## **Dust physico-chemical and optical properties**





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### **Thanks** to many colleagues, of course, but above all to **Claudia Di Biagio** and Jean-François Doussin (LISA) Yves Balkanski (LSCE)















Adapted from V. Grassian

# The variability issue (1/3)



- **Size distribution**: ~ 200 nm 20(0) μm
- Shape
- Mineralogical composition

Clays (illite, kaolinite, smectites, chlorite...), dolomite), iron oxydes, others....

quartz, feldspaths, carbonates (calcite,





## The variability issue (2/3)





adapted from Laskin et al. (2005)

AMMA (2006), unpublished



## The variability issue (3/3)



<sup>\*</sup> fine fraction only

# Advancing the prediction of optical properties and links to mineralogy



Scanza et al., ACP, 2014 but also Colarco et al., JGR, 2014; Journet et al., ACP, 2014; Perlwitz et al, ACP, 2015a; 2015b





### **Approaches**





Space

CINS

### **Approaches**





CINS

## The CESAM simulation chamber



- Controlled conditions (RH, T, irradiation, gaz mixtures)
- Simultaneous measurements of physicochemical and optical properties
- Long aerosol lifetime (> 24h for submicron particles)

Stainless-steel, 4.2 m<sup>3</sup> volume

Wang et al., AMT, 2011



### A number of challenges

# Generating mineral dust aerosols from natural soils





**RED-DUST** 



### It seem to works...







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## **Particle lifetime and concentration**











### Quantifying the dust aerosol mineralogy

### Total mass = Mass of crystalline minerals + Mass of iron oxydes + Mass of amorphous material

Total mass = gravimetrically or by elemental analysis (XRF, PIXE) or by conversion of particle size distribution Mass of crystalline minerals = X-ray diffraction (XRD) Mass of iron oxydes = X-ray absorption at near-edge spectroscopy (XANES) Mass of amorphous material = TEM counting of non-diffracting particles

### Clay (aluminosilicate)

### **Gypsum (calcium sulphate)**

### **Diatomite (amorphous silicate)**









### Quantifying the dust aerosol mineralogy

Mass of crystalline minerals = X-ray diffraction (XRD)

### **Identification**

Quantification



XRD analysis of low-mass samples (< 600 μg) Caquineau et al. (1997) Calibration of non-clay minerals (Klaver et al., 2011) Clay minerals based on difference with total mass

### Quantifying the dust aerosol mineralogy



## **Realism of the generated aerosol**

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# A laboratory-based project targeting the absorption properties of mineral dust

wavelength soil mineralogy size distribution





## Laboratory simulation experiments



# **Representation of the global soil** mineralogical composition





Soil mineralogical database, Journet et al. (2014)



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### **Selected soils for experiments**





Yellow rectangles: desert areas by Ginoux et al. (2012)



### Choice of samples from soil bank



Dots = selected soils



Box and Whiskers = global soils (Journet et al., 2014)



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### Soil and aerosol mineralogy





# Representing the size distribution



## **Modification of the IR spectrum**



liso



# **Complex refractive index: imaginary part** *RED-DUST*





## Synthesis and comparison





# A relatively narrow range of values, little if no spectral dependence for the real part



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# Advancing the prediction of optical properties and links to mineralogy



Scanza et al., ACP, 2014 but also Colarco et al., JGR, 2014; Journet et al., ACP, 2014; Perlwitz et al, ACP, 2015a; 2015b

### Imaginary CRI = f (mineral content) ?





### **UV/Vis CRI vs mineralogy**





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# **IR CRI vs mineralogy**







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# Ways forward

- Account the size-dependent mineralogy when calculating the optical properties

   uses an aerodynamic aerosol classifier
- Account particle asphericity when measuring particle size (Huang et al., 2021; Formenti et al., ESSD, submitted)





# CLImate relevant processing of Mineral Dust by volatile Organic compounds







## **One limiting factor**

Availability of soil samples for laboratory experiments





# Thank you!





### **RED-DUST**

### Relationship between SSA with iron and iron oxide content



### Dependence on the coarse size fraction

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## **Experimental simulations**



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authentic mineral dust generated from natural soils



Tunisian dust, 15% carbonates, <1% iron oxydes



### **Representation of the size**





\* depends on composition too



#### Geometric sizing - Morphology, Mineralogy

Off-line method; requires sampling of dust particles on appropriate substrates; samples are analysed by microscopy techniques.

#### Aerodynamic sizing - Mechanical behaviour

On-line methods; accelerates particles in an air stream and measures properties, which are proportional to the aerodynamic diameter.

#### Light Scattering - Optical / radiative properties

On-line method; measures the light scattered by particles at a given instrument geometry; scattered light is related to particle size for a calibration particle of spherical shape and known refractive index.

#### **Optical Inversion Methods - Radiative properties**

Off-line method; retrieves the effective diameter of an aerosol from remote sensing data on spectral extinction and angular scattering information, delivers column-integrated data.

adapted from Petzold (2008)







# **Aerosol sizing**

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**Merkus**, 2009

Huang, Y., Adebiyi, A. A., Formenti, P., & Kok, J. F. (2021). Linking the different diameter types of aspherical desert dust indicates that models underestimate coarse dust emission, Geophys. Res. Lett., 48, e2020GL092054, https://doi.org/10.1029/2020GL092054



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**Figure 1. Conversion factors linking the four different diameter types of aspherical dust.** Shown are (a) size-dependent conversions between the geometric and optical diameters, and size-invariant conversions between (b) the projected area-equivalent and geometric diameters, and between (c) the aerodynamic and geometric diameters. All three plots account for dust asphericity using the globally representative shape distributions detailed in Section 2.1. In panel (a), the OPC wavelength is taken as 780 nm, the scattering angle range is , and the real part of dust refractive index is 1.52. When the imaginary part increases from to and , the intersection optical diameters of red lines and the 1:1 reference line decrease from , to , and . Sensitivity tests of the conversions between the geometric and optical diameters to real and imaginary dust refractive indices, wavelength, and scattering angle range are shown in Supplementary Figures S1-S4, respectively.



Figure 2. Scattering cross sections Csca as a function of particle diameter for the OPC considered in this paper moderatelyabsorbing mineral dust (CRI = 1.53 - 0.0032i). The brown line represents Csca calculated by Mie theory assuming homogeneous particles while the orange line represents Csca calculated g to Huang et al. (2021) assuming homogeneous al ellipsoid particles.

Formenti, et al, submitted to Earth System Science CNIS Data, essd-2021-292, 2021

for Climate Services





## A reference size distribution dataset







Also reported by Foner and Ganor, 1992; Ganor et al., 1998; Chabas and Lefevre, 2000; Formenti et al., 2001; Falkovich et al., 2001; Sobanska et al., 2003; Levin, 2005



# Variabilité des propriétés optiques confirmé par les observations de terrain



### Domaine spectrale SW SSA (500 nm) ~ 0.75-1.0



# Domaine spectrale LW SSA (10 µm) ~ 0.2-0.6

## **Dust extinction spectra**



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## Synthesis and comparison



A relatively narrow range of values, little if no spectral dependence

