



# ***Dust predication activies at the Egyptian Meteorological Authority***

***A. S. Zakey, Z. Salah, A. El-Awadi, and F. El-Ashmawy***

# Introduction

- Sever Dust Storms (SDS) have a major impact on the air quality and climate of north Africa and Mediterranean. *We must find a solution to reduce the impacts of dust storm on the society and environment.* May be the land use change is the best solution for dust mitigation ???
- In general, the arid and semi-arid area of MENA, the SDS are occurs frequently. In this study we address the following scientific question: *What is the impacts of land-use change on SDS frequencies and intensities?*
- As well, the detection of the dust source region and monitoring of the dust plume from its primary outflow to final deposition will be taken into accounts.
- The main idea of integrated dust storm monitoring and modeling system should be described as a key research areas, including new dust modules (*More understandable Processes*) and techniques in satellite remote sensing and system integration.

# What Controls Dust Emissions

## MODELED NATURAL EMISSIONS:

- Preferential sources
- Vegetation cover / type
- Surface wind speed
- Soil particle size
- Soil moisture

## HUMAN IMPACTS:

### Soil surfaces

- Cultivation
- Overgrazing
- Deforestation
- Roads, construction, military activity

### Climate

- Meteorology (winds, precip)
- Natural vegetation

## **Some recent advances in regional dust modeling**

- Analyzed wind fields or nudged runs to facilitate model-data intercomparison, improve dust source parameterizations.
- Role of 'hot spots' and vegetation for dust emission.
- Observations: Major field campaigns to analyze properties of transported dust (ACE-Asia, PRIDE, SHADE).
- New satellite assessments and instrumentation.

## **Major remaining problems**

- Subgridscale winds, initial vertical transport.
- Soil properties in particular in 'hot spots', parameterization of crusting.
- Relation of dust mineralogy to source soil properties.
- Further data compilations, observations near source regions.
- Percentage of dust from agricultural sources ('anthropogenic dust')?

**Most uncertainties in dust cycle modeling are due to dust emissions**

# Dust Scheme in RegCM

## Key question :

- Dust climatic forcing and impacts are still uncertain (IPCC).
- Potential importance of dust for weather forecast over North and West Africa, China and Atlantic.
- Global climatic impact of dust ( impact on SST variability, Paleoclimate).
- Dust aerosol and regional climatic response are still difficult to represent in GCM and RCM.

# RegCM-CHEM Model Core

- **Dynamics:**
  - MM5 Hydrostatic (Grell et al 1994)
  - Non-hydrostatic (MM5 or WRF, in progress)
- **Radiation:**
  - CCM3.6.6 (Kiehl 1996)
  - RRTM (in progress)
- **Large-Scale Clouds & Precipitation:**
  - SUBEX (Pal et al 2000)
- **Cumulus Convection:**
  - Grell (1993) + FC80 Closure
  - Anthes-Kuo (1977)
  - MIT/Emanuel (1991)
  - Betts-Miller (1993)
  - STRACO (in progress)
- **Boundary Layer:**
  - Holtslag (1990)
- **Nesting:**
  - Numerous GCM/Reanalysis Interfaces
  - One-way nesting
- **Biogenic Emiss:**
  - MEGANE (Twffic, 2010)
- **Tracers/Aerosols:**
  - Qian et al (2001) - sulfur chem.
  - Solmon et al (2005) - BC/OC chem.
  - Zakey(2006) dust module
  - Zakey(2008) Sea-salt
  - Zakey(2010) DMS and sulafte from Ocean
  - Shalaby (2010) - gas-phase chem.
- **Land Surface:**
  - BATS1e (Dickinson et al., 1993)
  - SUB-BATS (Giorgi et al., 2003)
  - CLM (Dai et al., 2003, Dai & Bi, in progress)
  - IBIS (Foley; Winter in progress)
- **Ocean Fluxes:**
  - BATS1e (Dickinson et al., 1993)
  - Zeng et al (1998)
  - Air-Sea Coupling (MITogcm, OASIS coupler, in progress)
- **Computations:**
  - User-Friendly
  - Multiple Platforms
  - Parallel Code





# ***RegCM-CHEM: Chemistry Core***

## **Chemistry:**

Condensed CBM-Z gas-phase chemistry (Zaveri and Peters, 1999).

## **Solver:**

Radical balance method (RBM) by (Sillman et al., 1991) and (Barth et al., 2002)

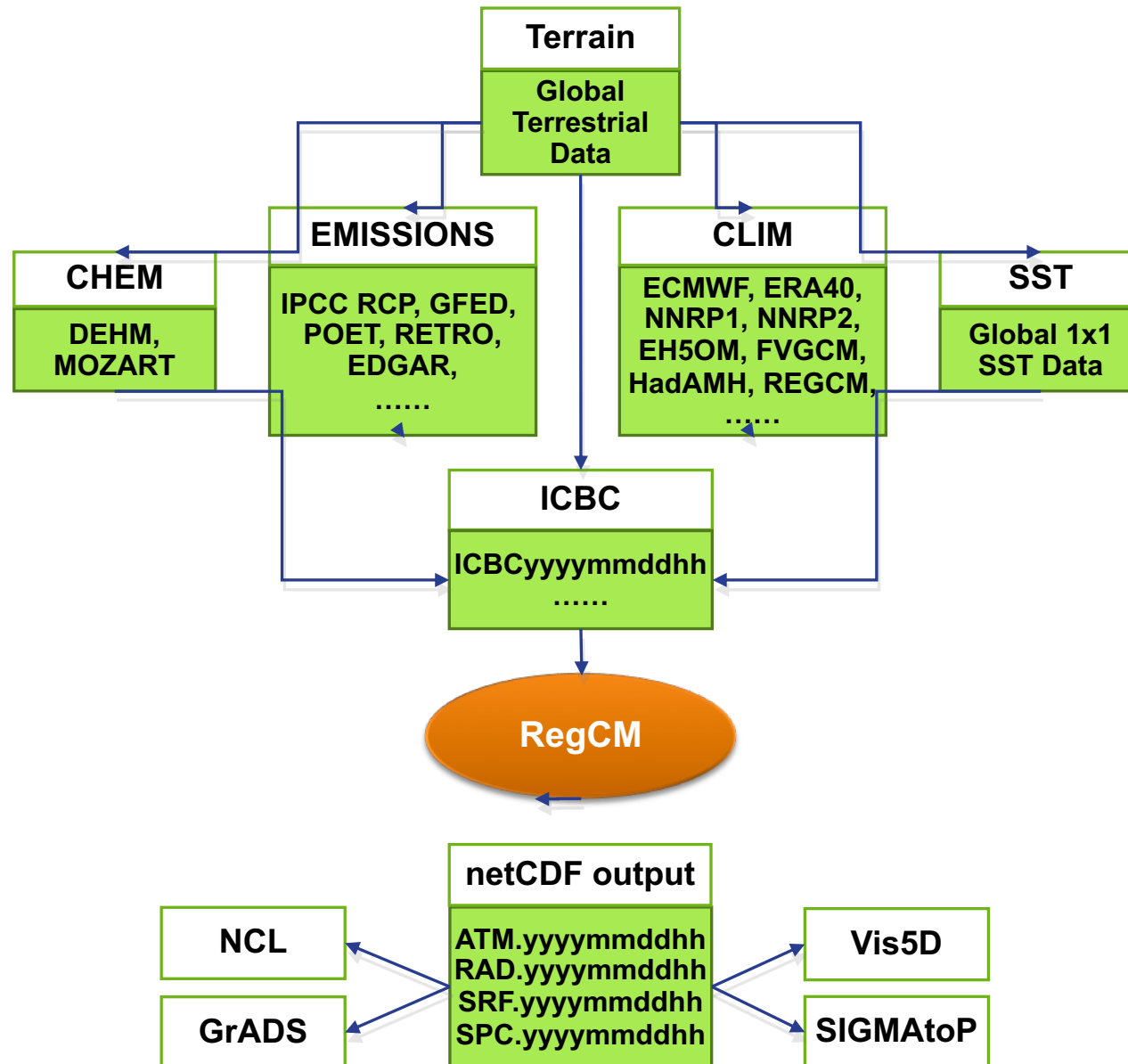
## **Photolysis rates:**

Tropospheric Ultraviolet-Visible Model (Madronich and Flocke, 1999) with cloud cover correction by (Chang et al., 1987)

## **Dry deposition:**

- “big leaf” multiple resistance model with aerodynamic, quasi-laminar layer, and surface resistance for 31 gaseous species.
- uptake resistance for vegetation, soil, water, snow and ice (20 land-use types).
- stomata and non-stomata resistances

# RegCM-CHEM: Modeling System Flow Chart





## Dust in RegCM

- General approach  $\longleftrightarrow$  Tracer model / RegDAM

$$\frac{\partial \chi}{\partial t} = \underbrace{-\bar{V} \cdot \nabla \chi}_{\text{Transport}} + \underbrace{F_H + F_V + T_{CUM} + S_\chi}_{\text{Primary Emissions}} - \underbrace{R_{w,ls} - R_{w,cum}}_{\text{Removal terms}} - D_{dep} + \underbrace{\sum Q_p - Q_l}_{\text{Physico-chemical transformations}}$$

Strongly dependent on the nature of the tracer

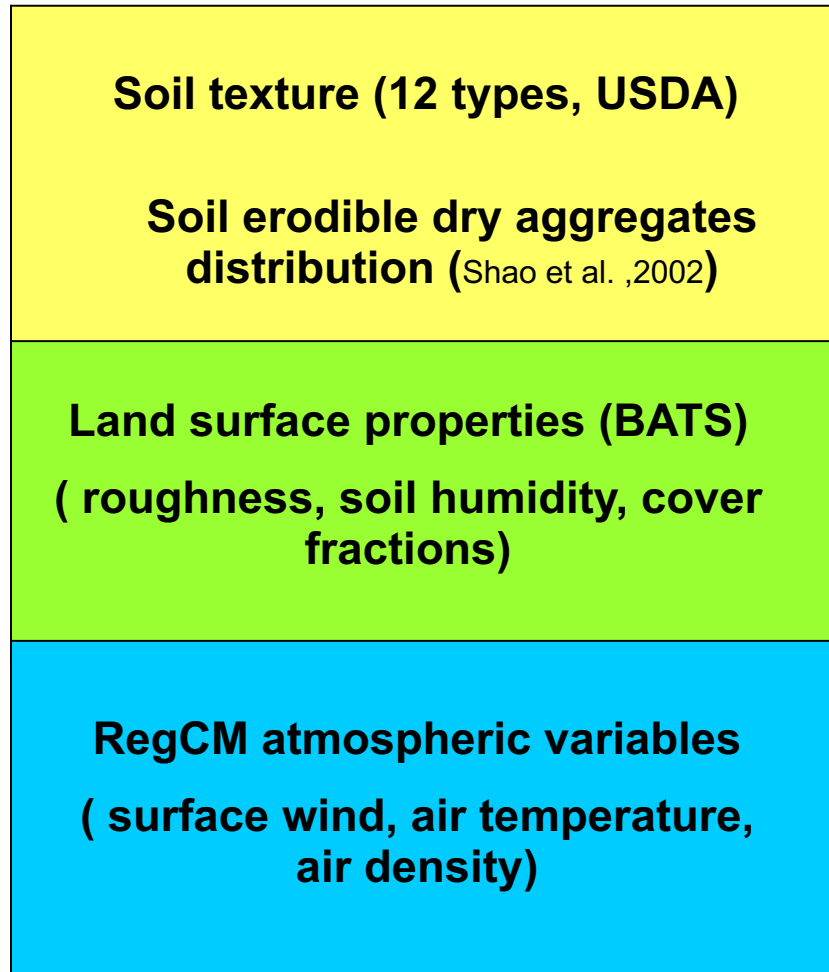
- Particles and chemical species considered.

$\text{SO}_2 \longleftrightarrow \text{SO}_4^{--}$		BC (soot)		OC (total organic carbon)		DUST (4 bins)			
Aqueous and gazeous conversion (Qian et al., 2001)		Hydrophilic (20% at emission)	Hydrophobic (80%at emission)	Hydrophilic (50%at emission)	Hydrophobic(50%at emission)	0.01-1 $\mu\text{m}$	1-2.5 $\mu\text{m}$	2.5-5 $\mu\text{m}$	5-20 $\mu\text{m}$



## Aerosol dust modeling in RegCM ( Zakey et al., 2006 ACP)

Input parameters



**DUST emission scheme**  
*Zakey et al., 2006*



**Saltation** (Marticorena et al. 1995)

Roughness and humidity correction



**Suspension**

**Sand-blasting** (Alfaro et al., 1997, 2001)



**Dust flux distribution**  
(3 log-normal emission modes)



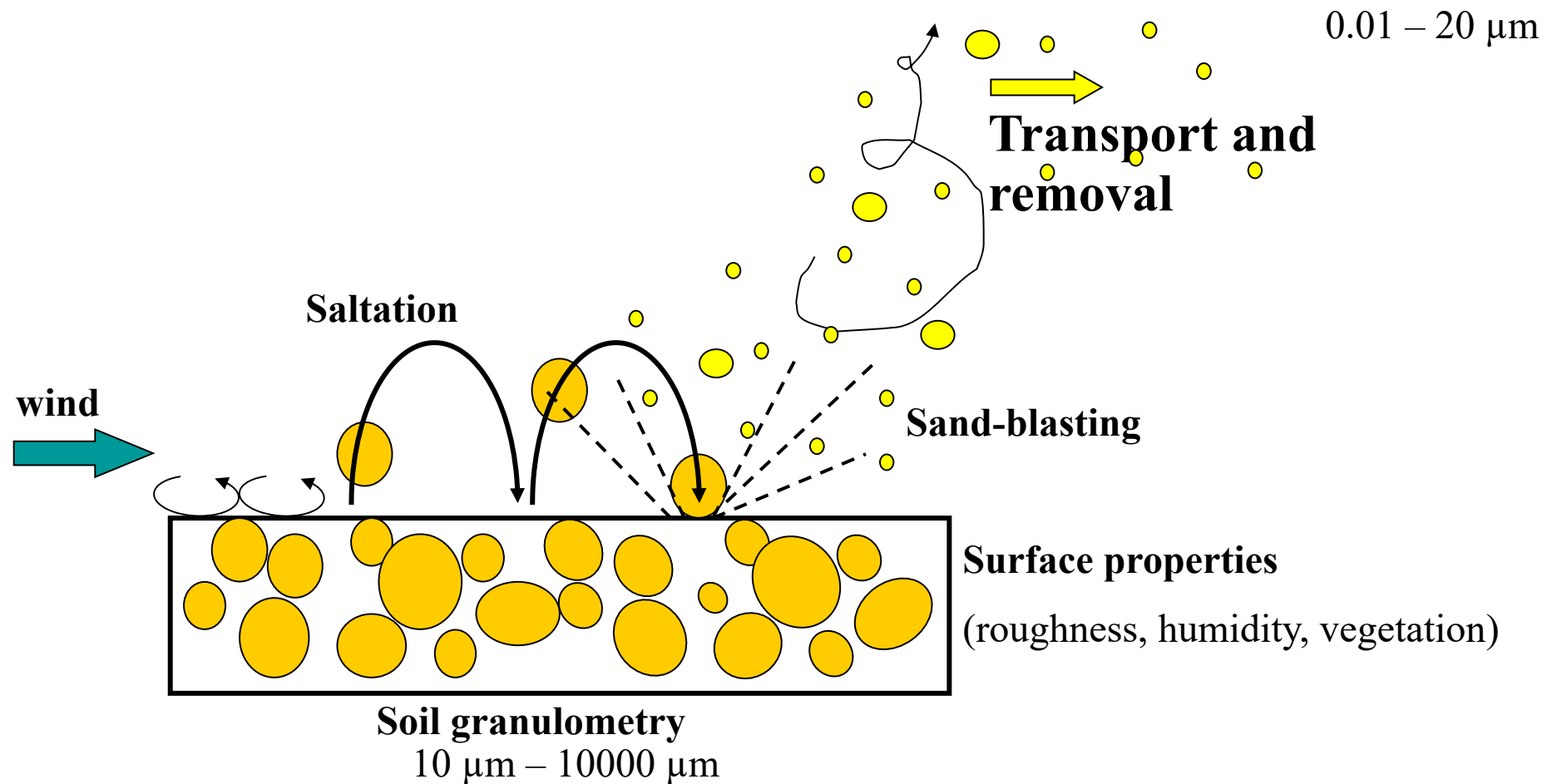
**Transport bins** (up to 12), usually 4



**AOP / radiation**

**Size dependent settling and surface deposition**





↕

**Zakey et al., 2006**

**RegCM**

Regional climate modeling approach

## Modeled dust emission

Dust Flux F: 
$$F = \alpha \cdot A_{eff} \cdot G \cdot (1 - A_{snow}) \cdot I_{\theta}$$

with

$$A_{eff} = 1 - (FPAR_{(max)} \cdot f_{shrub} + FPAR_{(monthly)} \cdot f_{grass}) \cdot \frac{1}{25}$$

$$G = \frac{\rho_a}{g} \cdot u_*^3 \cdot \sum_i \left[ \left( 1 + \frac{u_{*tri}}{u_*} \right) \left( 1 - \frac{u_{*tri}^2}{u_*^2} \right) \cdot s_i \right]$$

for  $u_* \geq u_{*tr}$

$\alpha$ : uplift constant (depending on texture)

$A_{snow}$ : relative snow cover

$I_{\theta}$ : soil moisture dependence

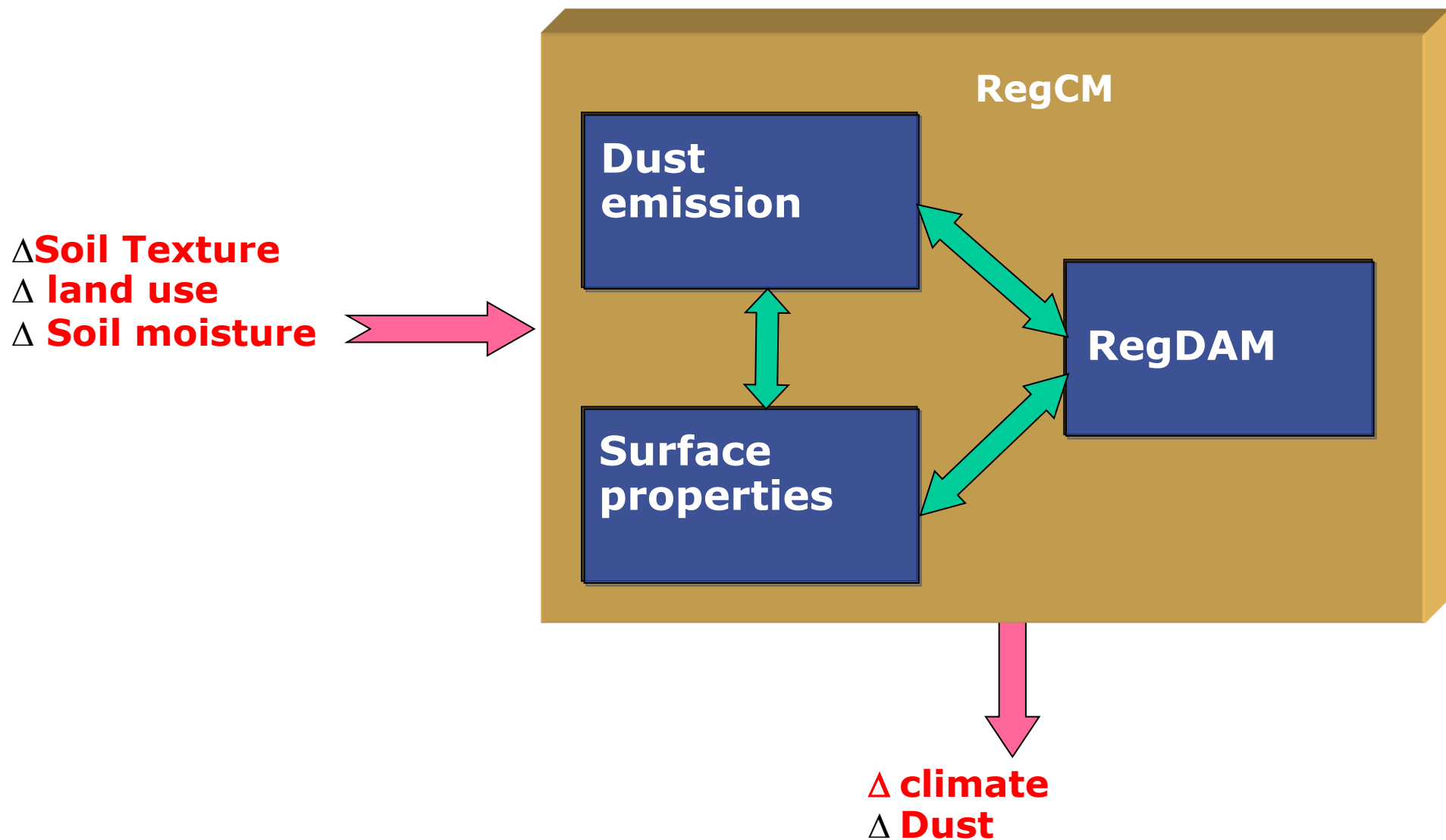
$\rho_a$ : air density

$g$ : gravitational constant

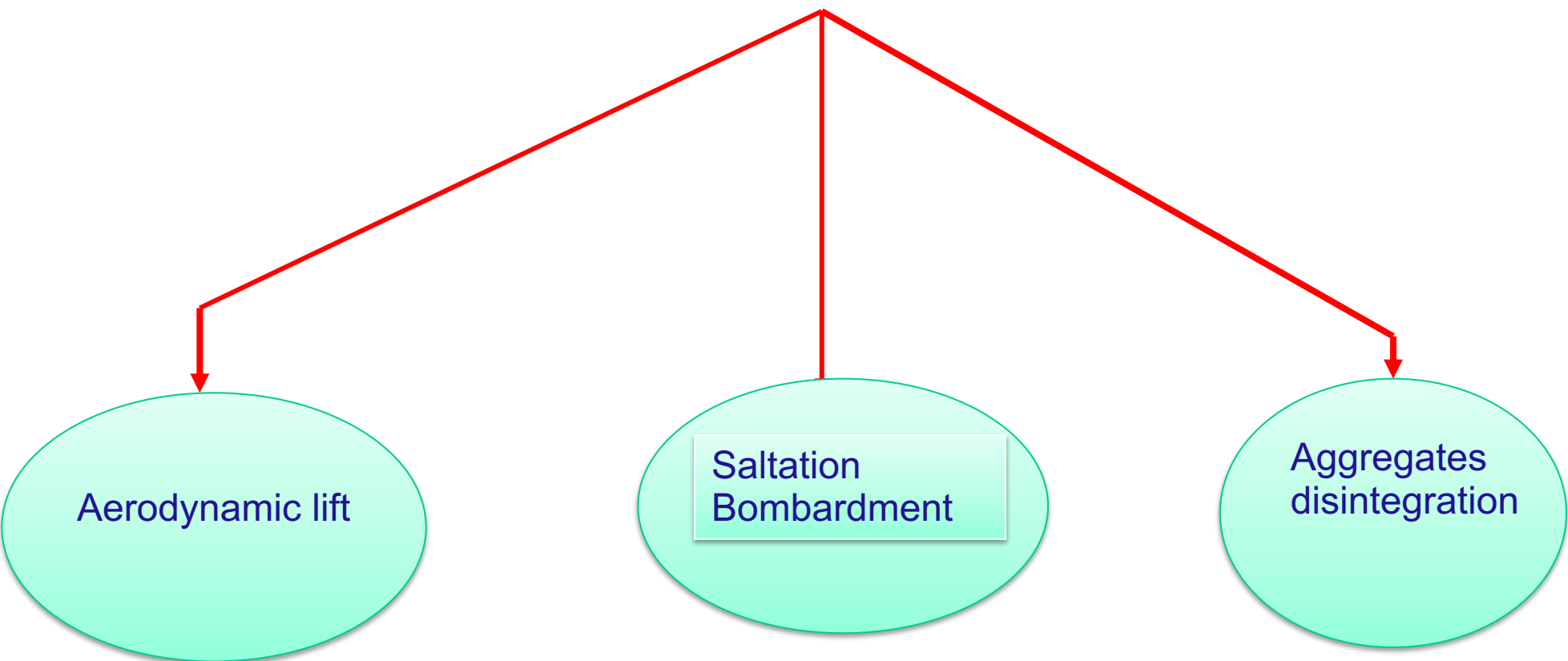
$u_*$ : wind shear velocity

$S_i$ : relative surface of particle size  $i$

## Approach: coupled Dust-climate RegCM (RegDAM)

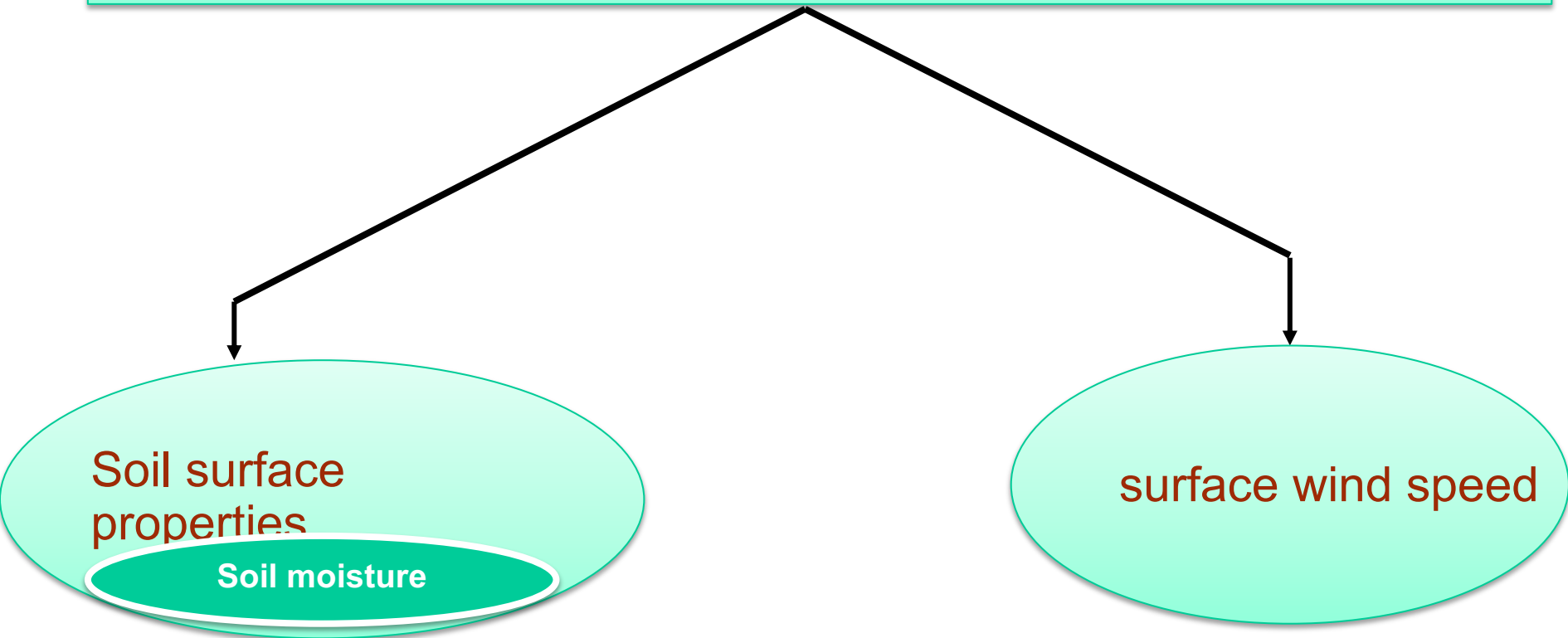


# Three Processes Control the dust emissions:





# Production of soil dust aerosols depends



$$\frac{dn(D_p)}{d(\log D_p)} = \sum_{i=1}^3 \frac{n_i}{\log \sigma_i \sqrt{2\pi}} \exp \left[ \frac{-(\log D_p / D_i)^2}{2(\log \sigma_i)^2} \right]$$

## Threshold friction velocity

Soil dust particles are mobilized only for wind speed greater than a threshold values (*Marticorena and Bergametti, 1995*).

The expected threshold friction velocity for the observed smallest size was:

$$u_{ts}^* = 0.129 K [1 - 0.0858 \exp\{-0.0617(\text{Re} - 10)\}] \quad \text{Re} > 10$$

$$= \frac{0.129 K}{(1.928 \text{Re}^{0.092} - 1)^{0.5}} \quad 0.3 < \text{Re} < 10$$

Where

$$\text{Re} = ad_s^X + b$$

$$K = \sqrt{\frac{2g\rho_p D_p}{\rho_a} \left( 1 + \frac{0.006}{\rho_p g (2D_p)^{2.5}} \right)}$$

## Threshold friction velocity Correction: *Drag partition*

*Marticorena and Bergametti(1995)* showed that the threshold wind friction velocity is strongly dependent on the roughness height. They derived a drag partition scheme from an approach developed by *Arya (1975)*. The drag partition parameterization is expressed by:

$$f_{\text{eff}} = \frac{u_{\text{tr}}^*}{u_{\text{ts}}^*} = 1 - \left[ \frac{\ln\left(\frac{z_0}{z_{0s}}\right)}{\ln\left[0.35\left(\frac{10}{z_{0s}}\right)^{0.8}\right]} \right]$$

Threshold friction velocity Correction: *Soil moisture*

The threshold friction velocity is also a function of soil moisture content, which has been parameterized (*Fécan et al., 1999*) as :

$$\frac{u_{tw}^*}{u_{tr}^*} = \left[ 1 + A \left( w - w' \right)^B \right]^{0.5} \quad \text{for } w > w'$$

$$= 1$$

$$\text{for } w < w'$$

$$w' = 0.0015(\%clay)^2 + 0.17(\%clay)$$

# Dust Flux

Vertical

Horizontal

$$F = m_d n N$$

Where :  $n$  is the number of saltation impacts per unit area and time;  $m_d$  the mass of dust particles

$$G_{tot} = E \frac{\rho_a}{g} u^{*3} \int_{D_{ss}} (1 + R)(1 - R^2) dS_{rel}(D_p) d(D_p)$$

$$R = \frac{u_{tw}^*}{u^*}$$

$S_{rel}$  is the relative surface covered by particles

## *The kinetic energy flux*

In the saltation and sandblasting process, the fine particles released either from saltating aggregates or from the surface depend on the individual kinetic energy (*Alfaro, and Gomes, 2001*).

$$e_c = \frac{\rho_p \pi}{12(D_p)^3 (20u^*)^3}$$

The kinetic energy flux  $dF_{kin}(D_p)$  of the saltating aggregates with diameters between  $D_p$  and  $D_p + dD_p$  is proportional to the corresponding horizontal saltation flux  $dF_h(D_p)$  (*Gillette and Stockton, 1986; Alfaro et al., 1997; Alfaro and Gomes, 2001*) as following:

$$dF_{kin}(D_p) = \beta dF_h(D_p)$$



## ***Total emitted dust flux***

The mass flux  $dF_{dust,i}(D_p)$  of the particles of the  $i^{th}$  dust mode released by the impacts of soil aggregates of size  $D_p$  are :

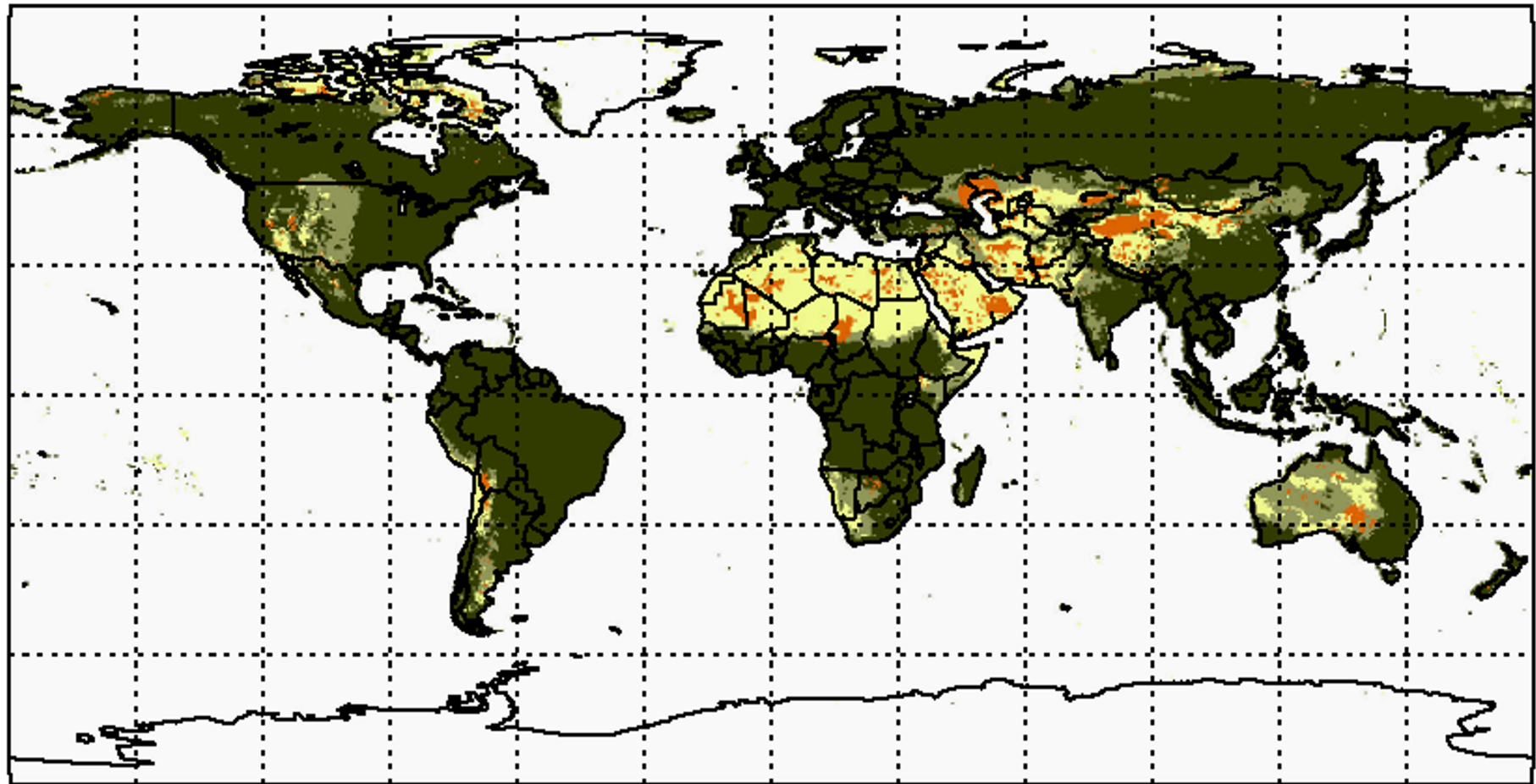
$$dF_{dust,i}(D_p) = \left(\frac{\pi}{6}\right) \rho_p \beta dF_h(D_p) \left[ \frac{p_i(D_p) d_i^3}{e_i} \right]$$

The total dust flux that represents fine transportable dust particles as the dust source strength is:

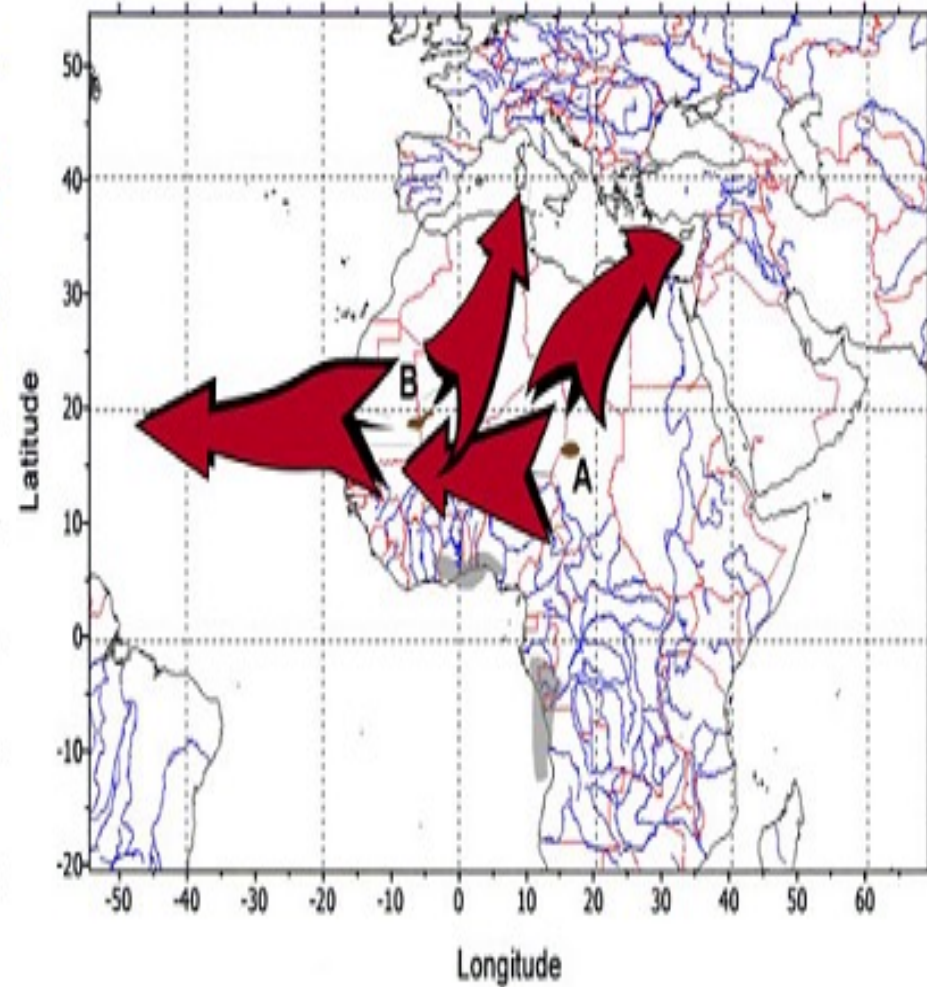
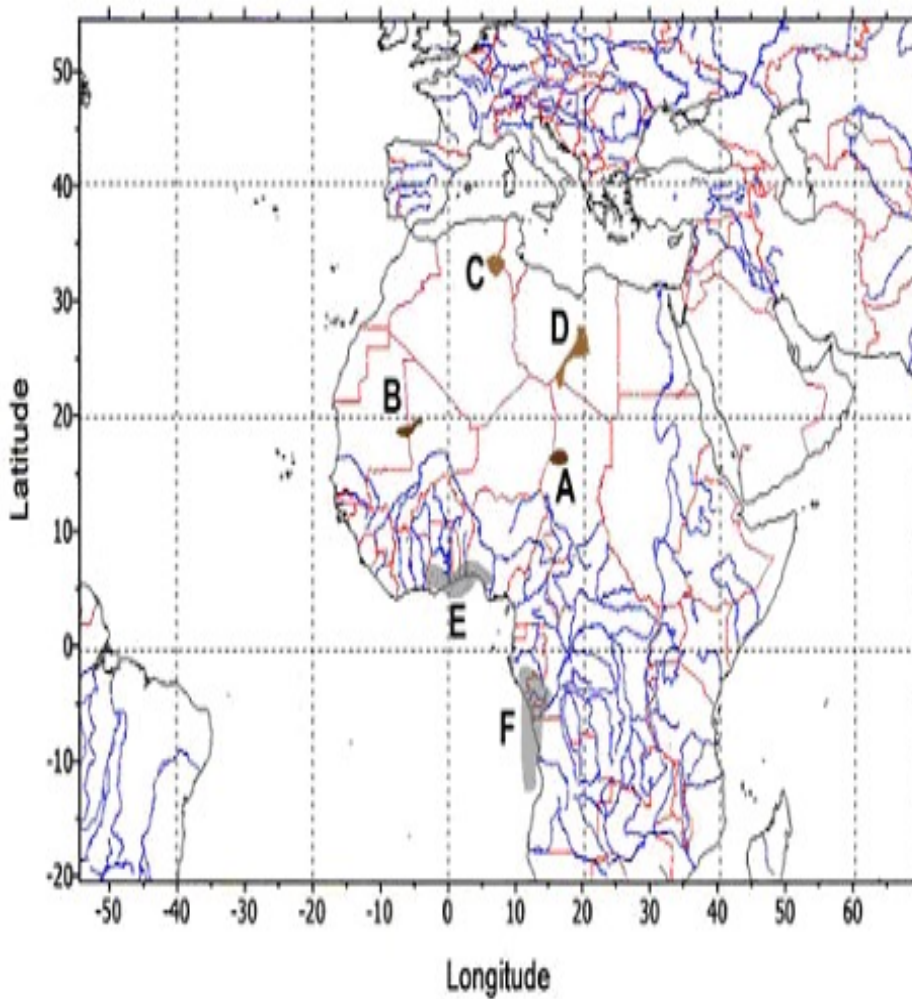
$$F_{total} = \left[ \sum_{i=1}^3 F_{dust,i} \right]$$

## Dust source areas in the RegDAM model

Dust Source Areas



**a) The main sources of dust emission    b) dust movements from its sources**

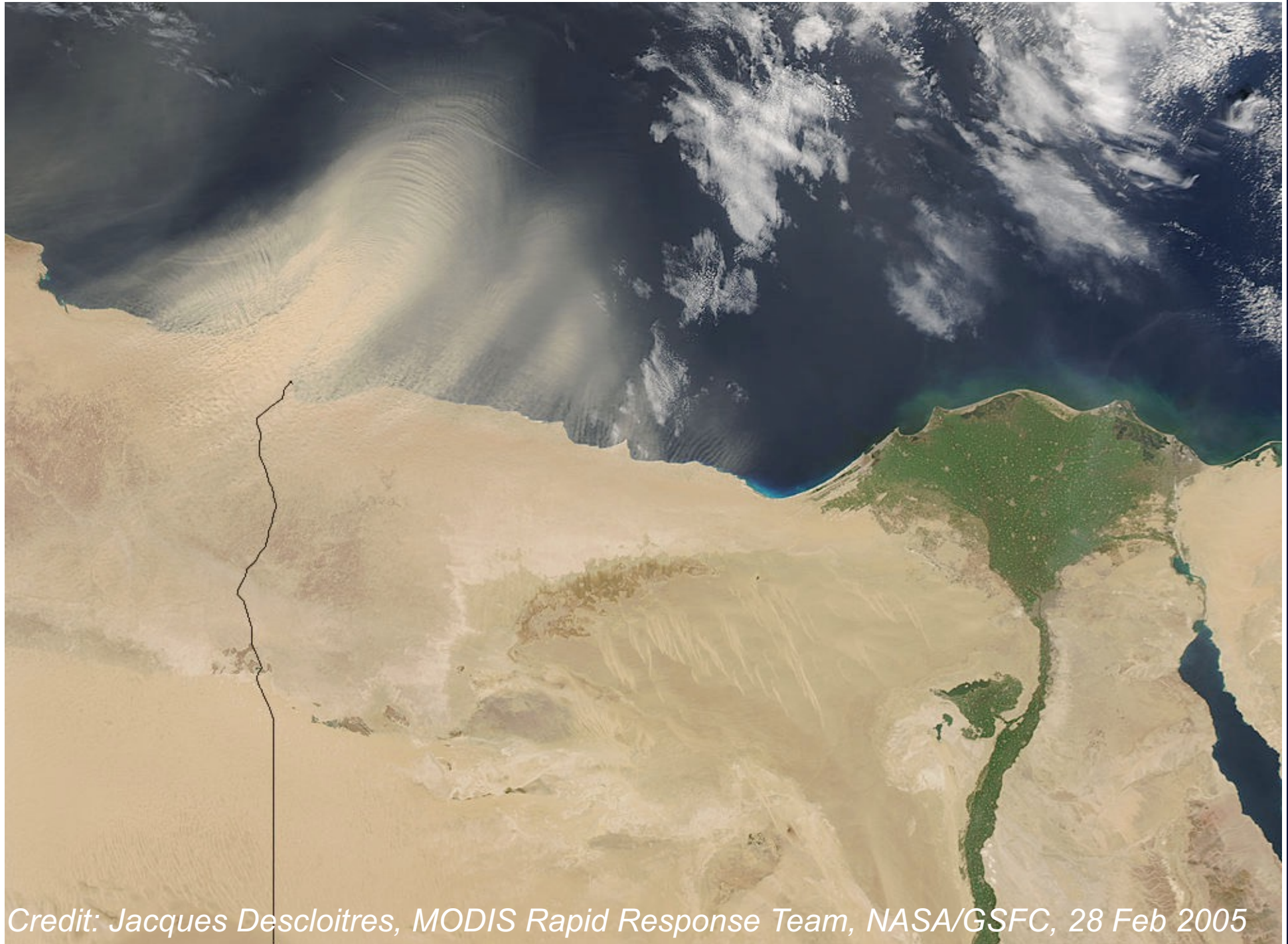




# Emissions: Dust

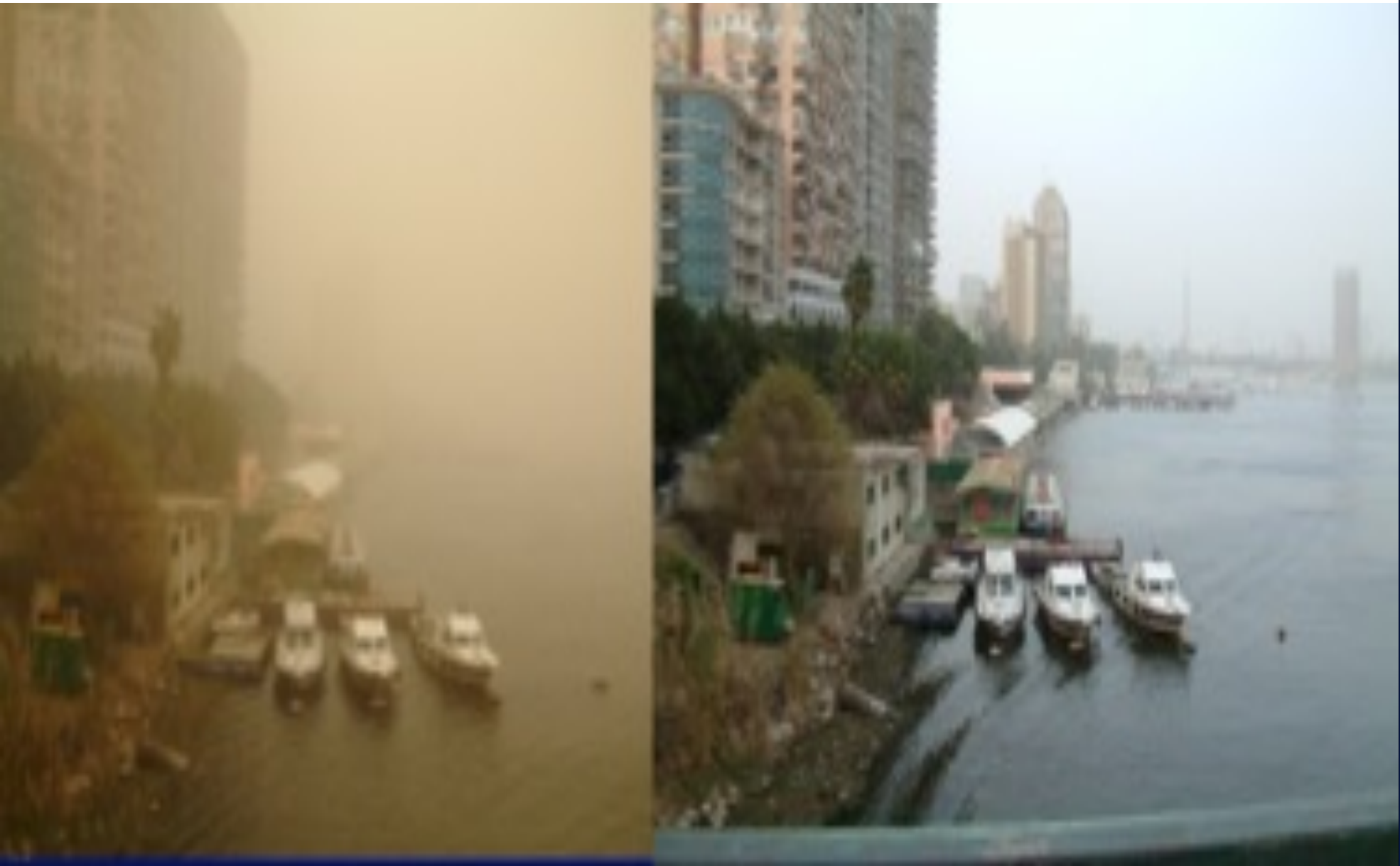


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**The Scientific Research Department**

*Credit: Jacques Descloitres, MODIS Rapid Response Team, NASA/GSFC, 28 Feb 2005*





High aerosol loading over Nile Delta and surrounding region as seen in the three day composite of MODIS Terra level-2 AOD (at 10 km spatial resolution) during October 1-3, 2008.

The vertical profile of feature mask obtained from CALIPSO on October 2, 2008 show aerosols at 2.5-6 km height.

