

A misty forest scene with a stream and trees. The image has a soft, ethereal quality with a light blue and grey color palette. The text "به نام خدا" is centered in the middle of the image in a white, elegant font.

به نام خدا

Monitoring of Dust Optical Characteristics Using Remote Sensing Instruments

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7th Training Course on WMO SDS-WAS Products, Ahvaz, Iran (10-14 Nov,2018)



This Presentation covers

- ❖ Introduction
- ❖ Impacts of dust storms
- ❖ Satellite remote sensing
- ❖ Remote sensing instruments
- ❖ Aerosol optical properties
- ❖ Satellite & ground data validation
- ❖ Variation in Aerosol Characteristics during Dust

What are aerosols and Dust?

Sources

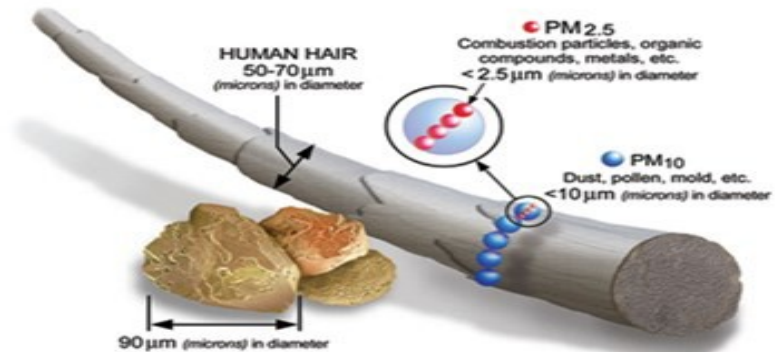
Natural

Anthropogenic

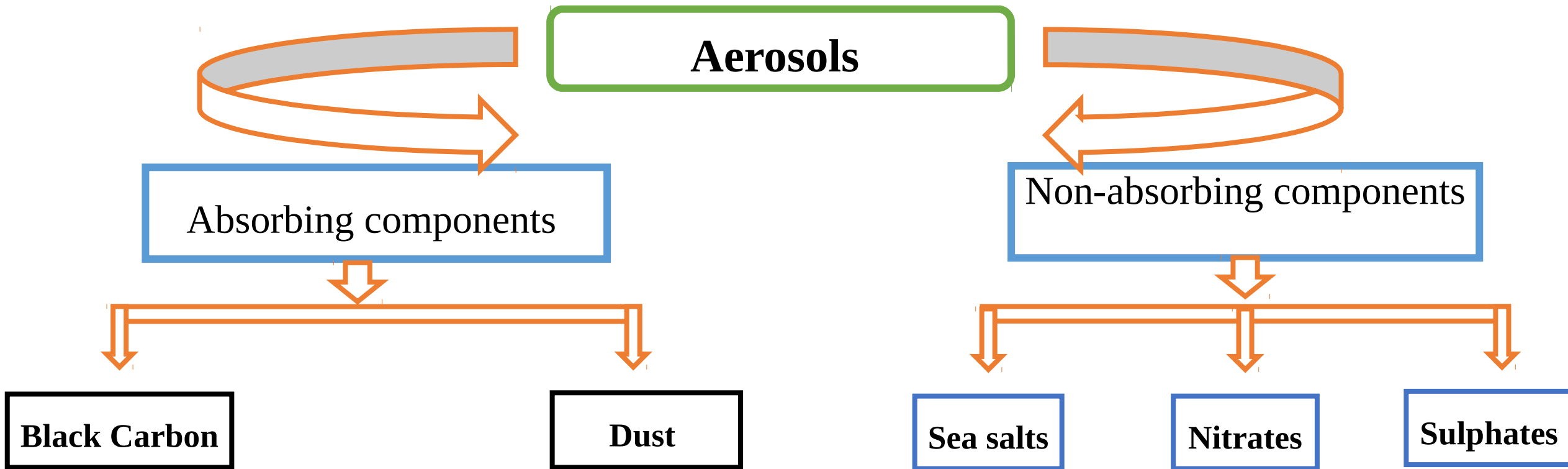
Sizes

PM 10
Coarse

PM 2.5
Fine



Classification of Atmospheric Aerosols



Aerosol Types

Dust aerosols

- Coarse particles probably consist of a mixture of quartz and clay minerals produced in arid regions



Biomass burning aerosols

- Fine particles mainly consist of black carbon and organic aerosols probably produced from natural or manmade fires



Urban/industrial aerosols

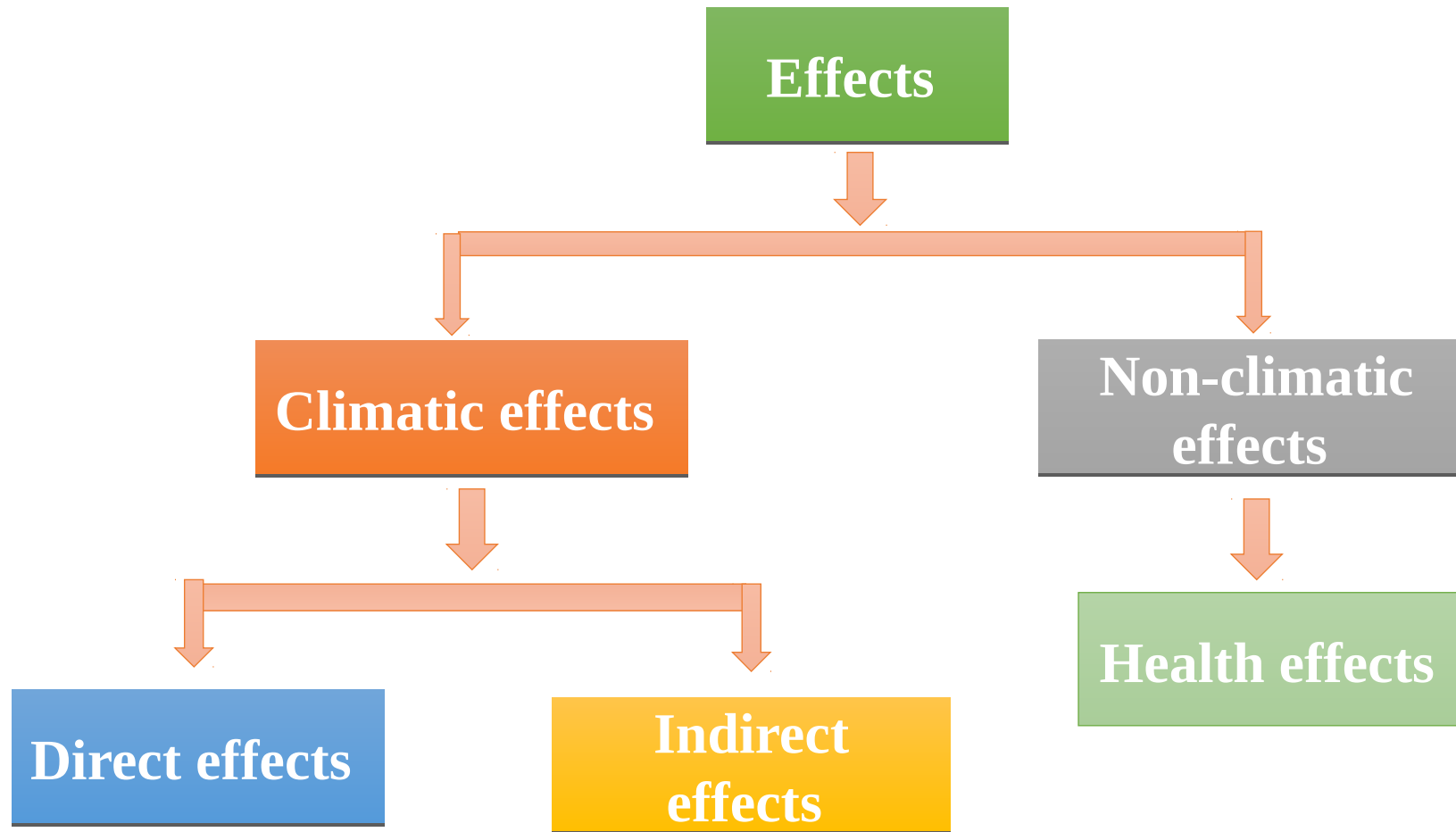
- Fine particles mainly consist of anthropogenic aerosols produced in urban and industrial regions



Dust characteristics

- Dust storms occasionally reach up to **1km | 5km | 10km height**, and are as thick as **100m | 2km | 5km**
- Over land, dust optical depth is typically around **0.1 | 0.5 | 1** or **2 | 5** for storms, in the visible range.
- Dust absorbs and scatters **infrared** radiation in the **Mie | Rayleigh | optical** region

Effects of Aerosols





Sources of Dust Emissions: 3 major types

Anthropogenic sources

(Dust from unpaved road)



Agricultural sources

(Dust from crop land)



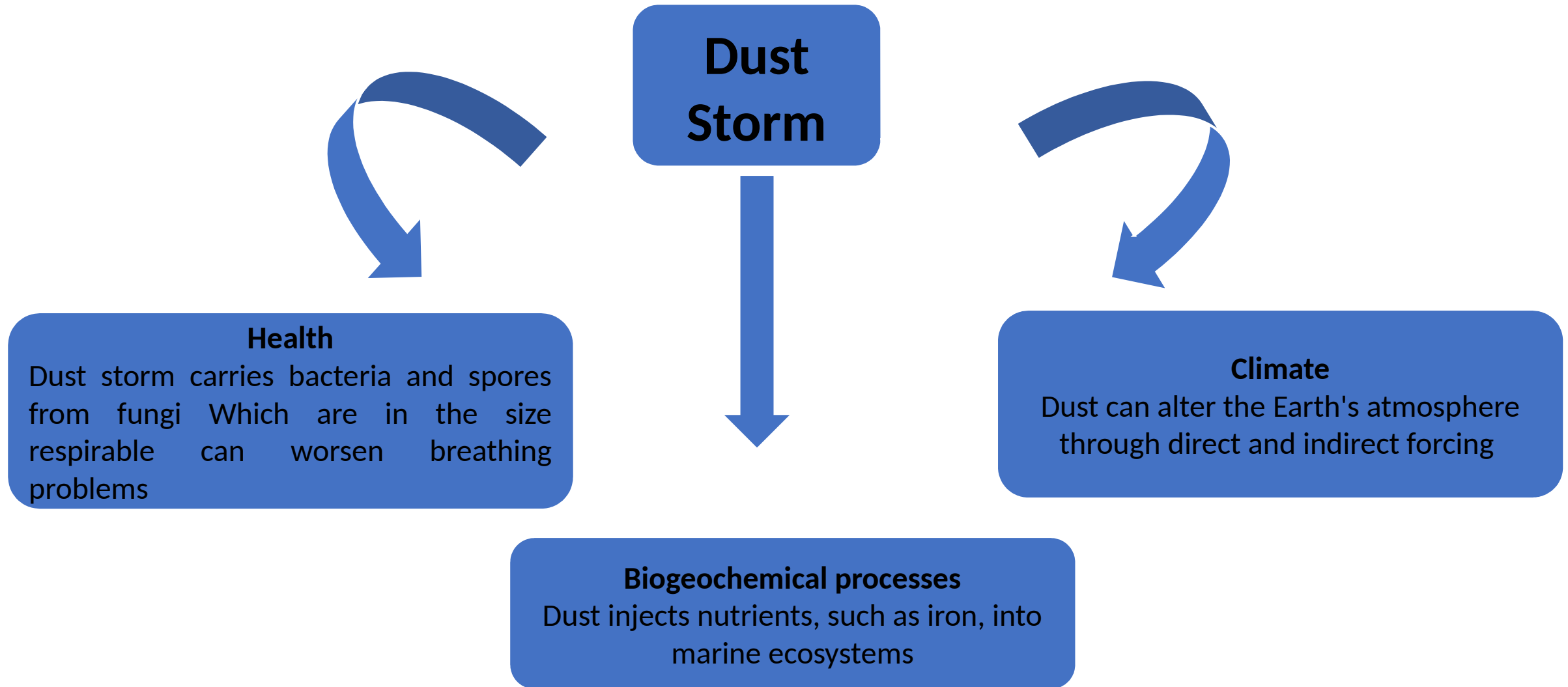
Natural desert sources



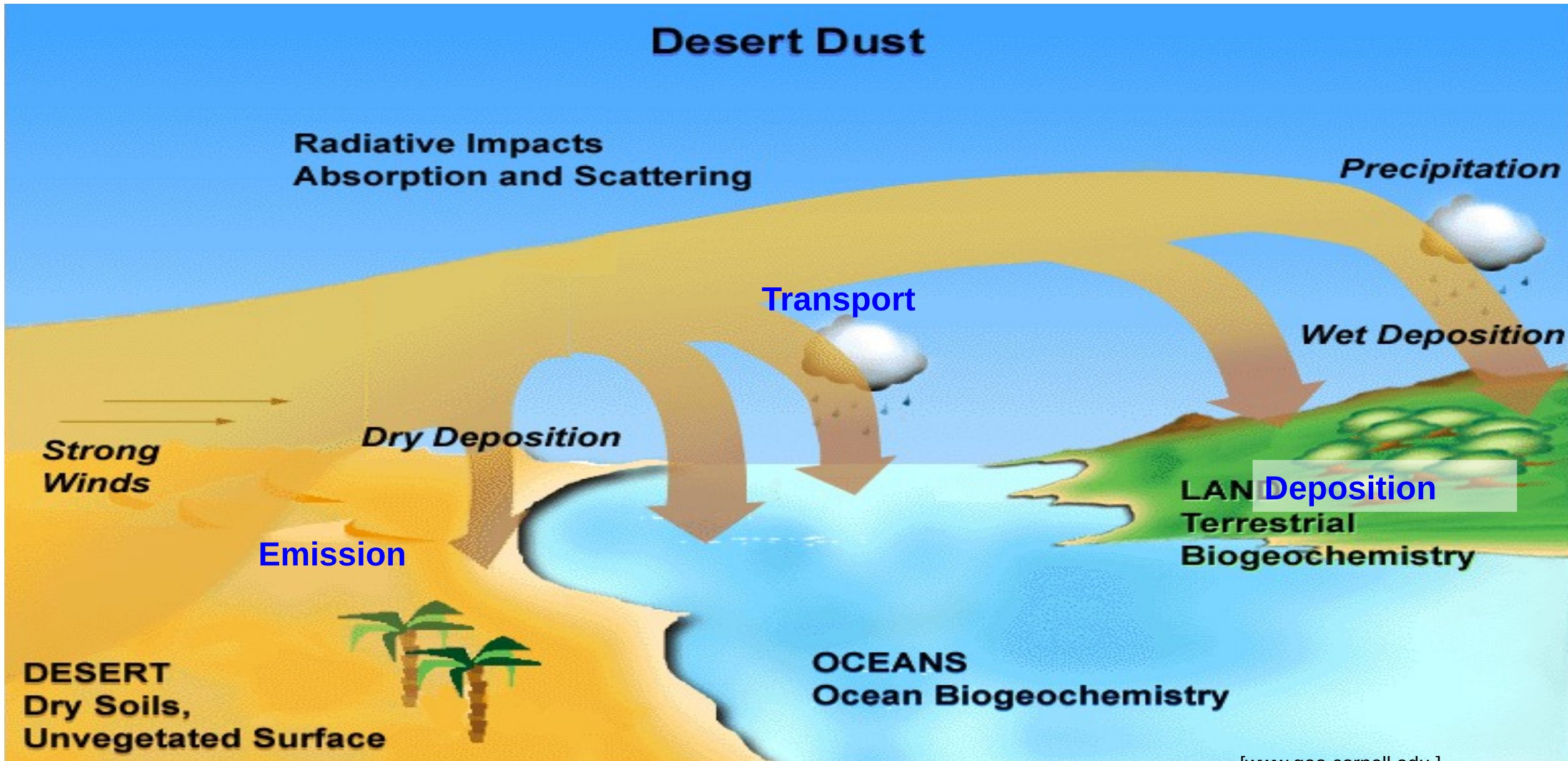
❖ World's Largest Deserts

Terra/MODIS True Colour View of Earth



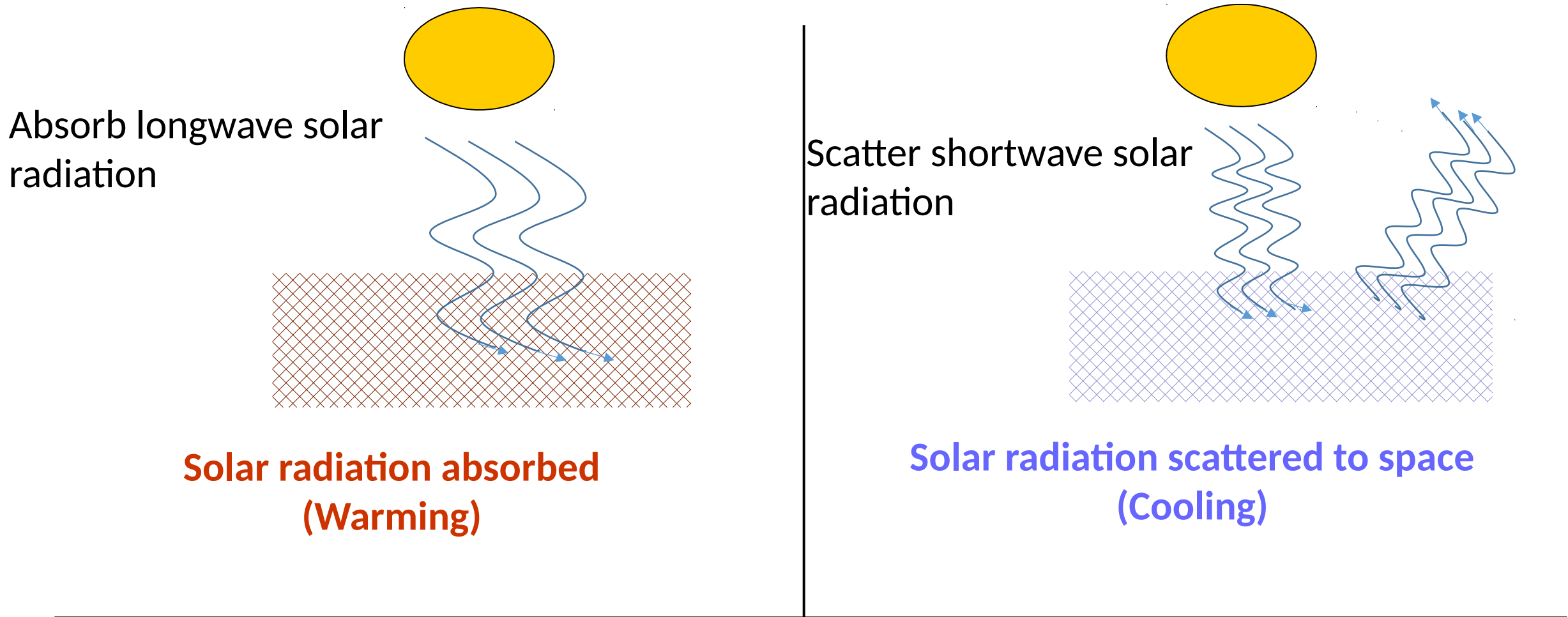


❖ Atmospheric Dust Cycle



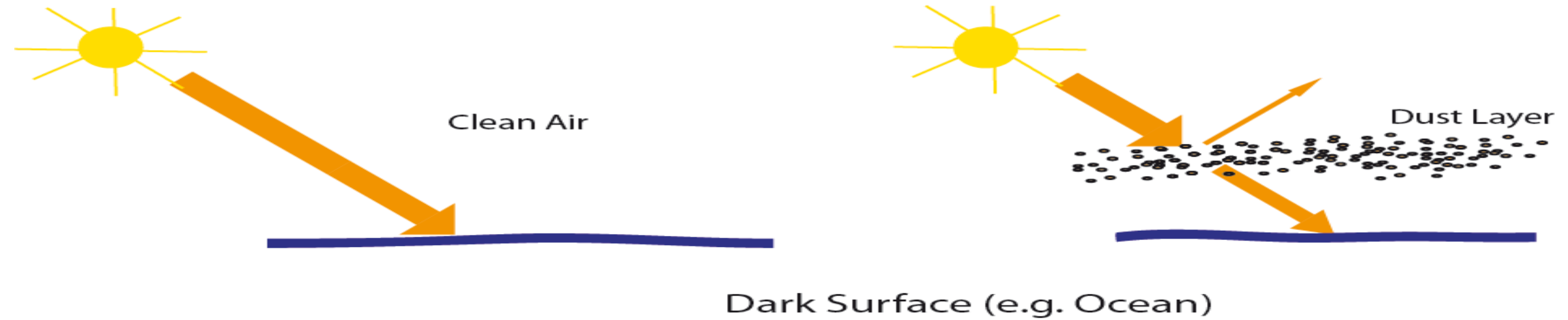
[www.geo.cornell.edu]

Direct radiative forcing by dust aerosols (Direct effect)

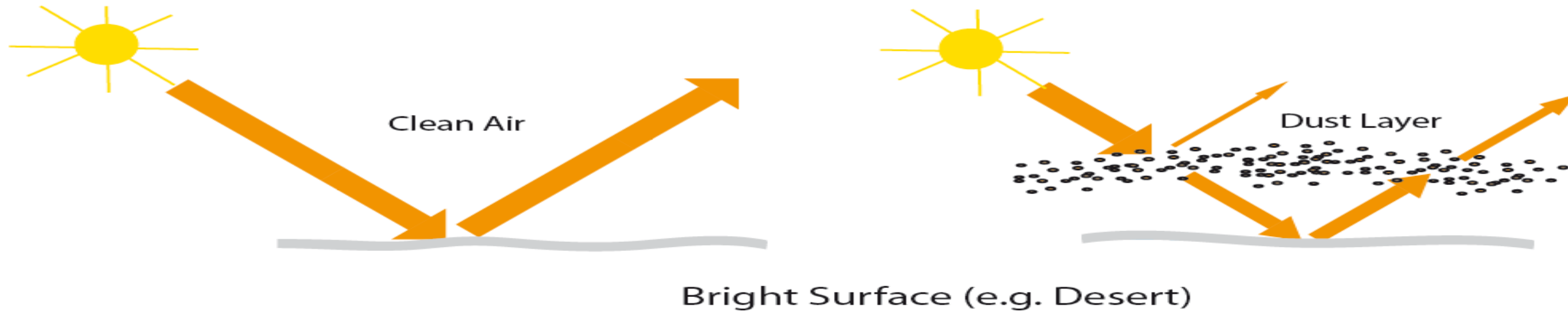


❖ Impacts of dust storm

Dust causes increased reflection of sunlight

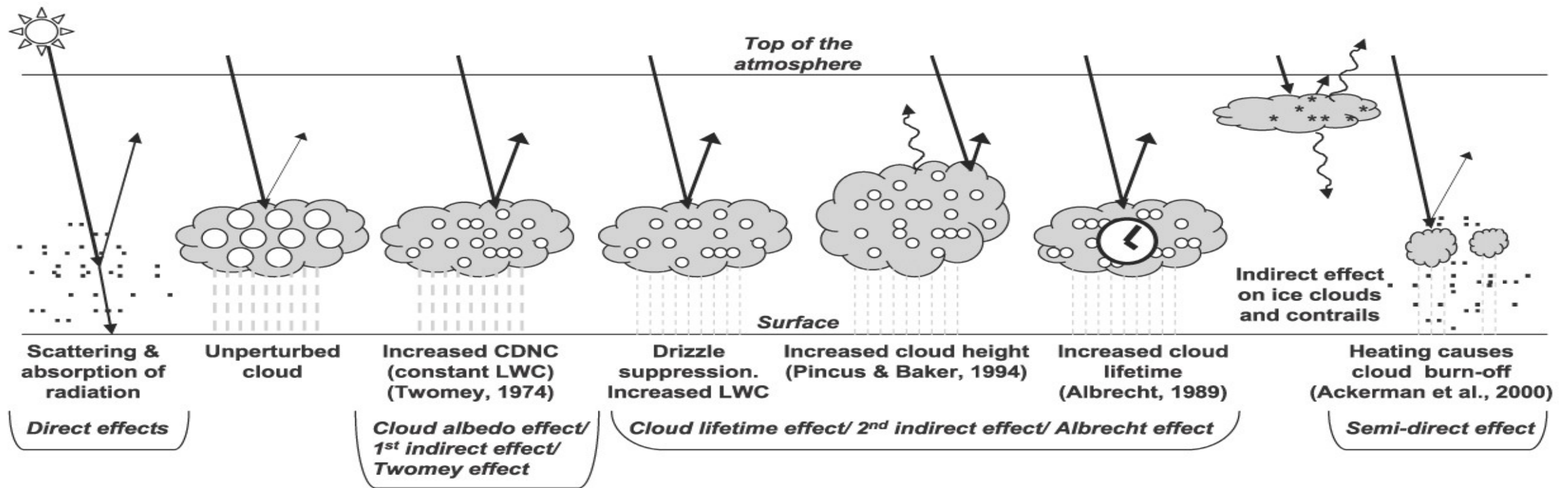


Dust causes decreased reflection of sunlight



Aerosol – Cloud Interactions

Indirect effect: aerosol-cloud effect



IPCC report, 2007

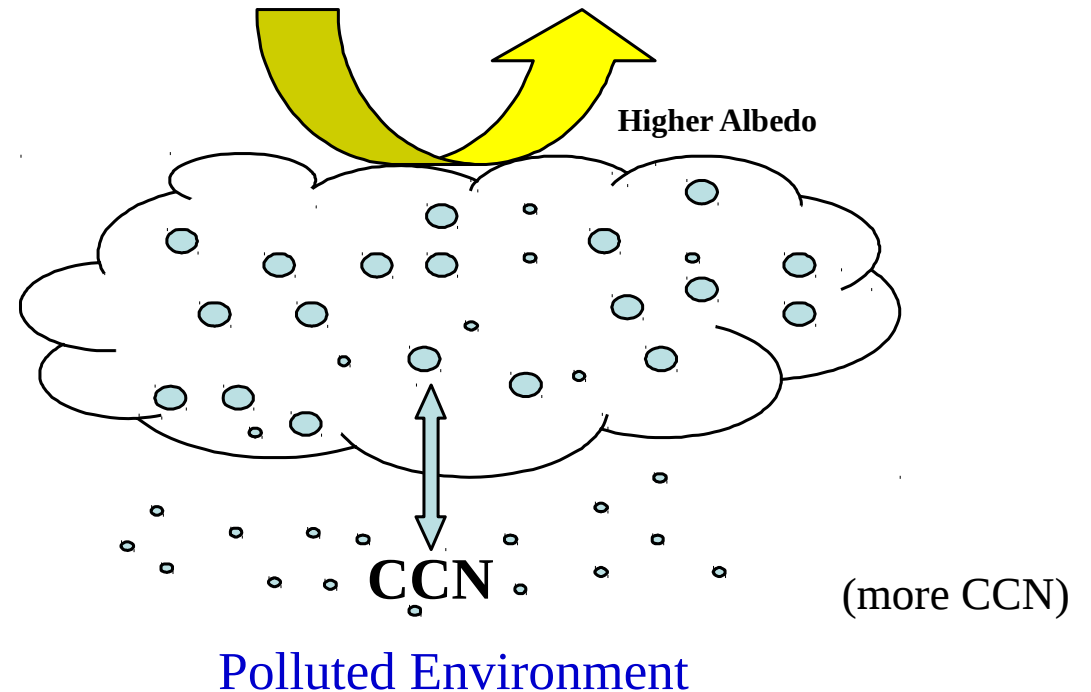
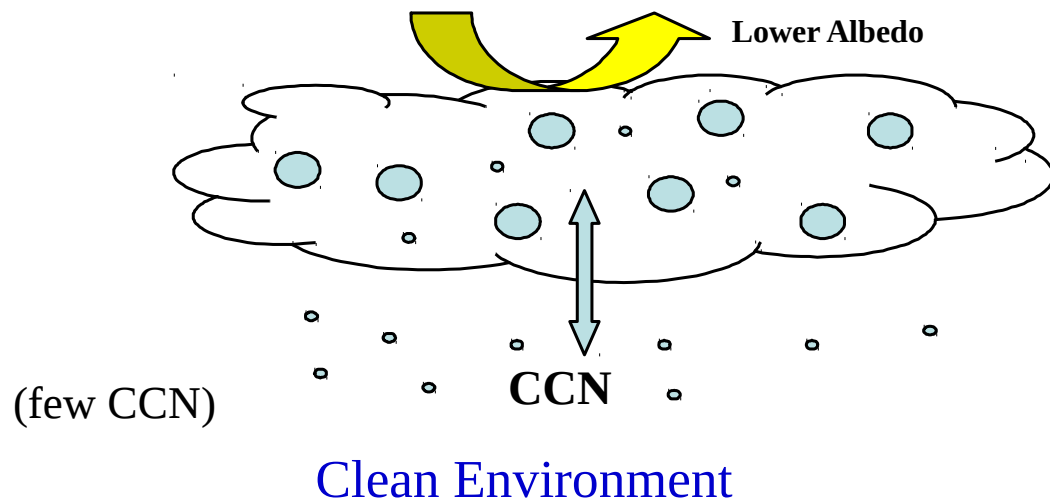
❖ Impacts of aerosol on cloud

How can humans affect clouds?

By changing CCN; cloud properties are a *strong* function of their concentration. This phenomenon is known as aerosol indirect effect.

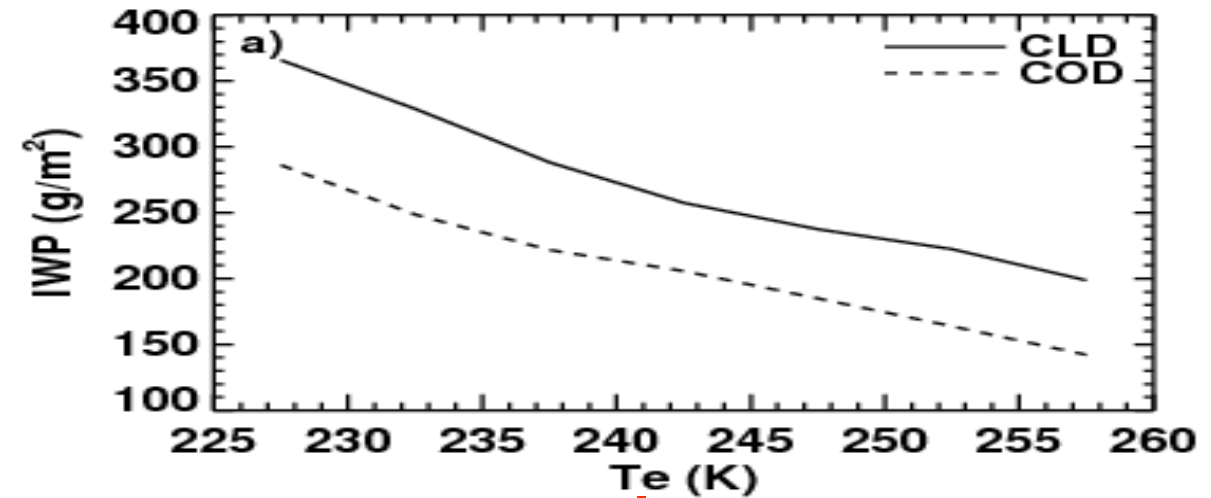
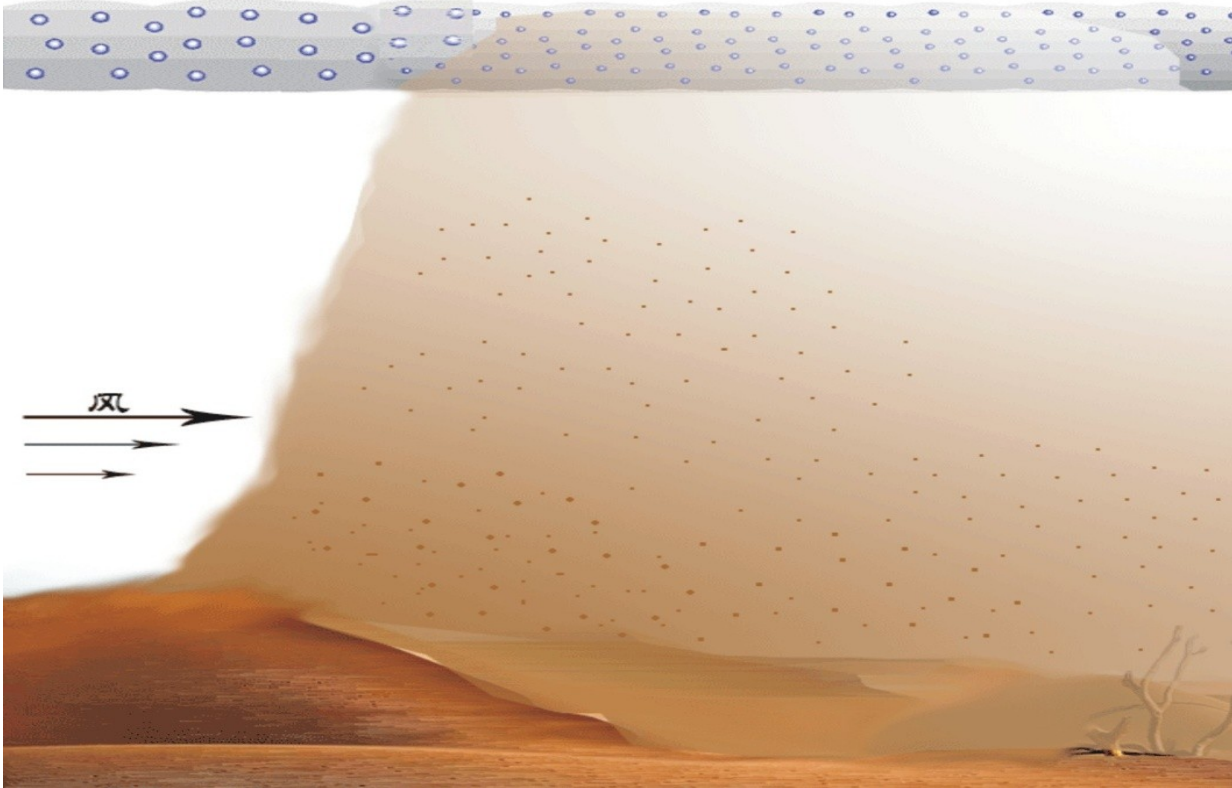
The aerosol indirect effect can lead to ***climatic cooling*** by:

- Increasing cloud reflectivity (albedo)
- Increasing cloud lifetime & coverage.



Climatic Effects of Aerosols

Dust indirect effect was evidenced by observation

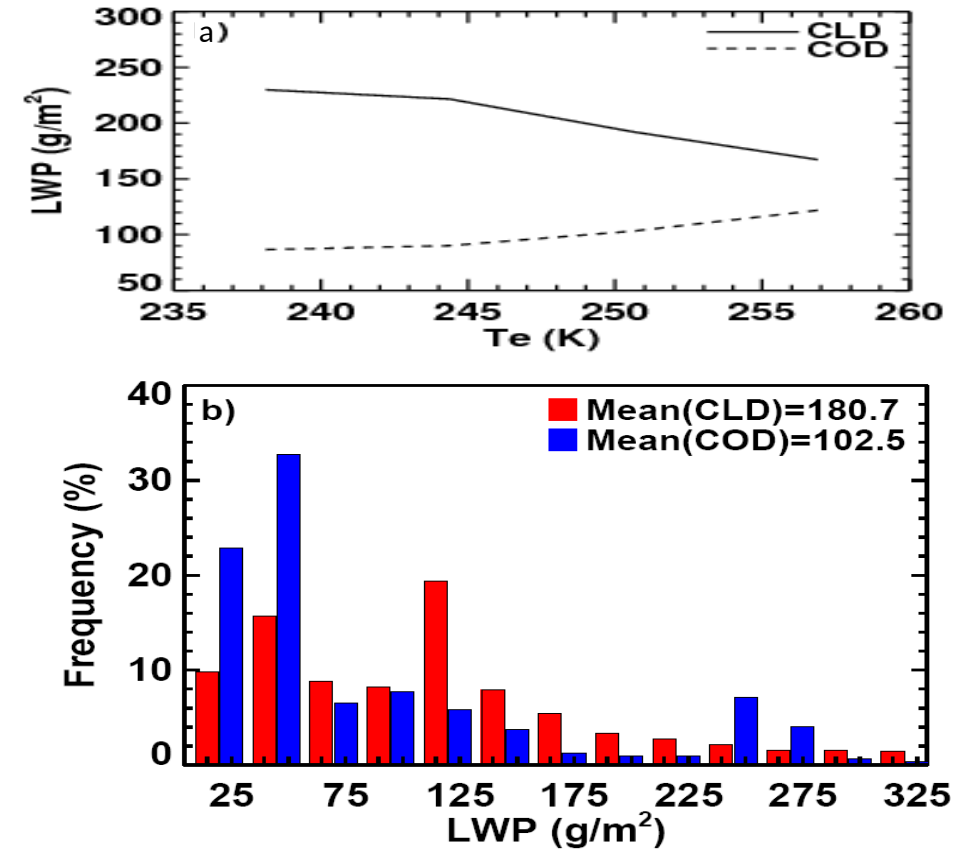
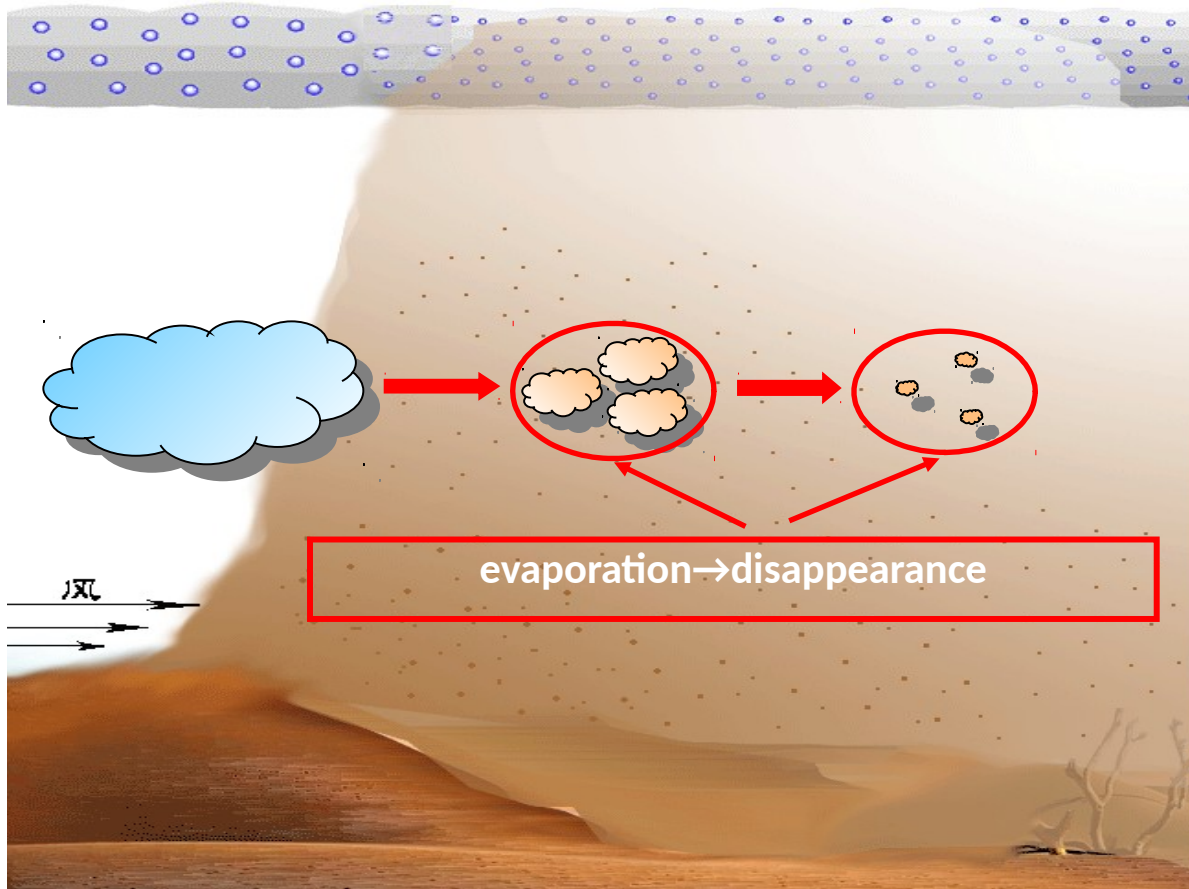


Huang et al., GRL, 2006a

Dust Aerosol, as Cloud condensation nuclei (CCN), can change cloud microphysical property, cloud fraction and cloud lifetime, thus can indirectly influence the radiation budget of the earth-atmosphere system.

Climatic Effects of Aerosols

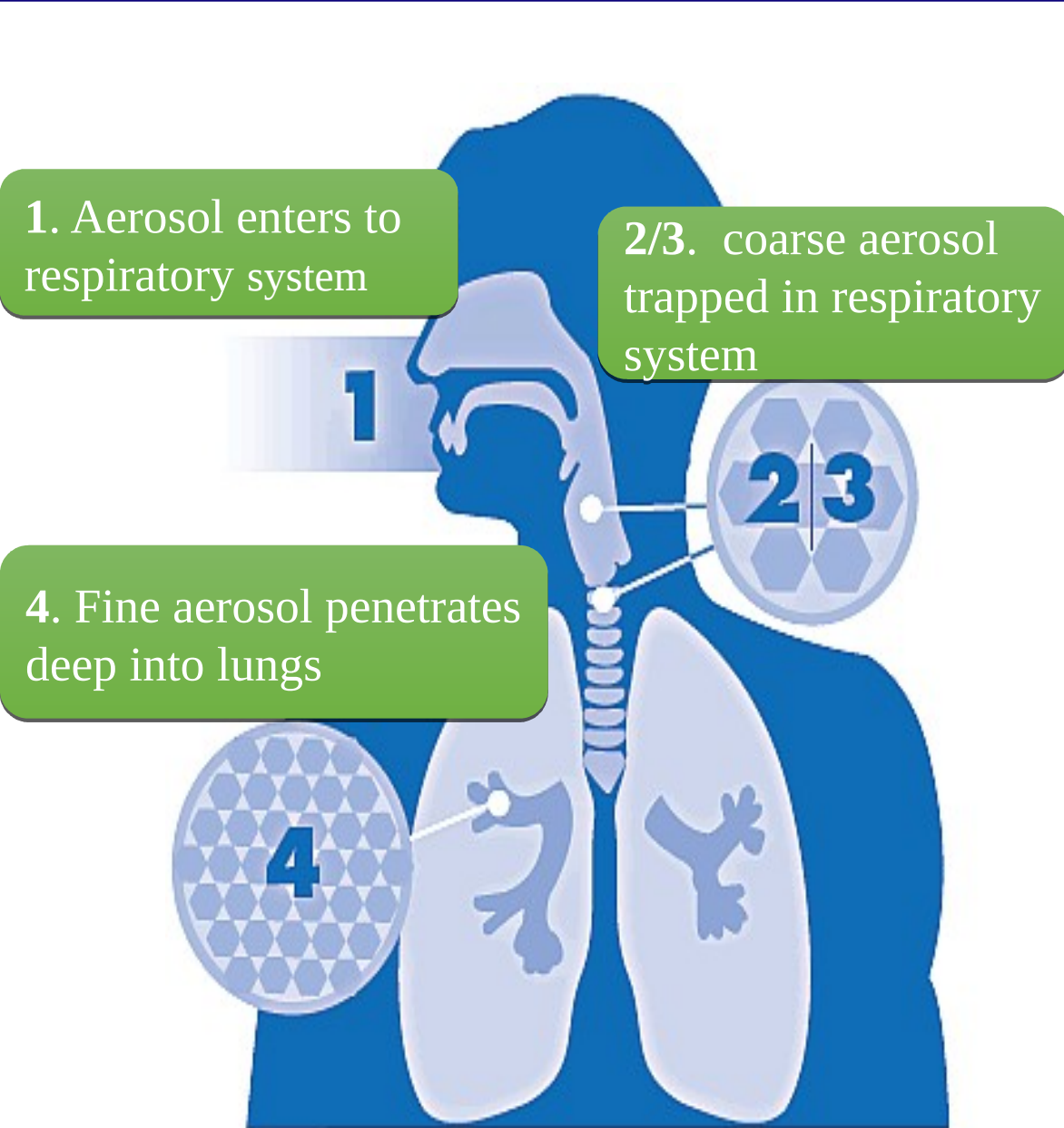
Semi-direct effect may contribute to arid climate



Huang et al., GRL, 2006b

Dust Aerosols' semi-direct effect can increase droplet evaporation, reduce liquid water path by 49.8%, lead to less precipitation.

Human Health Effect of Aerosols



Asthma



Pulmonary inflammation



Hay fever



Eyes infection



Respiratory symptoms



Skin allergy

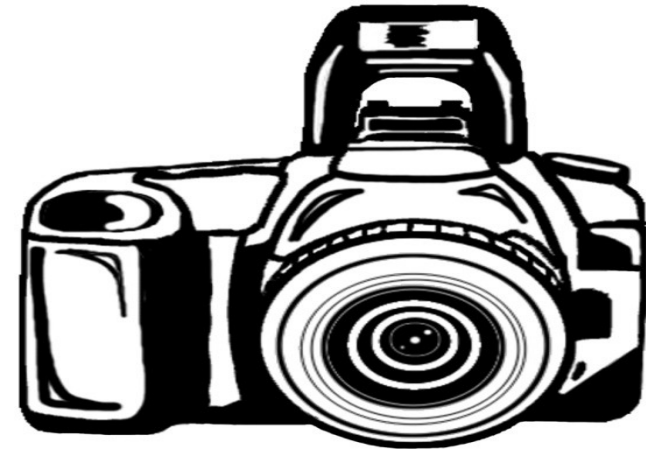


Remote Sensing

Satellite Remote Sensing

Basics of Satellite Remote Sensing

Collecting information about an object without being in direct physical contact with it.



- Remote Sensing:
 - The art and science of obtaining information about an object without physically contact between the object and sensor
 - The processes of collecting information about Earth surfaces and phenomena using sensors not in physical contact with the surfaces and phenomena of interest.
 - There is a medium of transmission involved i.e. Earth's Atmosphere.

Satellite Remote Sensing

Applied Sciences Program

Applications to Decision Making: Thematic Areas



**Agricultural
Efficiency**



Air Quality



Climate



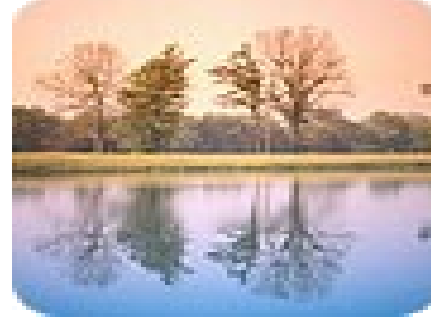
Disaster Management



Ecological Forecasting



Public Health

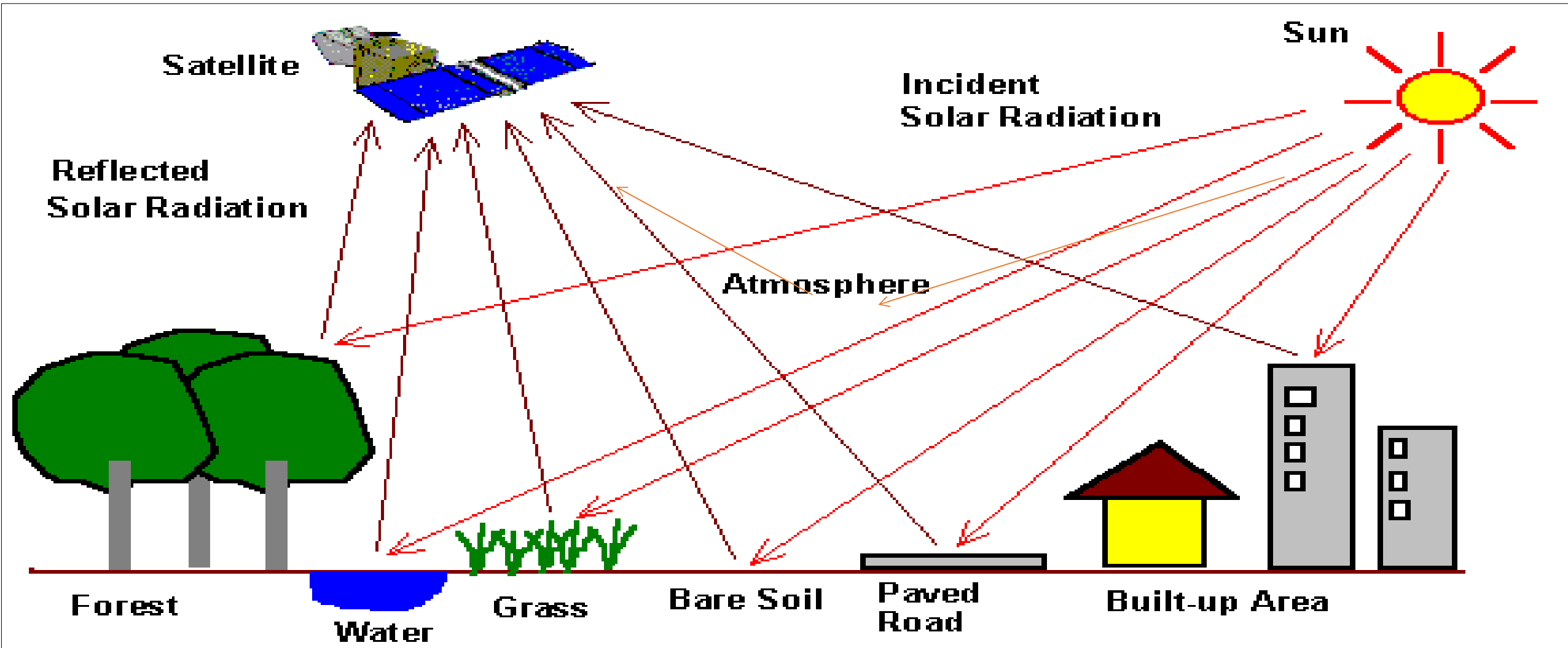


**Water
Resources**

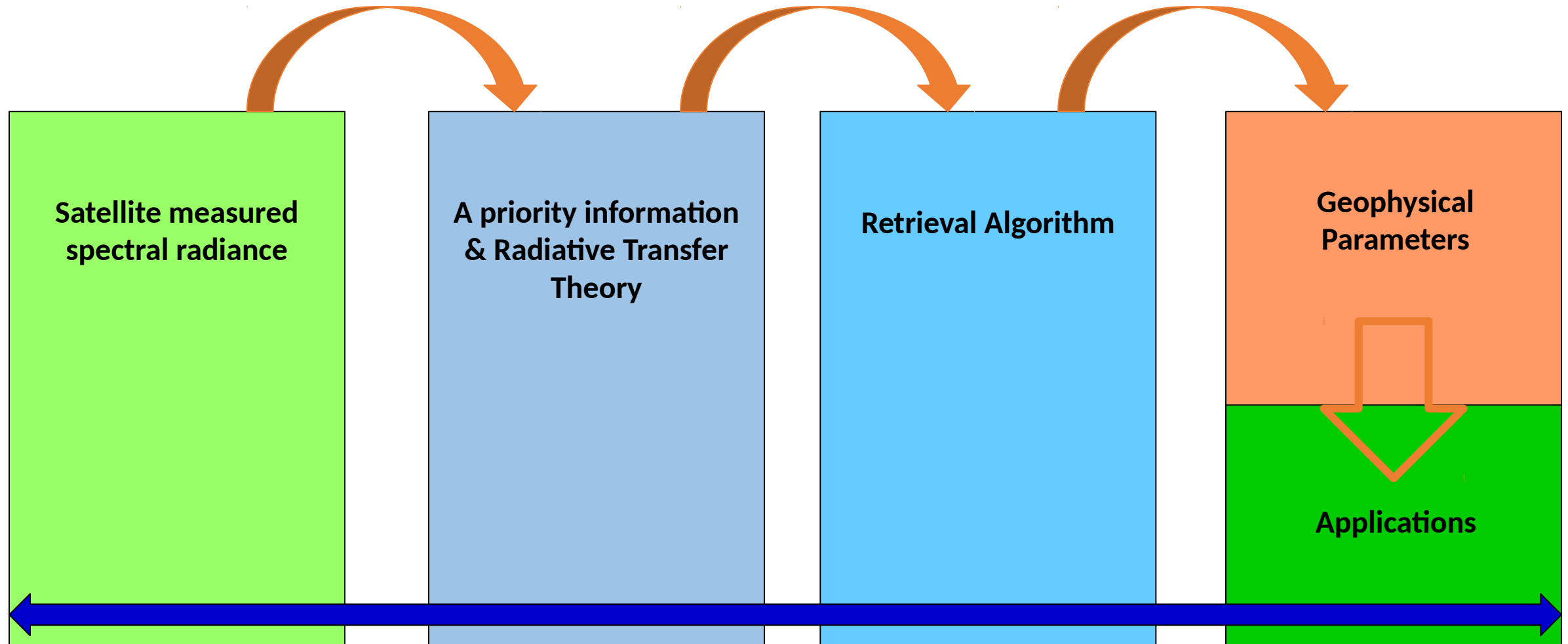


Weather

What does satellite measures ?

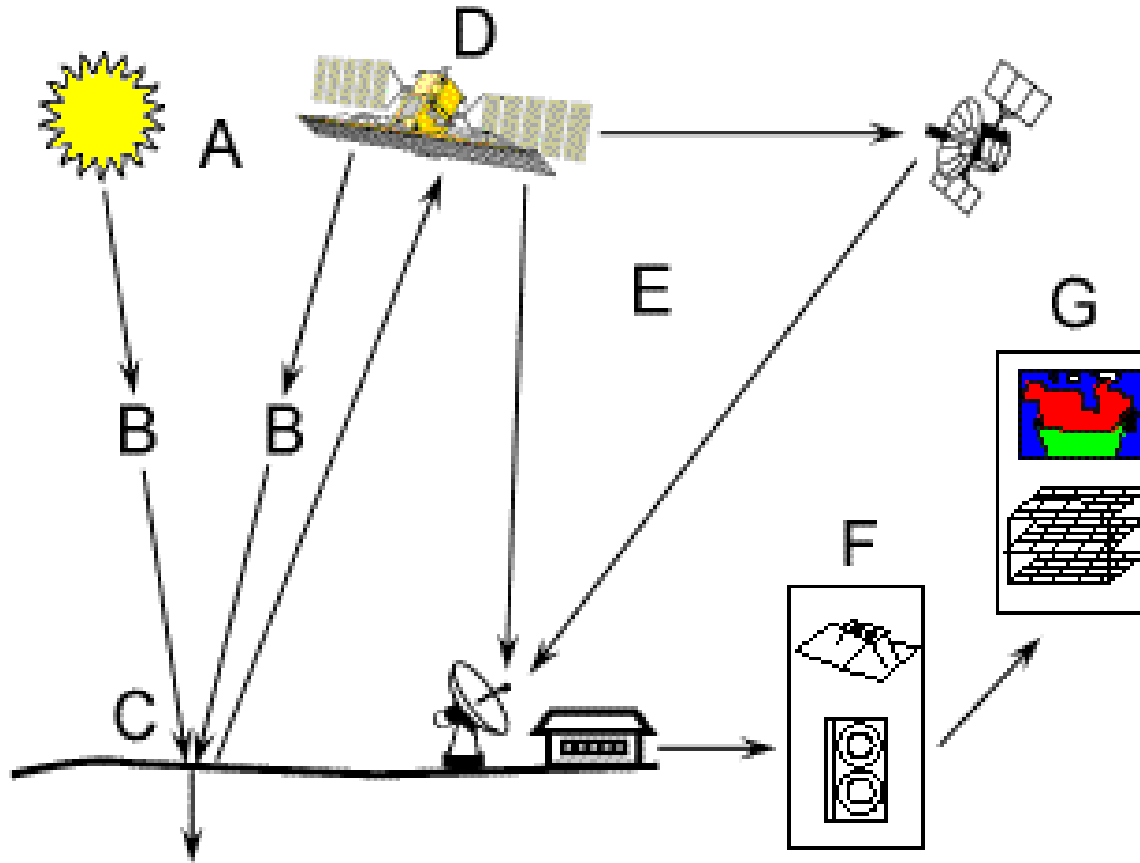


Remote Sensing Process



Satellite Remote Sensing

Remote Sensing Process Components



Energy Source or Illumination (A)

Radiation and the Atmosphere (B)

Interaction with the Target (C)

Recording of Energy by the Sensor (D)

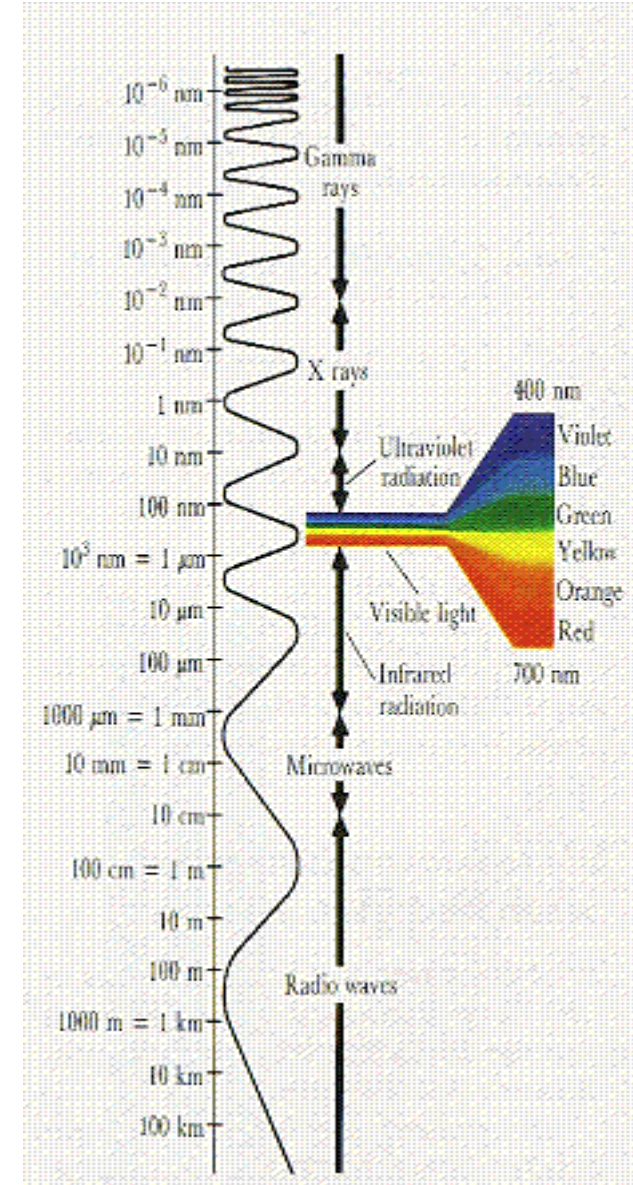
Transmission, Reception, and Processing (E)

Interpretation and Analysis (F)

Application (G)

Wavelength Ranges in Remote Sensing

- **UV:** some absorptions + profile information+ aerosols
- **VIS:** surface information (vegetation)+some absorptions+ aerosol
- **IR:** temperature information + cloud information+ water / ice distinction
- +many absorptions / emissions + profile information
- **MW:** no problems with clouds+ ice / water contrast+ surfaces + some emissions + profile information

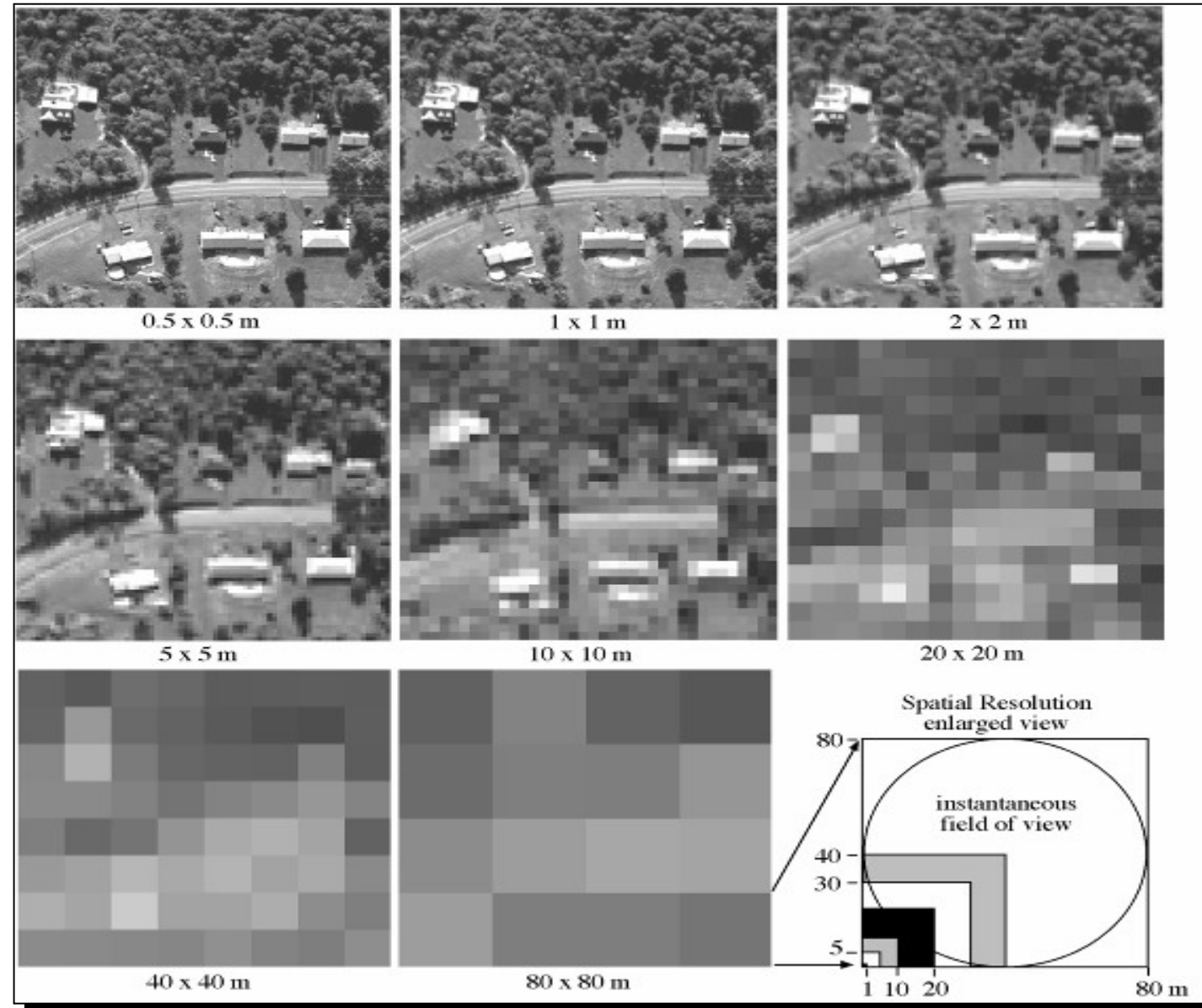


Resolution

- All remote sensing systems have four types of resolution:
 - Spatial
 - Spectral
 - Temporal
 - Radiometric

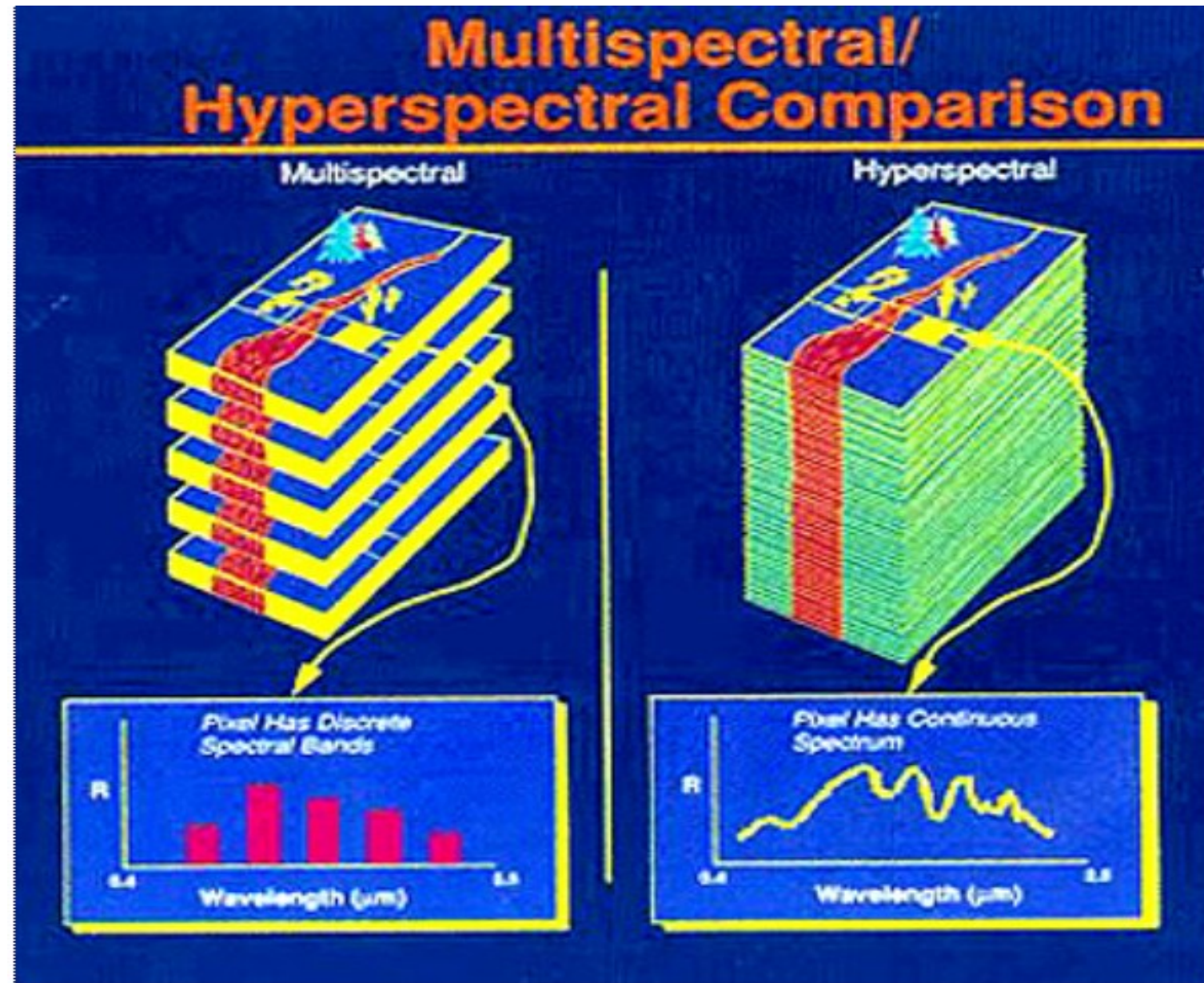
Fundamental in Remote Sensing: Spatial Resolution

- The earth surface area covered by a pixel of an image is known as spatial resolution
- Large area covered by a pixel means low spatial resolution and vice versa



Fundamental in Remote Sensing: Spectral Resolution

- Is the ability to resolve spectral features and bands into their separate components
- More number of bands in a specified bandwidth means higher spectral resolution and vice versa



Fundamental in Remote Sensing: Temporal Resolution

- Frequency at which images are recorded/ captured in a specific place on the earth.
- The more frequently it is captured, the better or finer the temporal resolution is said to be
- For example, a sensor that captures an image of an agriculture land twice a day has better temporal resolution than a sensor that only captures that same image once a week.

Fundamental in Remote Sensing: Radiometric Resolution

- Sensitivity of the sensor to the magnitude of the received electromagnetic energy determines the radiometric resolution
- Finer the radiometric resolution of a sensor, if it is more sensitive in detecting small differences in reflected or emitted energy

2-bit range
0 → 4

6-bit range
0 → 63

8-bit range
0 → 255

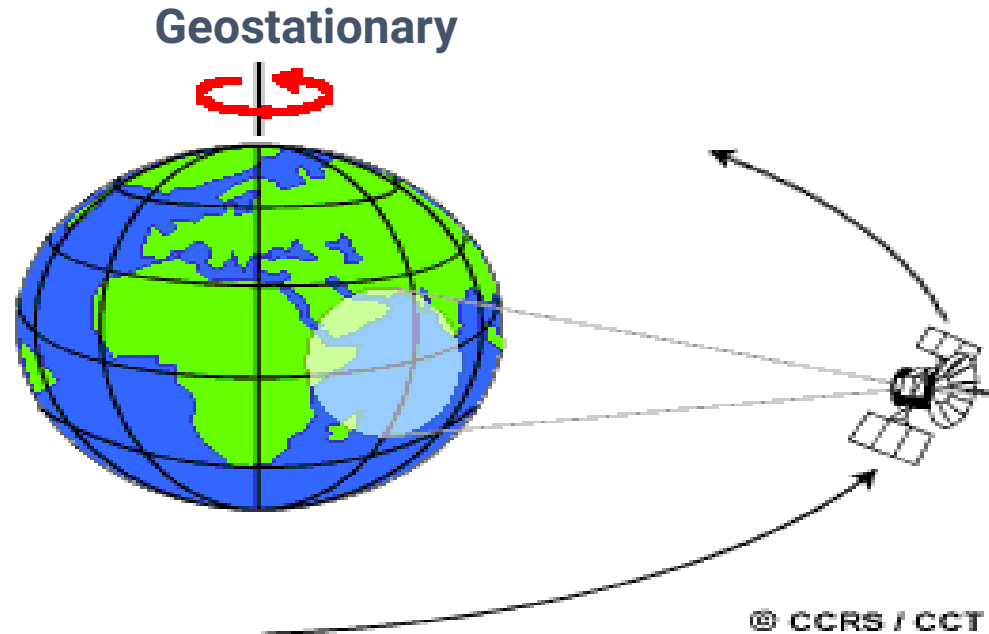
10-bit range
0 →



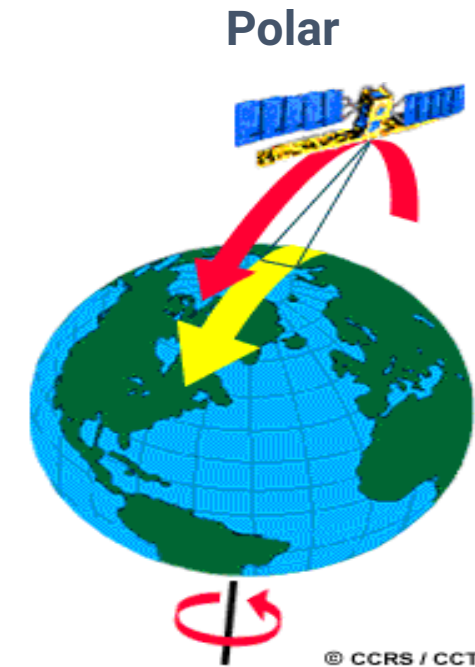
Some of the ways satellites/sensor can be classified

- **Orbits**
 - Polar vs Geostationary
- **Energy source**
 - Passive vs Active ...
- **Solar spectrum**
 - Visible, UV, IR, Microwave ...
- **Measurement Technique**
 - Scanning, non-scanning, imager, sounders ...
- **Resolution (spatial, temporal, spectral, radiometric)**
 - Low vs high (any of the kind)
- **Applications**
 - Weather, Ocean colors, Land mapping, Atmospheric Physics, Atmospheric Chemistry, Air quality, radiation budget, water cycle, coastal management ...

Common types of orbits

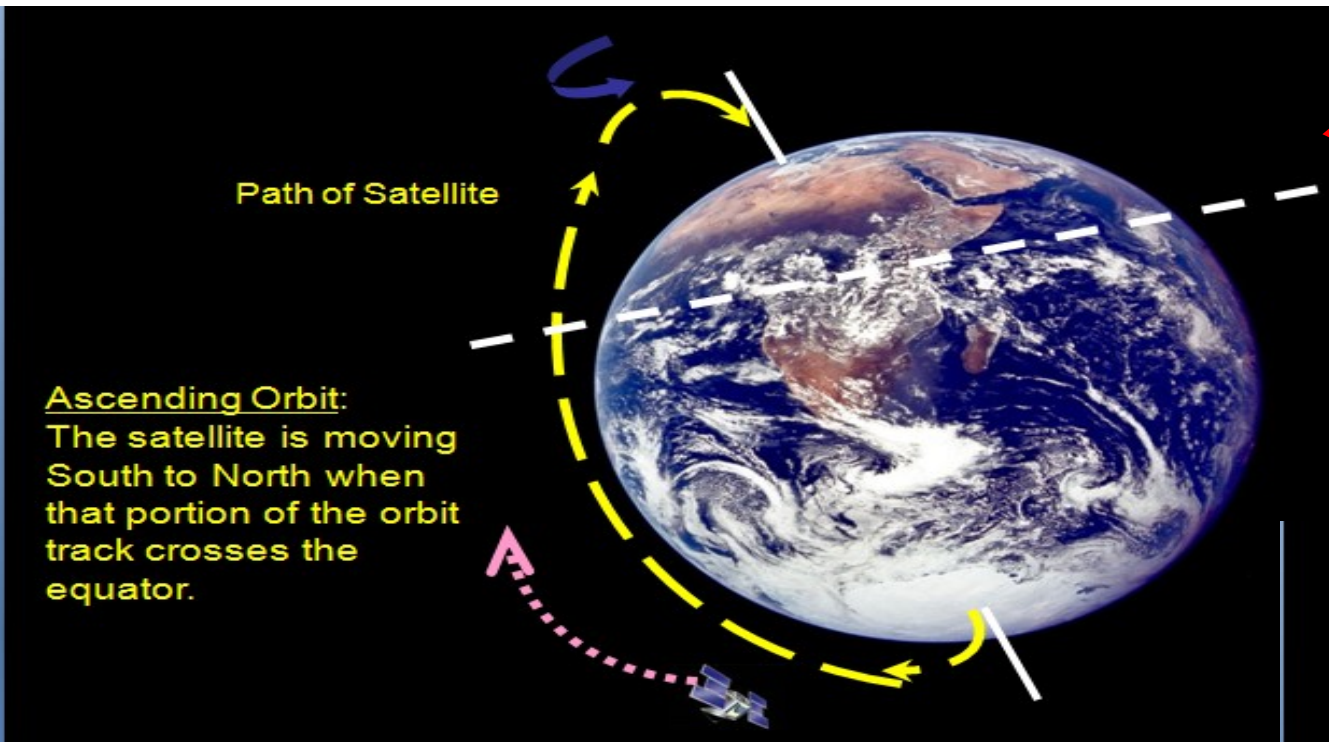


Geostationary orbit
An orbit that has the same Earth's rotational period
Appears 'fixed' above earth
Satellite on equator at ~36,000km



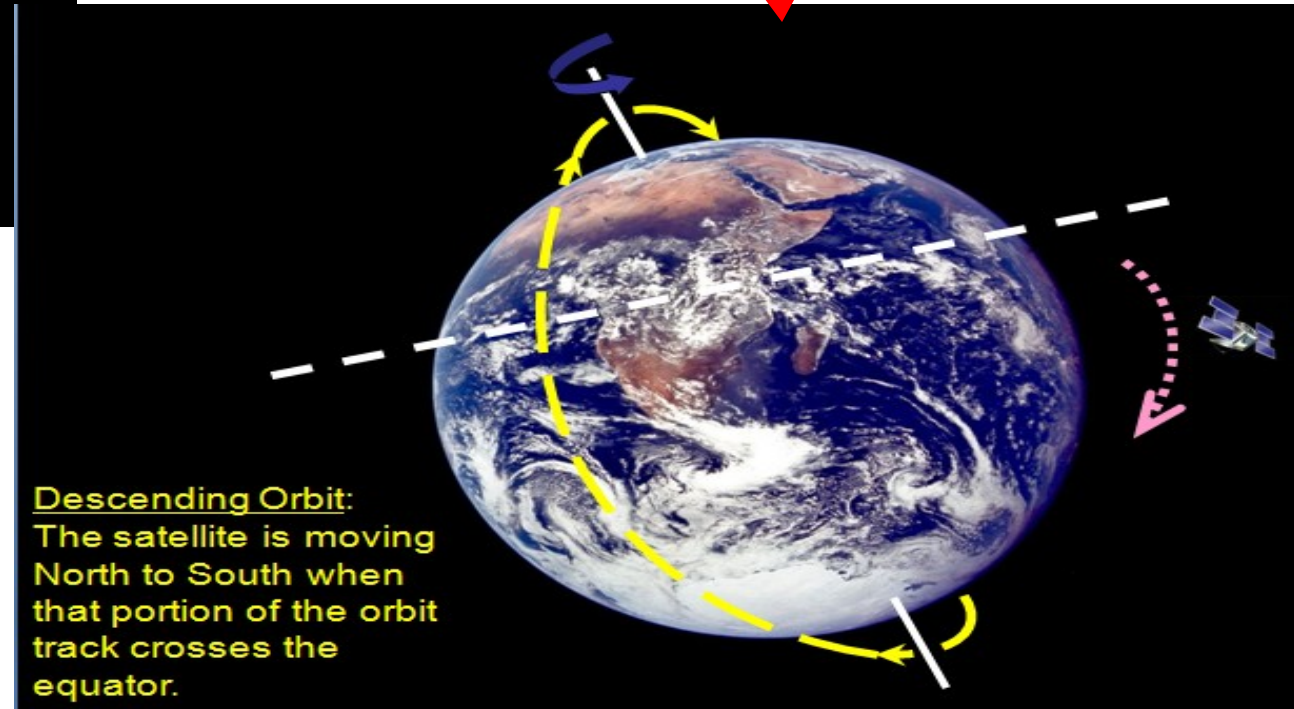
Polar orbiting orbit
fixed circular orbit above the earth, ~600-1000km in sun synchronous orbit with orbital pass at about same **local solar time** each day

Satellite/Sensor Classifications



Polar Orbits

**Ascending
vs
Descending**



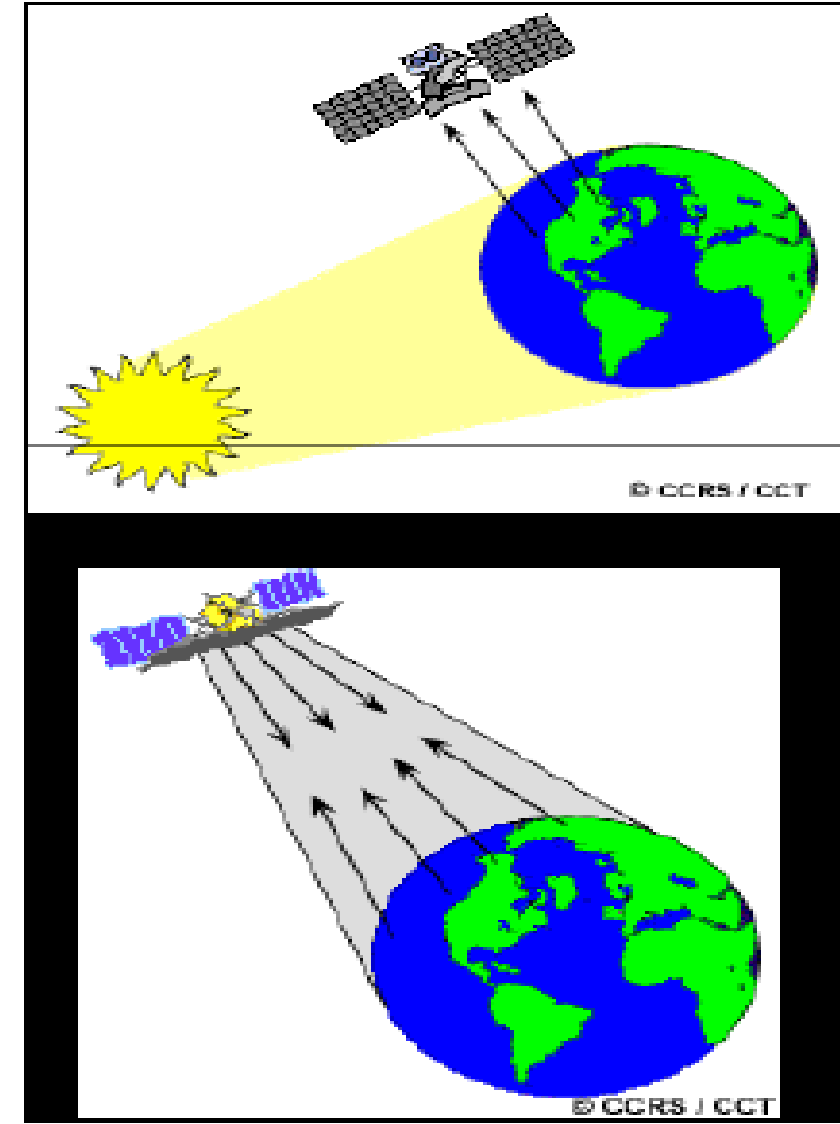
Satellite/Sensor Classifications

Passive Sensors: Remote sensing systems which measure energy that is naturally available are called passive sensors.

MODIS, MISR, OMI

Active Sensors: The sensor emits radiation which is directed toward the target to be investigated. The radiation reflected from that target is detected and measured by the sensor.

CALIPSO



Remote Sensing Instruments

Ground based instruments

Aerosol RObotic NETwork (AERONET)

Satellite based instruments

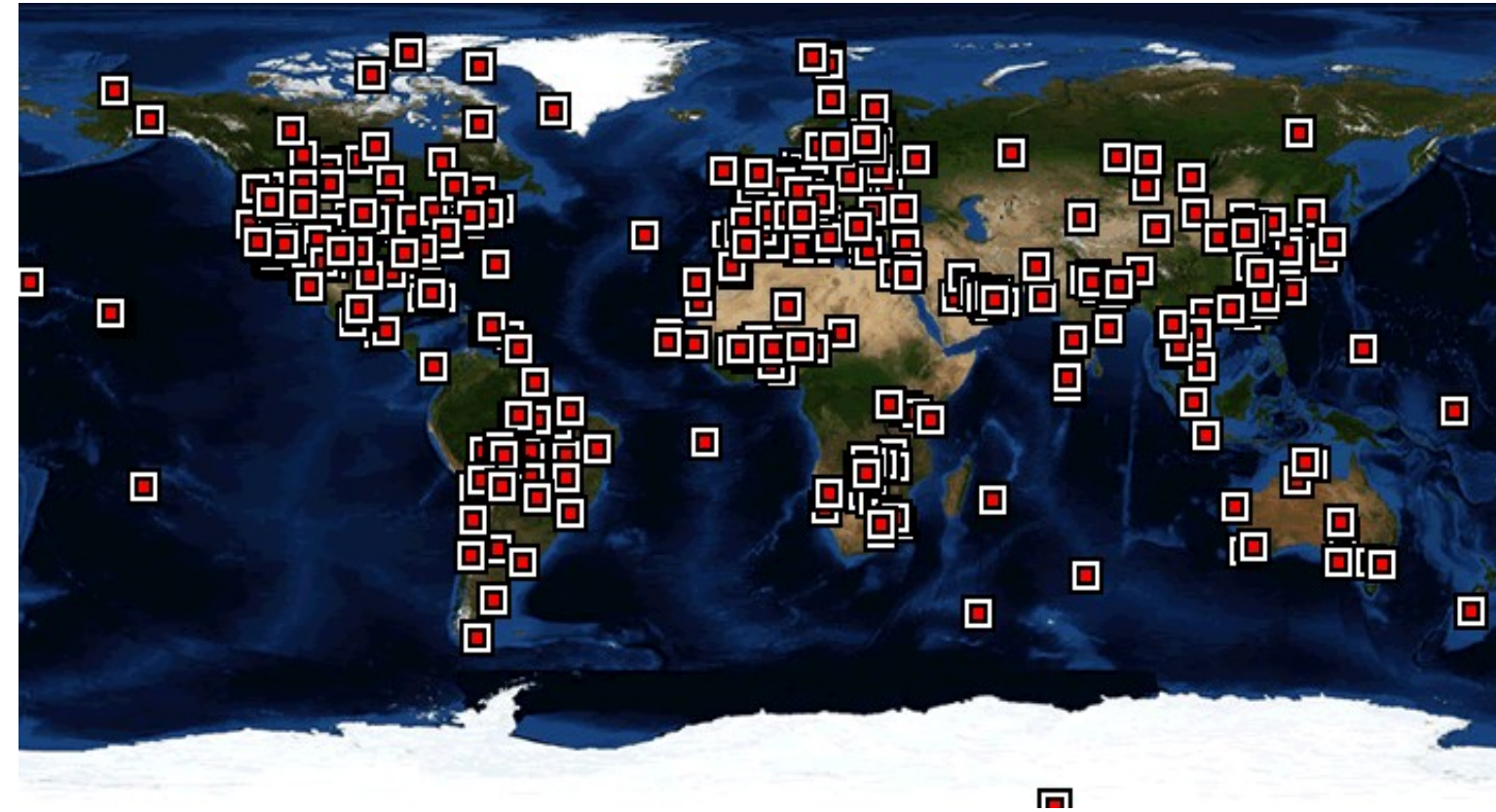
MODerate resolution Imaging Spectroradiometer (MODIS)

Ozone Monitoring Instrument (OMI)

Multangle Imaging SpectroRadiometer (MISR)

Cloud-Aerosol Lidar and Infrared Pathfinder Satellite Observations (CALIPSO)

AERONET/Sunphotometer



- AErosol RObotic NETwork
- Worldwide collection of sun photometers

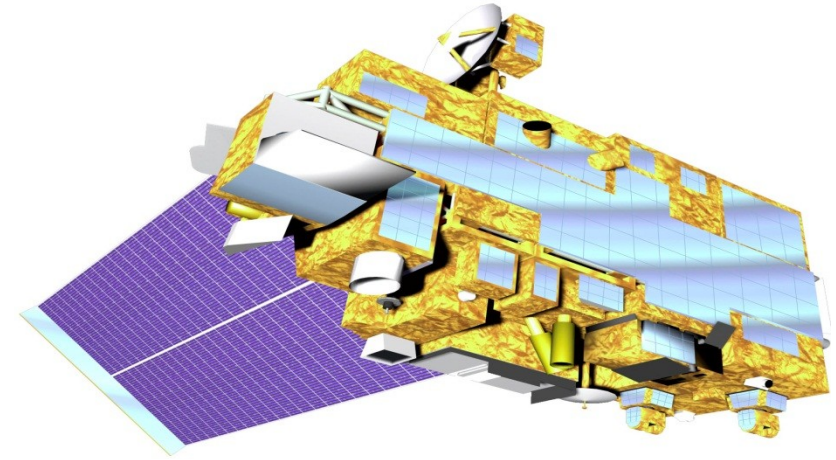


AERONET

- AERONET provides global observations of spectral aerosol optical depth (AOD), inversion products, and precipitable water.
- The direct sun measurements are made in eight spectral bands requiring approximately 10 seconds at wavelengths of 340, 380, 440, 500, 670, 870, 940 and 1020 nm
- The 940 nm channel is used for column water abundance determination
- Quality levels: Level 1.0 (unscreened), Level 1.5 (cloud-screened), and Level 2.0 (cloud-screened and quality-assured).

❖ Moderate Resolution Imaging Spectroradiometer (MODIS)

- NASA, Terra & Aqua
 - launched 1999, 2002
 - 705 km polar orbits, descending (10:30 a.m.) & ascending (1:30 p.m.)
- Sensor Characteristics
 - 36 spectral bands ranging from 0.41 to 14.385 μm
 - cross-track scan mirror with 2330 km swath width
 - Spatial resolutions:
 - 250 m (bands 1 - 2)
 - 500 m (bands 3 - 7)
 - 1000 m (bands 8 - 36)



Three Separate Algorithms

Land

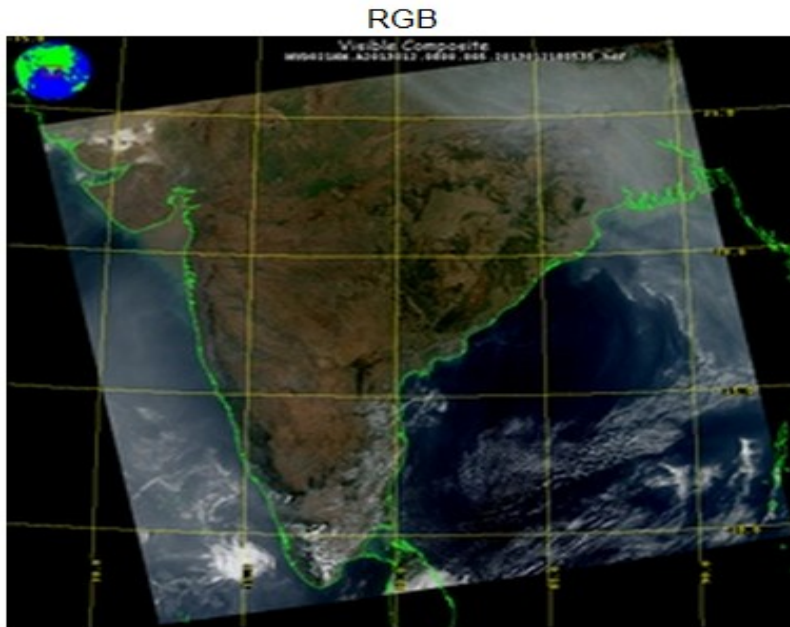
Ocean

Dark Target
(surface) –
limited to only
over dark vegetate surfaces

Deep Blue –
Used over bright
land surfaces

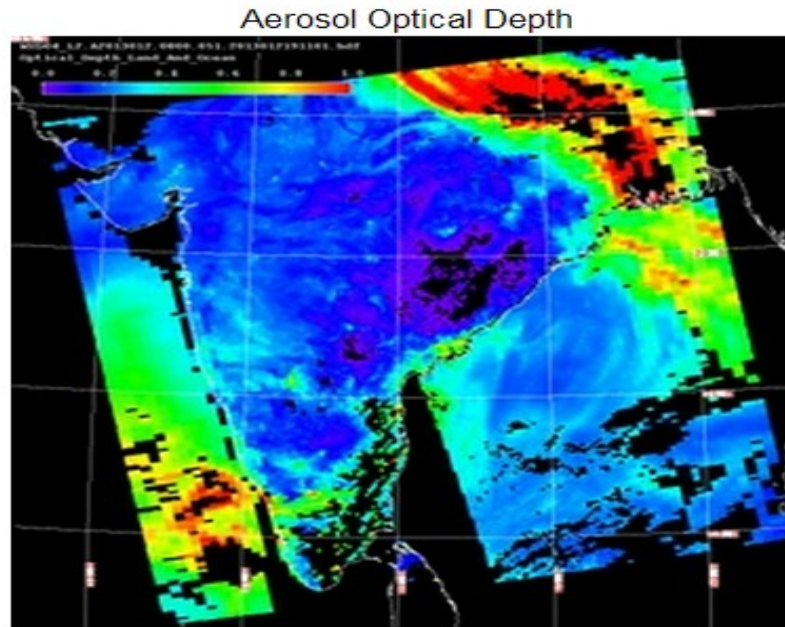
MODIS Levels of Data

January 12, 2013



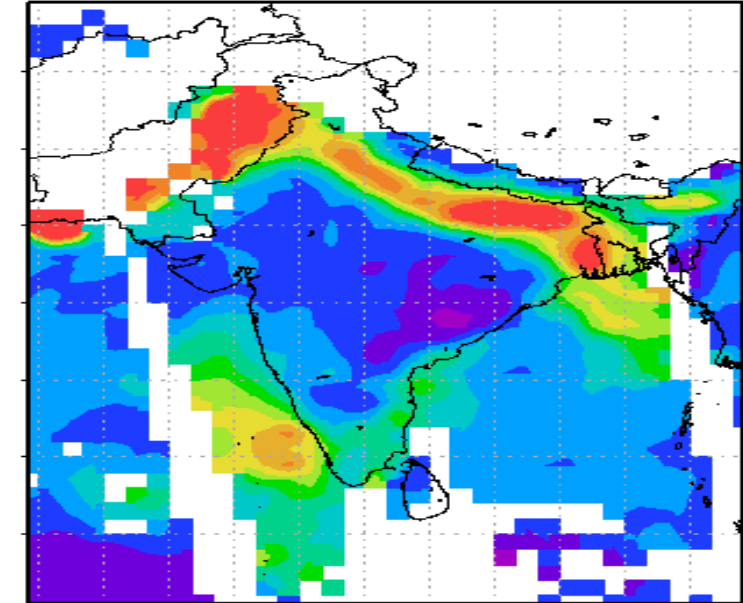
Level 1B

Calibration to Radiance



Level 2

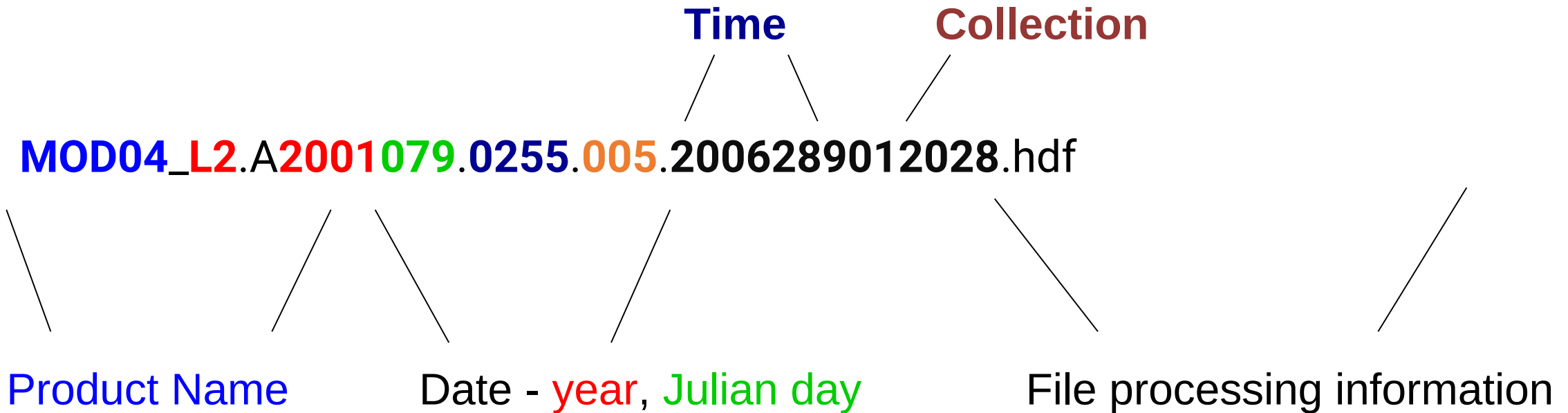
Aerosol Retrieval Algorithm



Level 3

Spatial & Temporal Averaging

Understanding a MODIS File Name



MODIS Aerosol Products

- **MODIS Level 2 data**

https://lpdaac.usgs.gov/about/news_archive/echo_announces_release_reverb

- **MODIS Level 3 data**

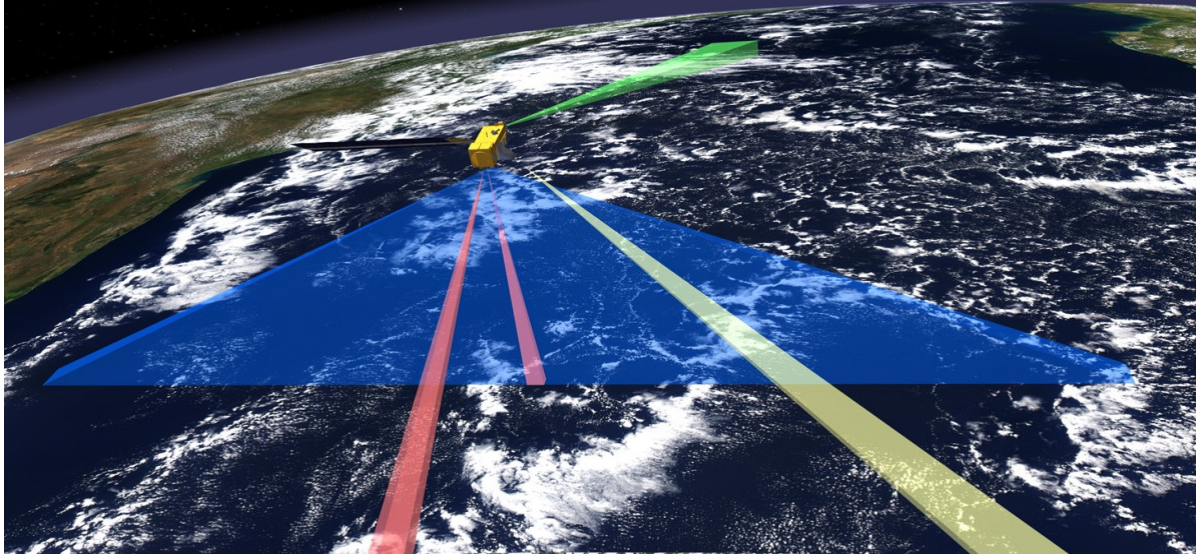
Giovanni – web tool for imagery visualization and analysis

<https://giovanni.gsfc.nasa.gov/giovanni>

- **MODIS Subsets**

<https://earthdata.nasa.gov/earth-observation-data/near-real-time/rapid-response/modis-subsets>

Ozone Monitoring Instrument



One of four sensors on the EOS-Aura platform (OMI, MLS, TES, HIRDLS)

An international project: Holland, USA, Finland

Launched on 07-15-04

Instrument Characteristics

- Nadir solar backscatter spectrometer
- Spectral range 270-500 nm (resolution~1nm)
- Spatial resolution: 13X24 km footprint
- Swath width: 2600 km (global daily coverage)

Retrieval Products

Column Amounts

- Ozone (O_3)
- Nitrogen Dioxide (NO_2)
- Sulfur Dioxide: (SO_2)
- Others

Aerosols

OMI data

OMI data site

<http://disc.sci.gsfc.nasa.gov/Aura/data-holdings/OMI>

Version 003 OMI Level 2, Level 2G, Level-3 and Climatology Products			
Short Name & Data Access			Product Description
Level-2 Orbital Swath (Nadir pixels 13x24 km)	Level-2G Global Binned (0.25x0.25 or 0.125x0.125 deg)	Level-3 Global Gridded (0.25x0.25 or 1x1 deg)	
Aerosols			
OMAERUV	OMAERUVG	OMAERUVd	OMI/Aura Near-UV Aerosol Optical Depth and single Scattering Albedo
OMAERO	OMAEROG	OMAEROe	OMI/Aura Multi-Wavelength Aerosol Optical Depth and single Scattering Albedo

[OMI-Aura_L2-OMAERUV_2011m1024t0521-o38692_v003-2011m1024t115317.he5](#)

Product name

YYYYmMMDDtHHMM

Cloud-aerosol lidar and infrared pathfinder satellite observations (CALIPSO)

Cloud-aerosol lidar and infrared pathfinder satellite observations (CALIPSO)

- The CALIPSO satellite was launched on April 28, 2006, with equator crossing times of about 13:30 and 01:30 and a 16-day repeating cycle
- It gives the distribution of aerosols and clouds in vertical atmospheric profiles on the global/regional scale
- It carries a Cloud-Aerosol Lidar with Orthogonal Polarization (CALIOP) instrument that operates at two wavelengths (532 nm and 1064 nm)
- CALIPSO satellite discriminates the different aerosol types including clean marine, dust, polluted continental, clean continental, polluted dust and smoke

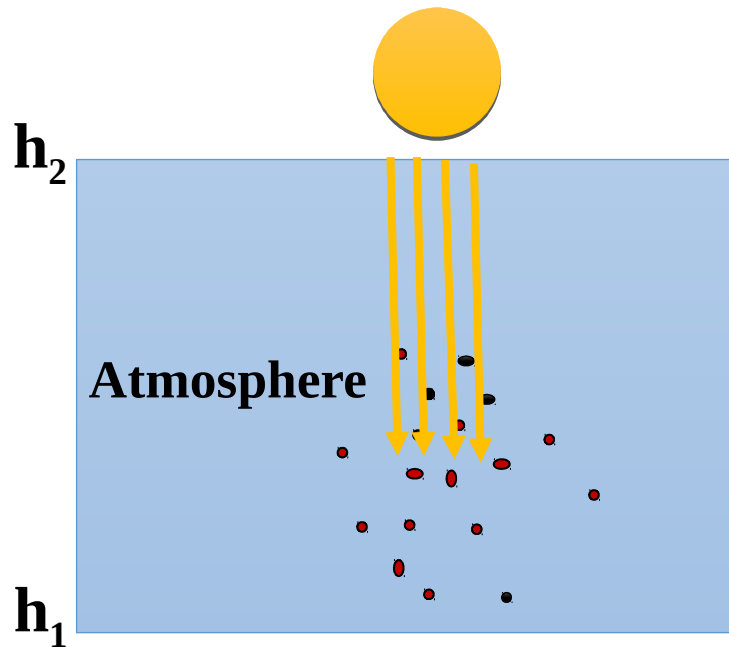
<https://www-calipso.larc.nasa.gov/>



Aerosol Optical Properties

Aerosol Optical Properties

Aerosol Optical Depth (AOD) and Angstrom Exponent (AE)



$$\text{AOD}(\lambda) = \int_{h_1}^{h_2} \beta_{\text{ext},\lambda}(\mathbf{h}) d\mathbf{h}$$

- **AOD** represents the total extinction of solar radiation caused by aerosol via **scattering** or **absorption**
- **AE** measures the aerosol size distribution

$$\text{AE} = -\frac{d \ln[\text{AOD}(\lambda)] / d \ln(\lambda)}{d \ln[\text{AOD}(\lambda)] / d \ln(\lambda)}$$

❖ AE is the size indicator of aerosol particle

Aerosol Optical Properties

Aerosol Index (AI)

AI is the ratio of measured to calculated radiances (at 360 nm)

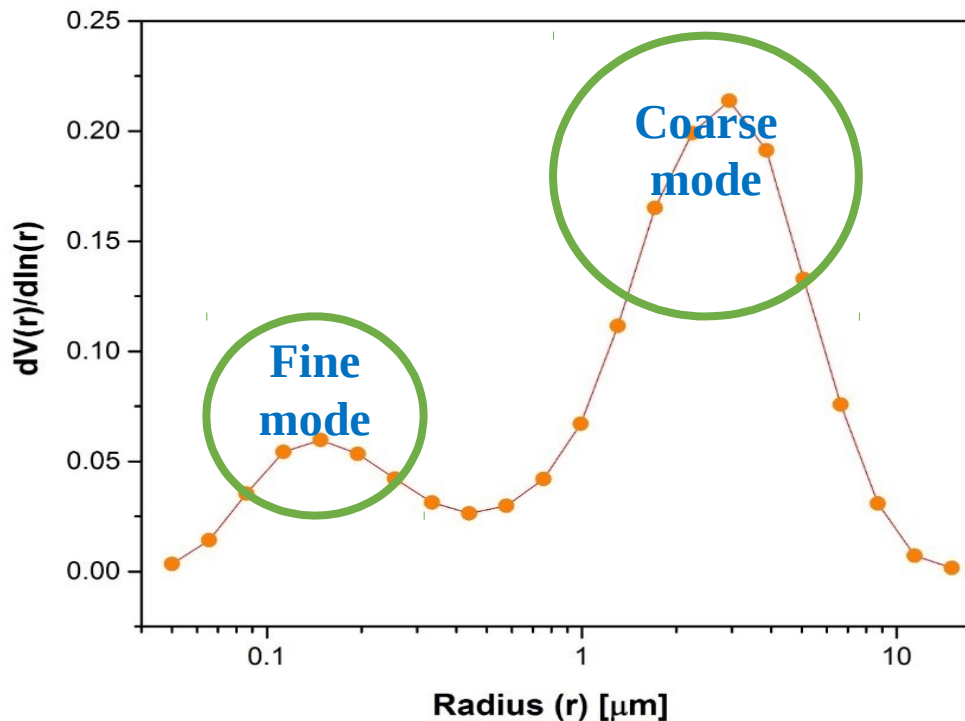
$$AI = 100 \log_{10} (I_{360} \text{ measured} / I_{360} \text{ calculated})$$

➤ **AI** separates the absorbing from non-absorbing aerosols in UV region

- ❖ $AI > 0$ for totally absorbing particles
- ❖ $AI < 0$ for purely scattering particles

Aerosol Optical Properties

Fine Mode Fraction (FMF) and Coarse Mode Fraction (CMF)



FMF/CMF of AOD is the fraction of fine/coarse mode AOD to the total AOD

$$\text{FMF} = \frac{\text{AOD}_{\text{fine mode}}}{\text{AOD}_{\text{total}}}$$

$$\text{CMF} = \frac{\text{AOD}_{\text{coarse mode}}}{\text{AOD}_{\text{total}}}$$

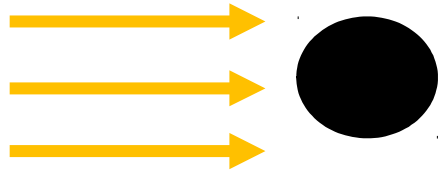
- ❖ Fine mode aerosols: 0.1 to 0.25 μm
- ❖ Coarse mode aerosols: 1.0-2.5 μm

➤ **FMF** gives information about the size distribution

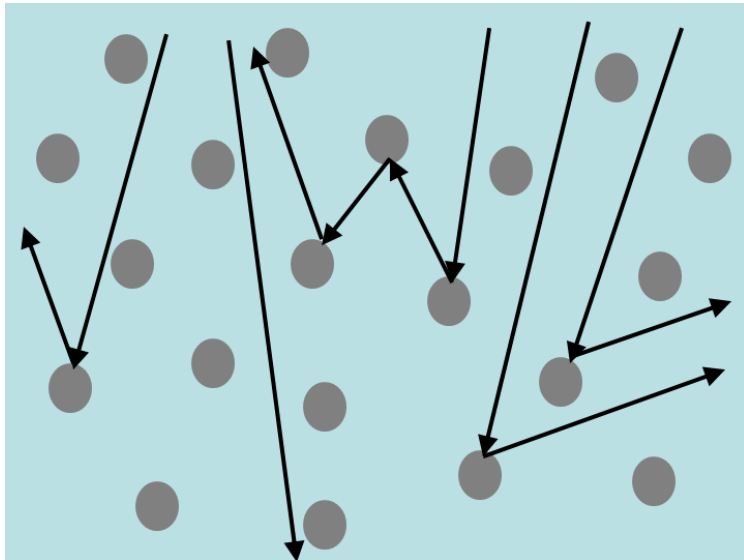
- ❖ FMF = 0 for single coarse mode particle
- ❖ FMF = 1 for single fine mode particle

Aerosol Optical Properties

Single Scattering Albedo (SSA)



Absorption



Scattering

SSA is the ratio of scattering efficiency (Q_{scat}) to total extinction efficiency (Q_{ext}):

$$SSA = \frac{Q_{scat}}{Q_{ext}} = \frac{Q_{scat}}{Q_{scat} + Q_{abs}}$$

➤ **SSA** is used to distinguish absorbing from non-absorbing aerosols

- ❖ SSA = 0 for totally absorbing (dark) particles
- ❖ SSA = 1 for purely scattering particles

Atmospheric Forcing

- In order to understand how the Earth's climate is changing, it is critical to make sense of each mechanism that causes warming or cooling in the atmosphere
- Each process that changes the balance of radiation coming into and going out of the Earth-Atmosphere system is known as atmospheric forcing
- The long wave (LW) and shortwave (SW) radiative forcing at the surface are important components of the Earth's radiation balance

Radiative Transfer Equations

- Once the solutions for the hemispheric intensities are known, it is easy to calculate hemispheric fluxes by performing the angular integrations
- Assumes that there is **no** dependency on ϕ in a plane-parallel atmosphere

Upward Fluxes

$$F_{\lambda}^{\uparrow} = 2\pi \int_0^{\pi} I_{\lambda}^{\uparrow}(\mu) \mu d\mu$$

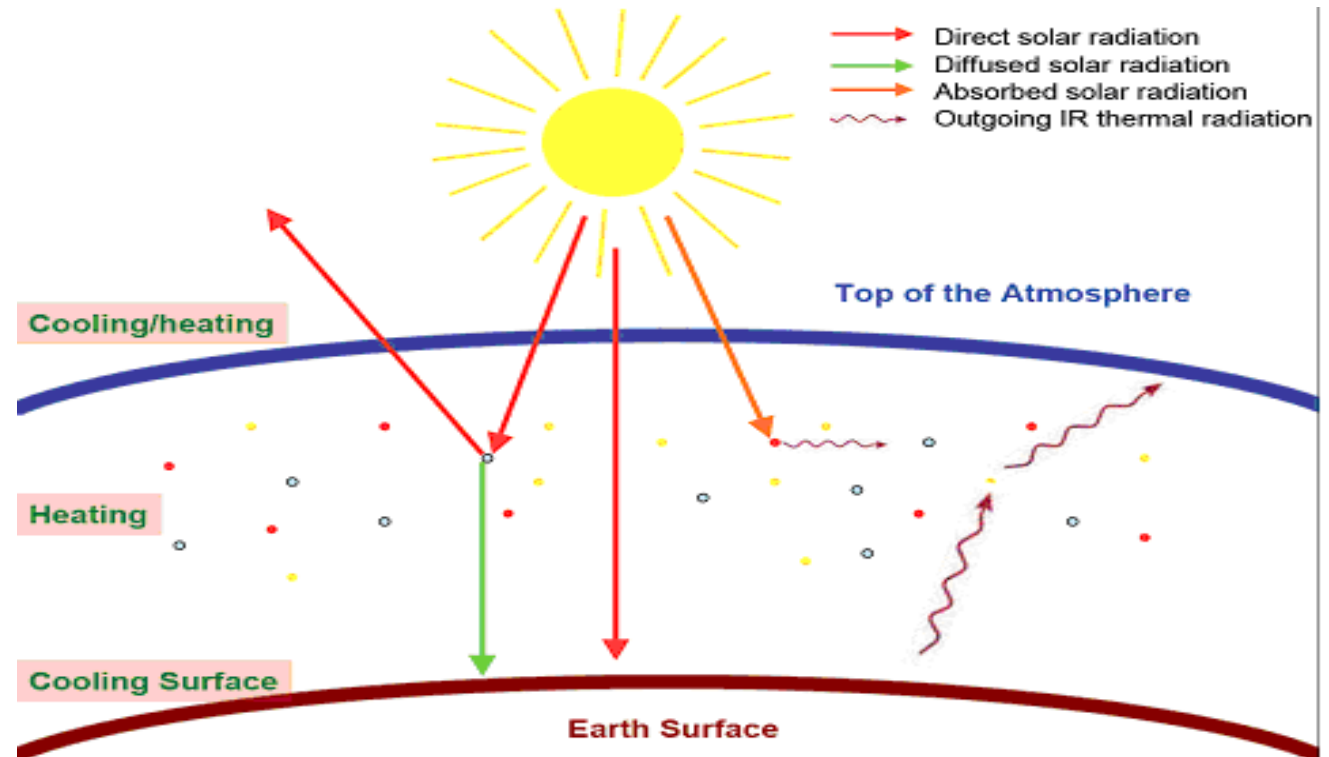
Downward Fluxes

$$F_{\lambda}^{\downarrow} = 2\pi \int_0^{\pi} I_{\lambda}^{\downarrow}(-\mu) \mu d\mu$$

The **total net flux** can be defined as:

$$F_{net} = F^{\uparrow} - F^{\downarrow}$$

$$\begin{array}{ll} F_{net} < 0 & \Rightarrow \text{cooling the atmosphere} \\ F_{net} > 0 & \Rightarrow \text{heating the atmosphere} \end{array}$$



$$\text{Heating rate} = \frac{\partial T}{\partial t} = \frac{g}{C_p} \frac{F_{net}}{\Delta P}$$

Aerosol Radiative Properties

Aerosol Radiative Forcing (ARF)

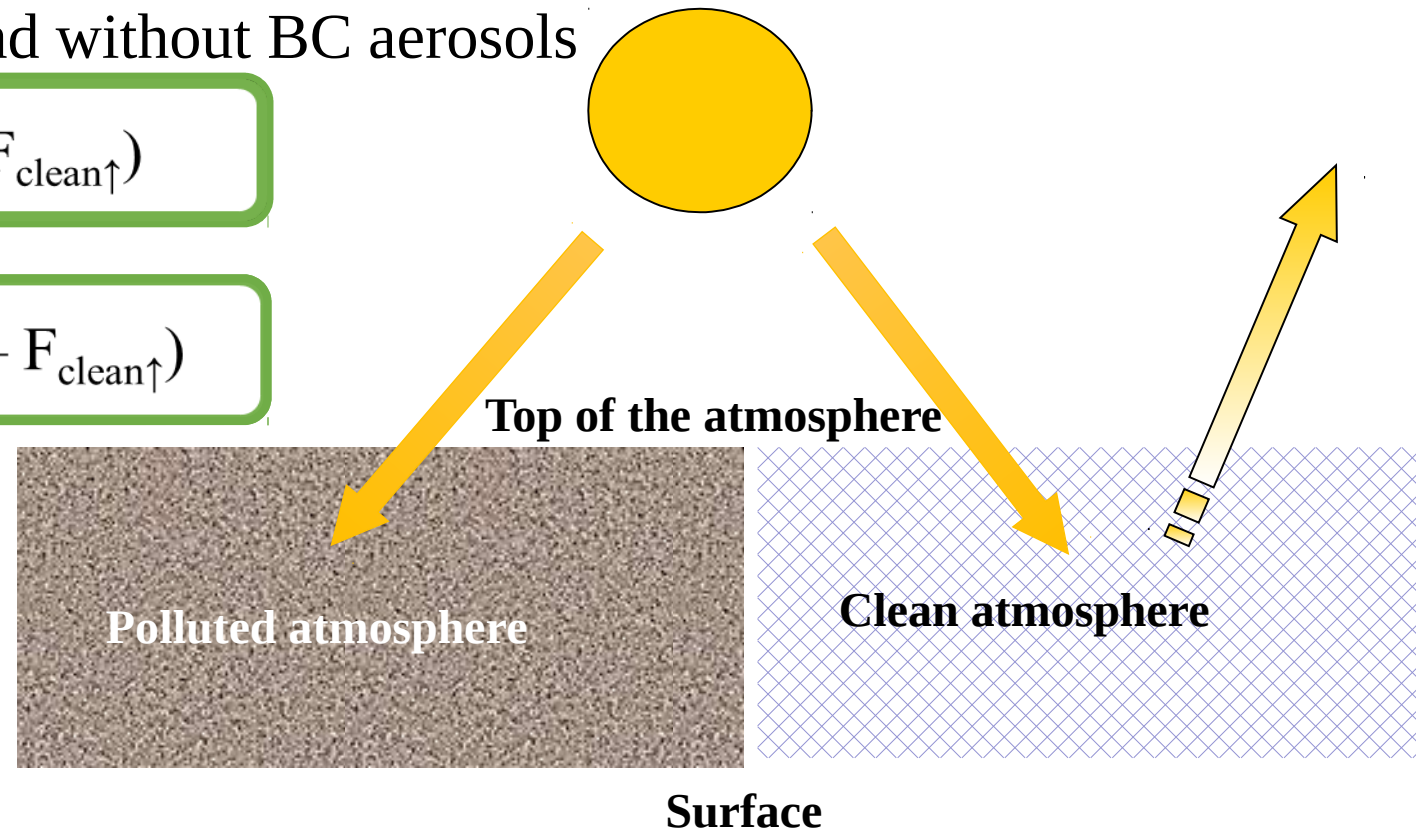
- The Aerosol radiative forcing at the TOA and the surface is obtained as the difference between the down and up fluxes for both with and without BC aerosols

$$F_{TOA} = (F_{aerosol\downarrow} - F_{aerosol\uparrow}) - (F_{clean\downarrow} - F_{clean\uparrow})$$

$$F_{Surface} = (F_{aerosol\downarrow} - F_{aerosol\uparrow}) - (F_{clean\downarrow} - F_{clean\uparrow})$$

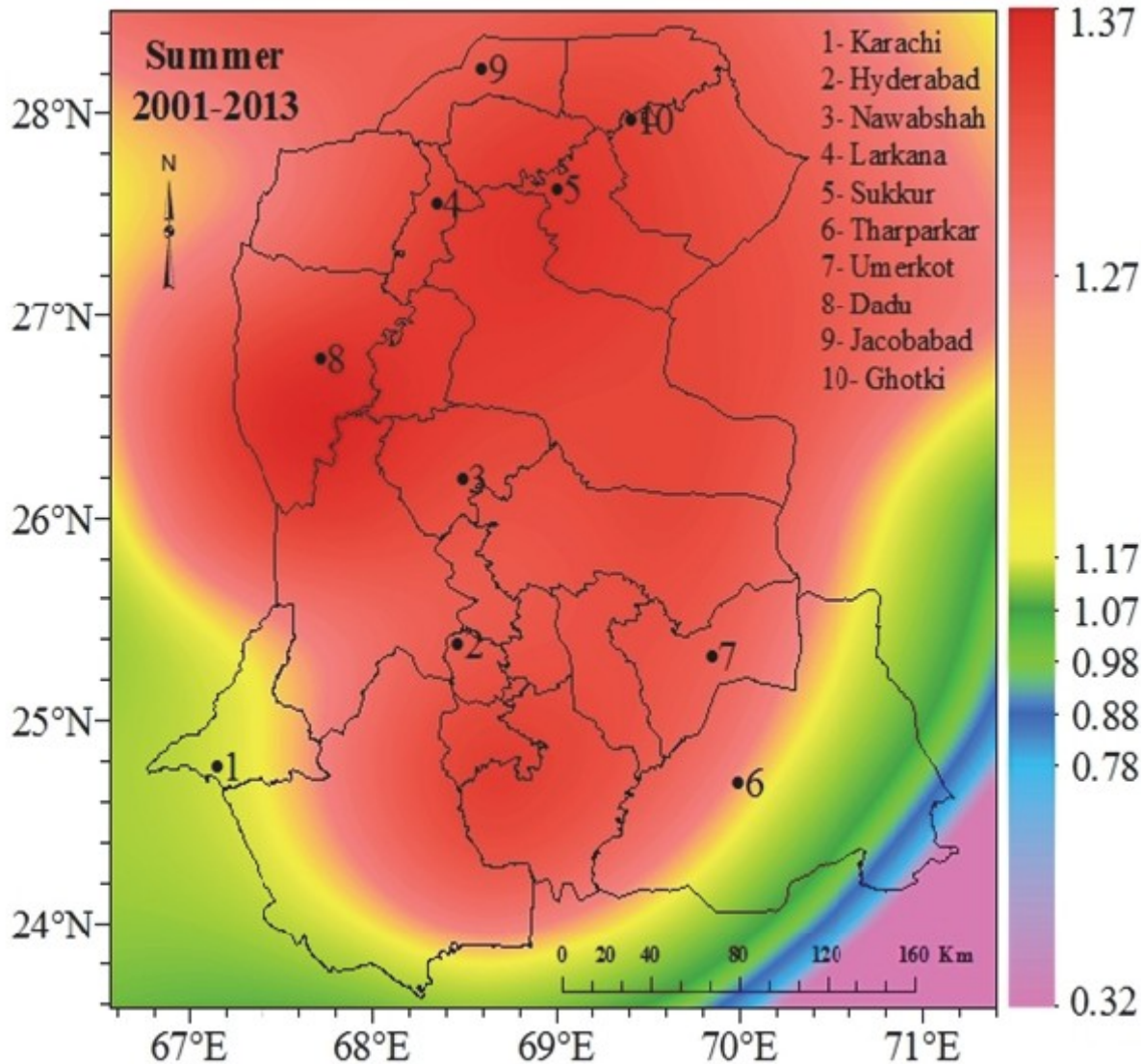
$$\Delta F_{Atmosphere} = F_{TOA} - F_{Surface}$$

$$\text{Heating Rate} = \frac{g}{C_p} \frac{\Delta F_{Atmosphere}}{\Delta P}$$



Analysis of Aerosol Optical Properties

MODIS Aerosol Optical Depth Product



Sharif et al., (2015)

MOD04 or MYD04

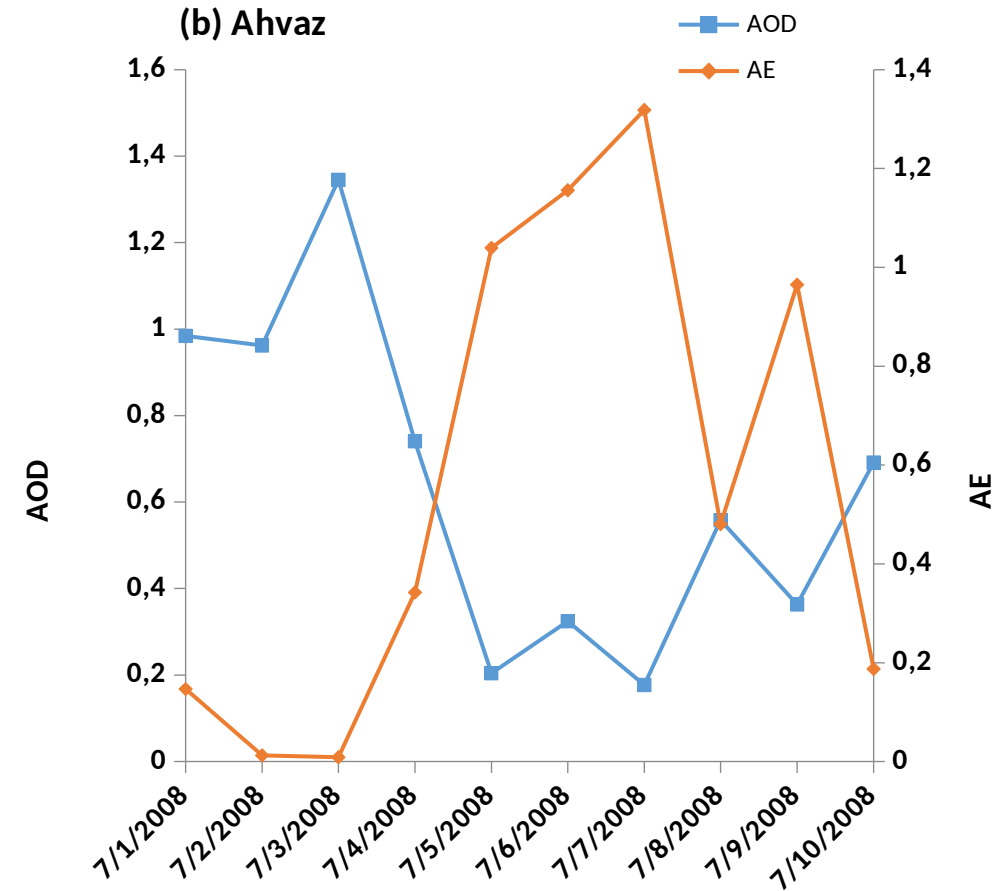
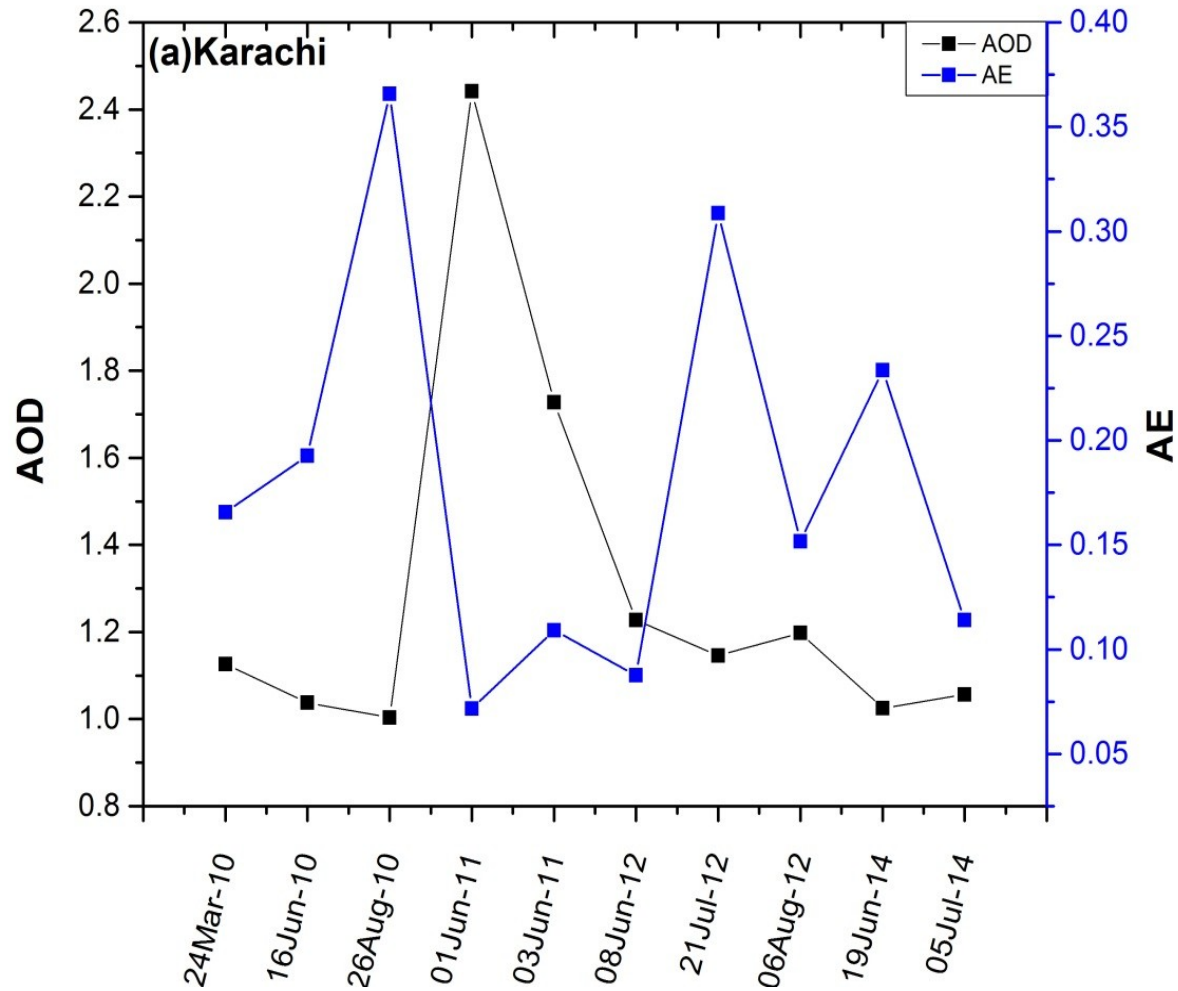
10 km – instantaneous

01 deg – daily, weekly, monthly

- At least two daytime overpasses - Terra and Aqua
- Sensitive to Boundary Layer
Industrial, smoke & dust aerosols
- Well validated over land

Aerosol Optical Properties

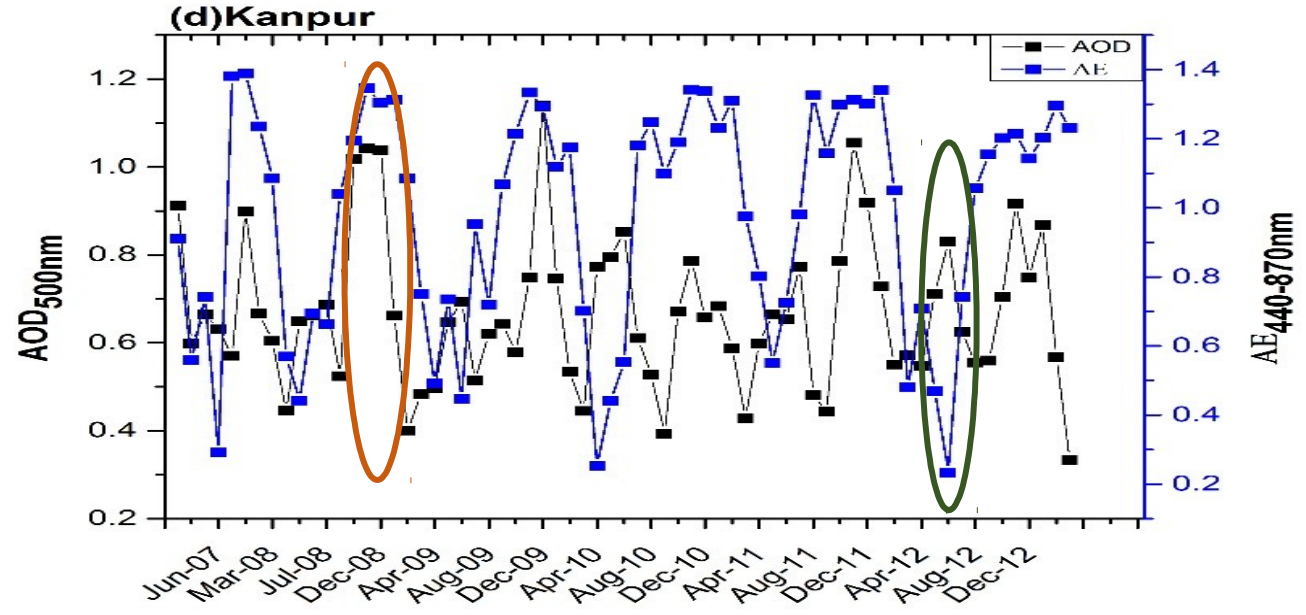
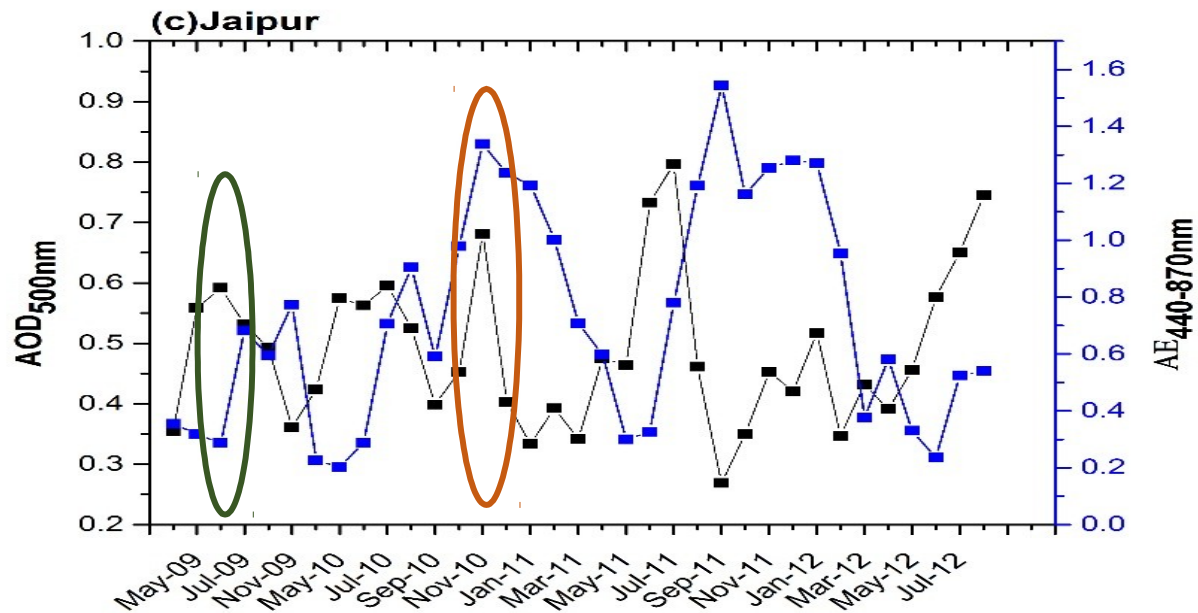
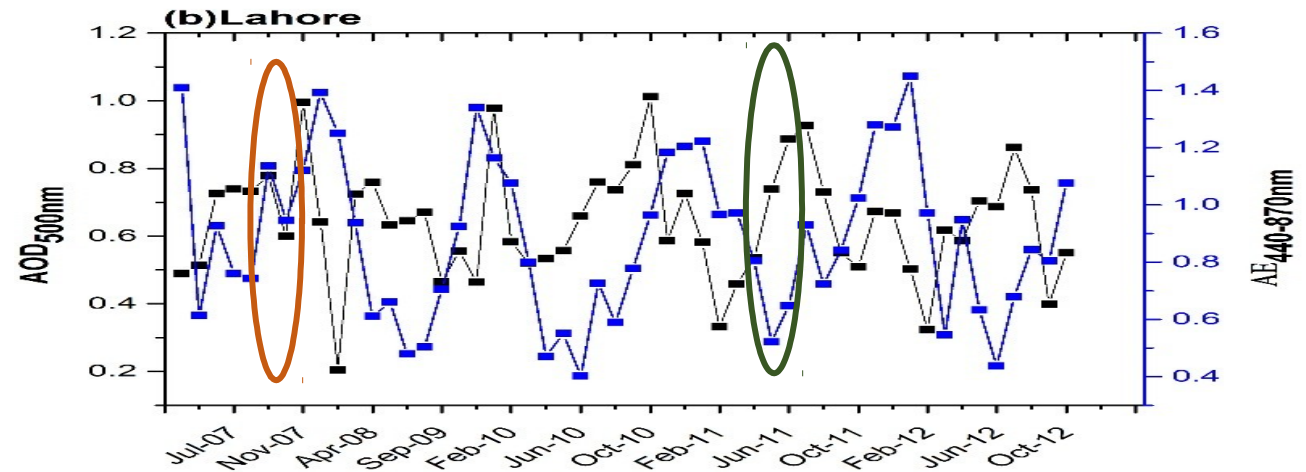
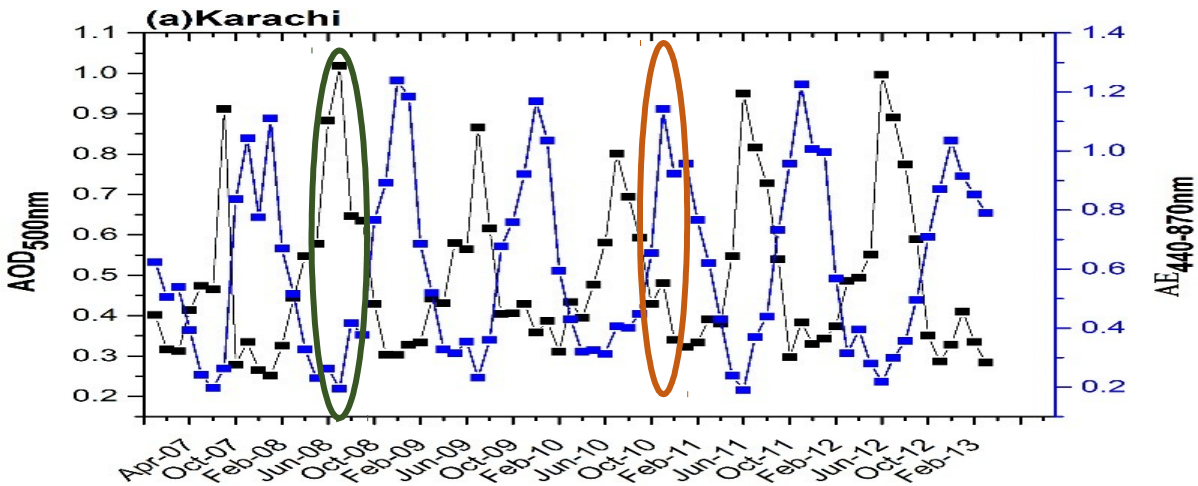
Aerosol Optical Depth (AOD) and Angstrom Exponent (AE) on Dusty Days



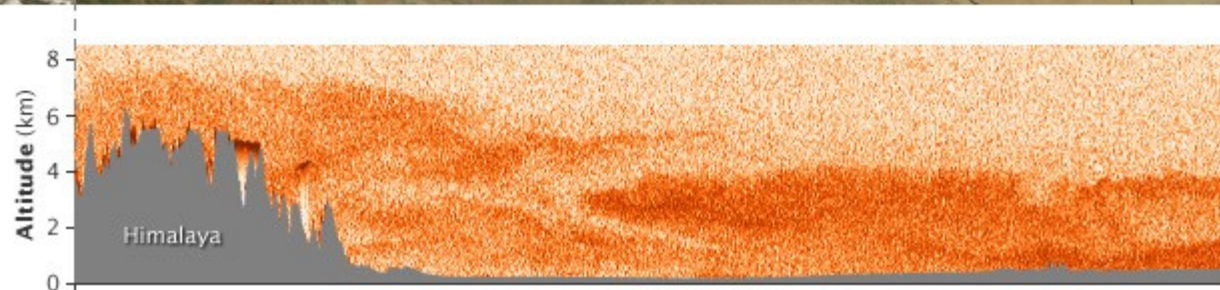
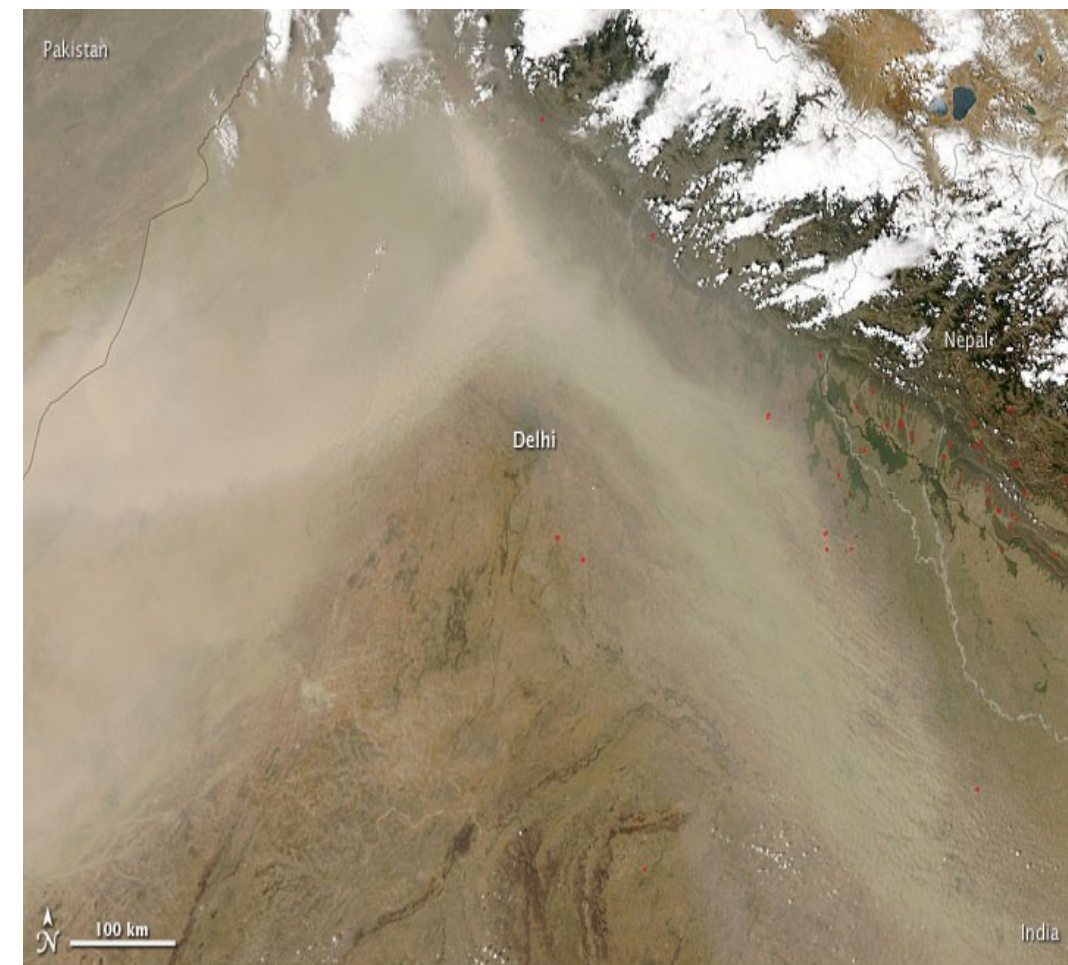
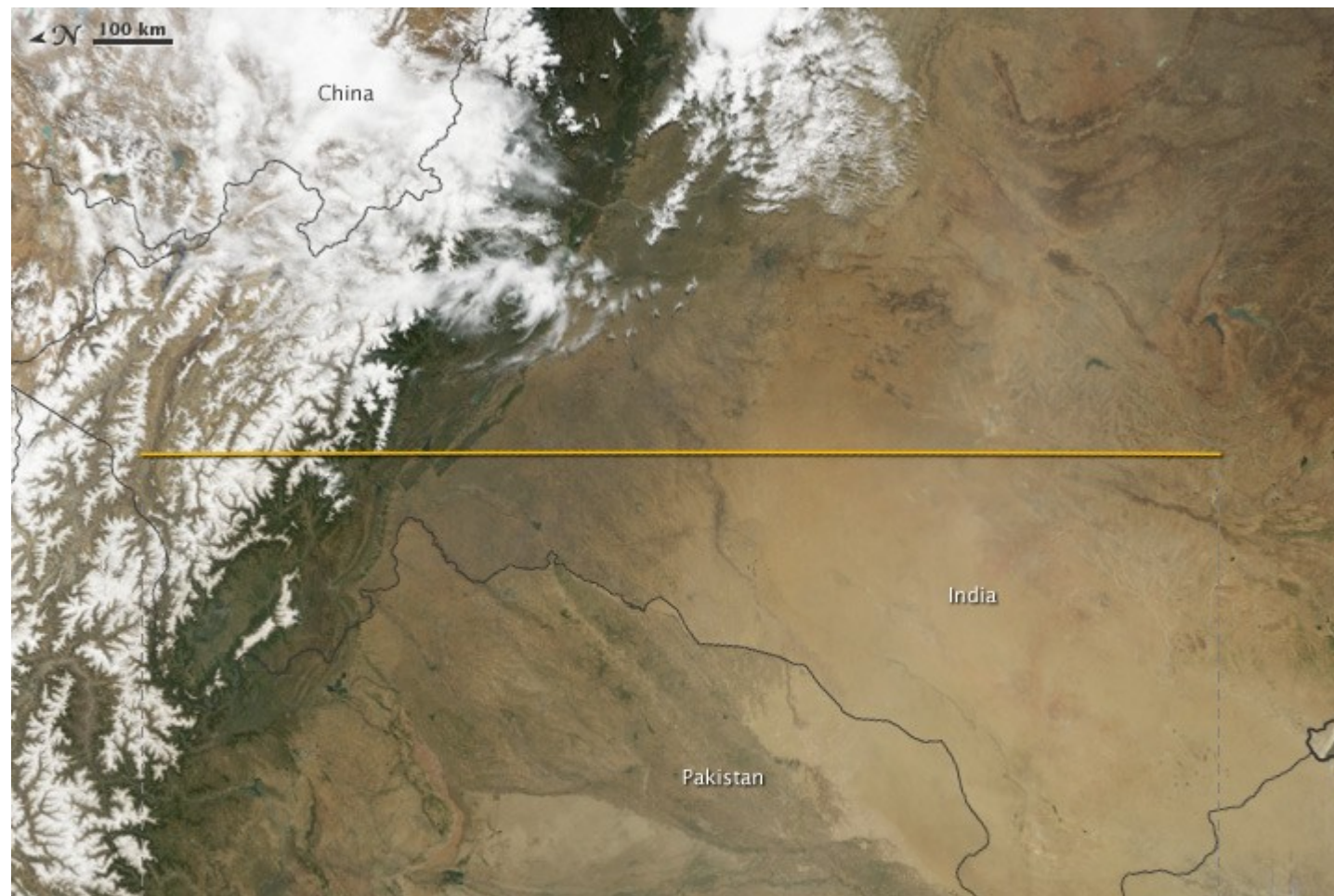
Monthly Averaged Variability of AOD and AE

High AOD with low AE during summer and pre-monsoon months

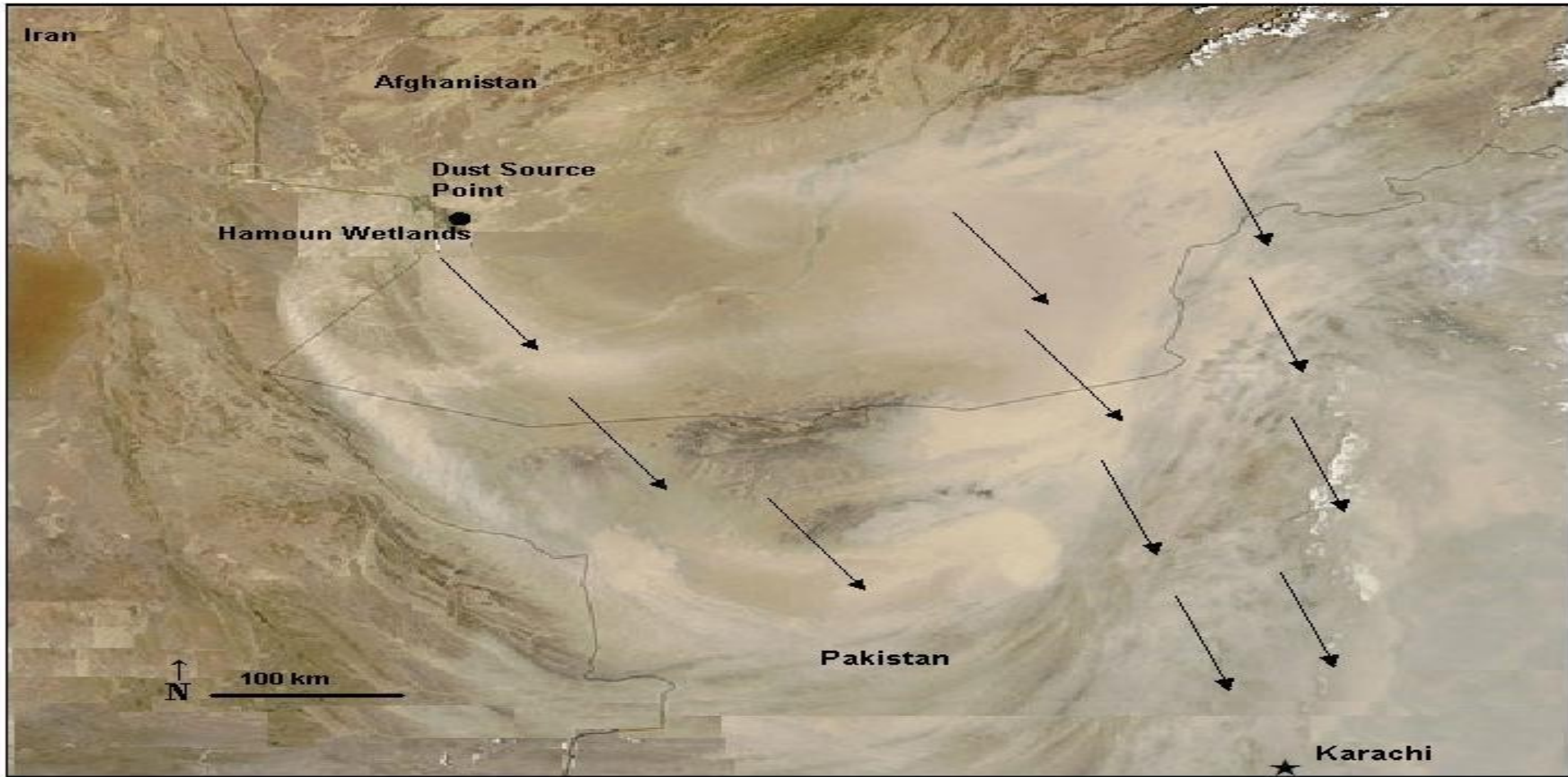
High AOD with high AE during winter and post-monsoon



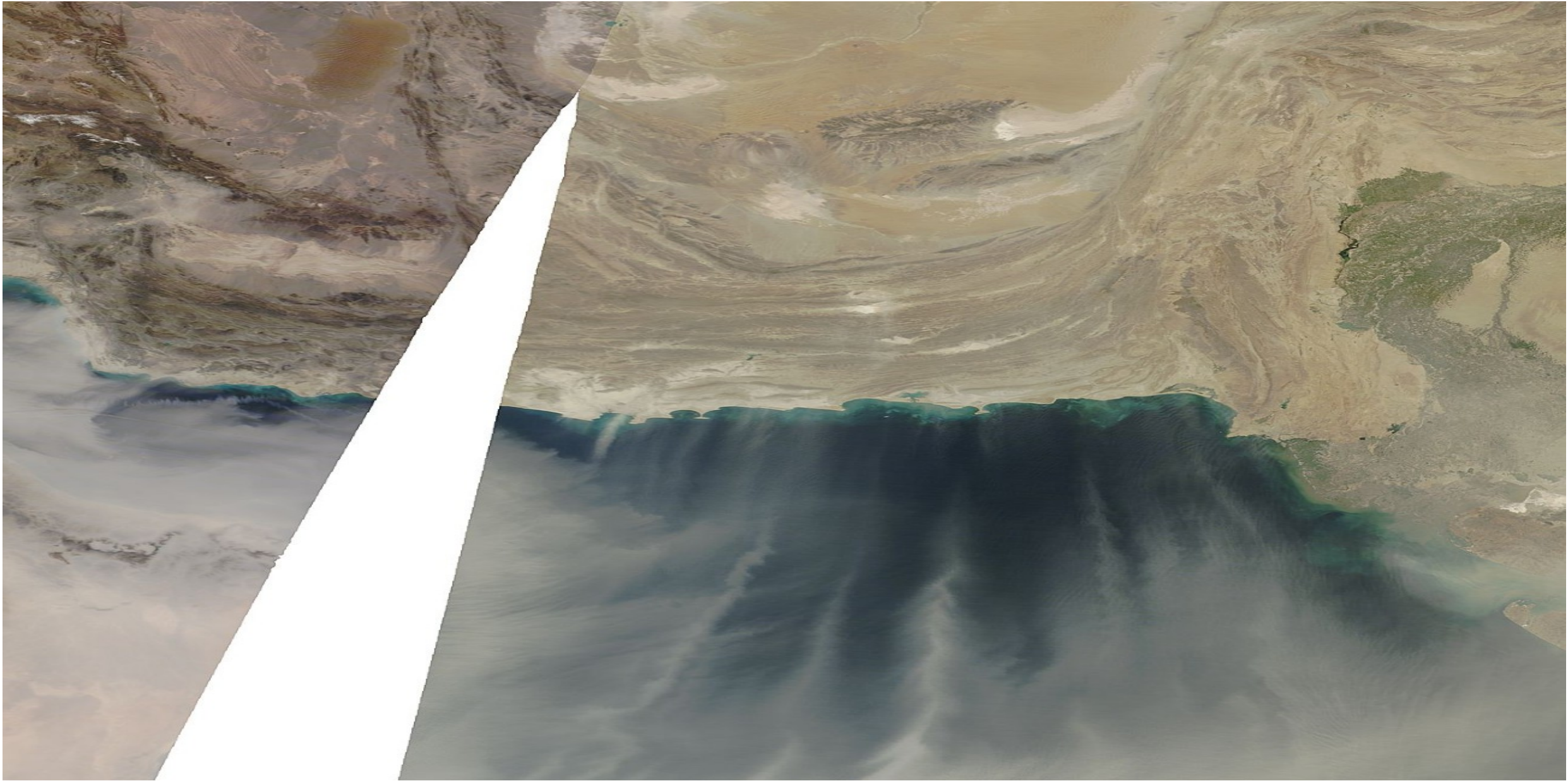
Evidence of Dust by MODIS in the Himalayan foothills



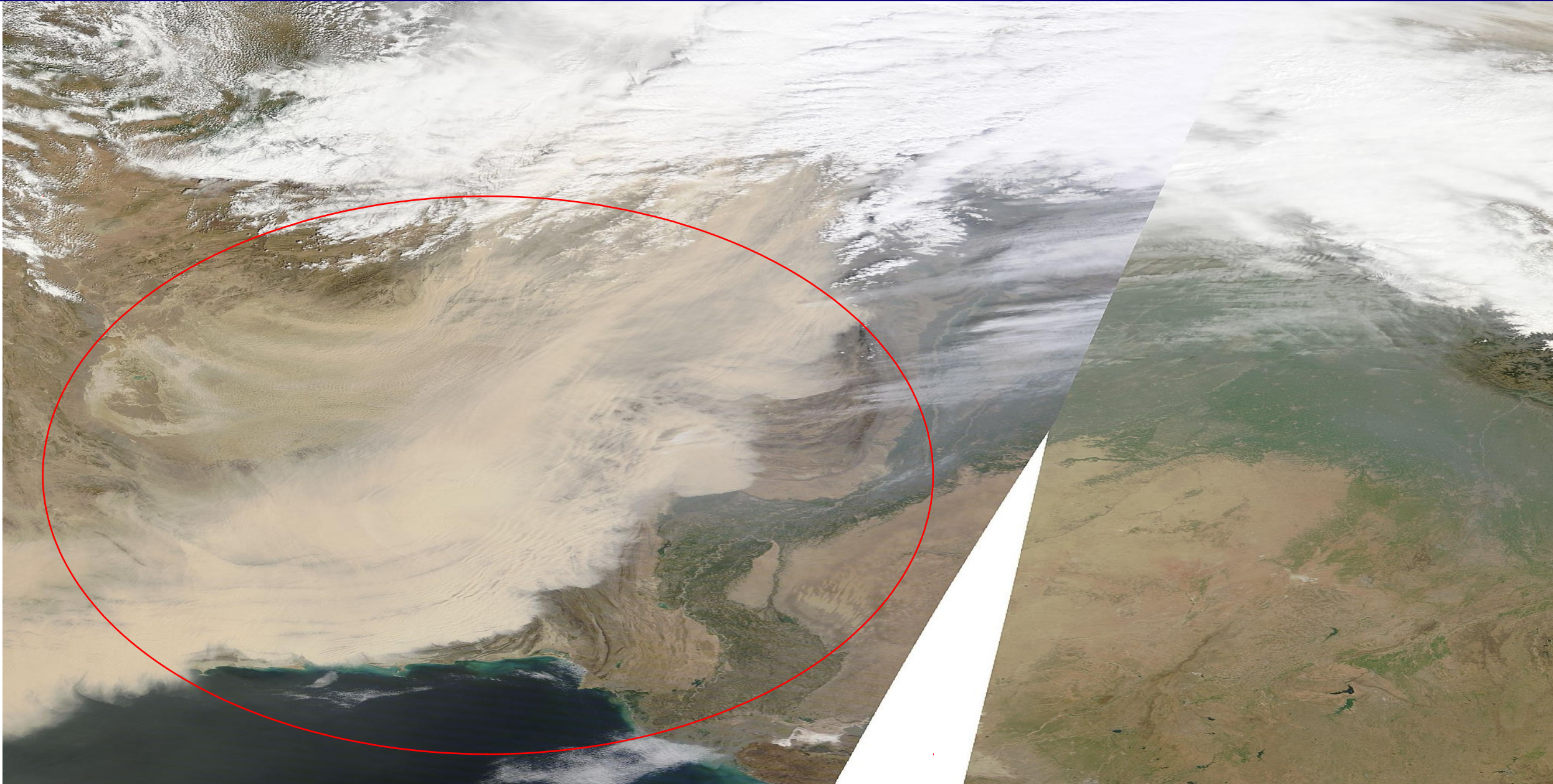
Dust Transport from Hamoun Wetlands



Dust Transport over the Arabian Sea



Dust over Iran & Pakistan



Doing More with Satellite Imagery (Climate studies)

If we understand the physics of how particular wavelengths interact with objects in the world we can create images to emphasize what we want to see

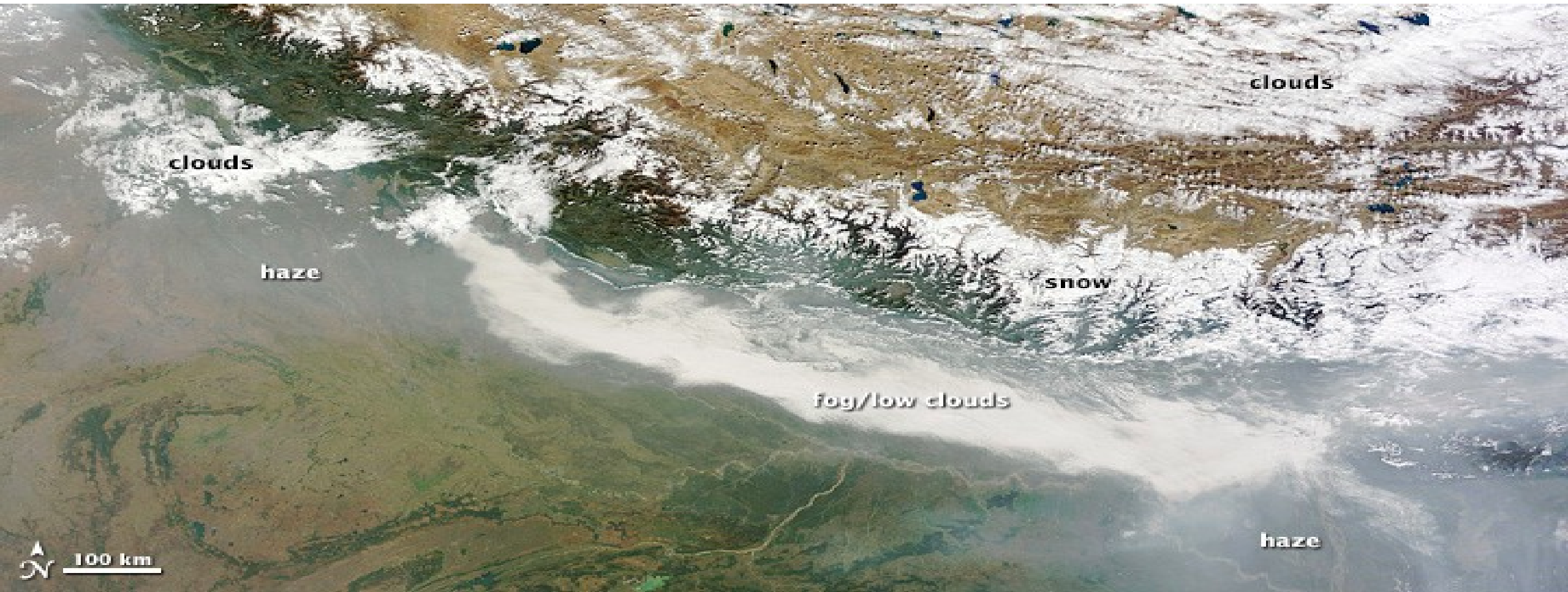


In visible imagery water is dark because it absorbs most of the energy.

Clouds are white because most of the incoming energy is reflected

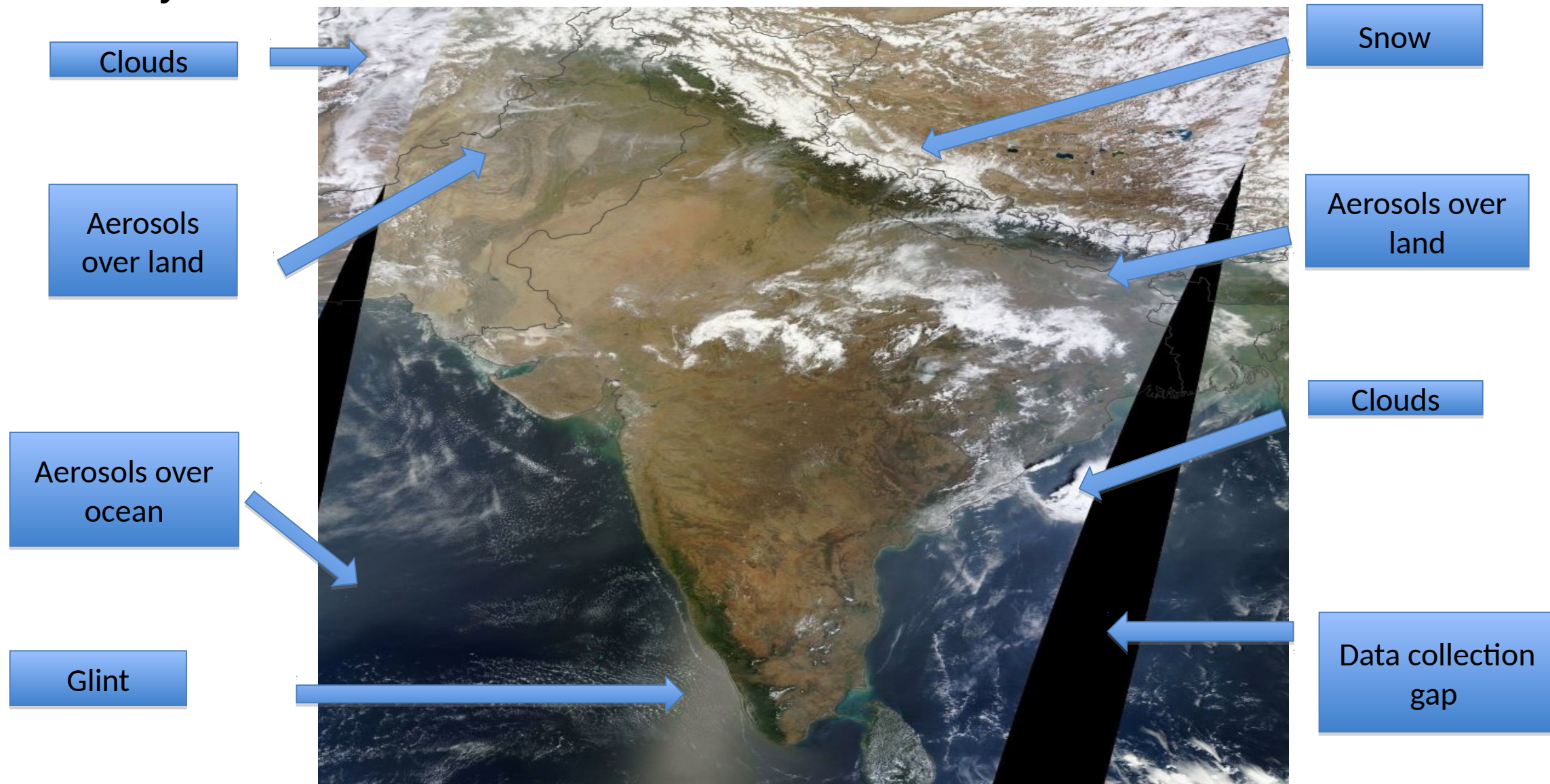
Pollution is hazy depending upon its absorptive properties

Features In True Color Image (Atmosphere)



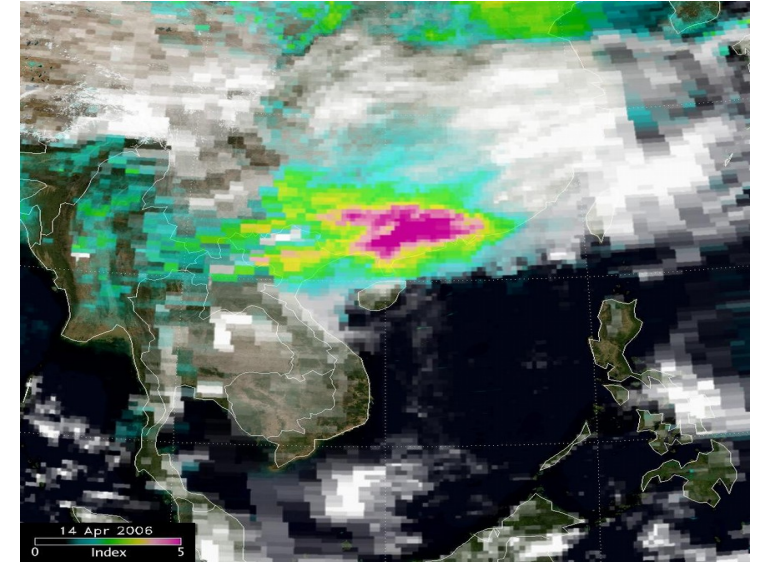
Satellite/Sensor identification feature

Identify Feature



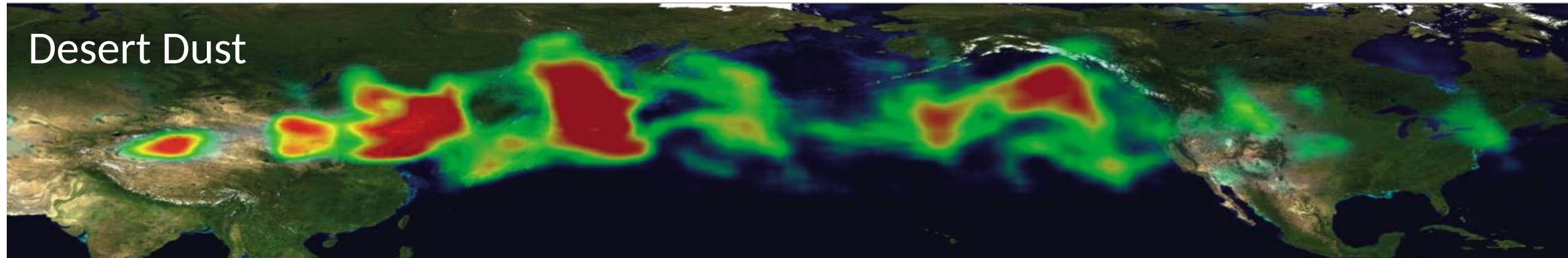
Aerosol Index

- Validation tool for transport models
- Separation of carbonaceous from sulfate aerosols
- Identification of aerosols above PBL (i.e., PBL aerosols are not detectable by AI)
- Tracking of aerosol plumes above clouds and over ice/snow

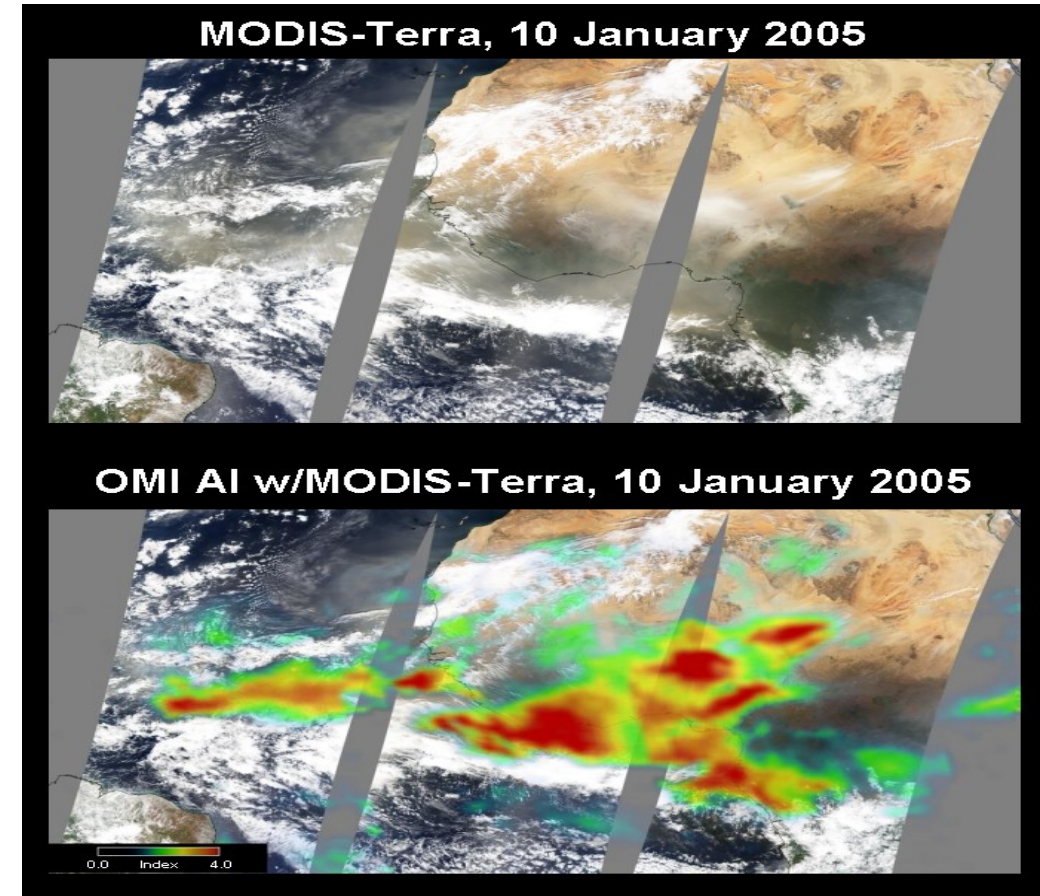
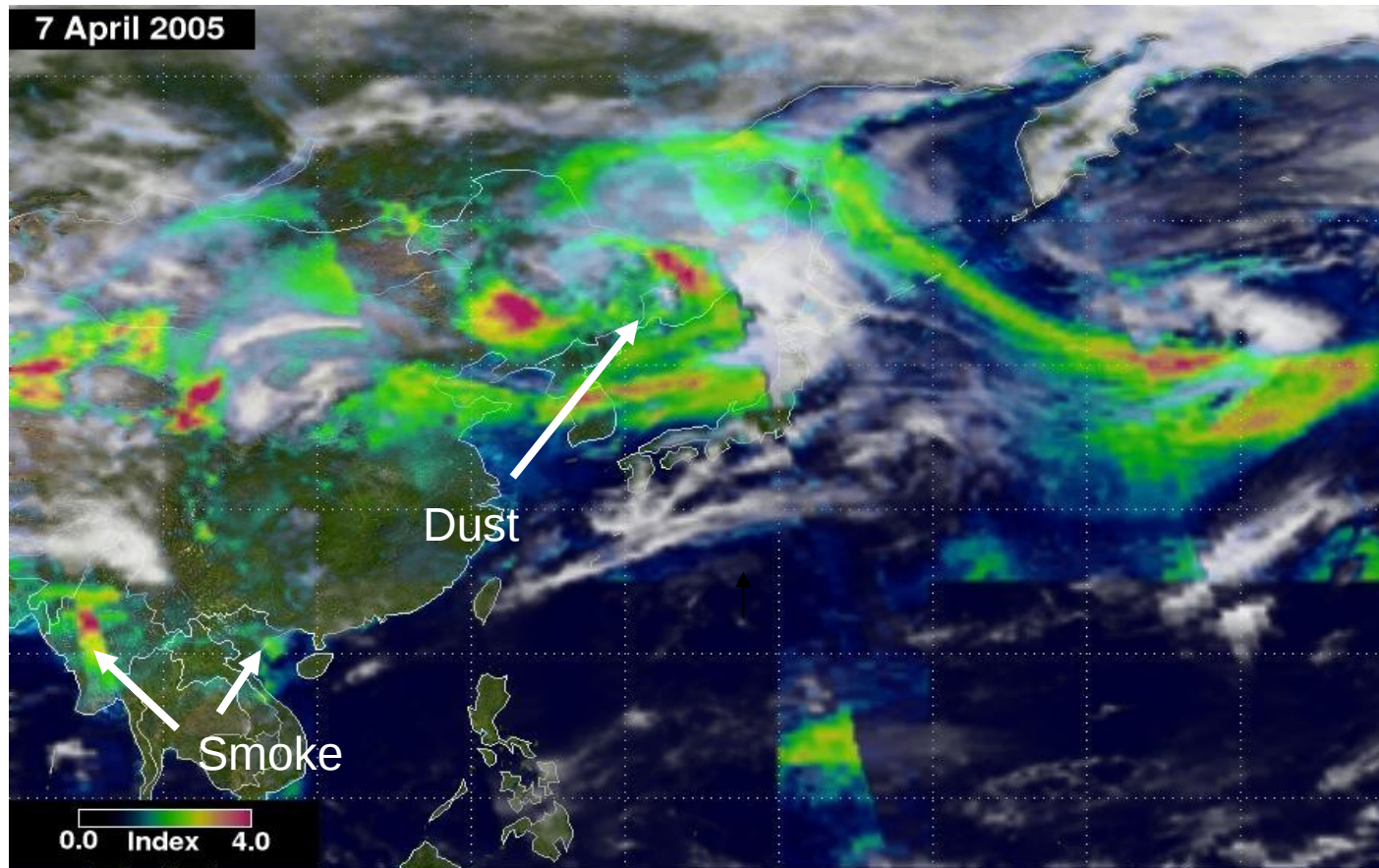


Aerosols over clouds:
April 14, 2006

Desert Dust



Aerosol Index



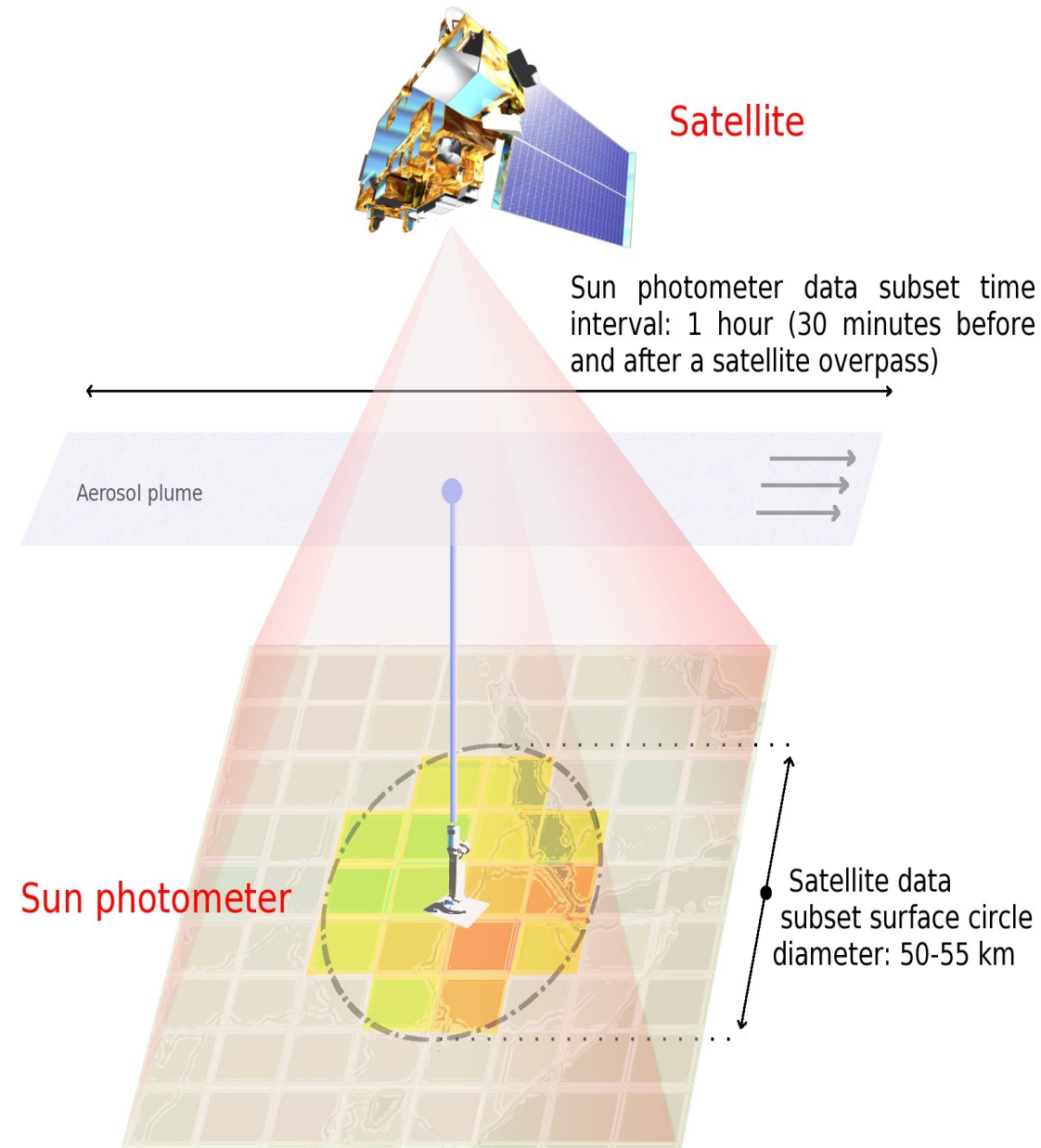
Aerosol Transport across the Oceans in terms of the Absorbing Aerosol Index

Satellite & Ground data validation

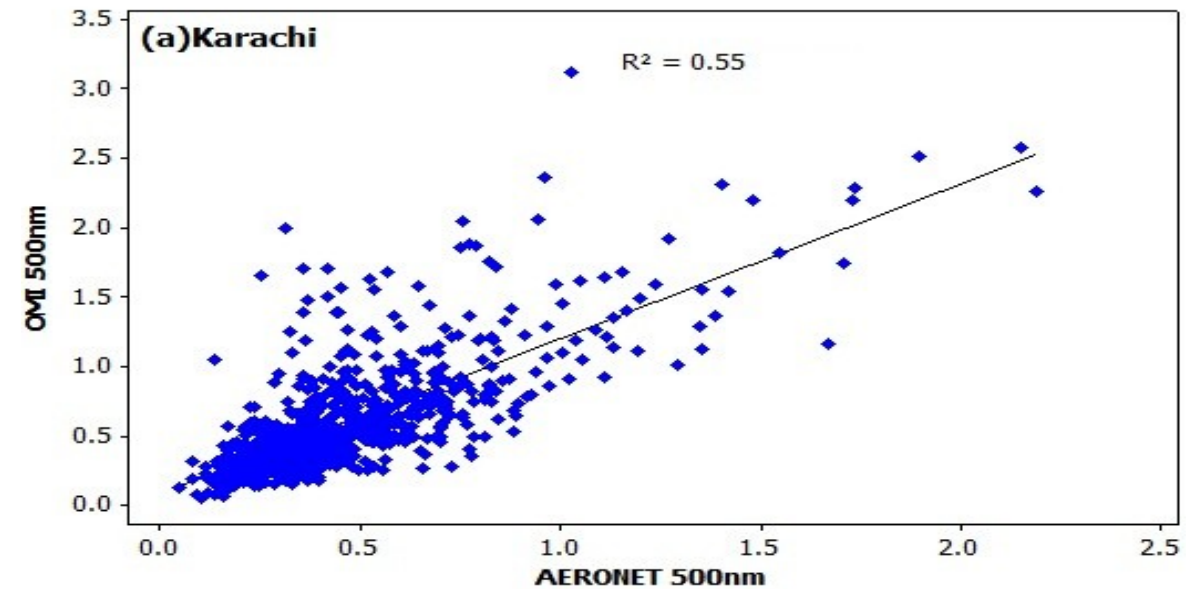
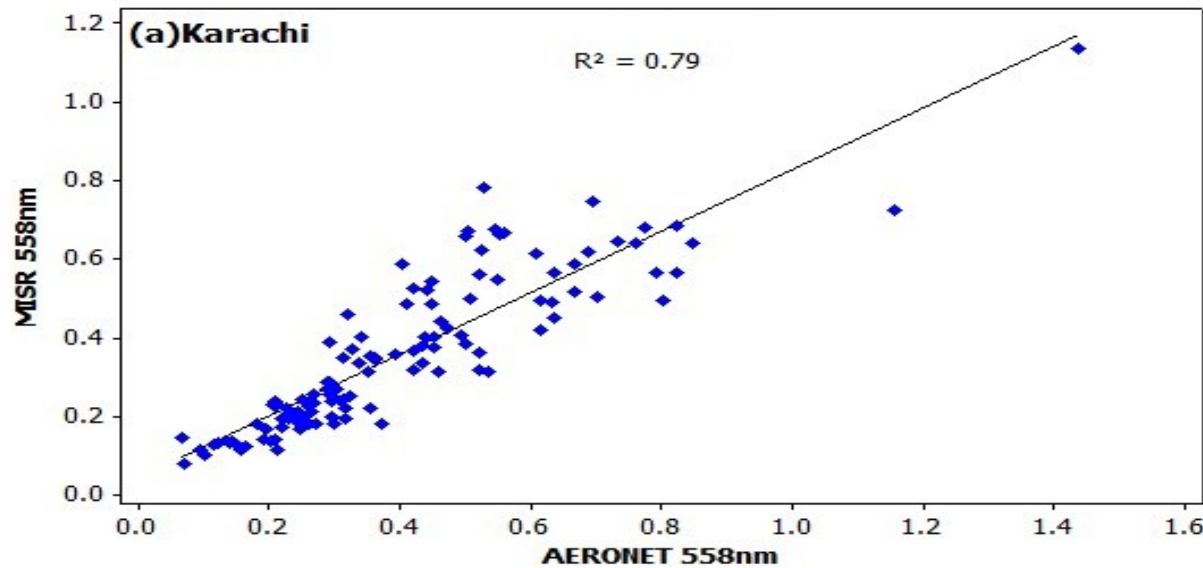
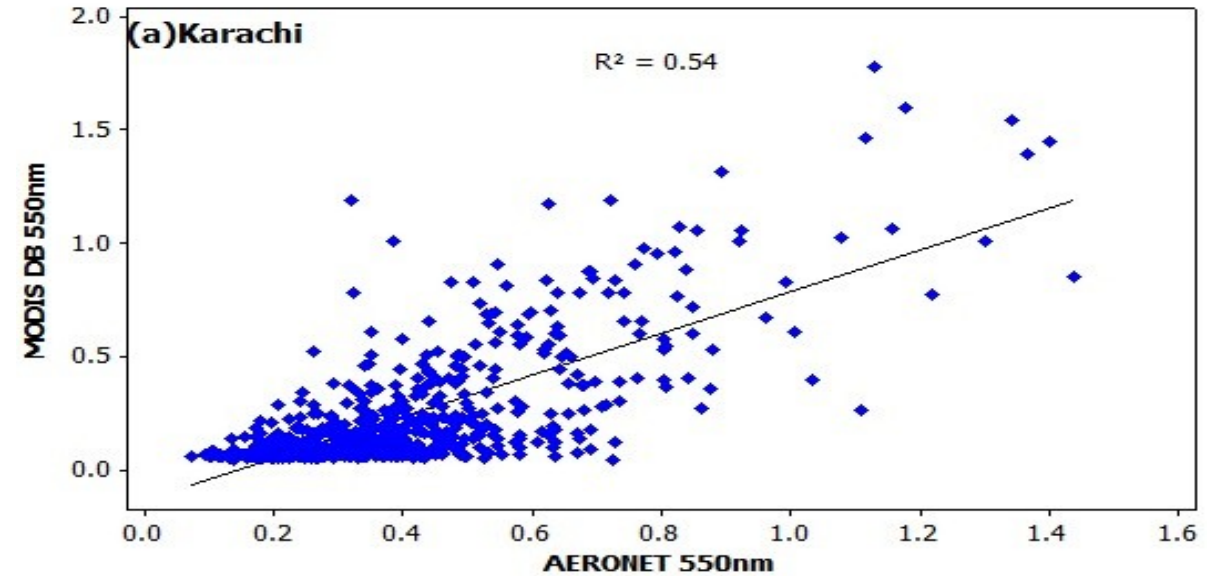
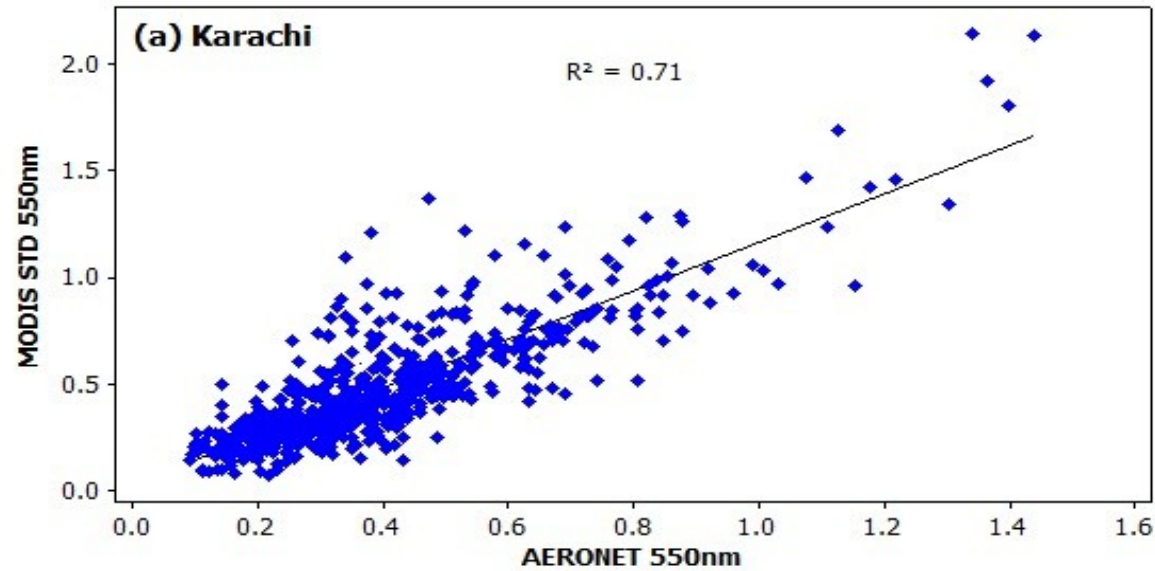
- ❖ An intercomparison of AOD values from different sensors is necessary if a long-term database for climatological studies is to be established
- ❖ To improve the accuracy and the coverage achievable with a single sensor
- ❖ AERONET AOD values (at 500 nm) were interpolated at common wavelength of 550 nm, using power law:

$$\text{AOD}_{550\text{nm}} = \text{AOD}_{500\text{nm}} \left(\frac{550}{500} \right)^{-\alpha}$$

where α is the (440–870 nm) Angstrom exponent

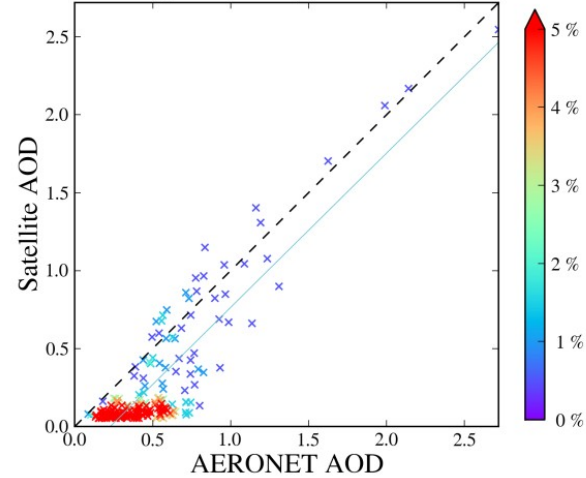


Intercomparison of Satellite and AERONET AOD



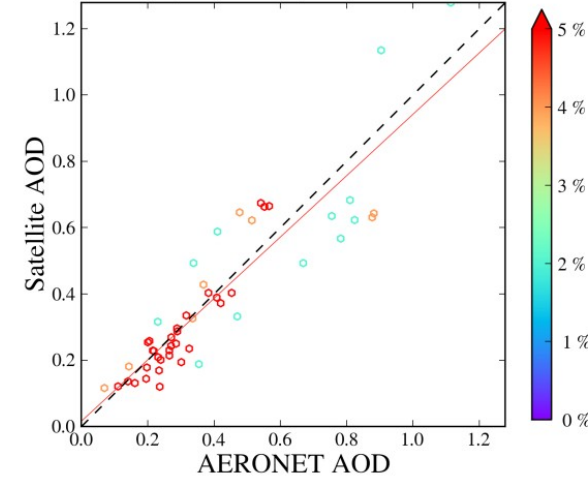
Intercomparison of Satellite and AERONET AOD

Land AOD-mean(Complete), '10-'13, seasons: All, QA: pre-filtered, at Karachi, outliers removed



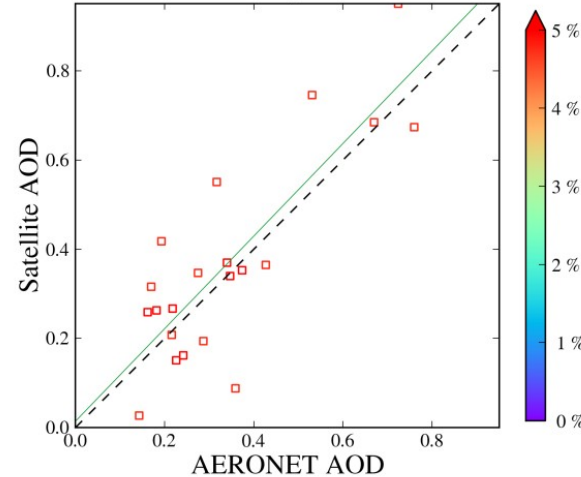
✕ AMODIS DB: $R^2=0.76$, $RMSE=0.30$, $y=0.99x-0.22$, $N=197$, outliers=2

Land AOD-mean(Complete), '10-'13, seasons: All, QA: pre-filtered, at Karachi, outliers removed



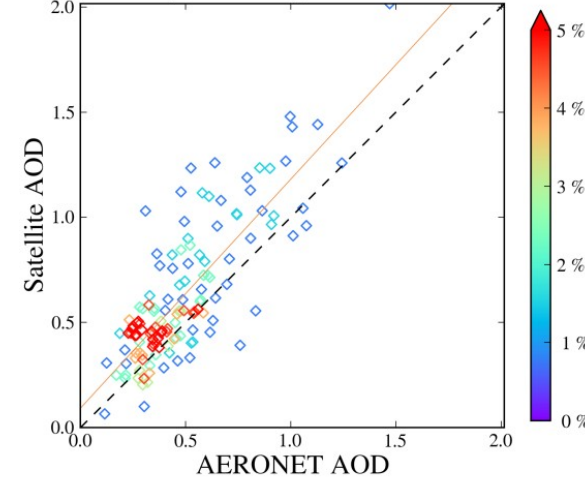
● MISR: $R^2=0.81$, $RMSE=0.11$, $y=0.93x+0.02$, $N=50$, outliers=2

Land AOD-mean(Complete), '10-'13, seasons: All, QA: pre-filtered, at Karachi, outliers removed



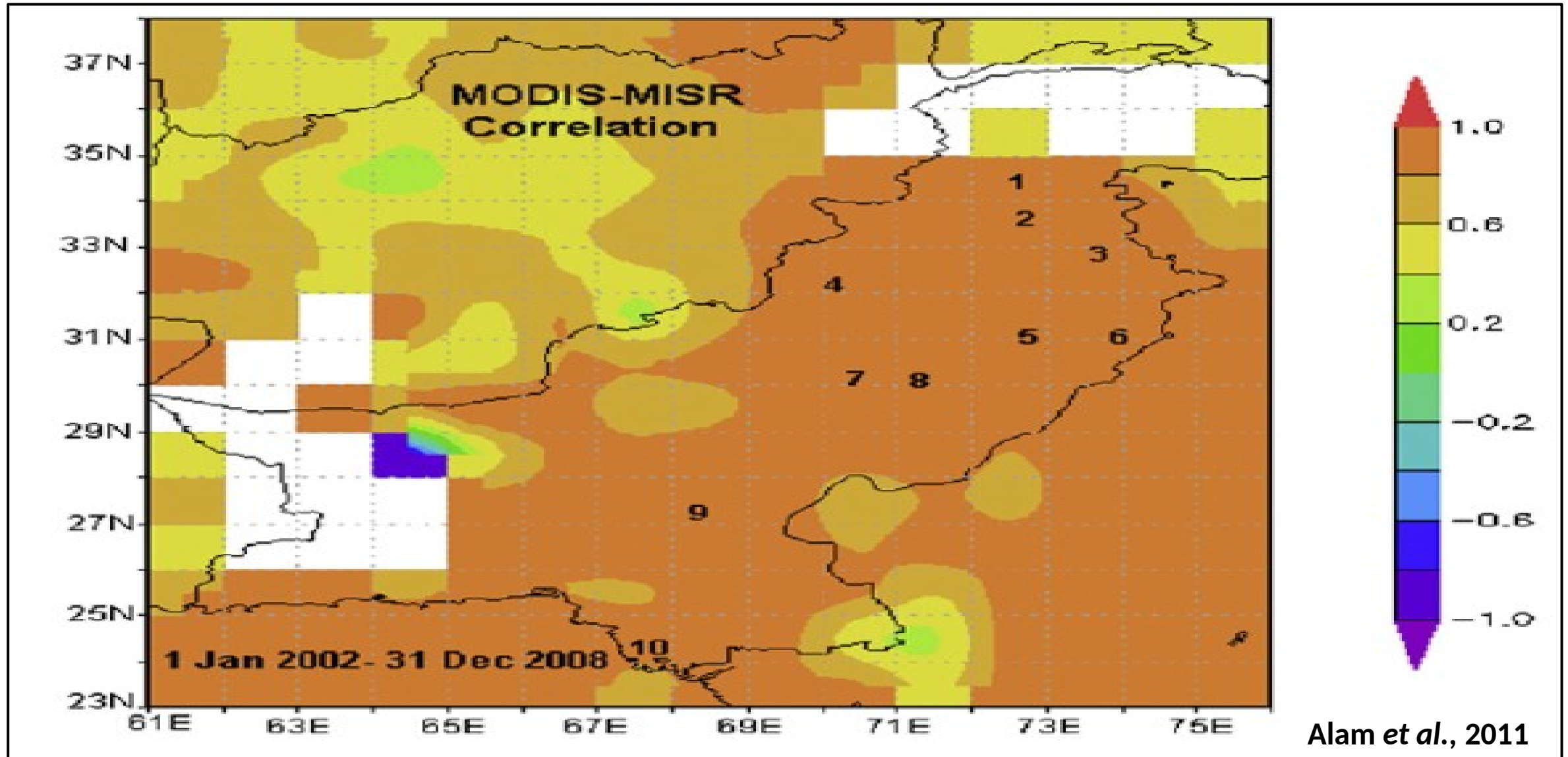
□ TMODIS: $R^2=0.68$, $RMSE=0.13$, $y=1.04x+0.01$, $N=21$, outliers=0

Land AOD-mean(Complete), '10-'13, seasons: All, QA: pre-filtered, at Karachi, outliers removed

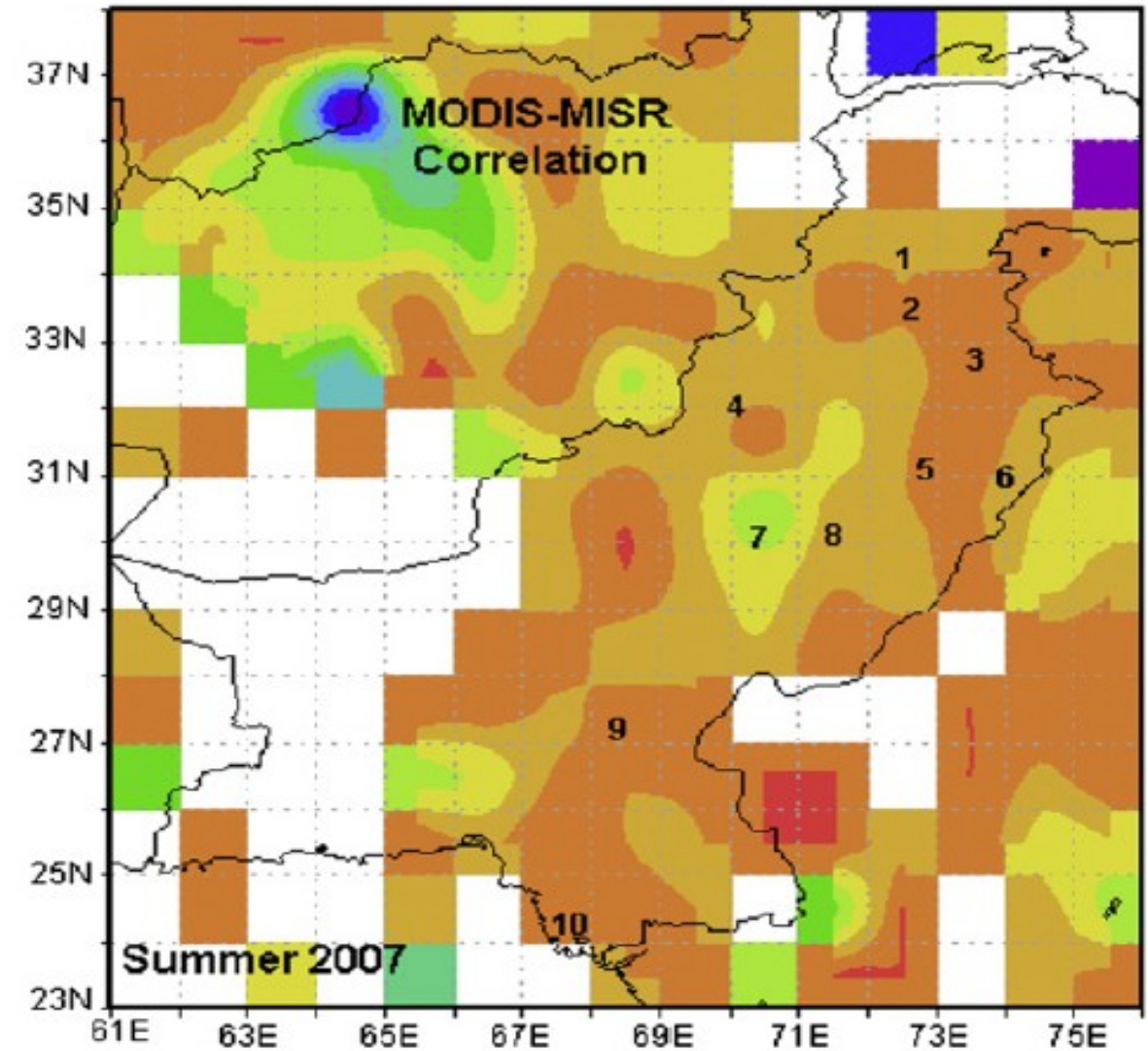
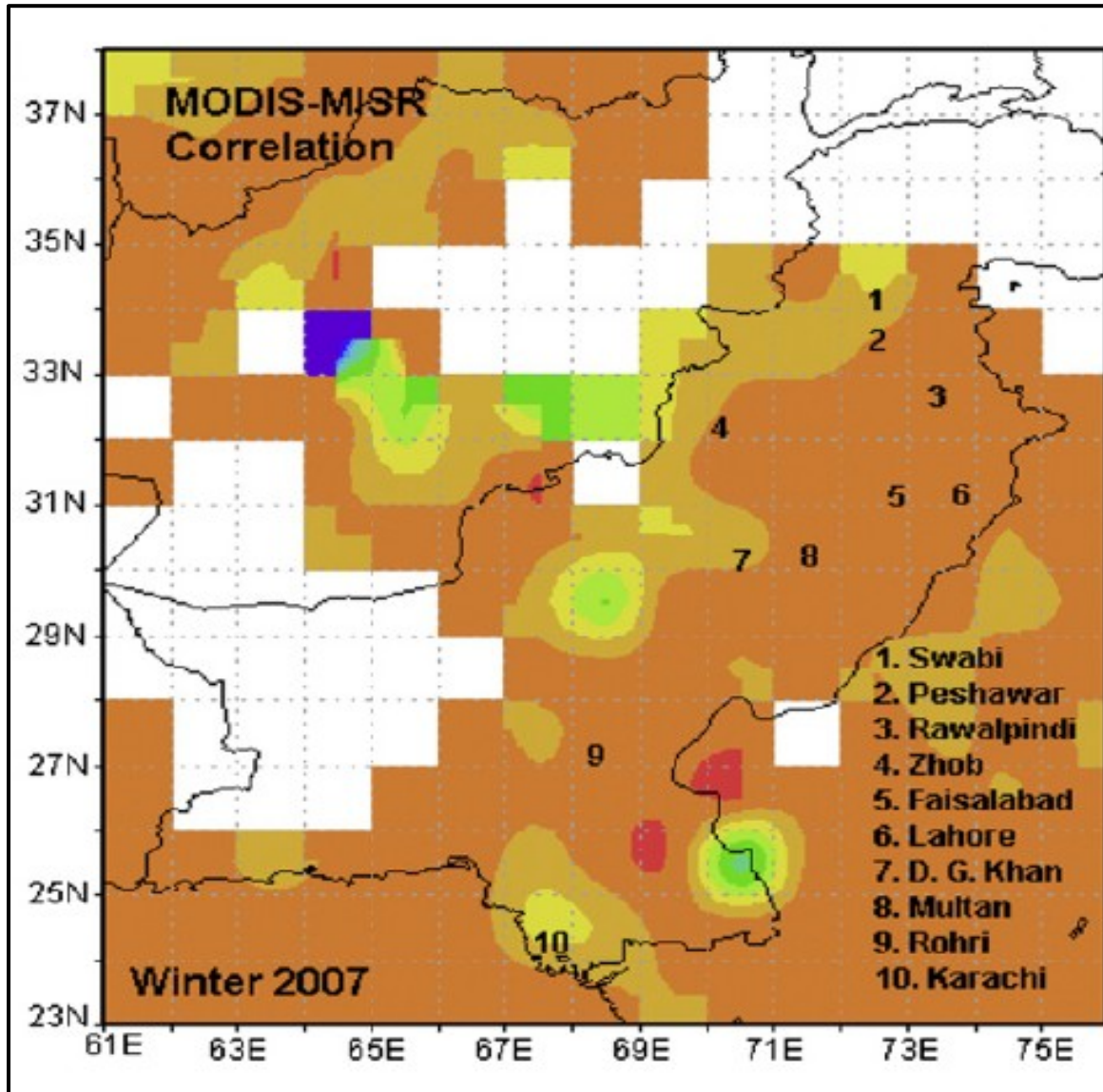


◇ OMI: $R^2=0.65$, $RMSE=0.24$, $y=1.09x+0.09$, $N=131$, outliers=3

MODIS vs. MISR

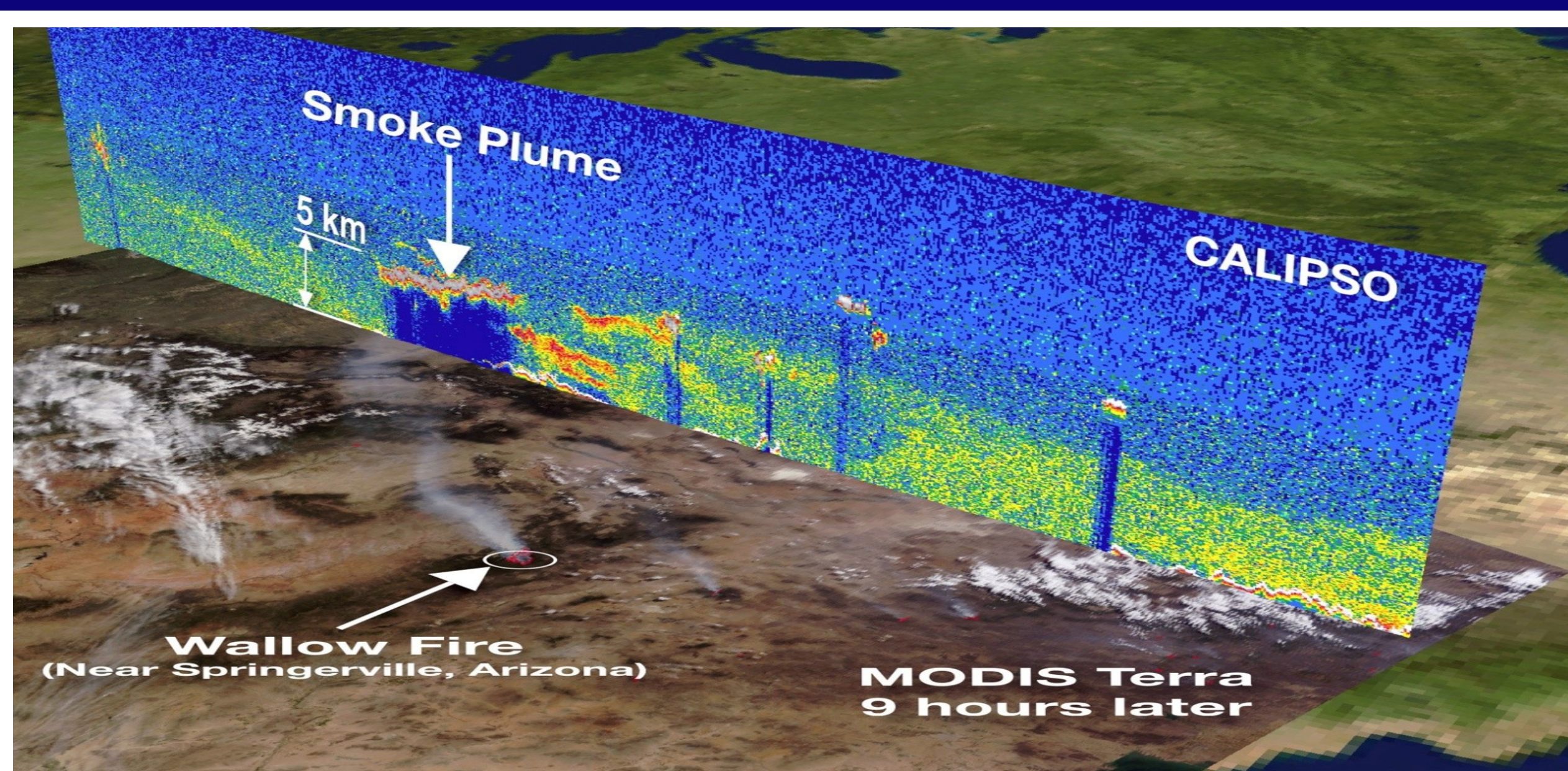


MODIS vs. MISR



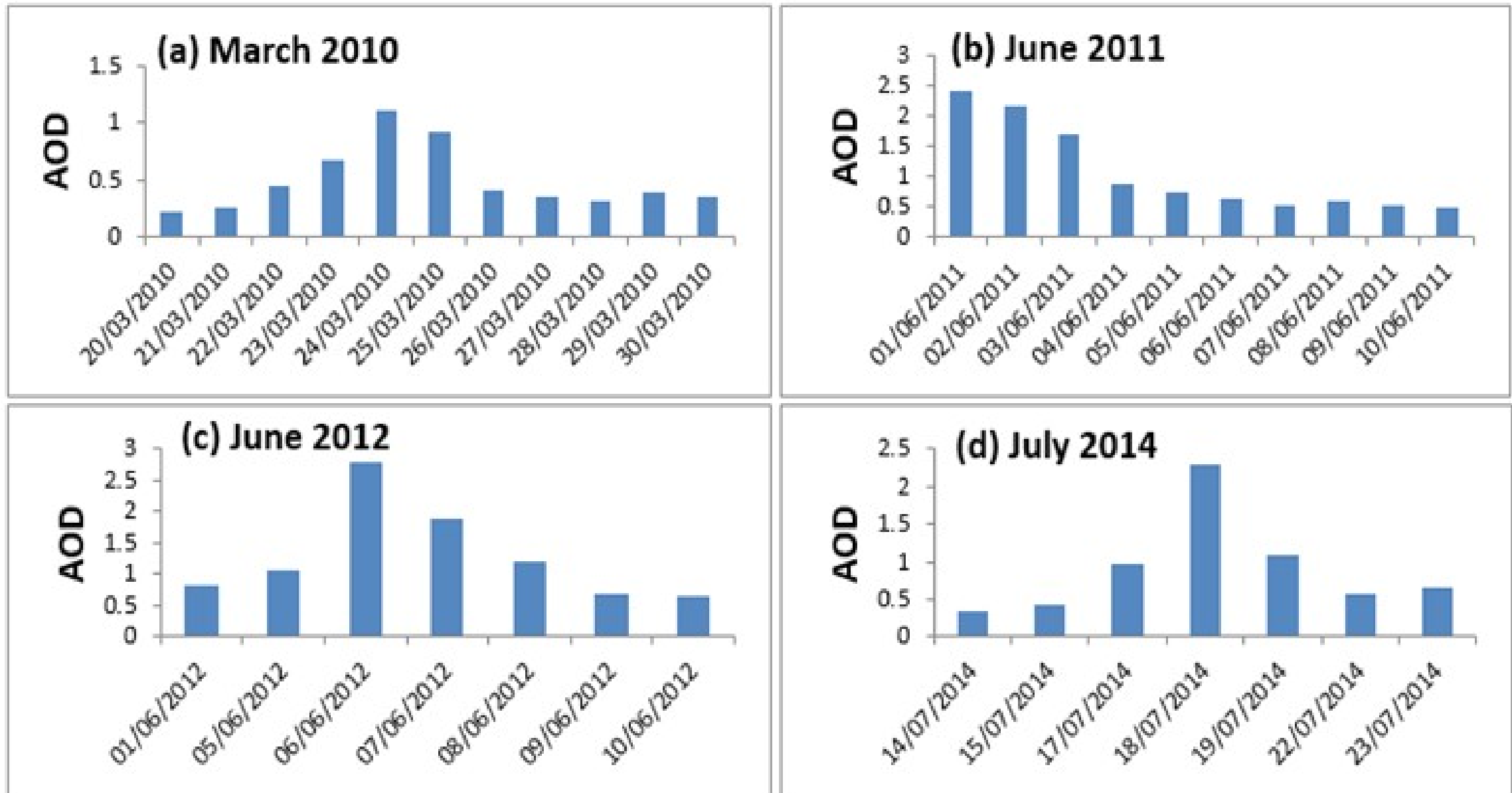
Alam et al., 2011

Satellite Data Validation

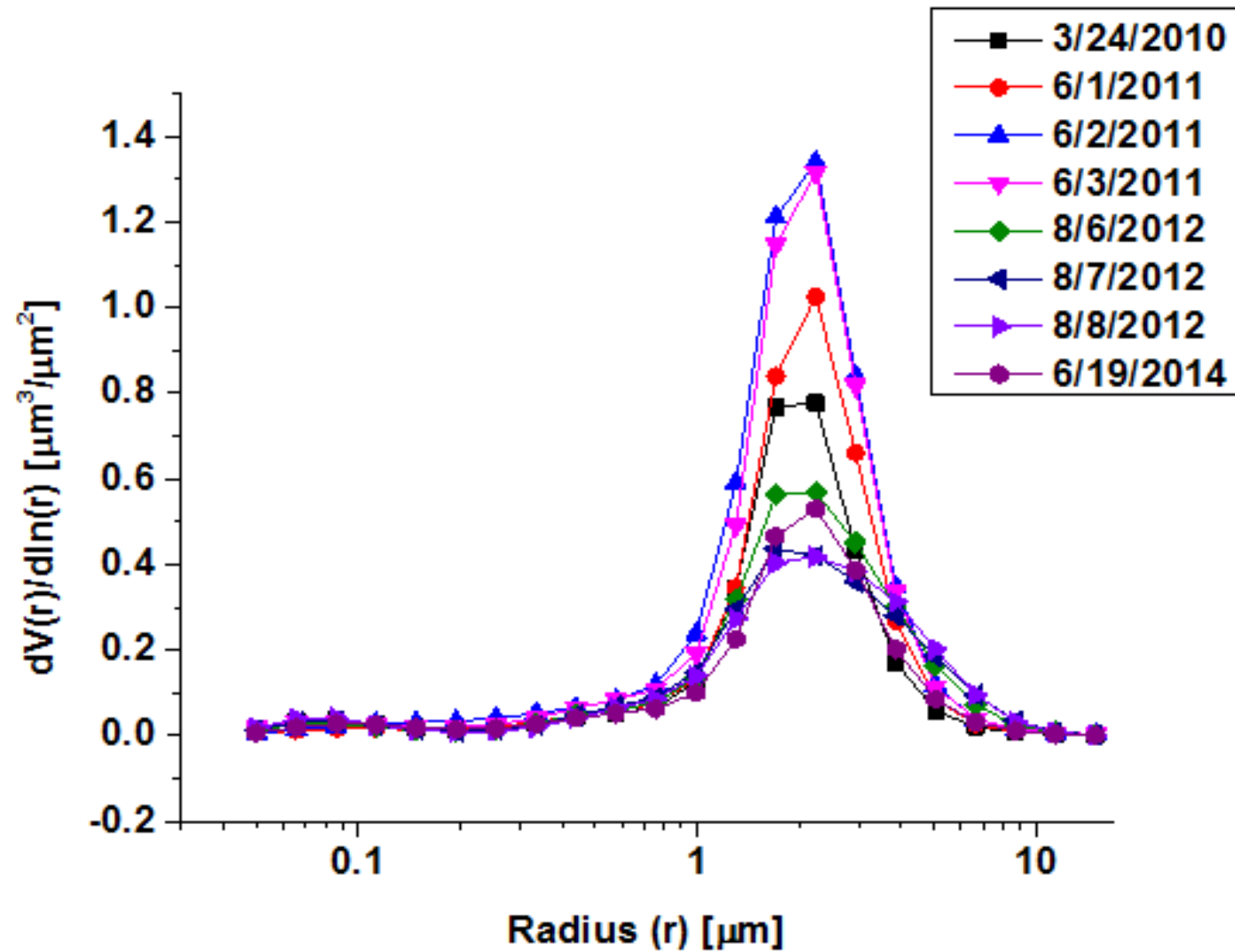


Dust Characteristics over Karachi

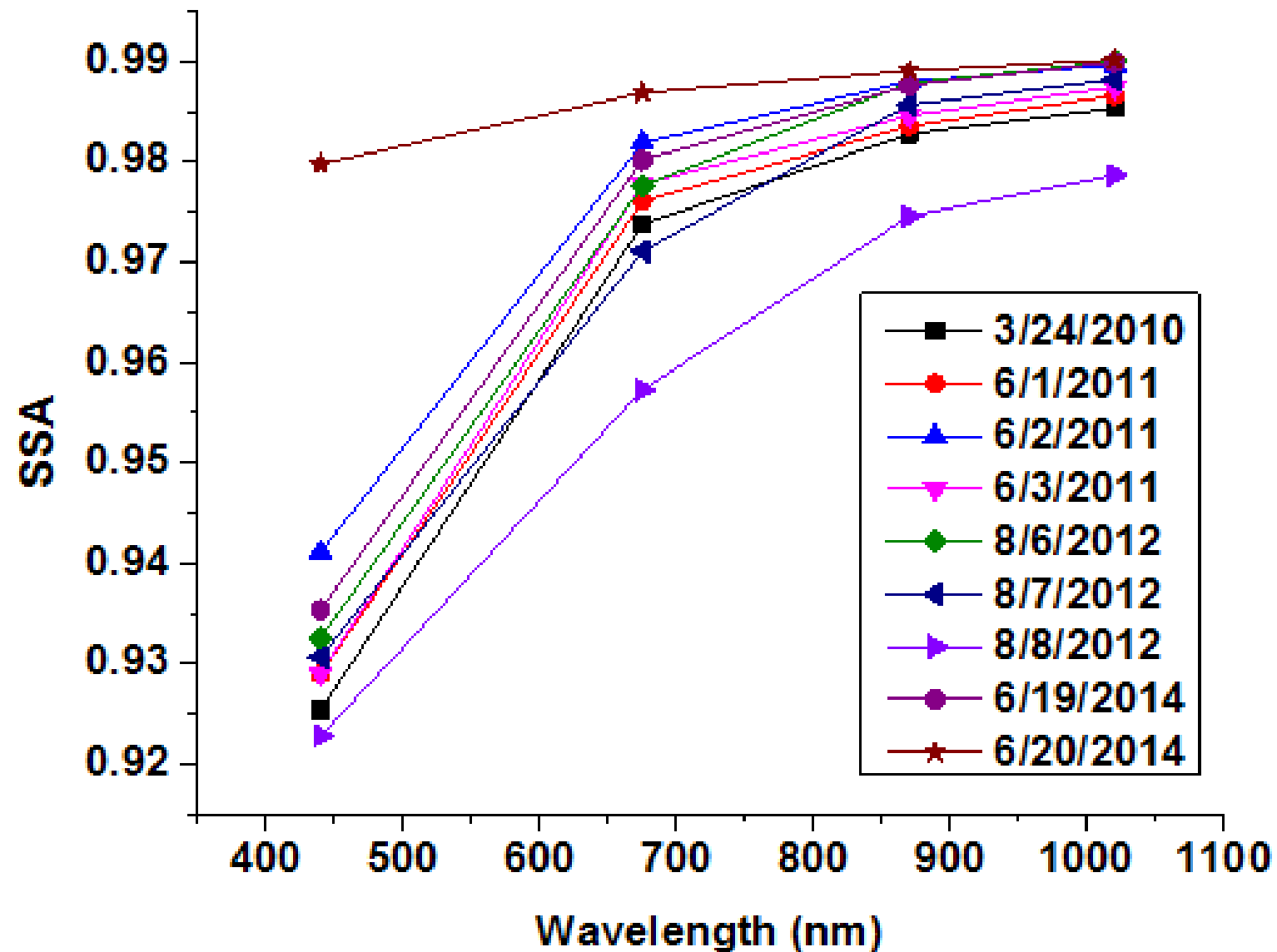
Variation in Aerosol Optical Depth during dusty and non dust days



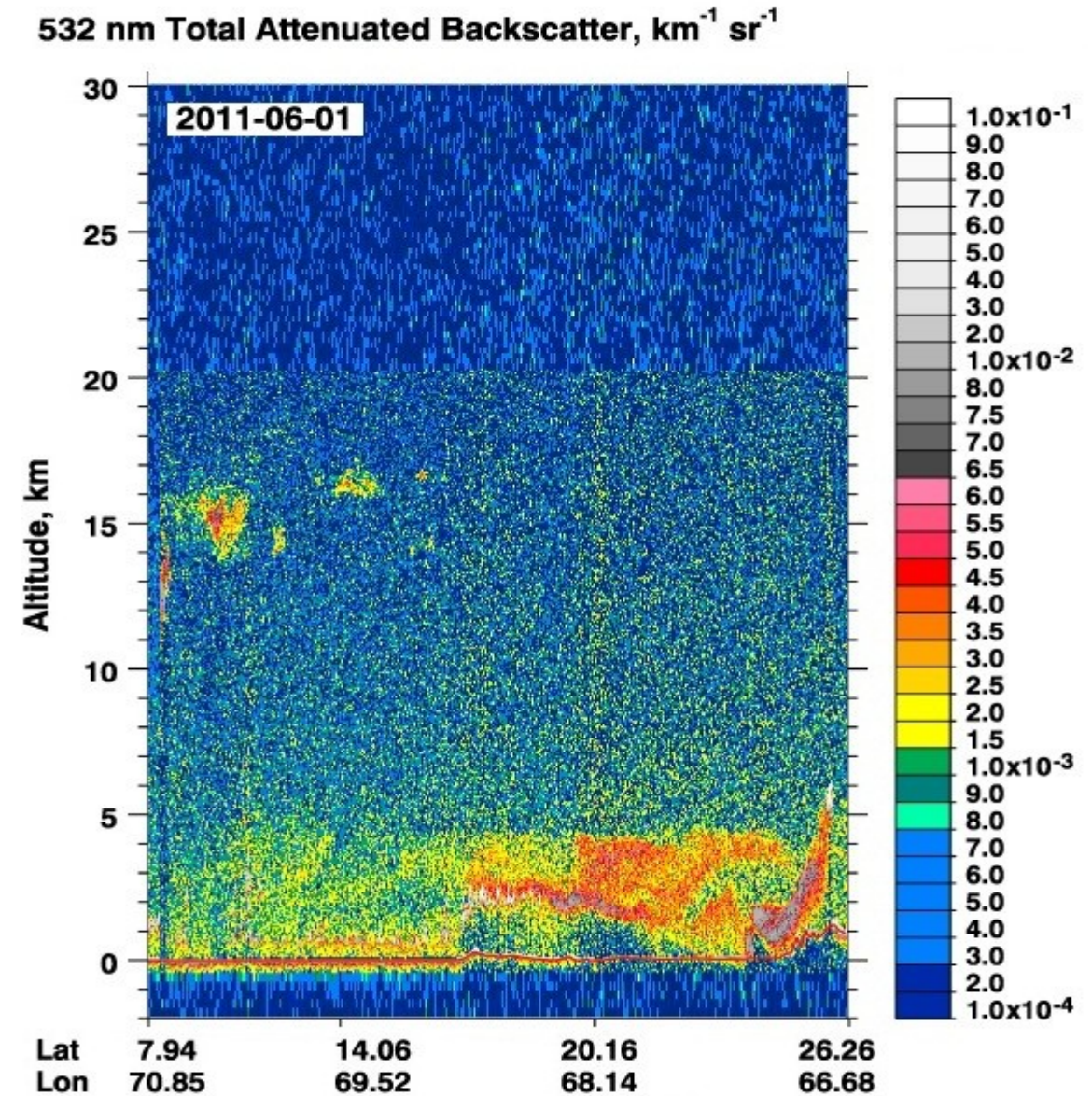
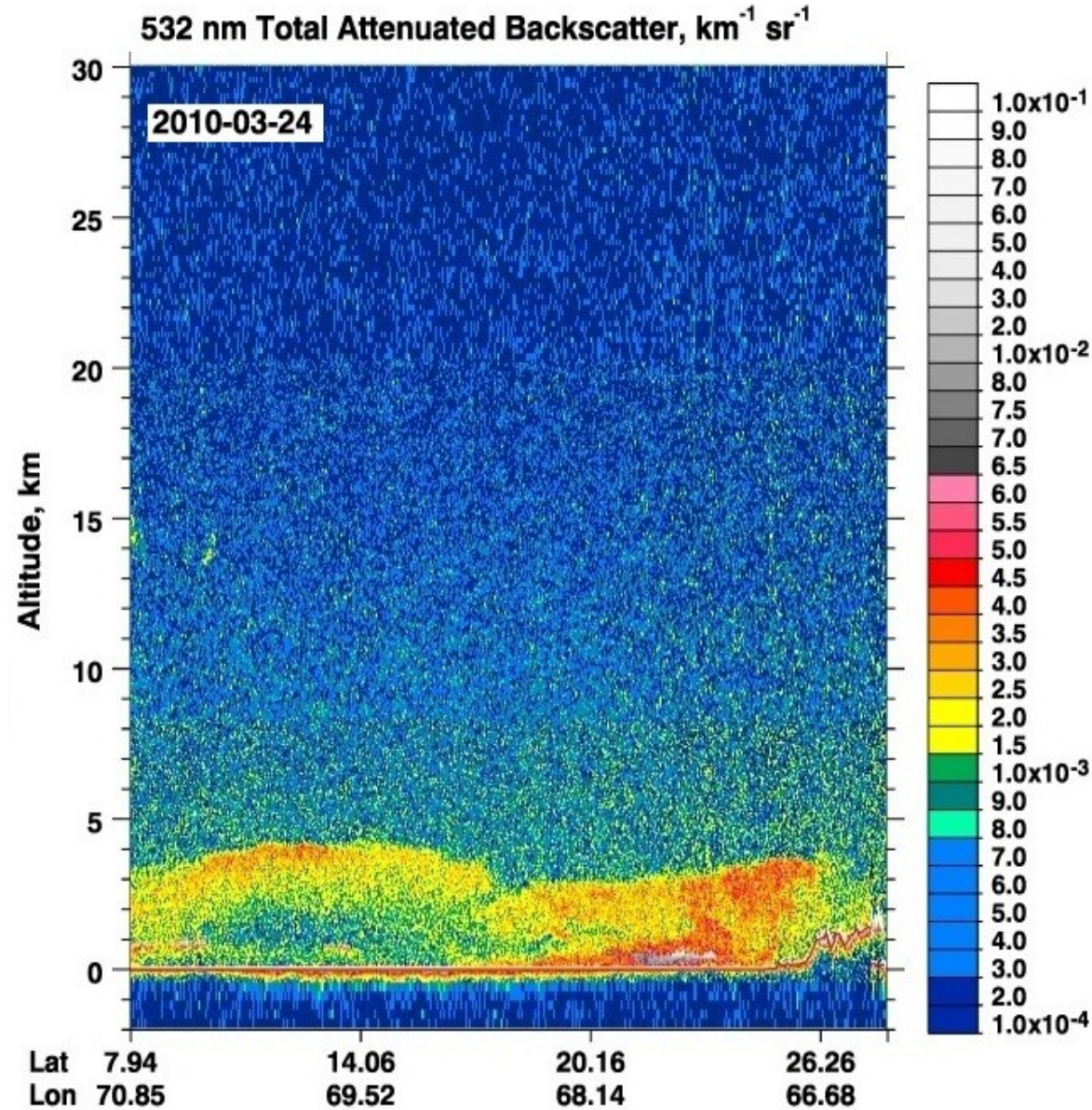
Volume Size Distribution



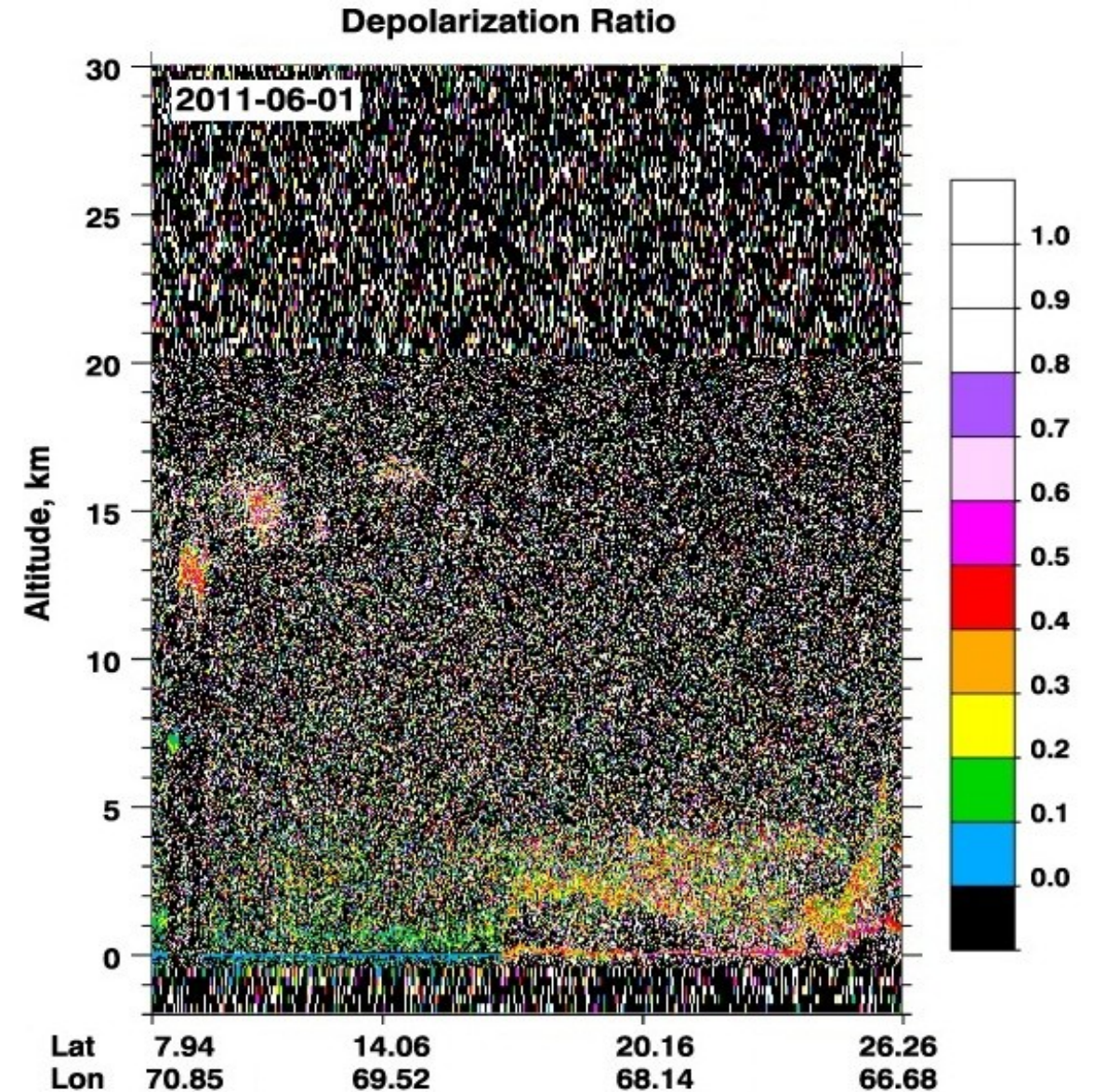
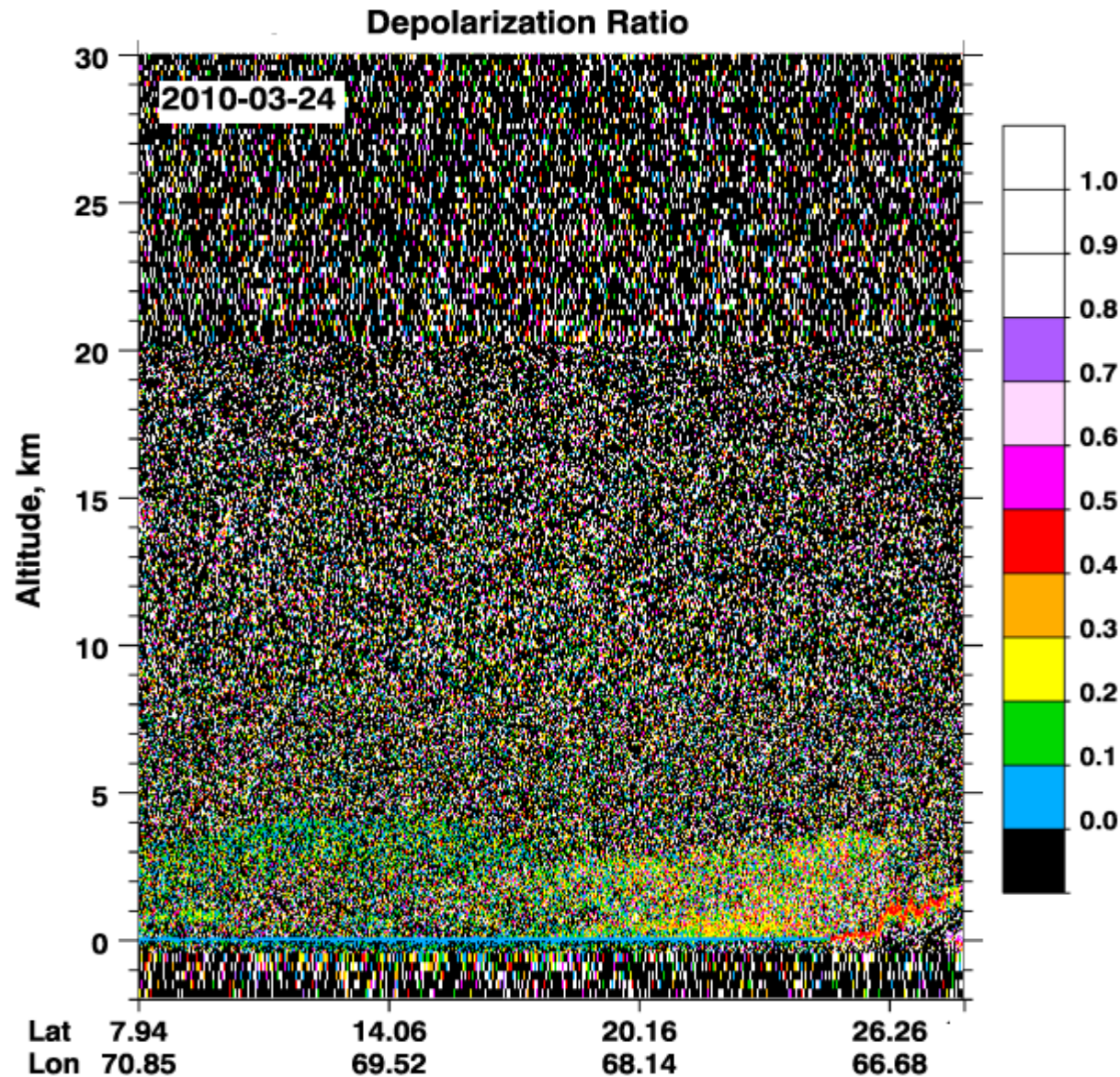
Single Scattering Albedo



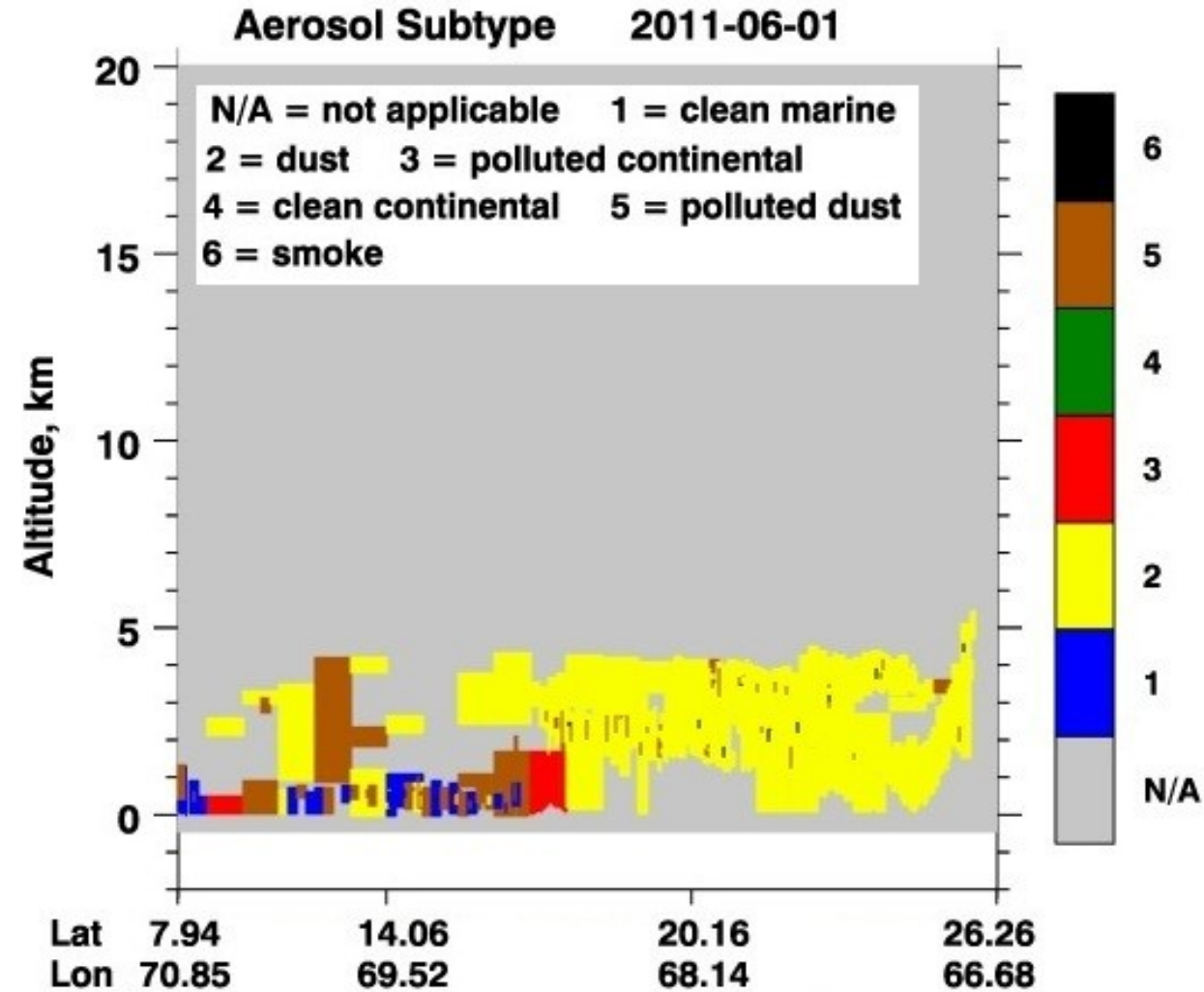
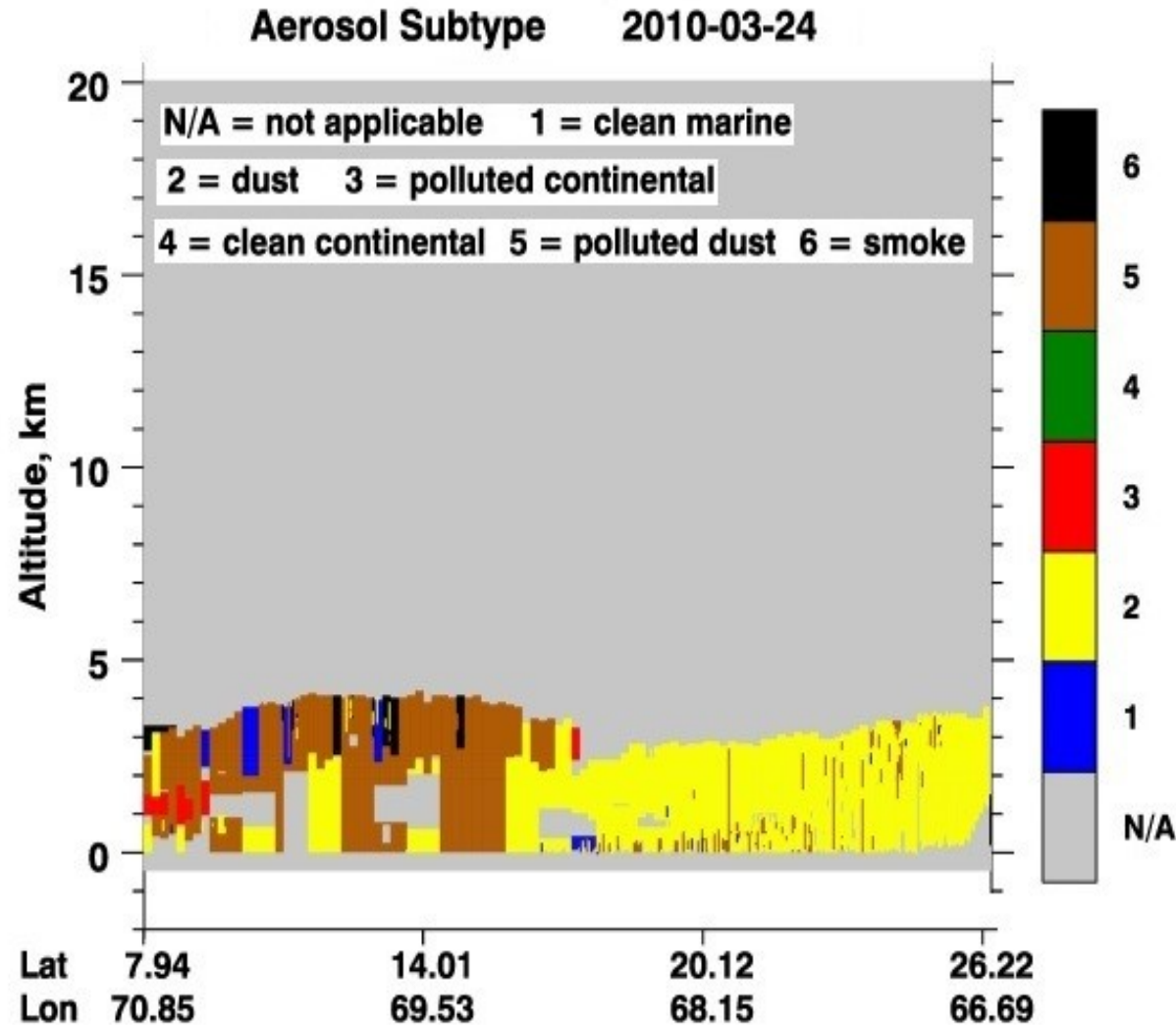
CALIPSO Total Attenuated Backscatter



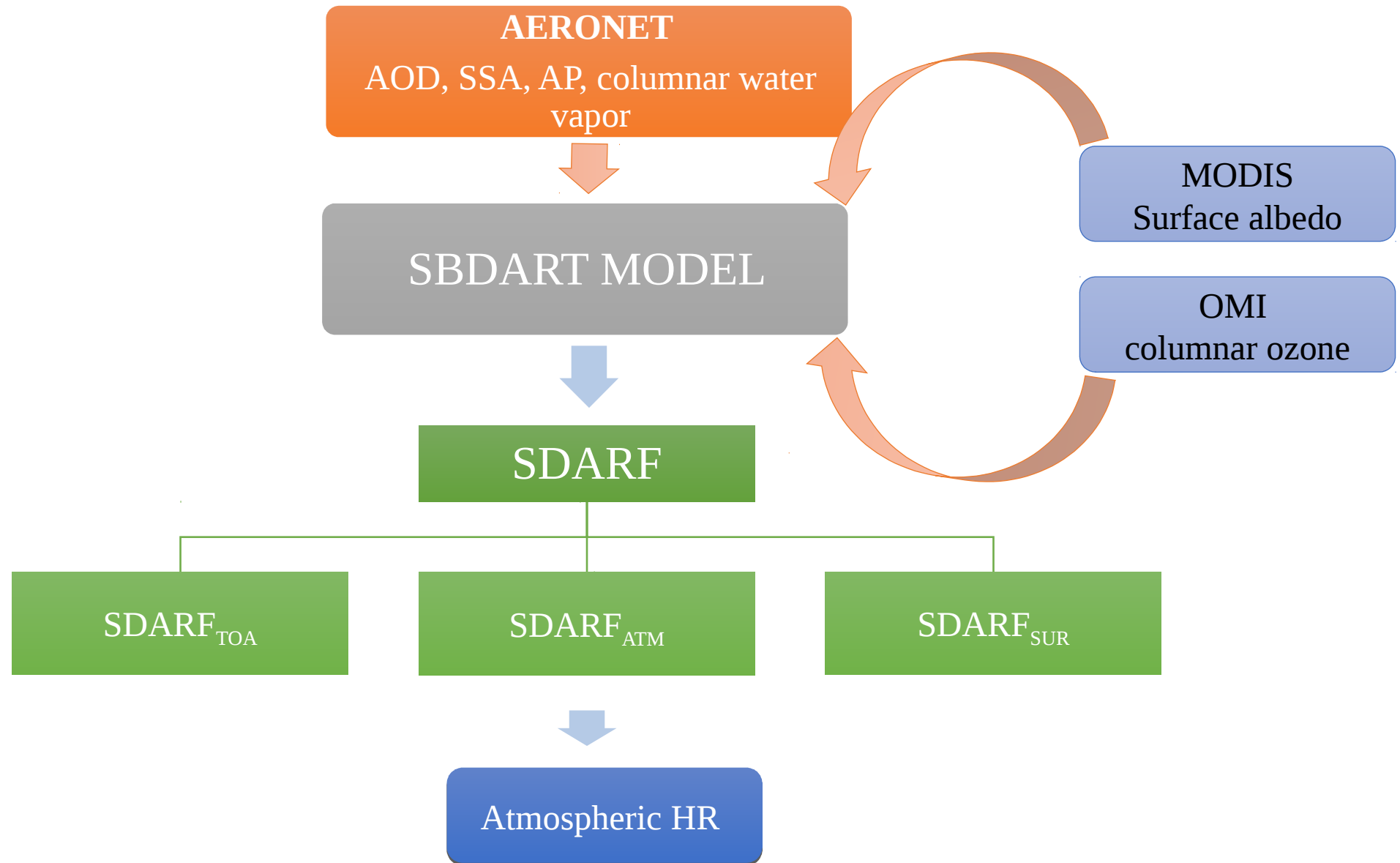
CALIPSO: Depolarization Ratio



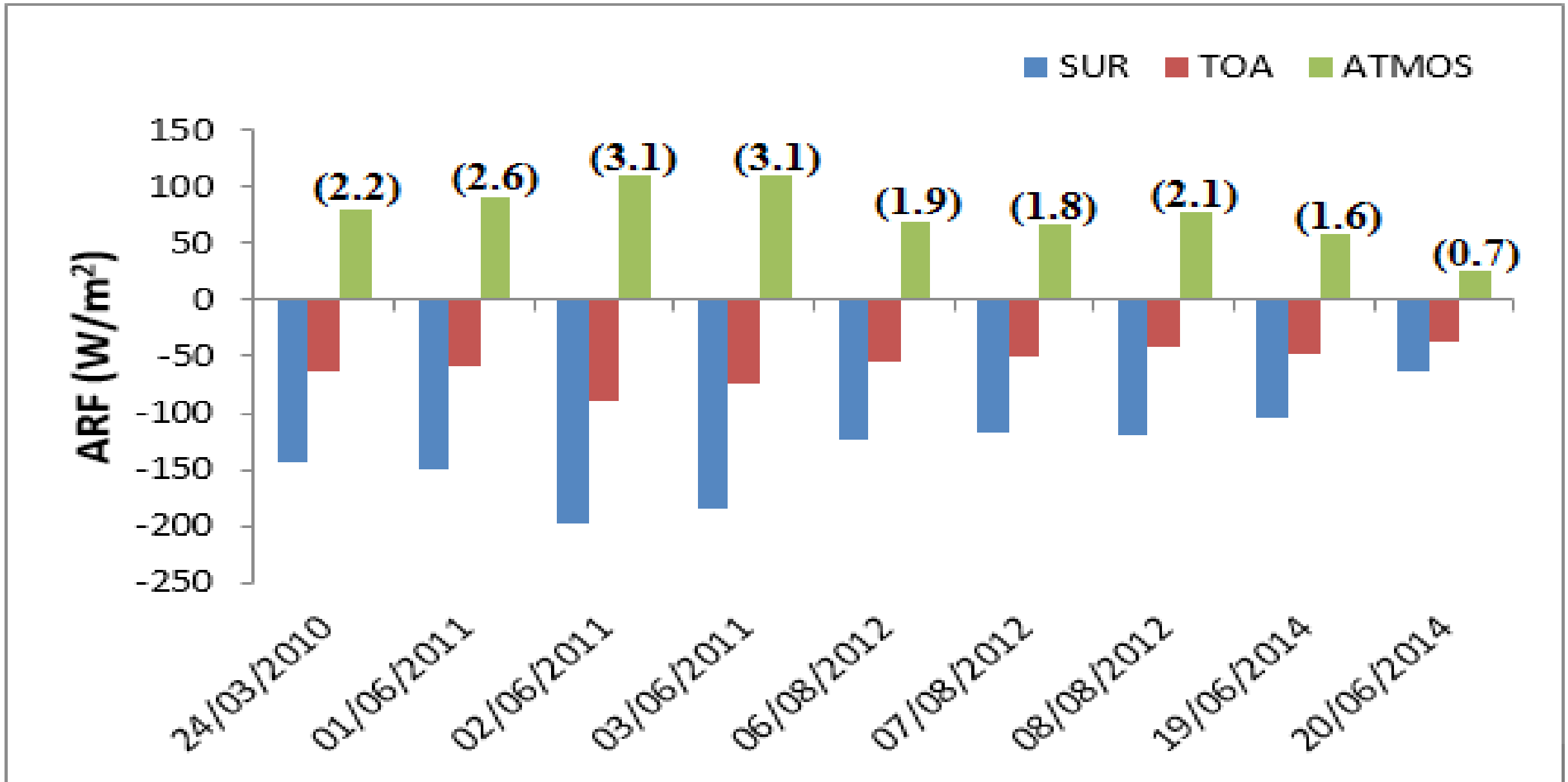
CALIPSO: Aerosol Classification



Schematic Diagram of Shortwave Radiative Forcing using SBDART Model



Aerosol Radiative Forcing and Heating Rates



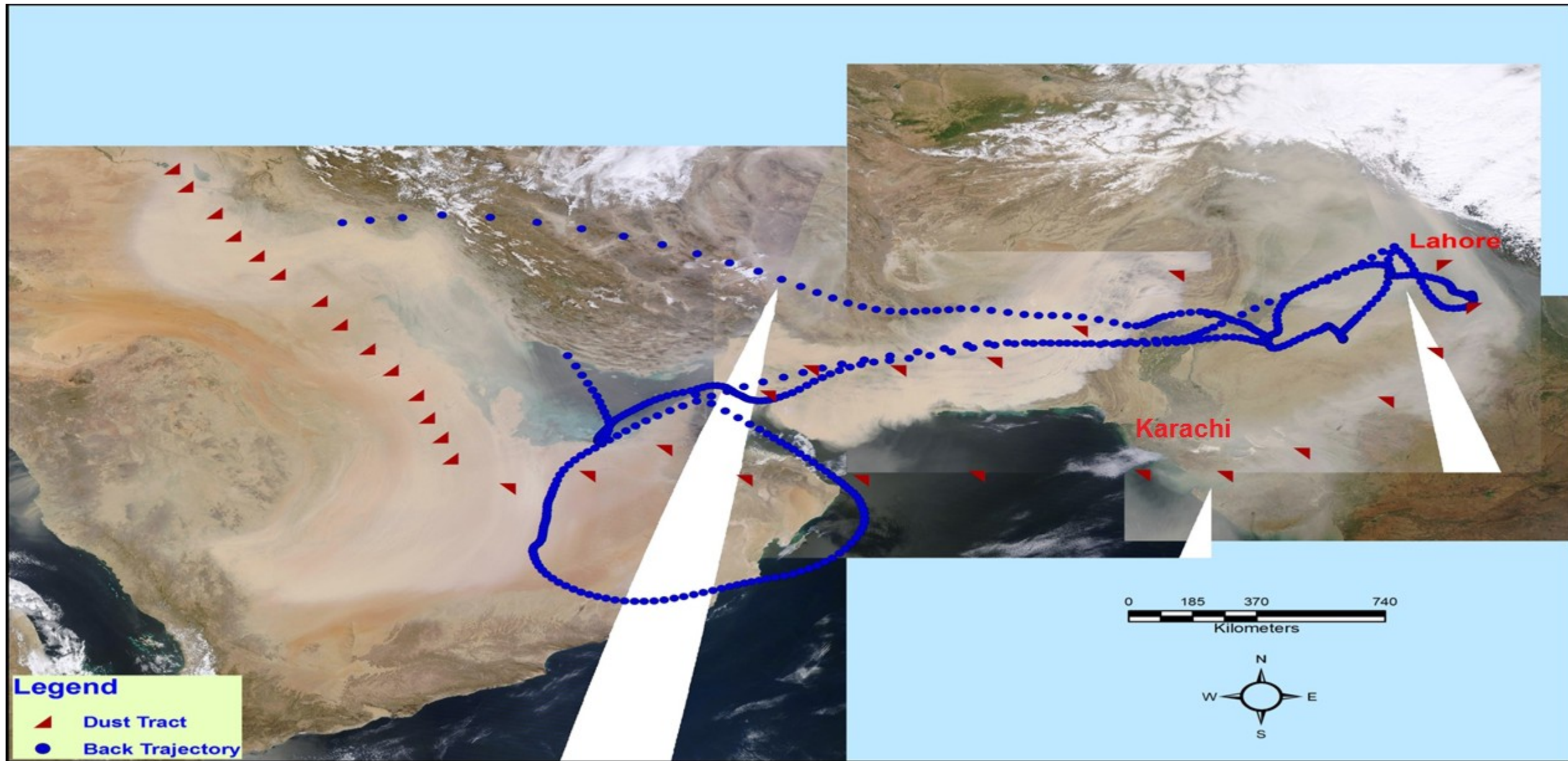
Dust Characteristics over the Middle East & Southwest Asia

Alam, K., Trautmann, T., Blaschke, T., Subhan, F. (2014). Changes in aerosol optical properties due to dust storm in the Middle East and Southwestern Asia. Remote Sensing of the Environment, 143, 216-227

❖ Dust in Middle East and Southwest Asia

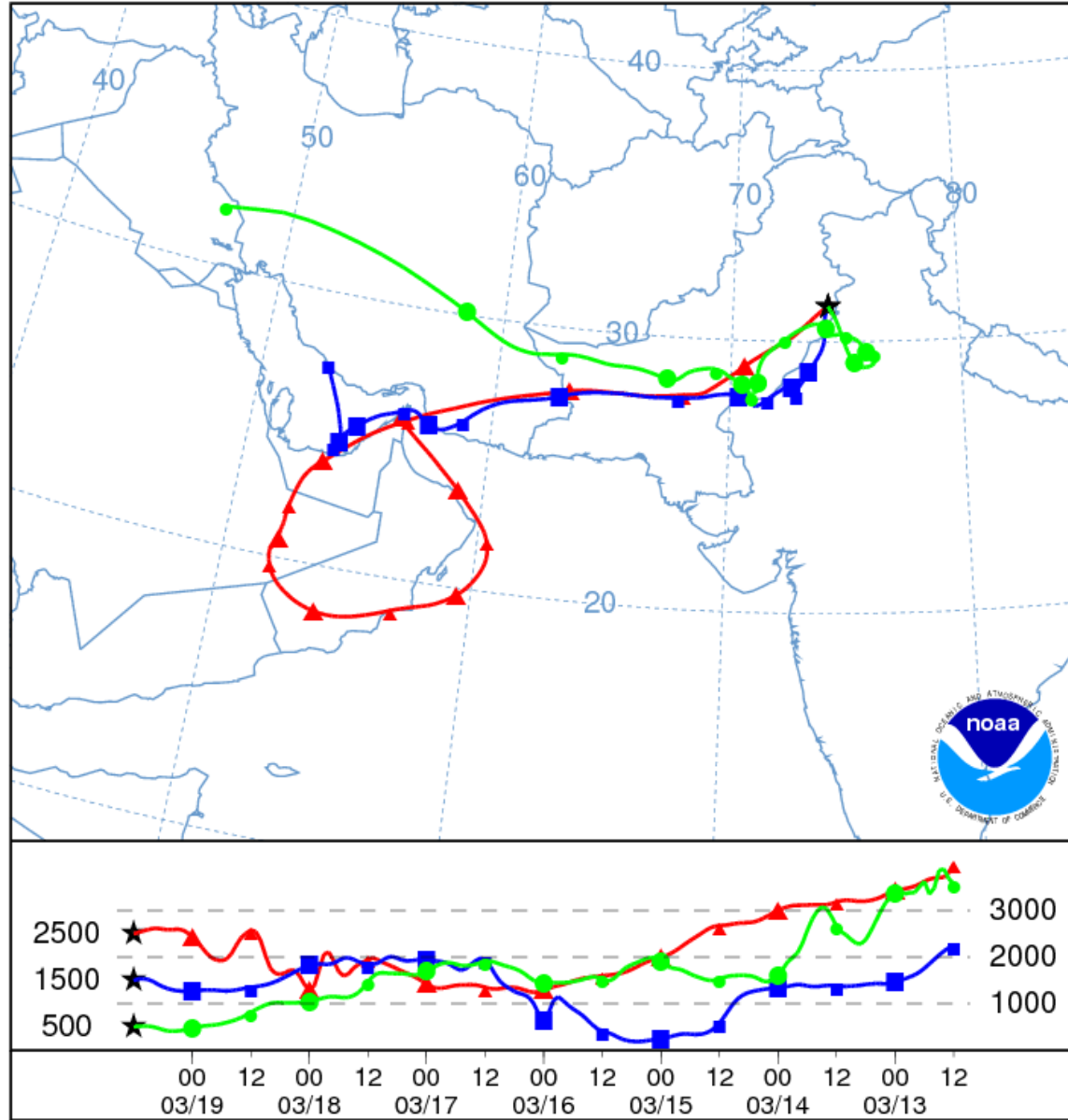


Formation of Dust events and Meteorological situation

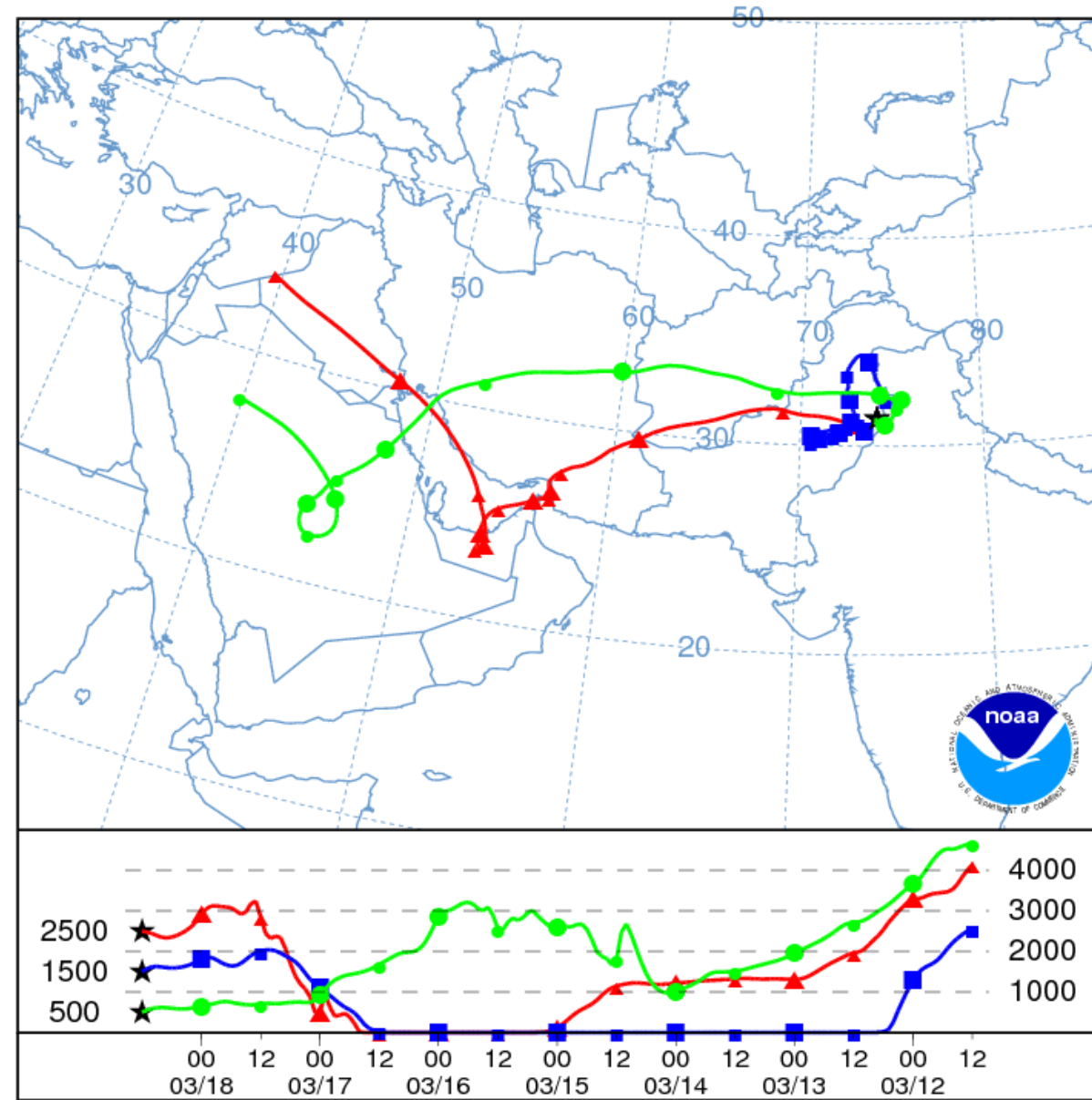


Aerosol Transport through HYSPLIT model

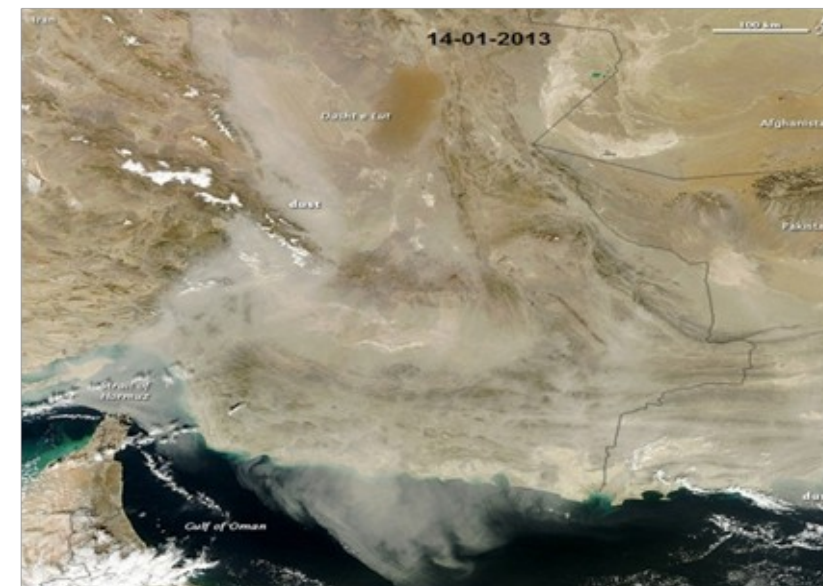
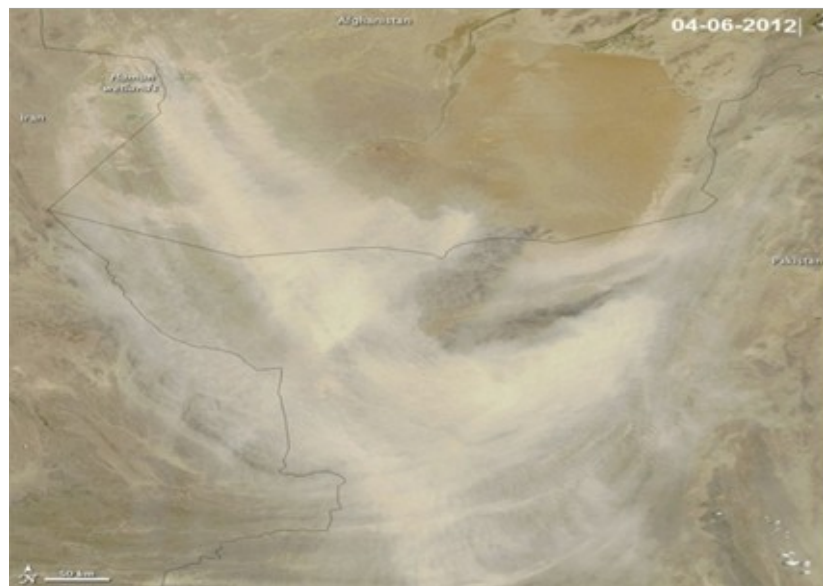
Source ★ at 31.30 N 74.20 E



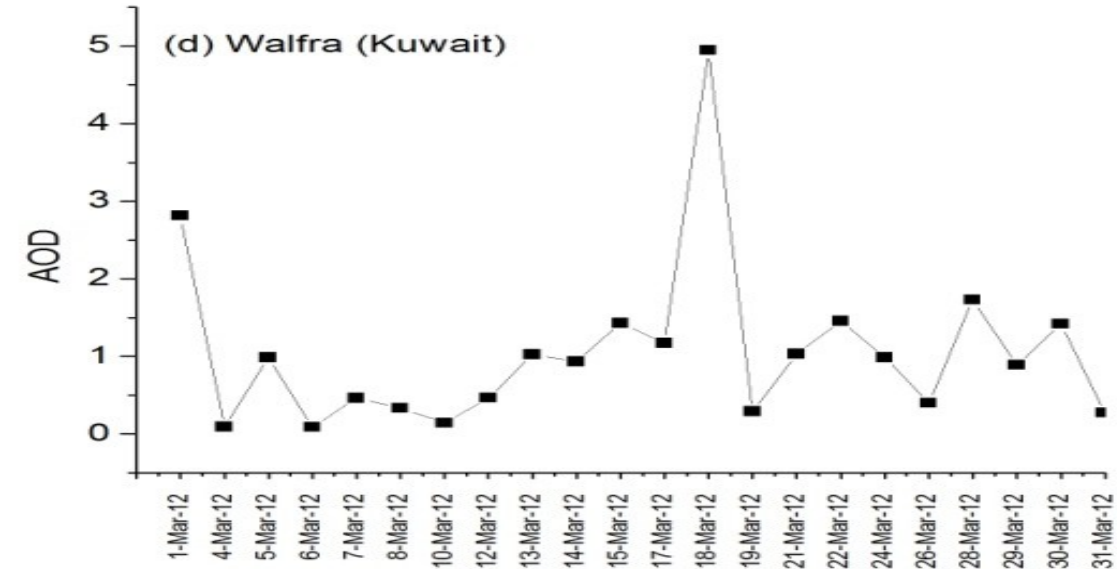
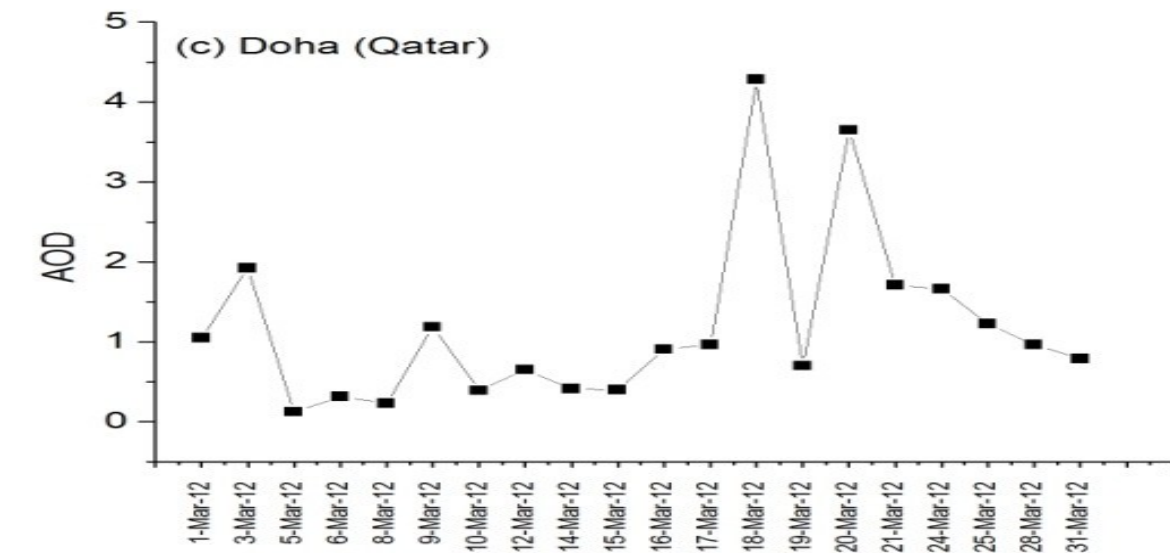
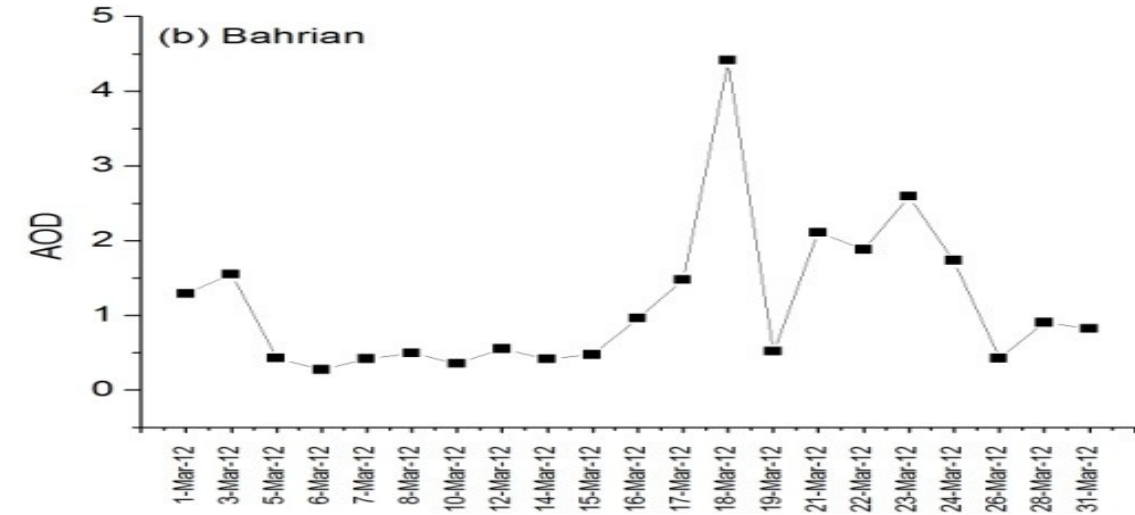
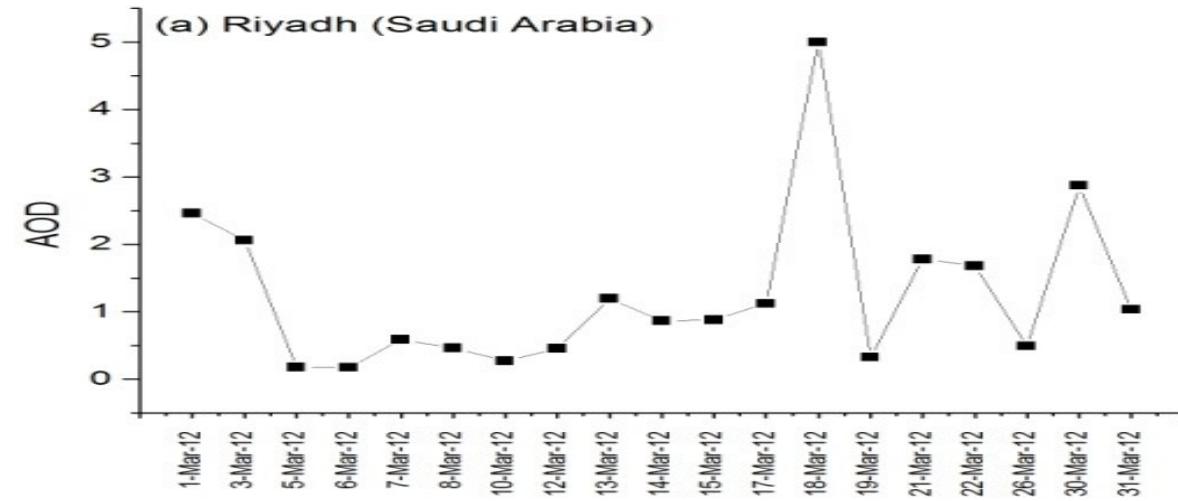
Source ★ at 31.30 N 74.20 E



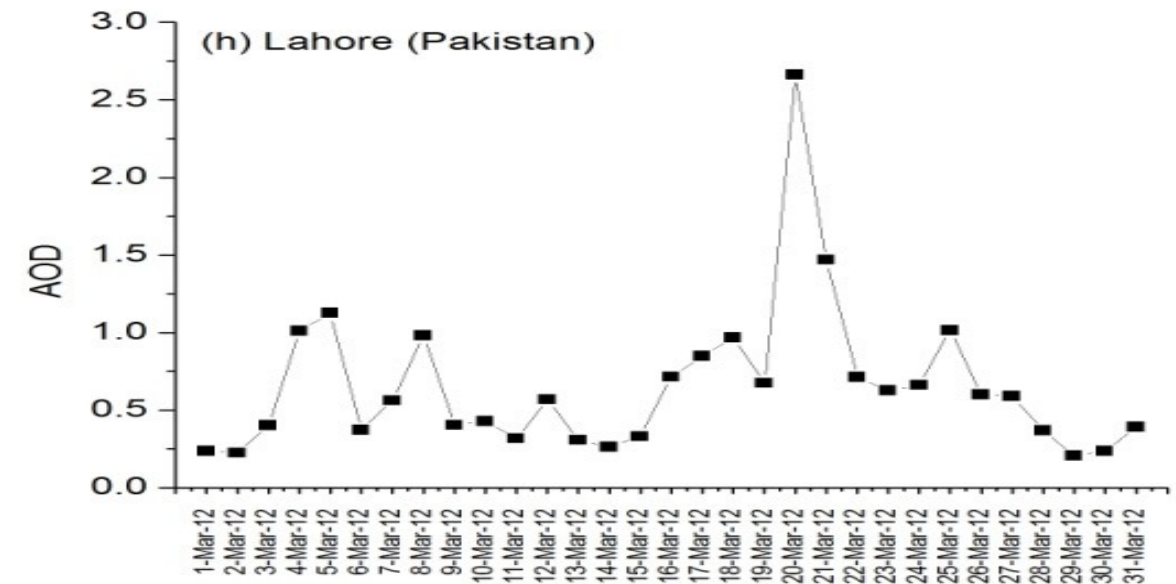
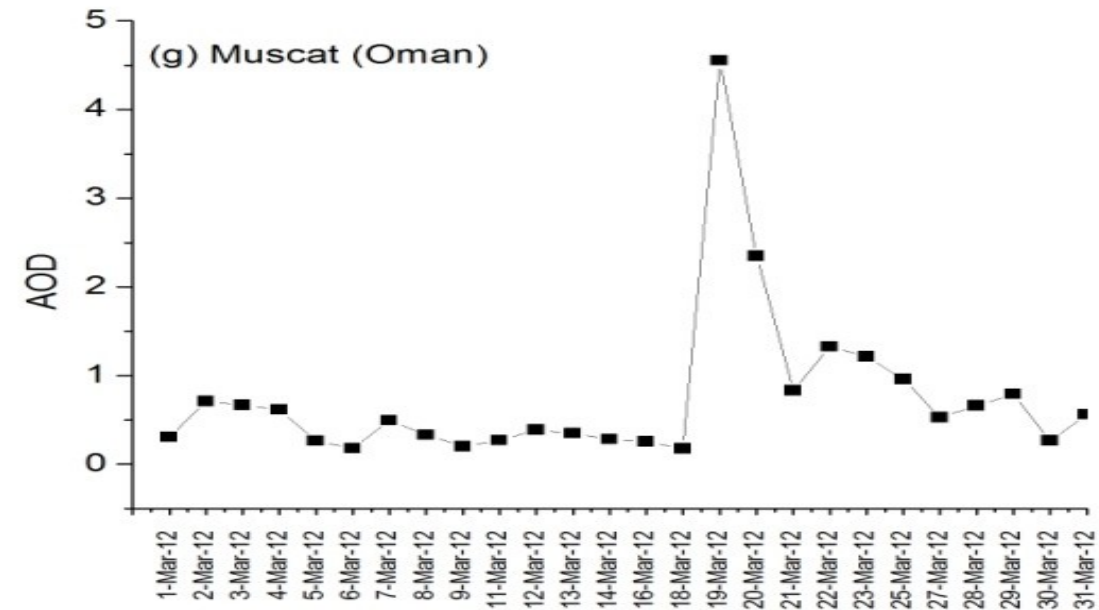
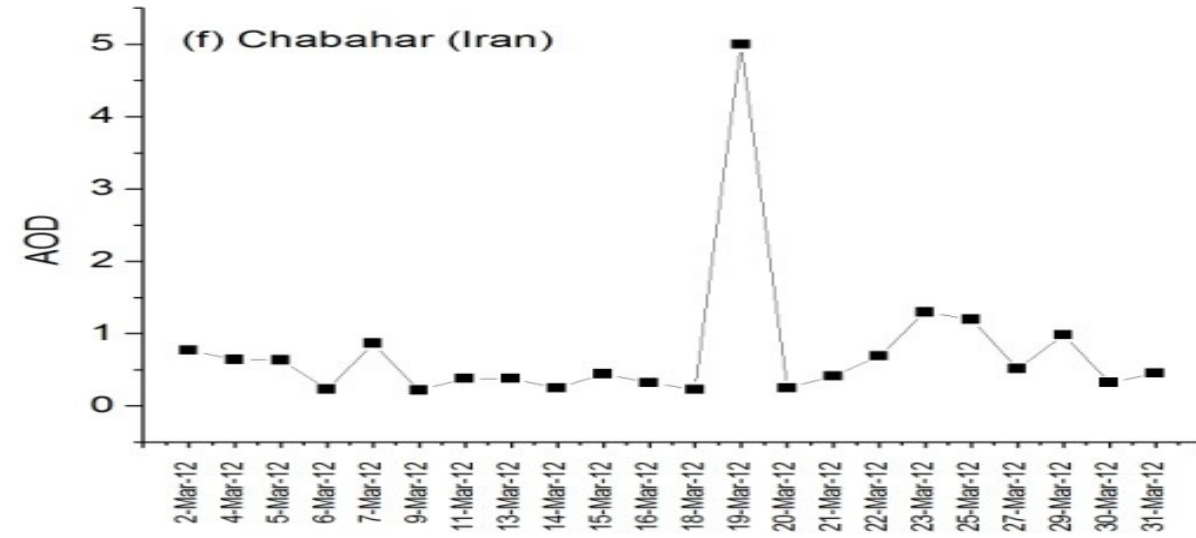
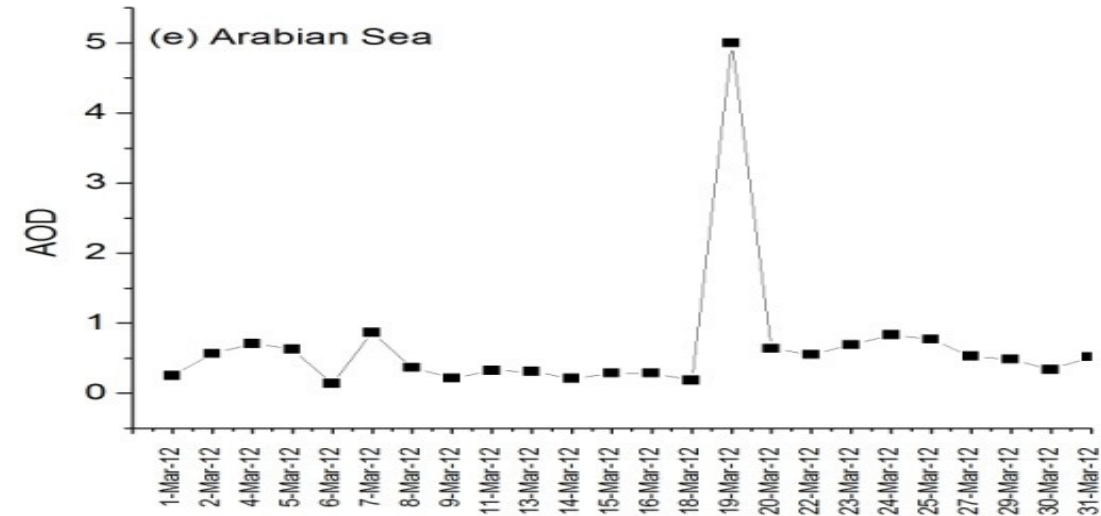
Dust detection through Satellite



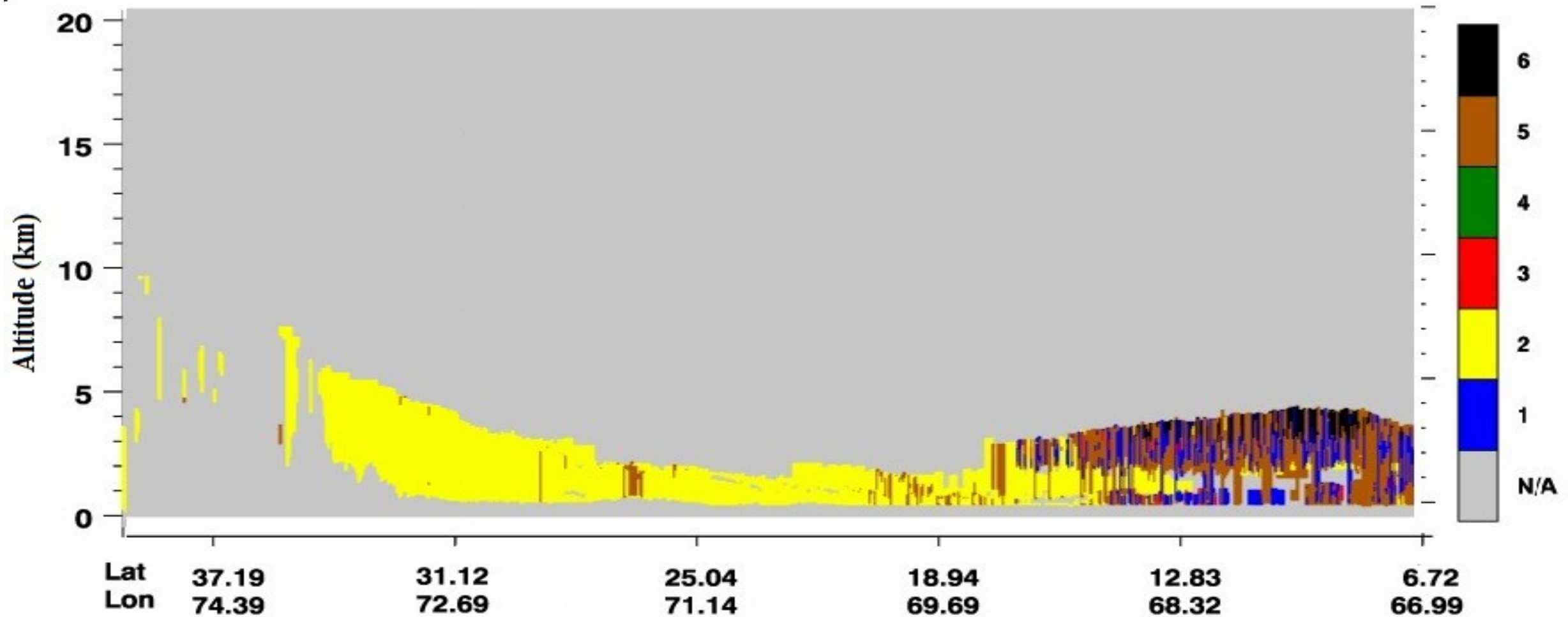
❖ Results & Discussion: Aerosol Optical Depth Variations



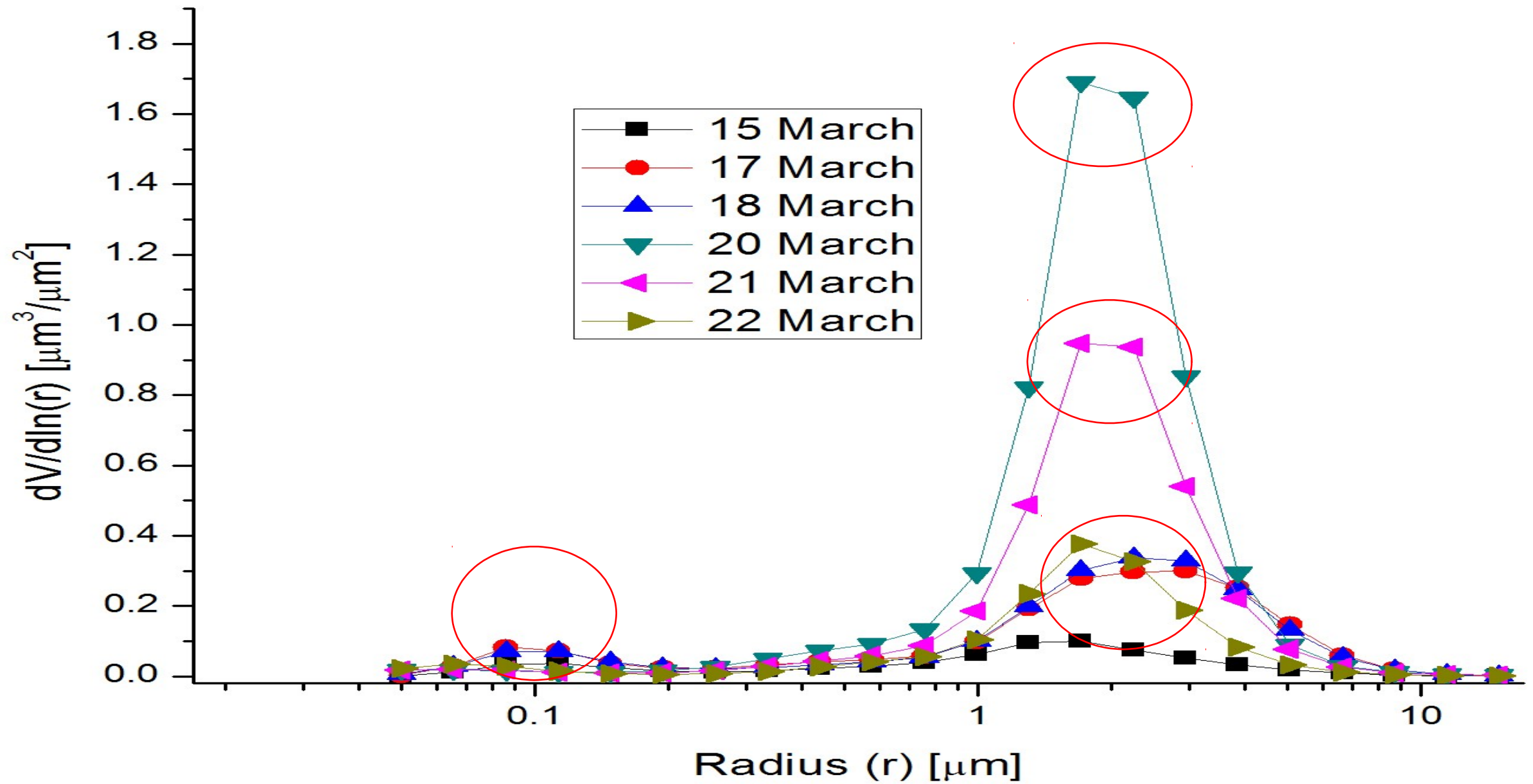
❖ Results & Discussion: Aerosol Optical depth variations

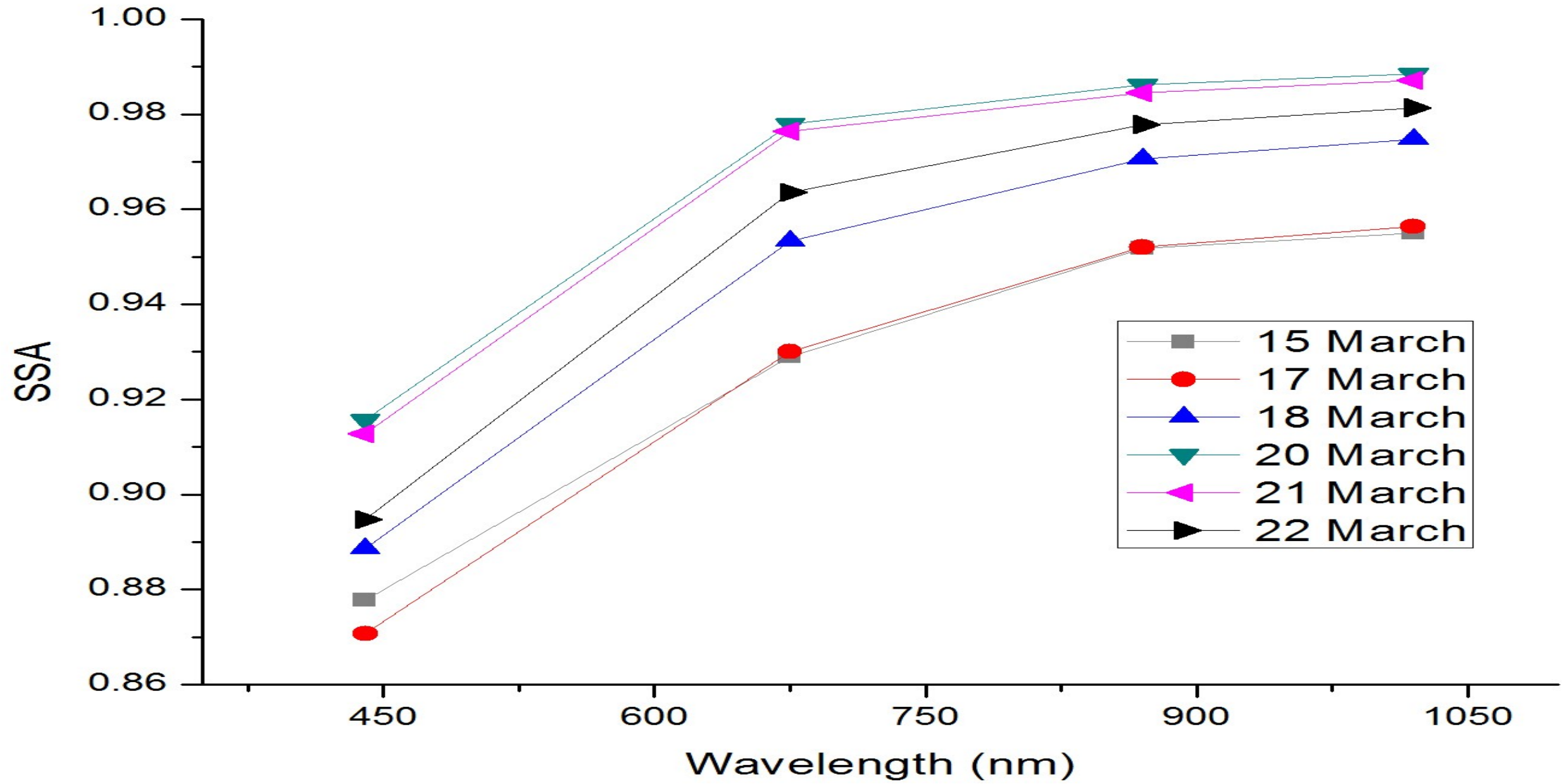


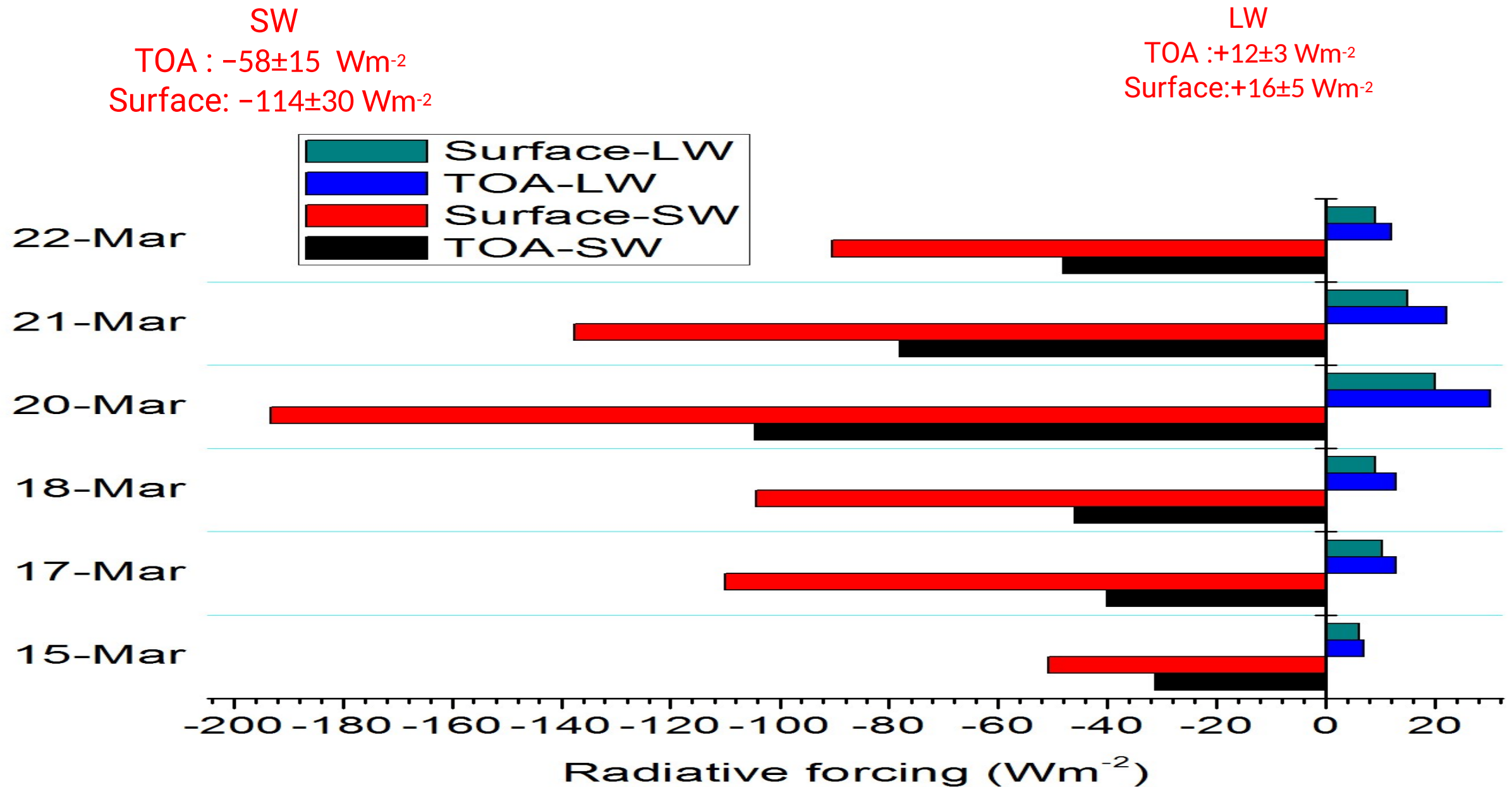
Aerosol sub-type classification



N/A = not applicable; 1= clean marine; 2 = dust; 3 = polluted continental; 4 = clean continental; 5 = polluted dust; 6 = smoke





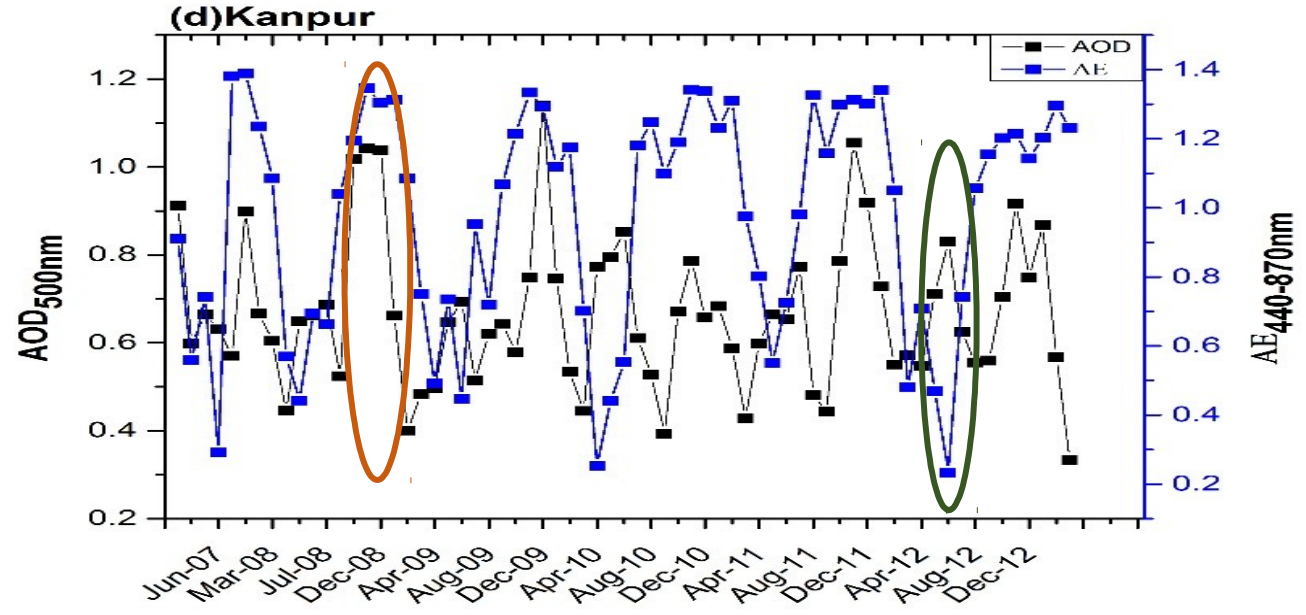
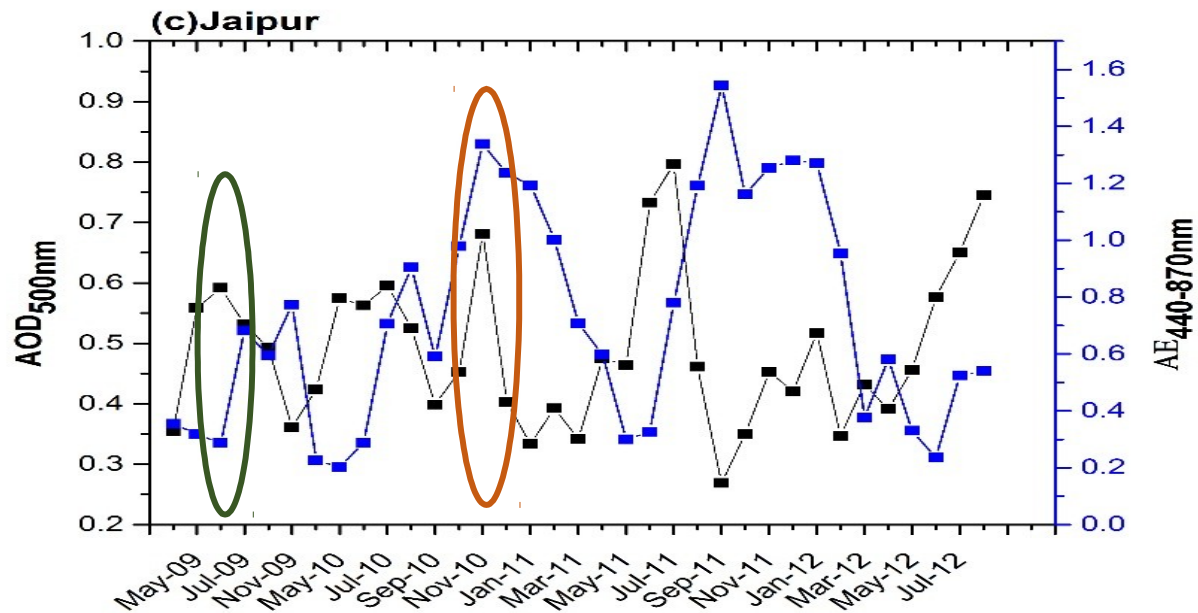
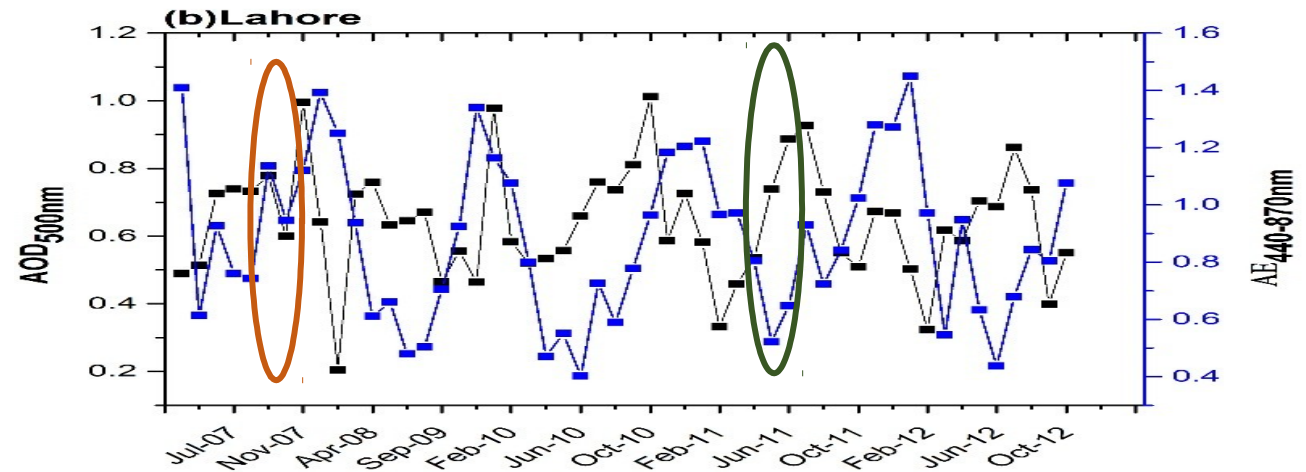
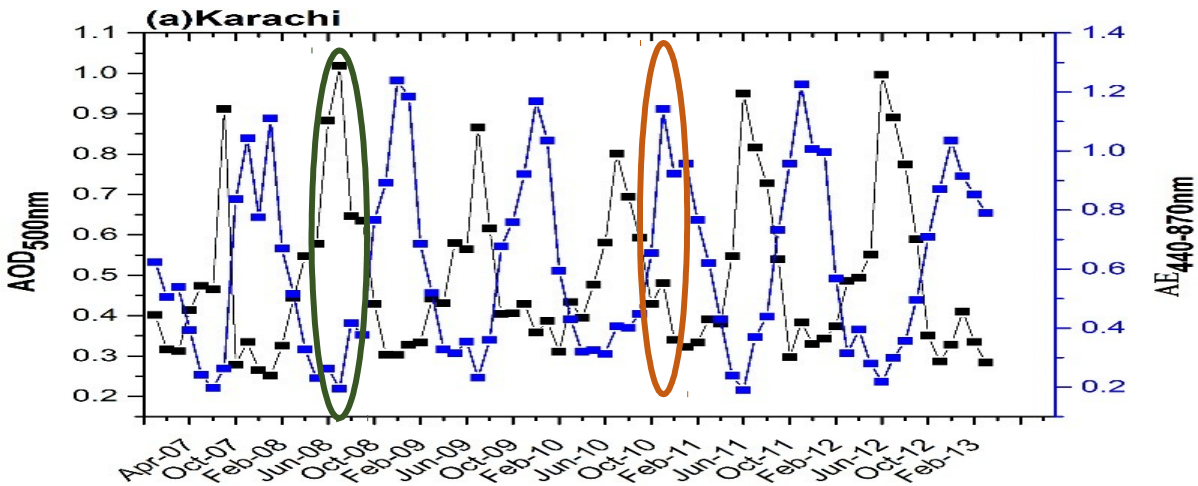


Classification of Aerosols

Monthly Averaged Variability of AOD and AE

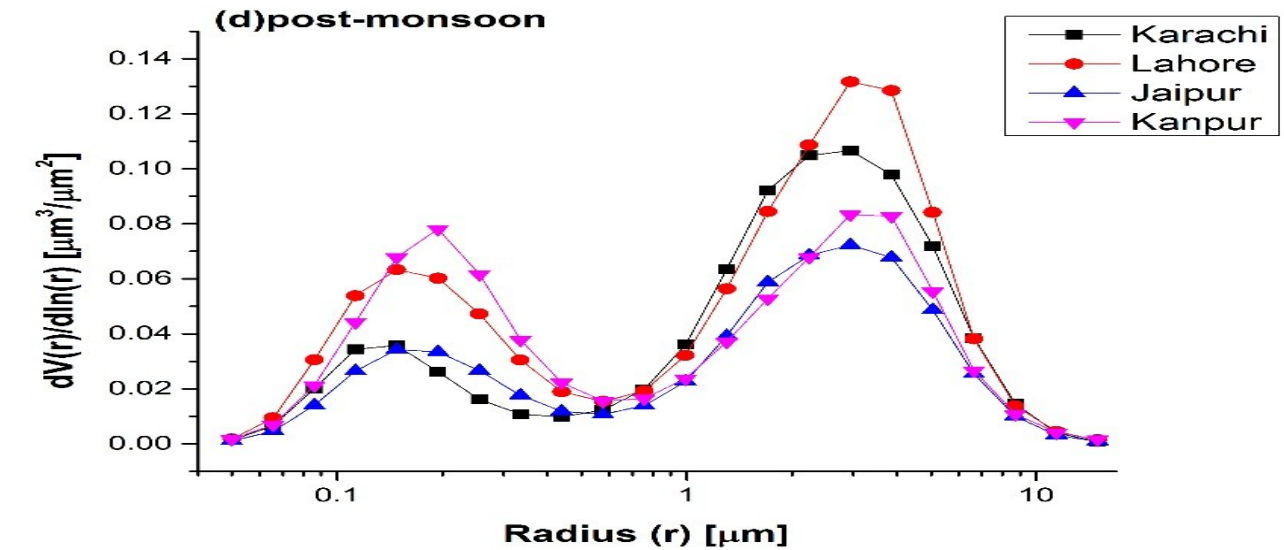
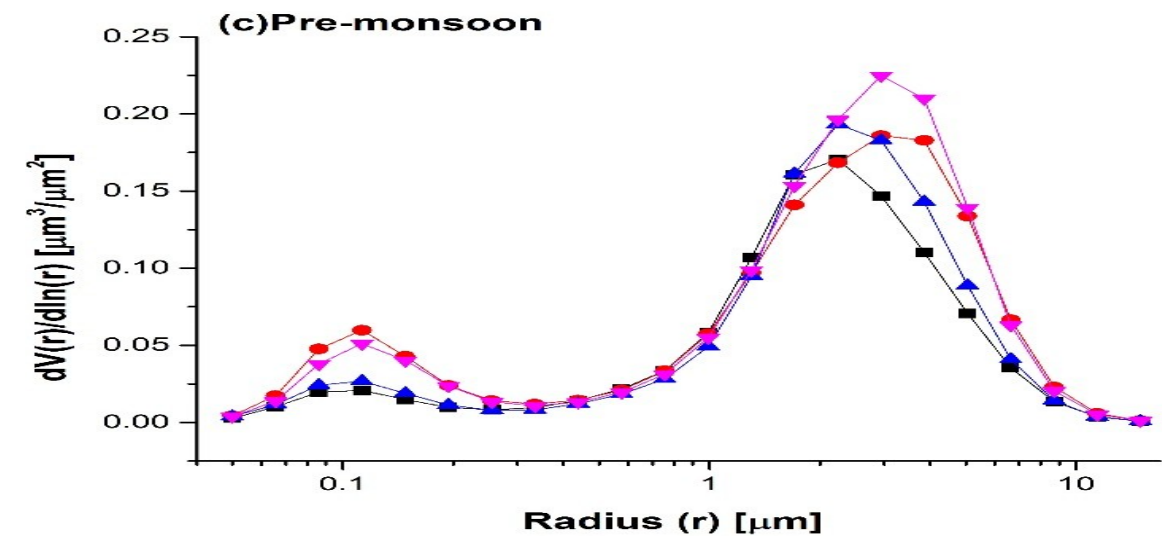
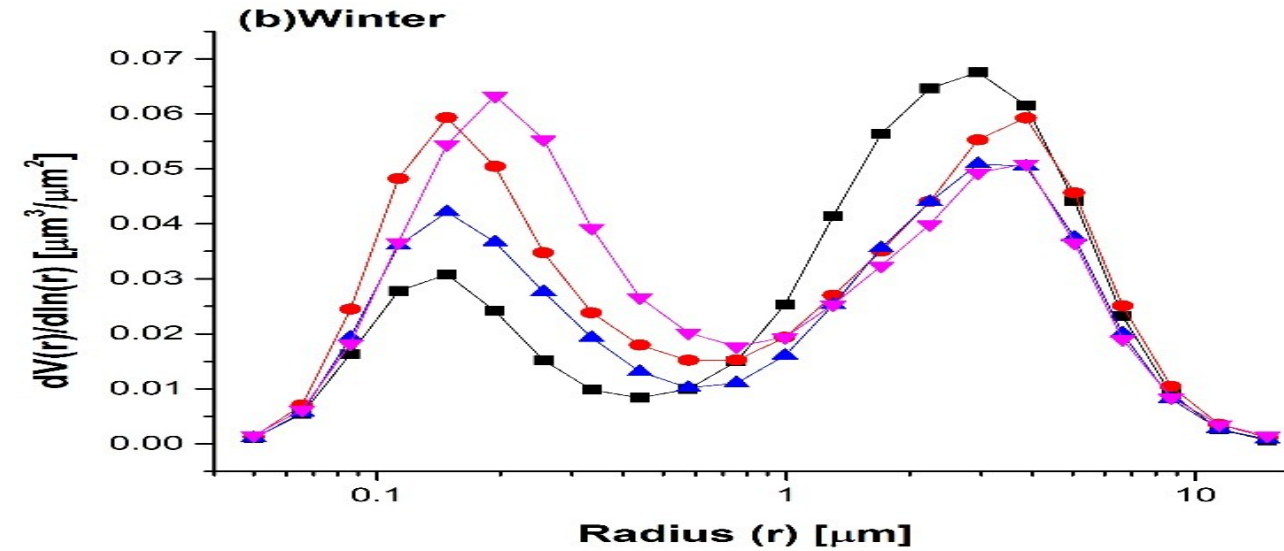
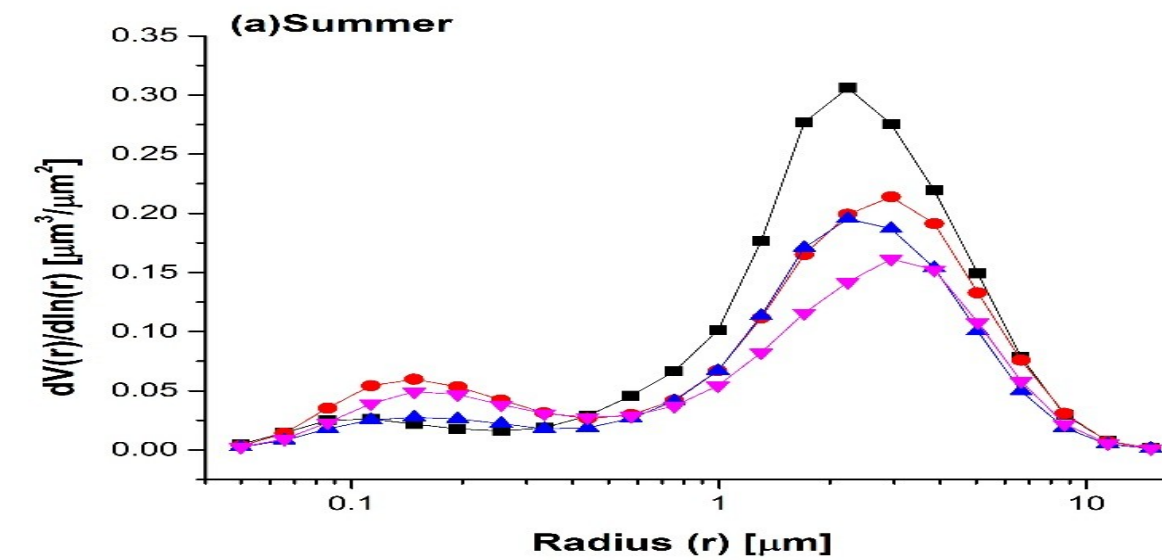
High AOD with low AE during summer and pre-monsoon months

High AOD with high AE during winter and post-monsoon



Seasonal Averaged Variability of AVSD

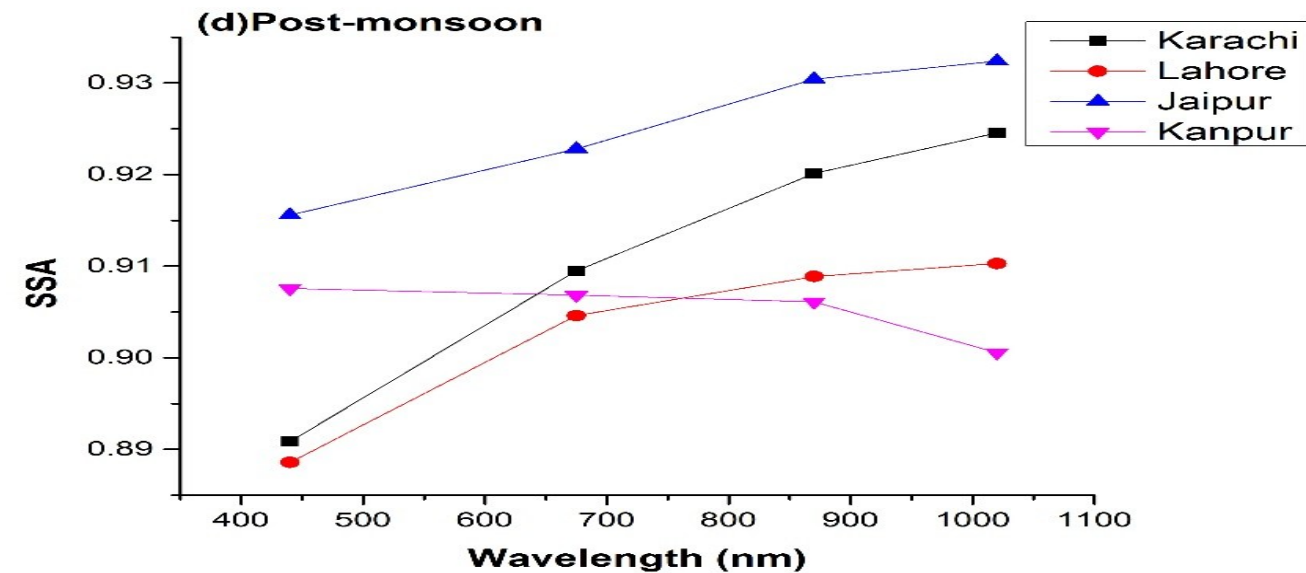
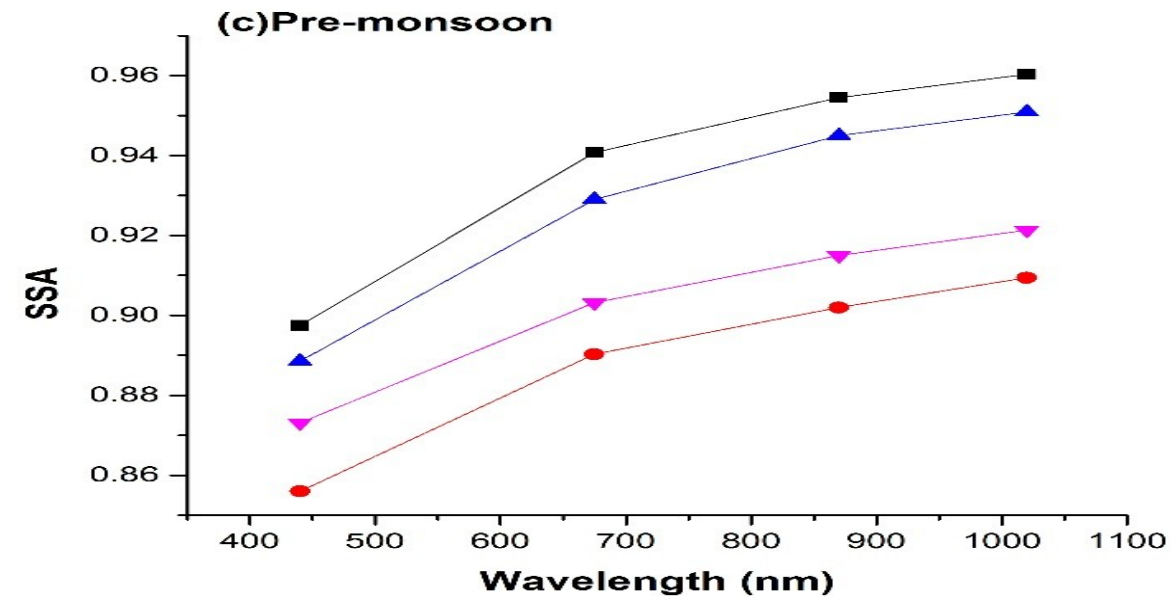
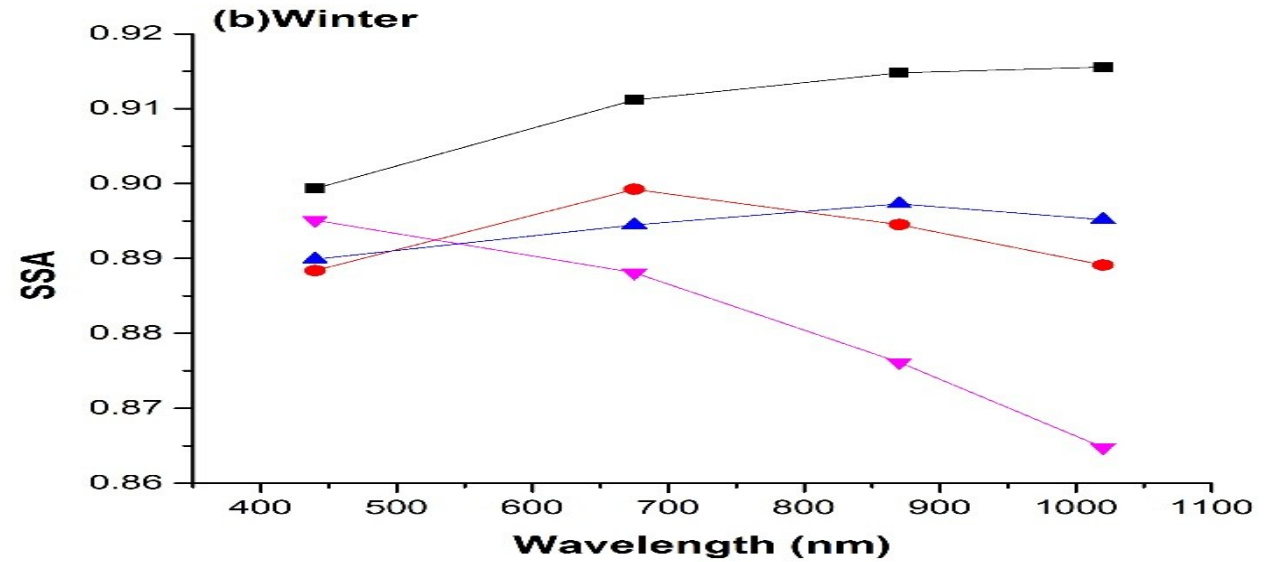
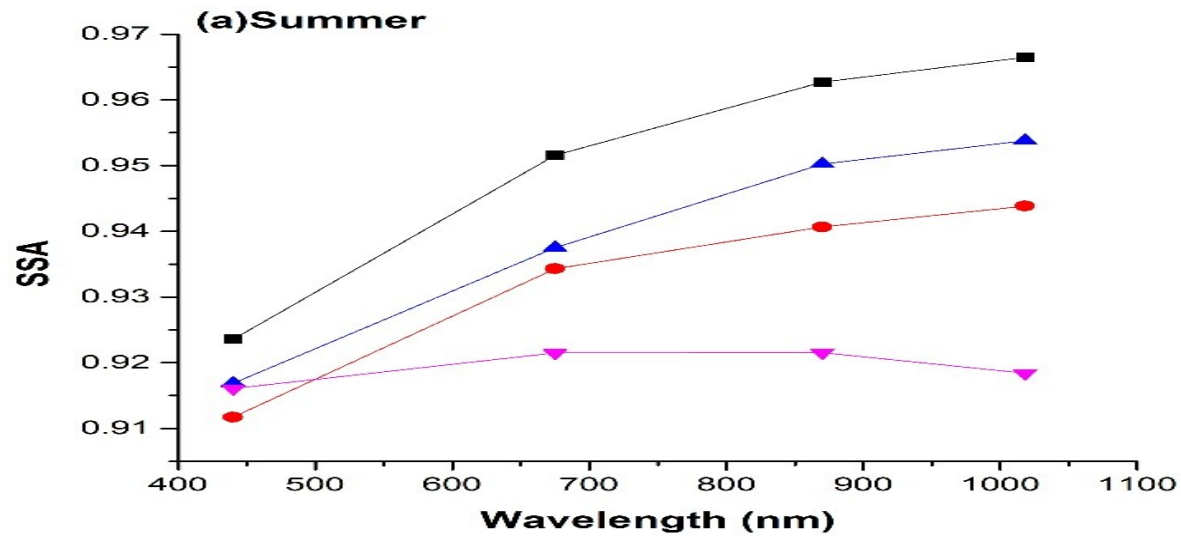
High coarse-mode peak in summer and pre-monsoon



Seasonal Averaged Variability of SSA

SSA increases with increasing wavelength

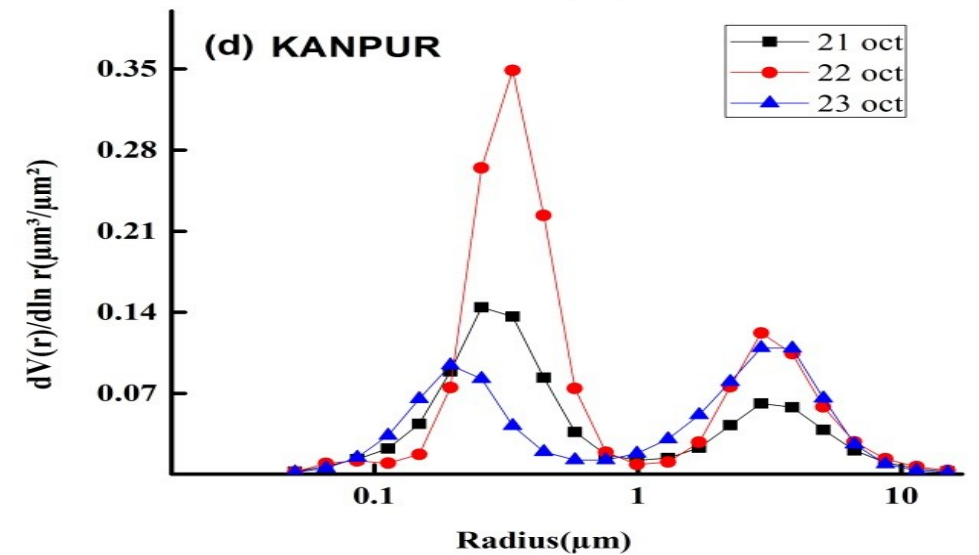
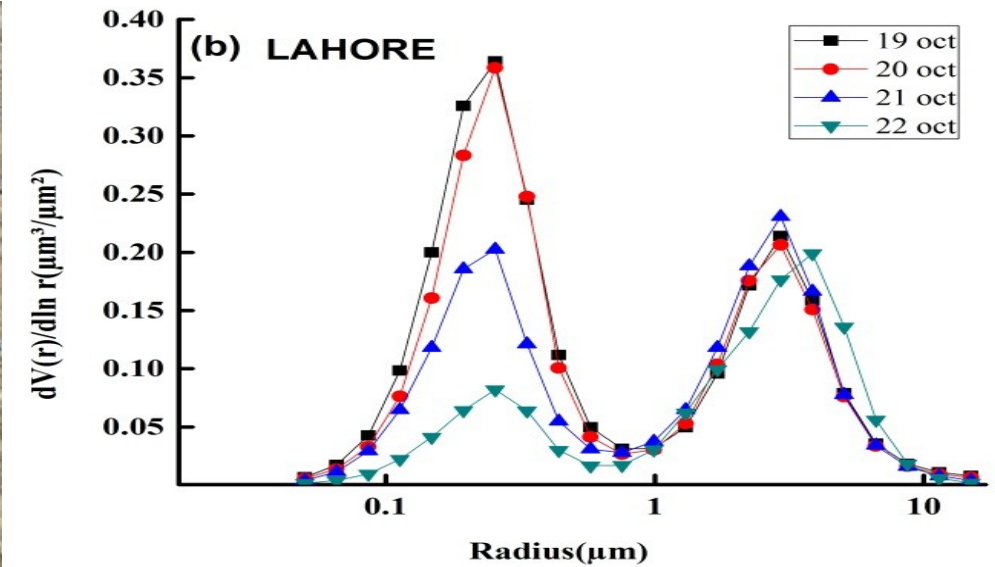
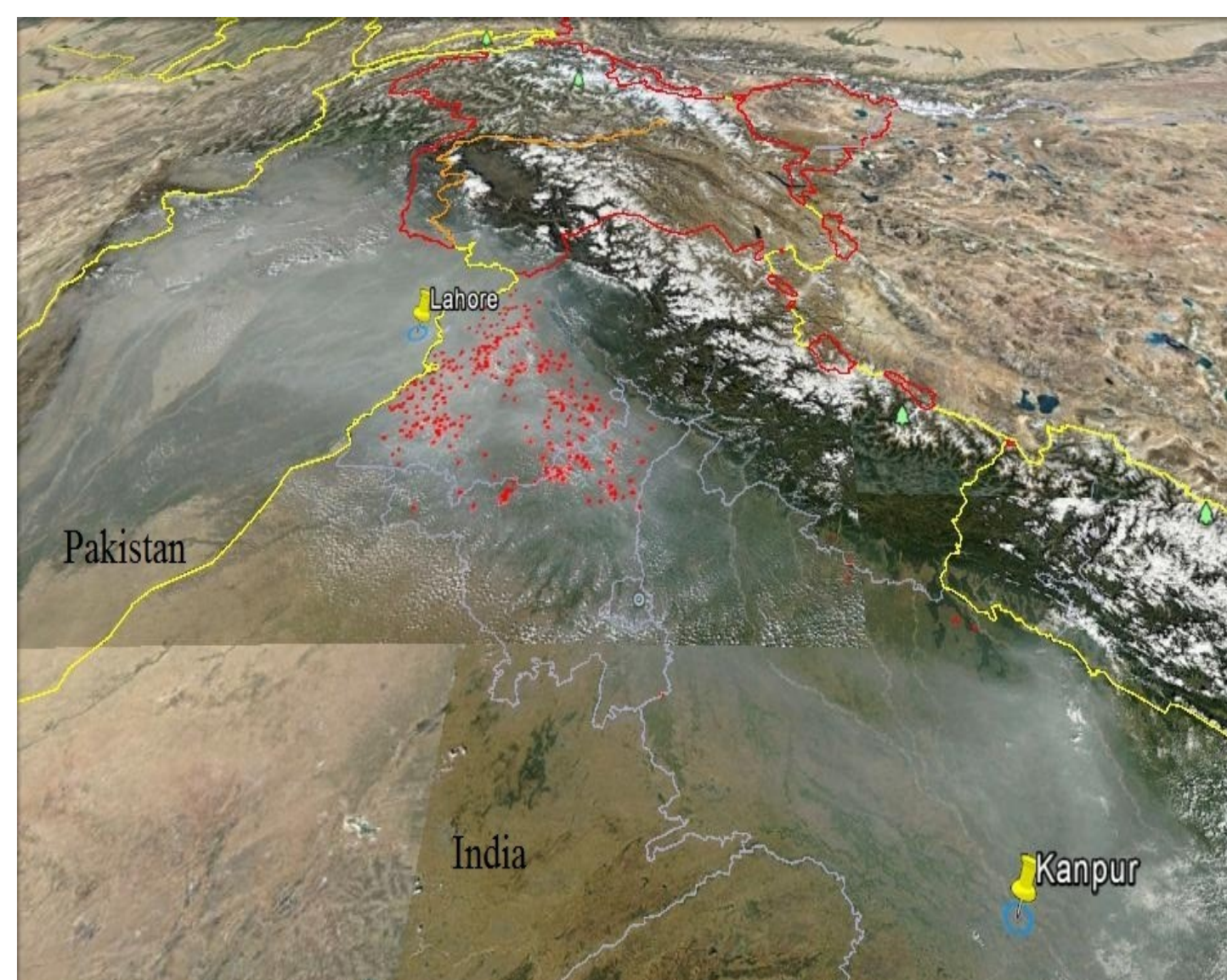
SSA values decreases with increasing wavelength



Seasonal Averaged Variability of SSA

Smoke/Haze

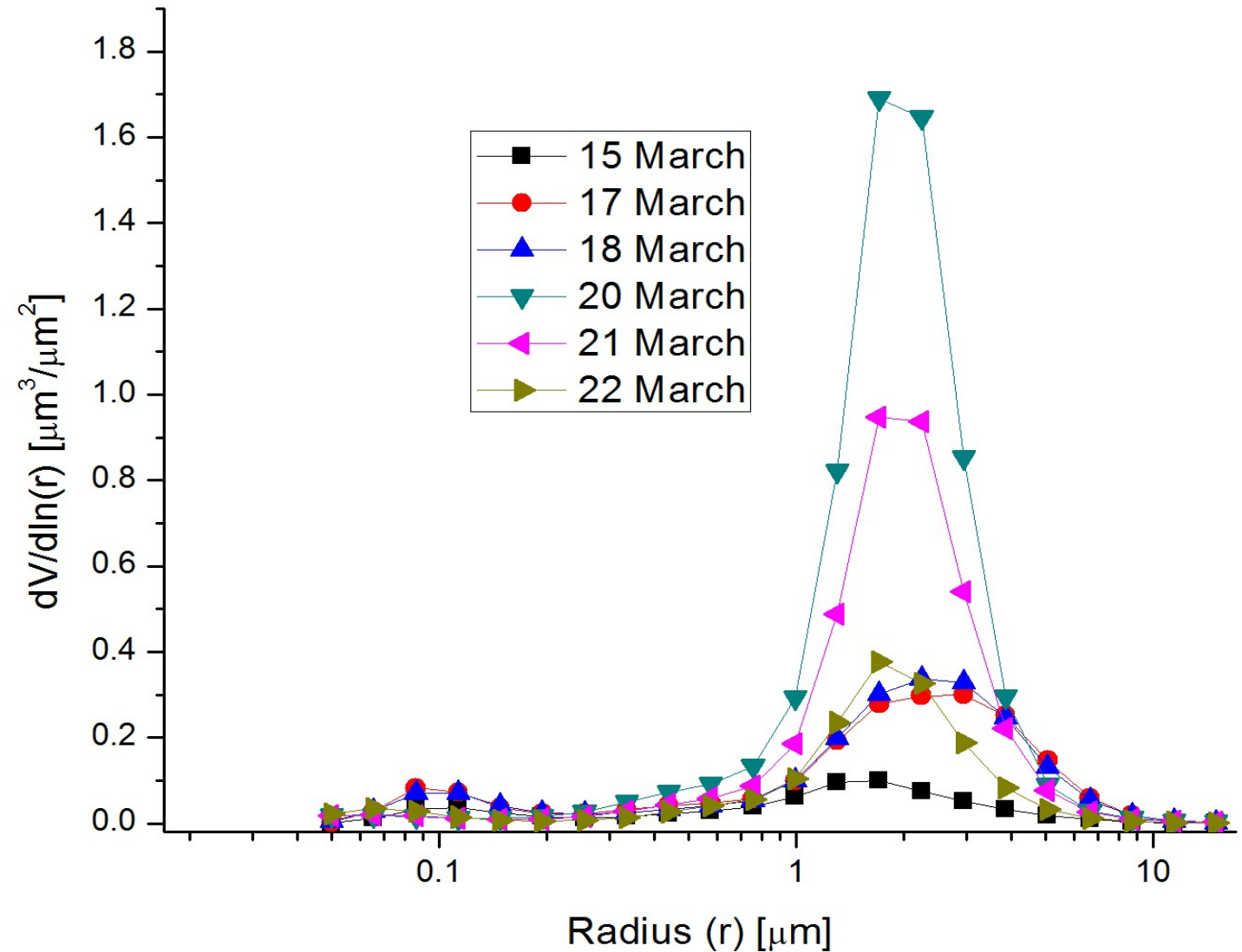
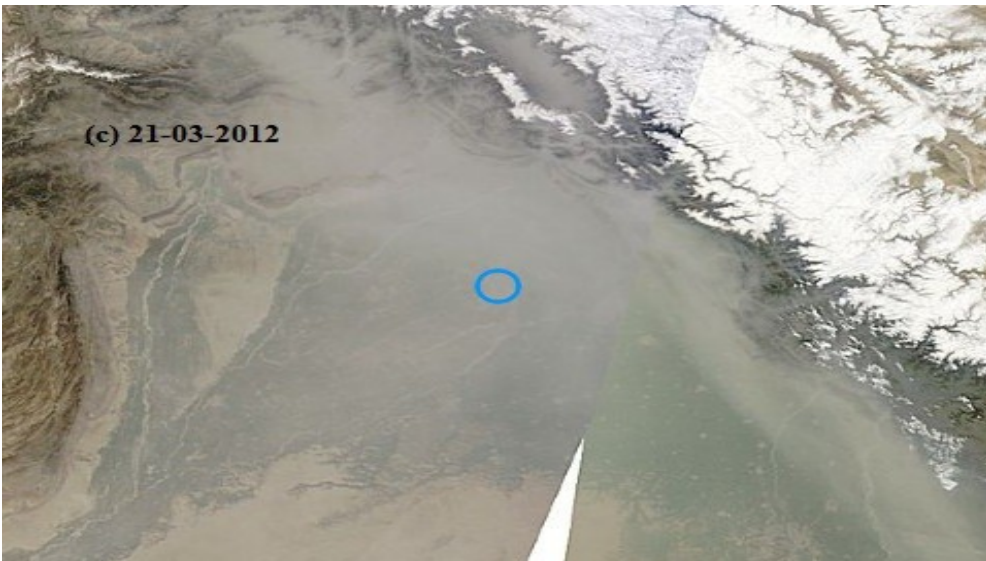
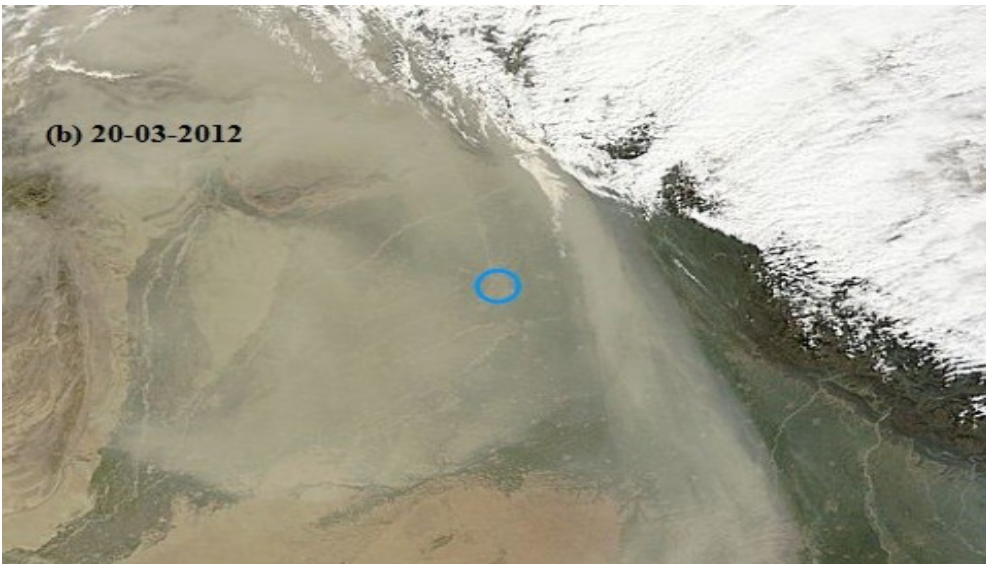
Volume size distribution



Seasonal Averaged Variability of SSA

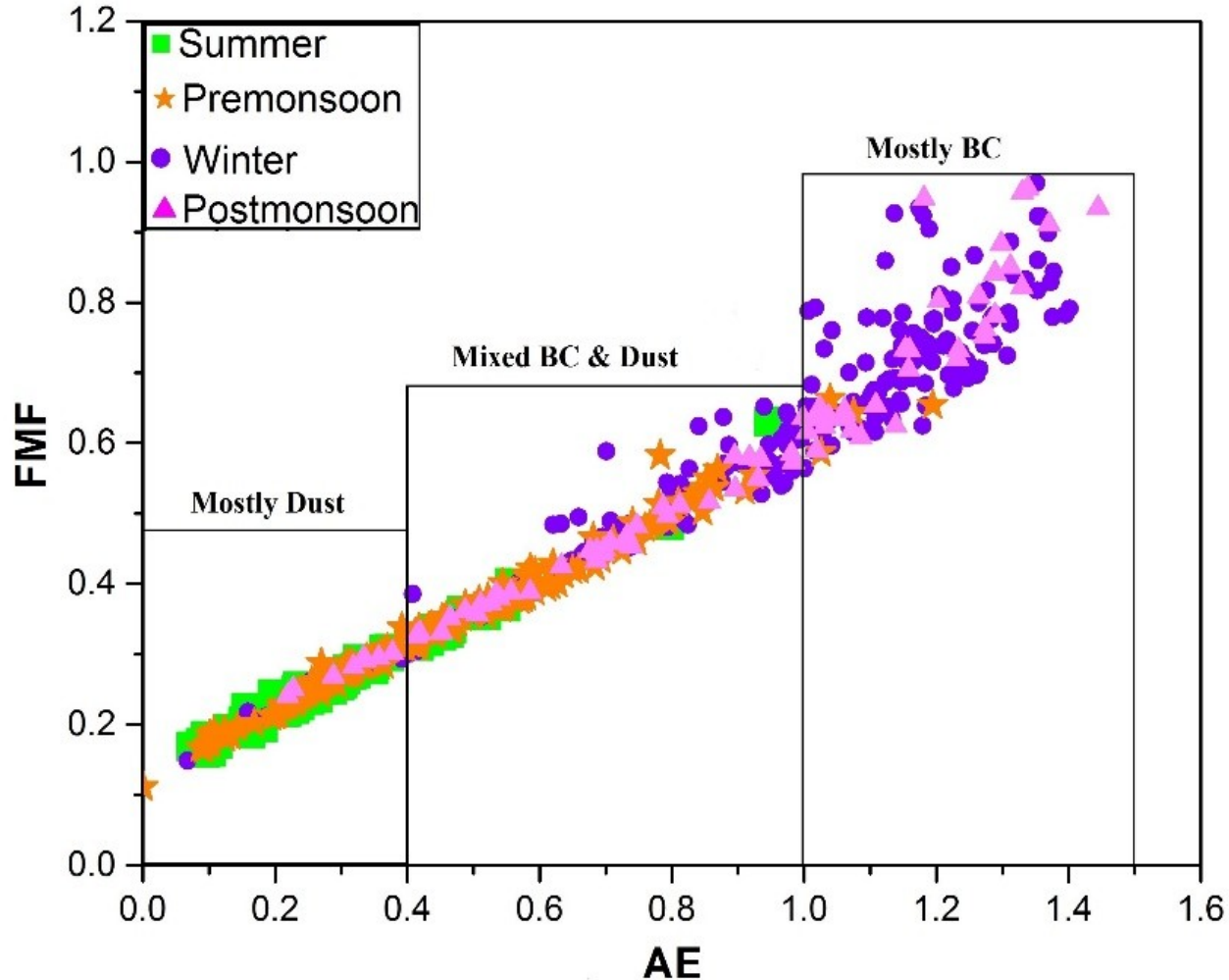
Dust

Volume size distribution



Classification of Aerosol

Relationship between FMF and AE

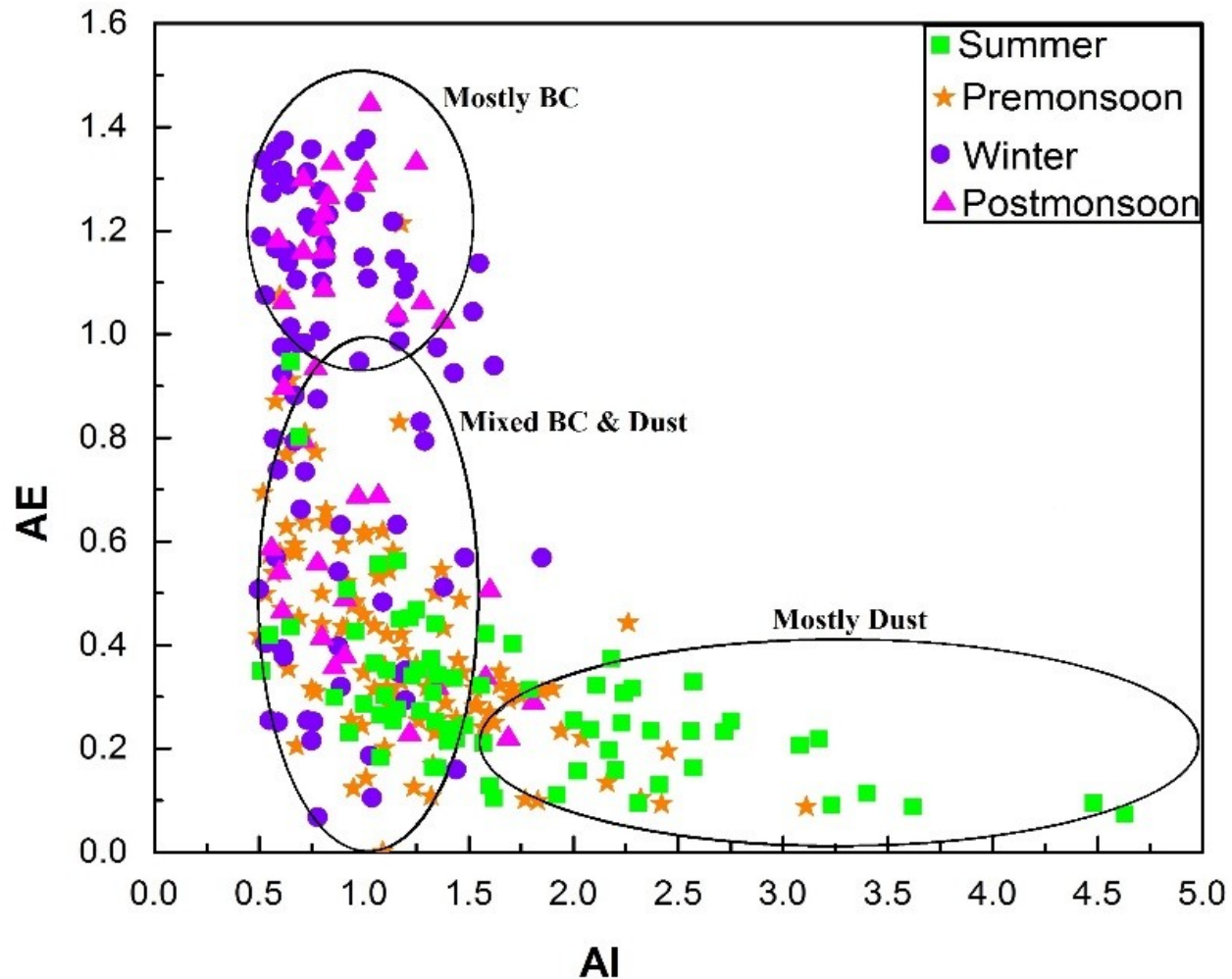


Mostly Dust \Rightarrow summer and premonsoon
Mostly BC \Rightarrow winter and postmonsoon

Aerosol types	FMF vs. AE	
Mostly Dust	$0.4 < \text{FMF} < 0.7$	$0.4 < \text{AE} < 1.0$
Mostly BC	$1.0 < \text{FMF} < 0.7$	$1.5 < \text{AE} < 1.0$
Mixed BC & Dust	$0.7 < \text{FMF} < 0.4$	$1.0 < \text{AE} < 0.4$

Classification of Aerosol

Relationship between AE and AI

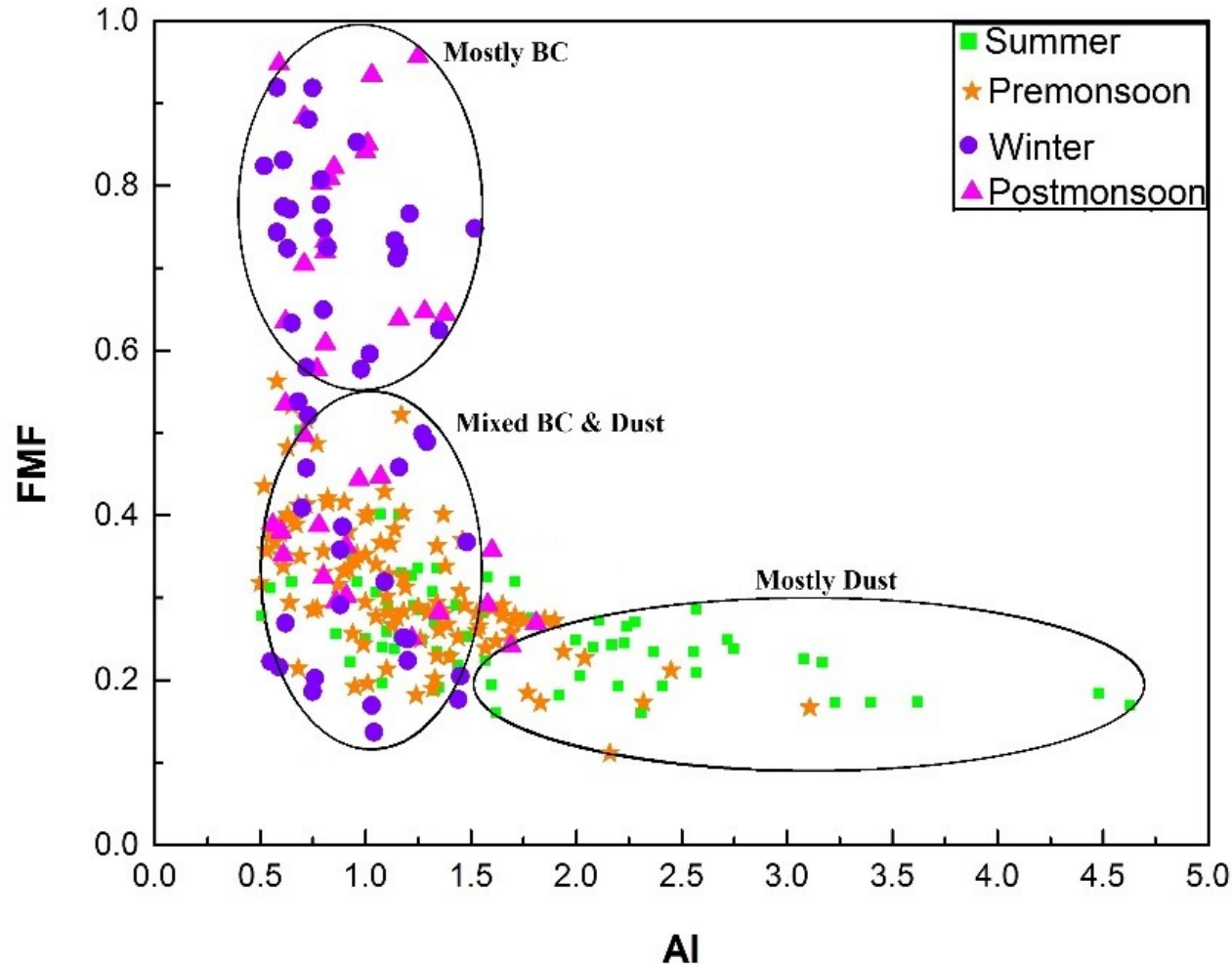


Mostly Dust \Rightarrow summer and premonsoon
Mostly BC \Rightarrow winter and postmonsoon

Aerosol types	AE vs. AI	
Mostly Dust	$0.4 < AE < 0.9$	$5.0 < AI < 1.57$
Mostly BC	$1.50 < AE < 0.92$	$1.52 < AI < 0.44$
Mixed BC & Dust	$1.0 < AE < 0.0$	$1.55 < AI < 0.5$

Classification of Aerosol

Relationship between FMF and AI

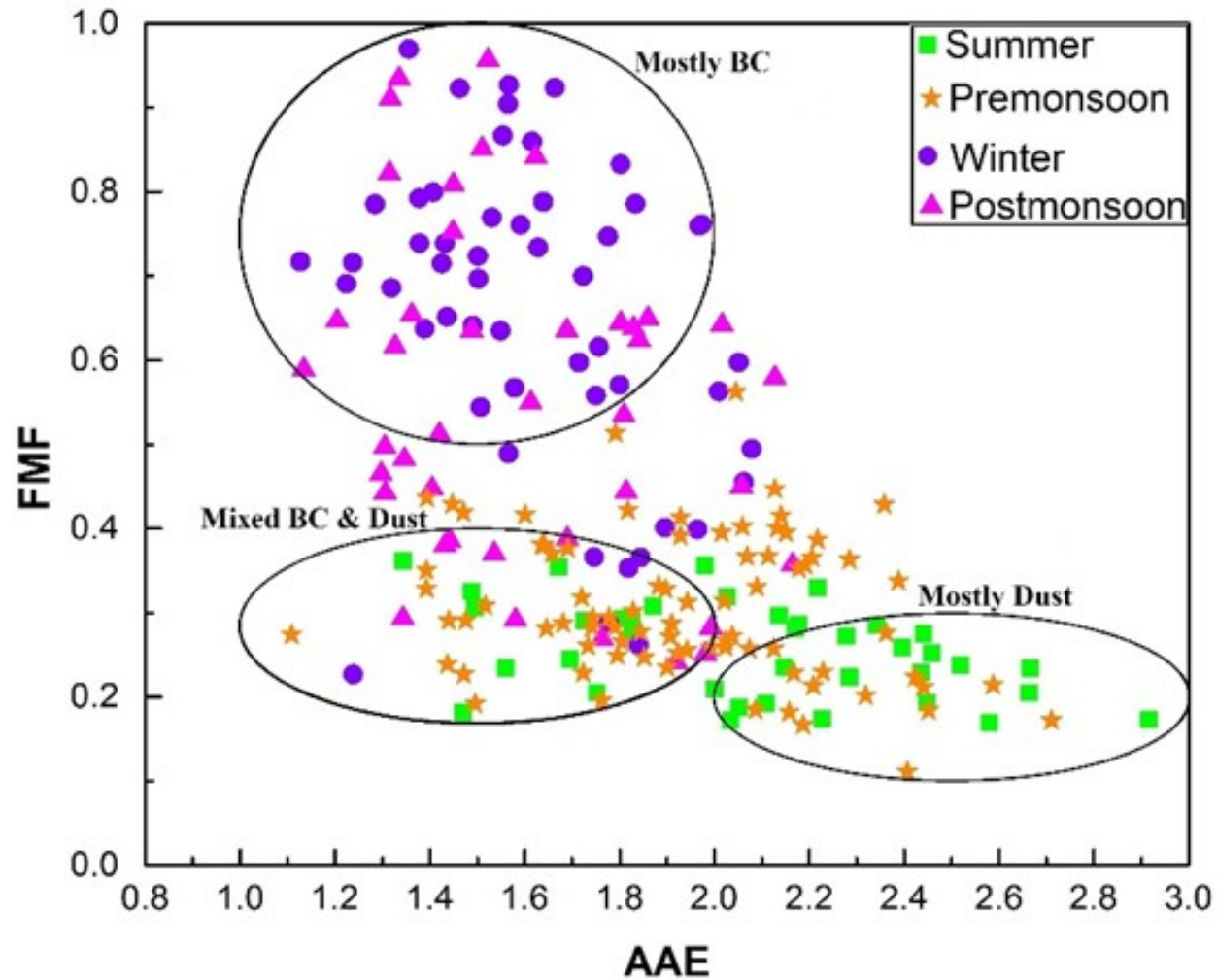


Mostly Dust \Rightarrow summer and premonsoon
Mostly BC \Rightarrow winter and postmonsoon

Aerosol types	FMF vs. AI	
Mostly Dust	$0.3 < \text{FMF} > 0.1$	$4.5 < \text{AI} > 1.50$
Mostly BC	$1.0 < \text{FMF} > 0.57$	$1.55 < \text{AI} > 0.4$
Mixed BC & Dust	$0.55 < \text{FMF} > 0.1$	$1.55 < \text{AI} > 0.5$

Classification of Aerosol

Relationship between FMF and AAE

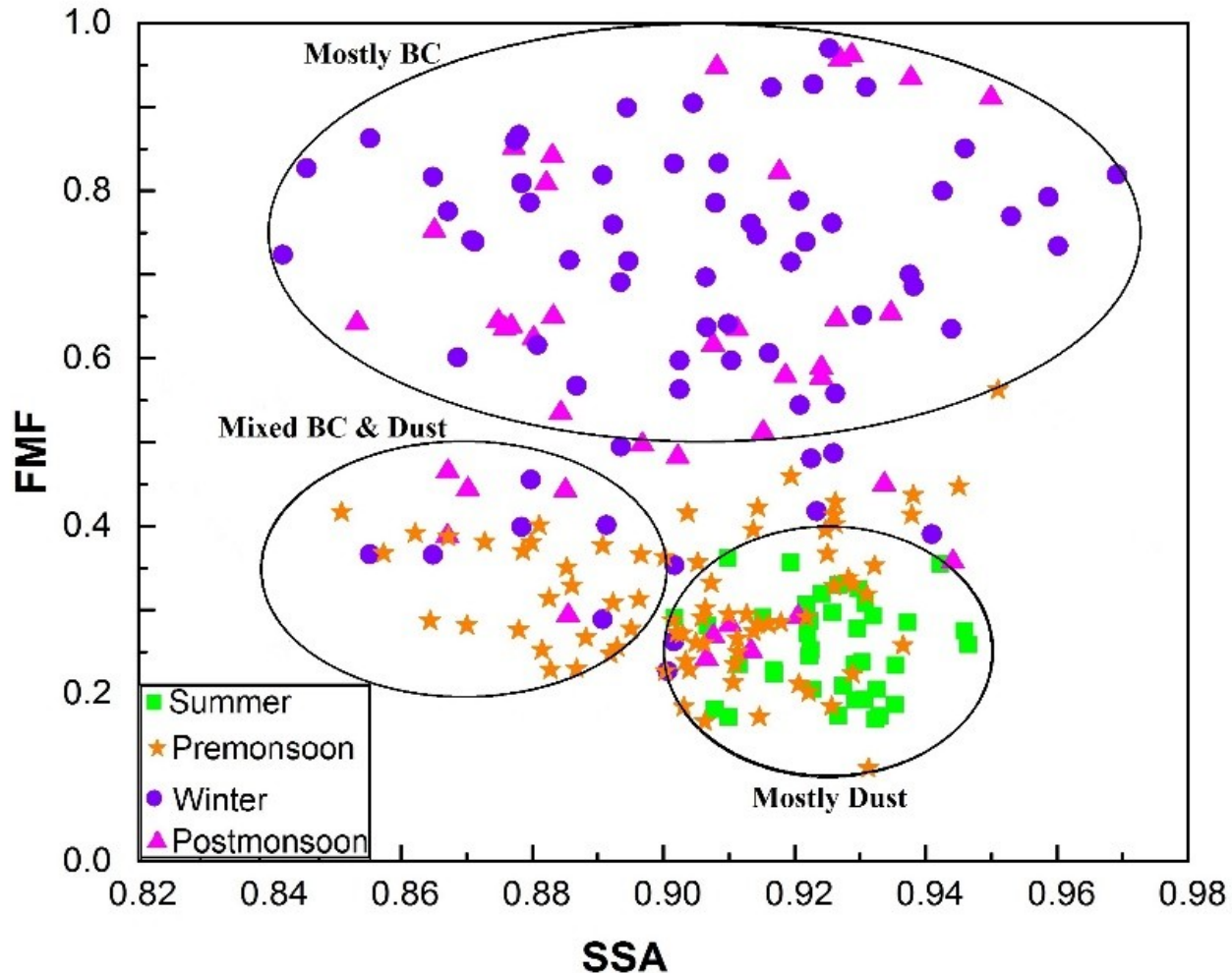


Mostly Dust \Rightarrow summer and premonsoon
Mostly BC \Rightarrow winter and postmonsoon

Aerosol types	FMF vs. AAE	
Mostly Dust	$0.3 < \text{FMF} < 0.5$	$3.0 < \text{AAE} < 2.0$
Mostly BC	$1.0 < \text{FMF} < 0.5$	$2.0 < \text{AAE} < 1.0$
Mixed BC & Dust	$0.4 < \text{FMF} < 0.1$	$2.0 < \text{AAE} < 1.0$

Classification of Aerosol

Relationship between FMF and SSA



Mostly Dust \Rightarrow summer and premonsoon
Mostly BC \Rightarrow winter and postmonsoon

Aerosol types	FMF vs. SSA	
Mostly Dust	$0.4 < \text{FMF} > 0.1$	$0.95 < \text{SSA} > 0.90$
Mostly BC	$1.0 < \text{FMF} > 0.5$	$0.91 < \text{SSA} > 0.84$
Mixed BC & Dust	$0.4 < \text{FMF} > 0.2$	$0.90 < \text{SSA} > 0.84$

The background features a gradient from light blue at the top to a deeper blue at the bottom. Several long, slender green leaves are scattered across the frame, some appearing to be part of a plant branch. The leaves have a vibrant green color with visible veins.

**با تشکر از
توجه شما**