

3rd Training Course on WMO SDS-WAS products (satellite and ground observation and modelling of atmospheric dust)
Muscat-Oman, December 8-12, 2013



Ground observations of mineral dust

Emilio Cuevas (ecuevasa@aemet.es)

&

Sergio Rodríguez (srodriguezg@aemet.es)

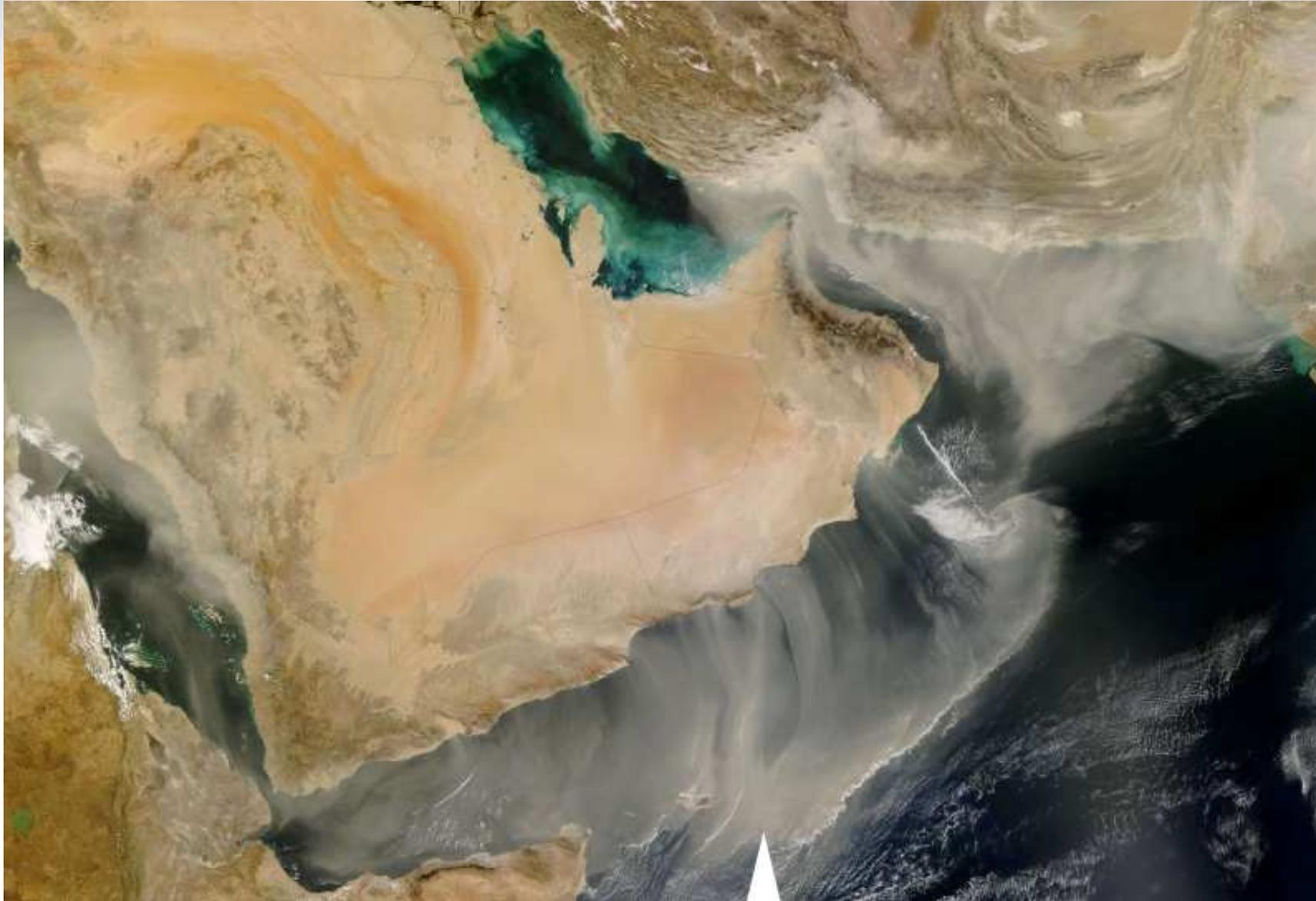
AEMET, Spain

Index

- Aerosols and dust background
- In-situ dust characterization
- In-situ dust estimations (Visibility)
- Ground based remote sensing
- Recommended ground-based observations in Middle East
- Summary

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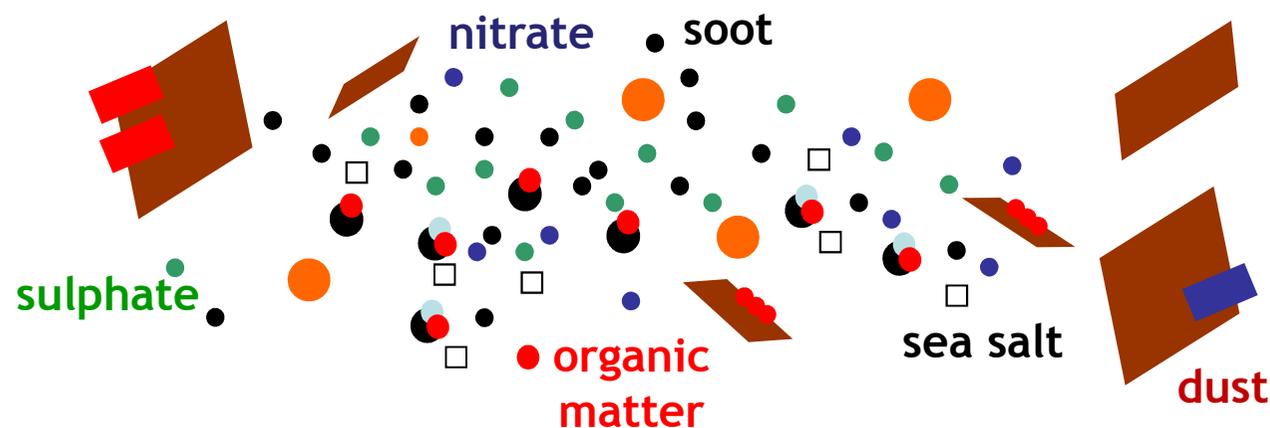
- **Aerosols and dust background**
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Images acquired from the MODIS instruments on both the Aqua and Terra satellites. The images were captured on February 22, 2008.

Aerosols: solid & liquid matter suspended in a gas.

Size 0.001 to 100 μm ($1 \mu\text{m} = 10^{-6} \text{ m}$) = 1 nm (10^{-9} m) to 100 μm (10^{-6} m).



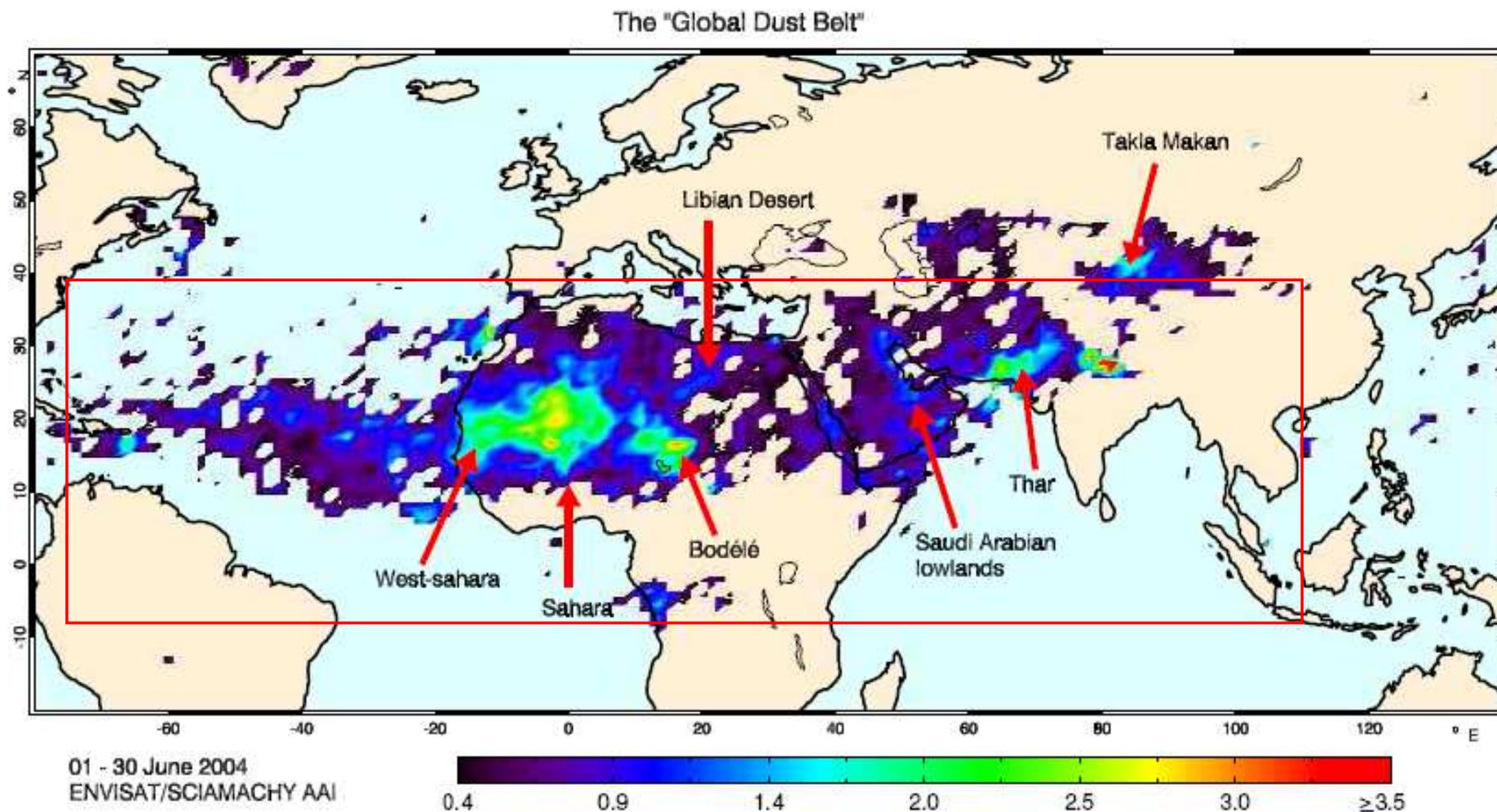
Mineral dust:

Small fragments of soil / crust of the Earth.

One of the most abundant aerosol in the Earth.

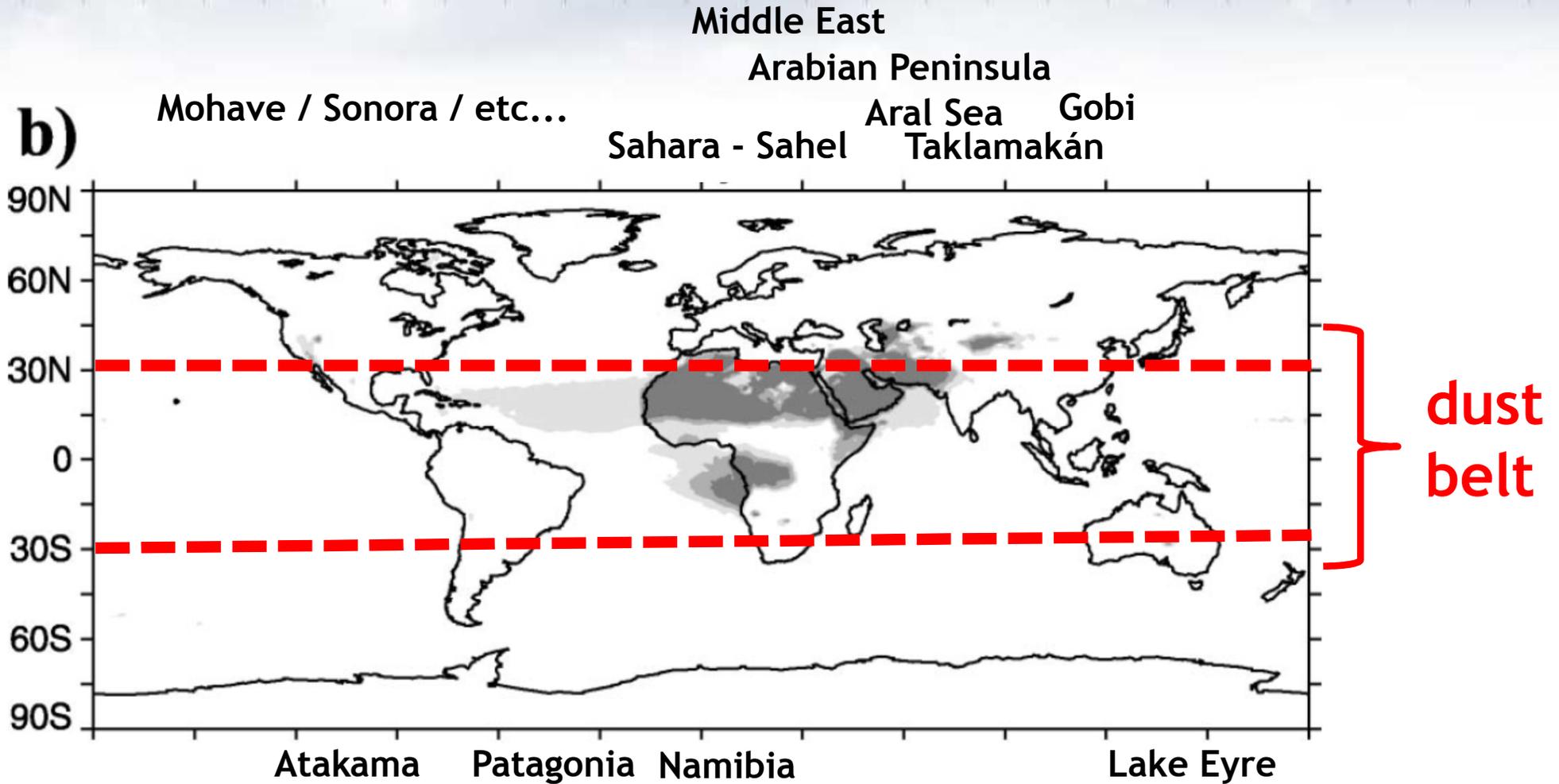
Mineral dust is one of the most important tropospheric aerosols on the global scale

The global distribution is very heterogeneous



Sahara, Sahel, Arabian Peninsula, Thar desert (Middle East), Aral Sea (Central Asia), Taklamakan desert (China), Gobi Desert (China/Mongolia), Lake Eyre Basin (Australia)

(de Graaf, 2006)



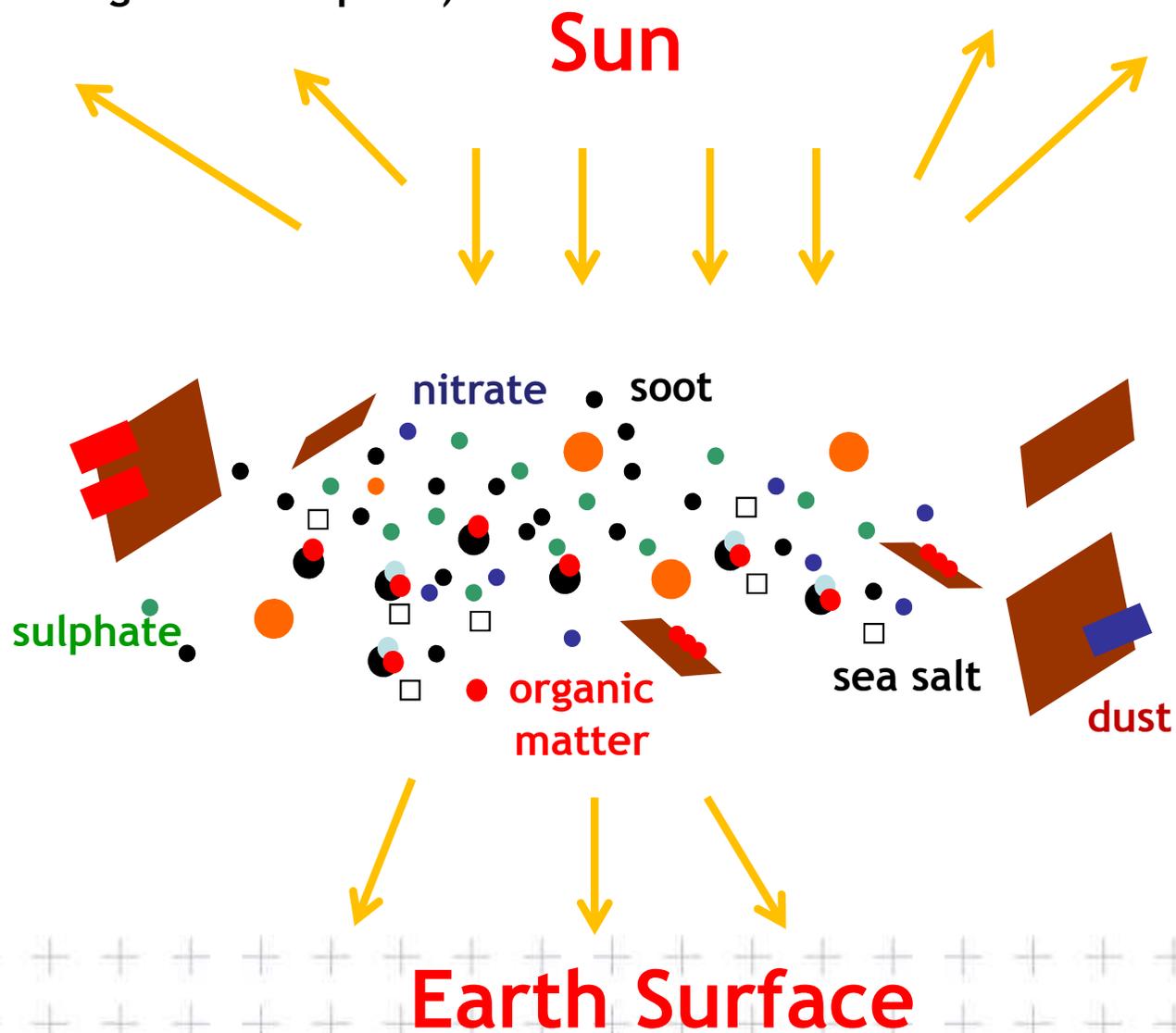
ENVIRONMENTAL CHARACTERIZATION OF GLOBAL SOURCES OF ATMOSPHERIC SOIL DUST IDENTIFIED WITH THE NIMBUS 7 TOTAL OZONE MAPPING SPECTROMETER (TOMS) ABSORBING AEROSOL PRODUCT

Joseph M. Prospero,¹ Paul Ginoux,² Omar Torres,³ Sharon E. Nicholson,⁴ and Thomas E. Gill⁵

Reviews of Geophysics, 40, 1 / February 2002

Aerosols & Climate

Direct effects: direct interaction between dust and radiation (scattering and absorption)



Aerosols & Climate

Direct effects: direct interaction between dust and radiation (scattering and absorption)

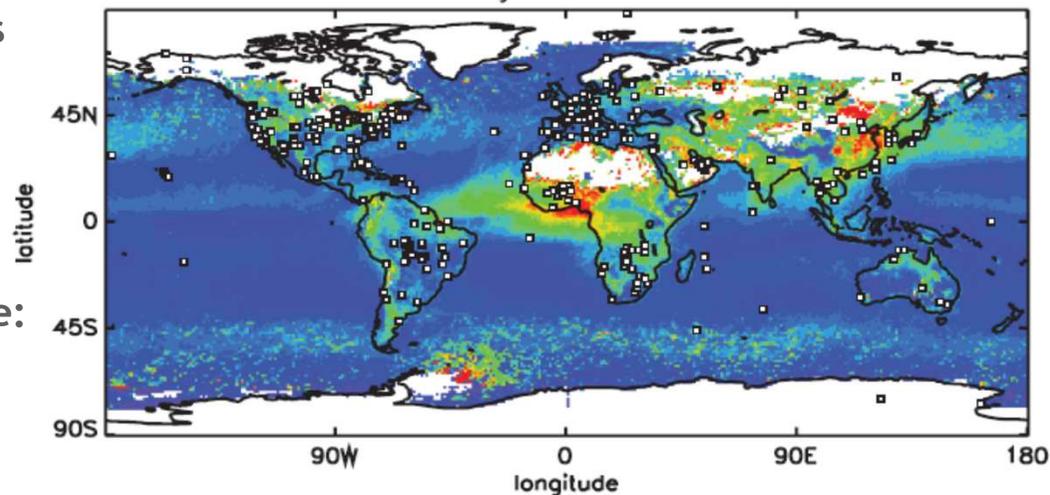
A total direct aerosol radiative forcing combined across all aerosol types can now be given for the first time as $-0.5 \pm 0.4 \text{ W m}^{-2}$, with a **medium-low** level of scientific understanding

The direct radiative forcing for individual species remains less certain and is estimated from models to be:

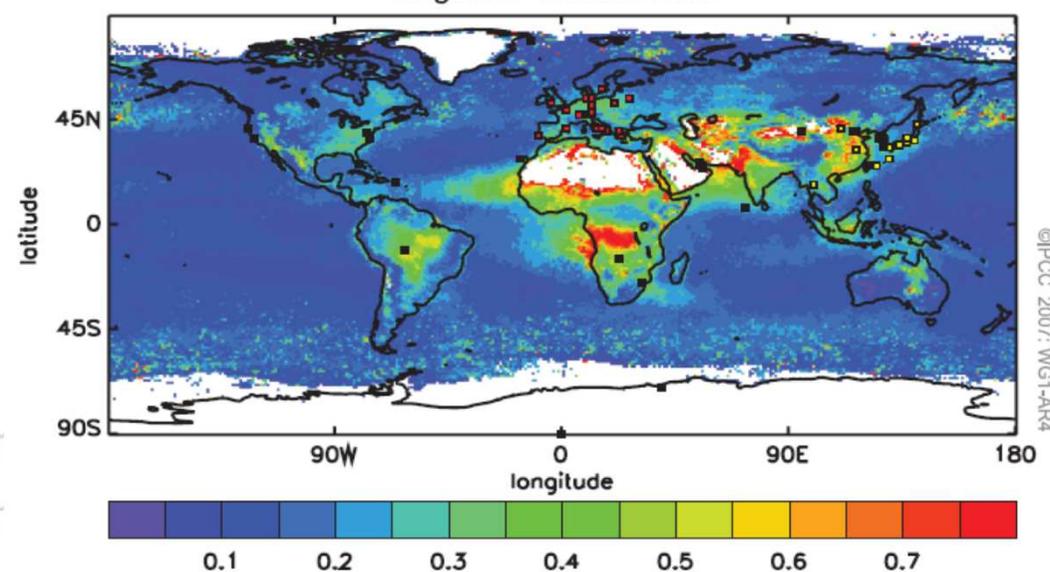
- $-0.4 \pm 0.2 \text{ W m}^{-2}$ sulphate
- $-0.05 \pm 0.05 \text{ W m}^{-2}$ fossil fuel organic carbon
- $+0.2 \pm 0.15 \text{ W m}^{-2}$ fossil fuel black carbon
- $+0.03 \pm 0.12 \text{ W m}^{-2}$ biomass burning
- $-0.1 \pm 0.2 \text{ W m}^{-2}$ for nitrate
- $-0.1 \pm 0.2 \text{ W m}^{-2}$ for mineral dust

TOTAL AEROSOL OPTICAL DEPTH

January to March 2001

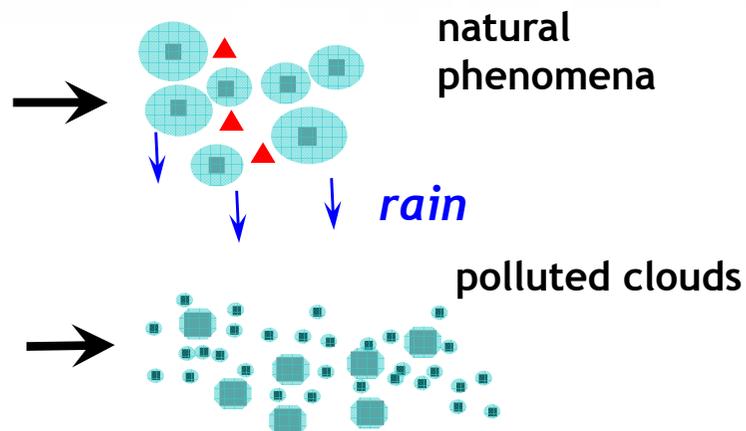
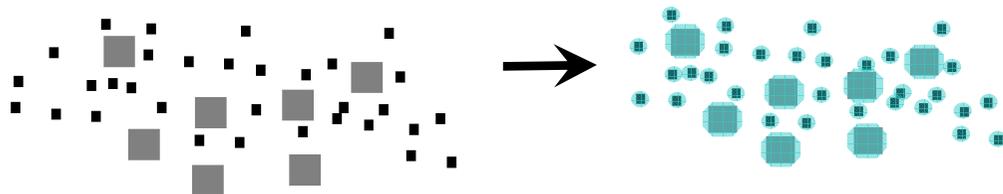


August to October 2001



Aerosols & Climate

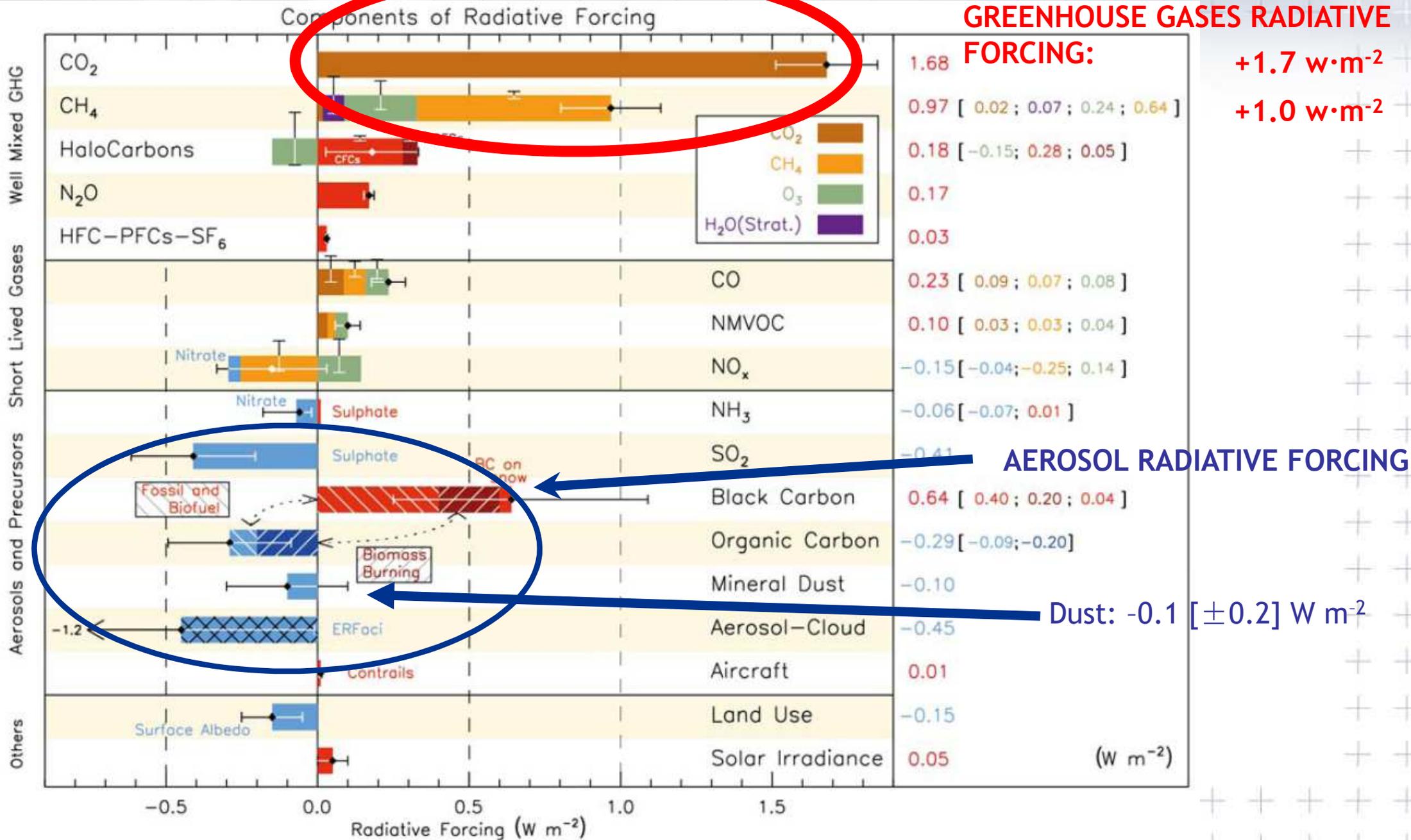
Indirect effects: change in the optical properties of clouds due to interaction with anthropogenic-aerosols



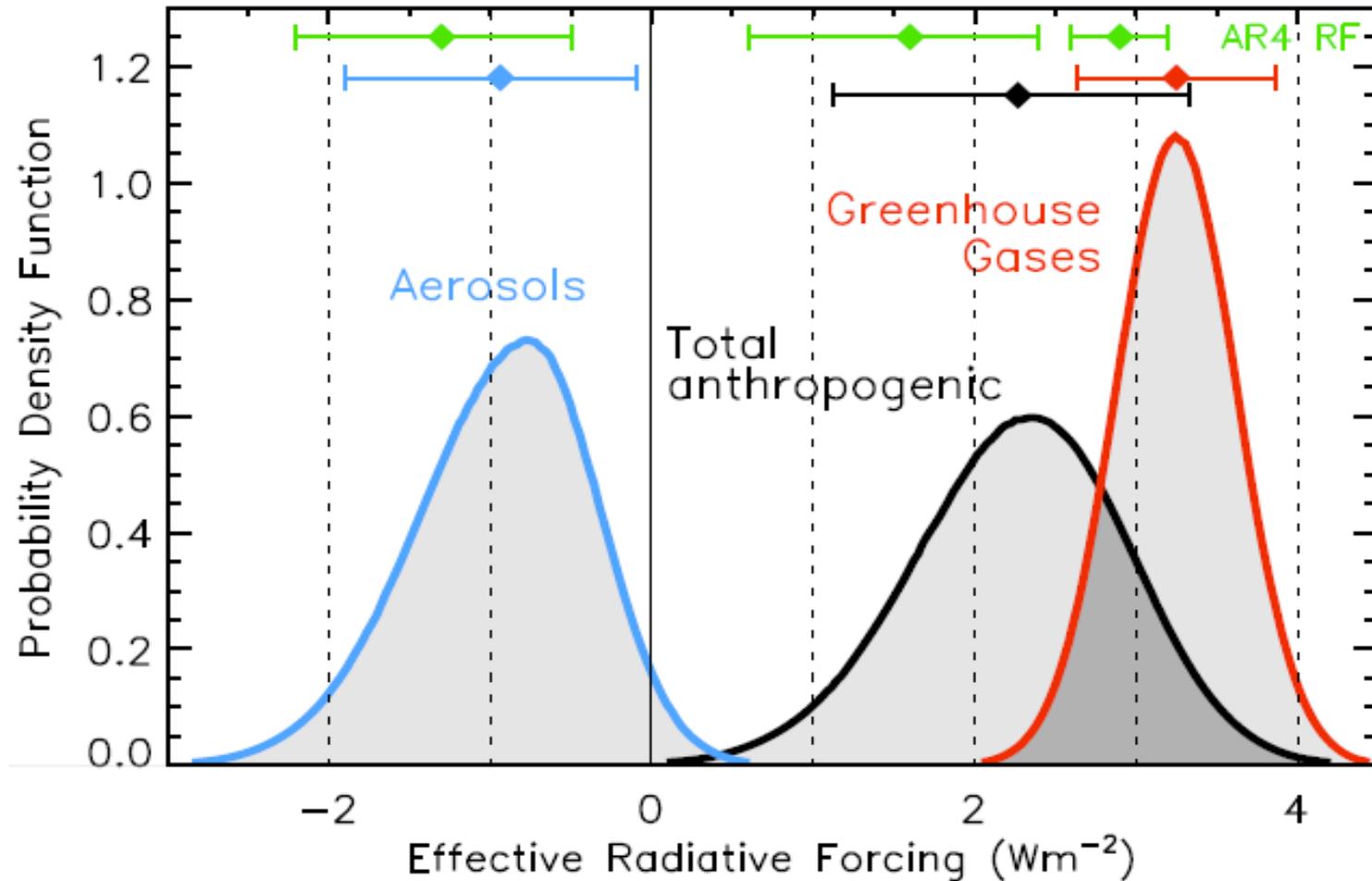
inhibited rain

{ increase in life time of clouds
change of optical properties
change of optical properties

Anthropogenic aerosols effects on water clouds cause an indirect cloud albedo effect (referred to as the first indirect effect in the TAR), which has a best estimate for the first time of -0.7 $[-0.3$ to $-1.8]$ $W m^{-2}$, with a **low level** of scientific understanding



Today's uncertainty in the total anthropogenic climate forcing is to a great extent caused by the large aerosol uncertainty



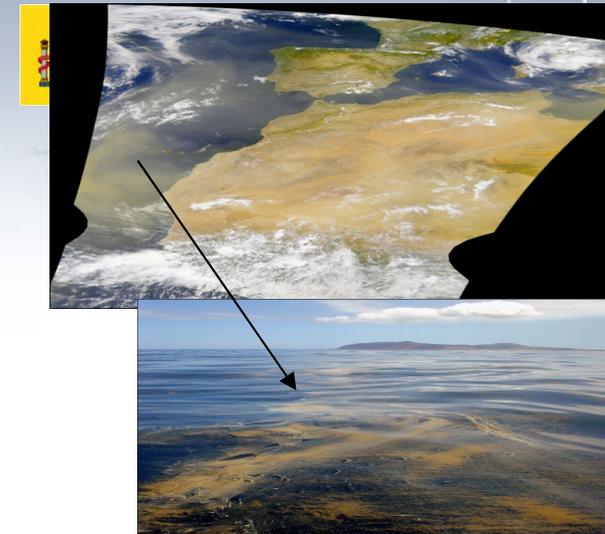
Background



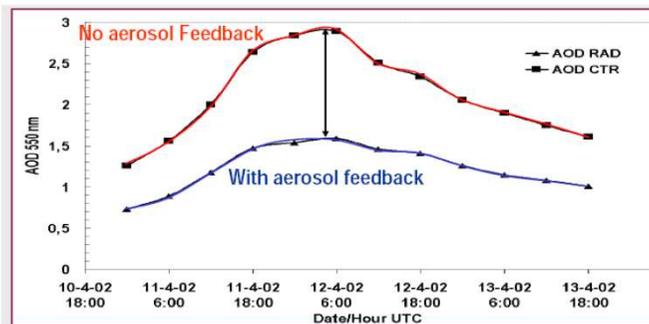
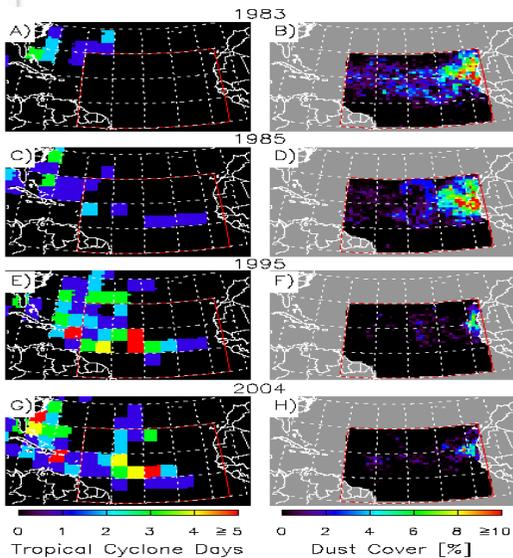
Human Health (Asthma, infections, Meningitis in Africa, Valley Fever in the America's)



Agriculture (negative & positive impacts)



Marine productivity (negative & positive impacts)



Improved Weather and Seasonal Climate prediction

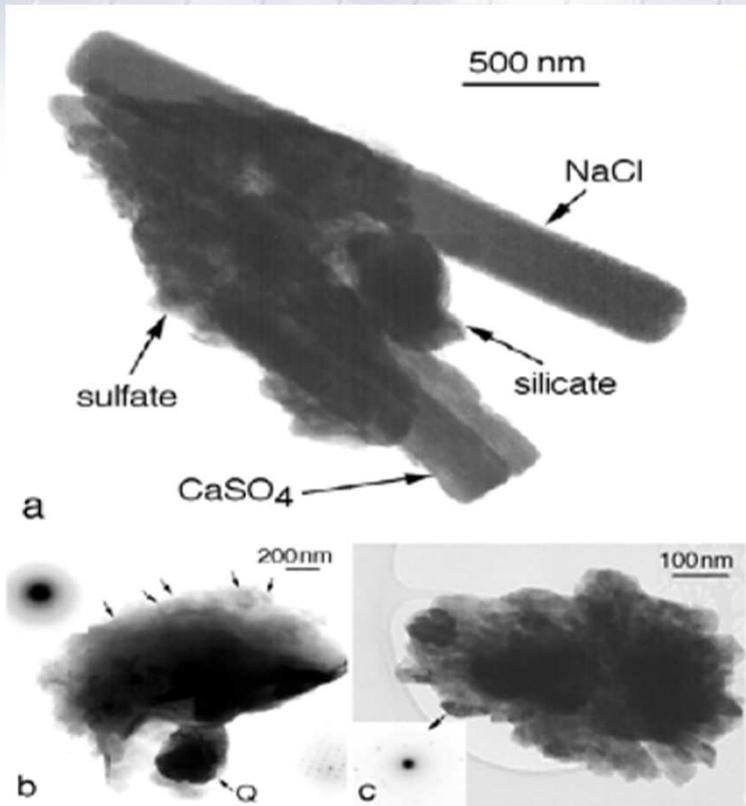
Industry (Semi-conductor, etc.)

Energy (Thermal solar energy)



Aviation (air disasters)
Ground Transportation

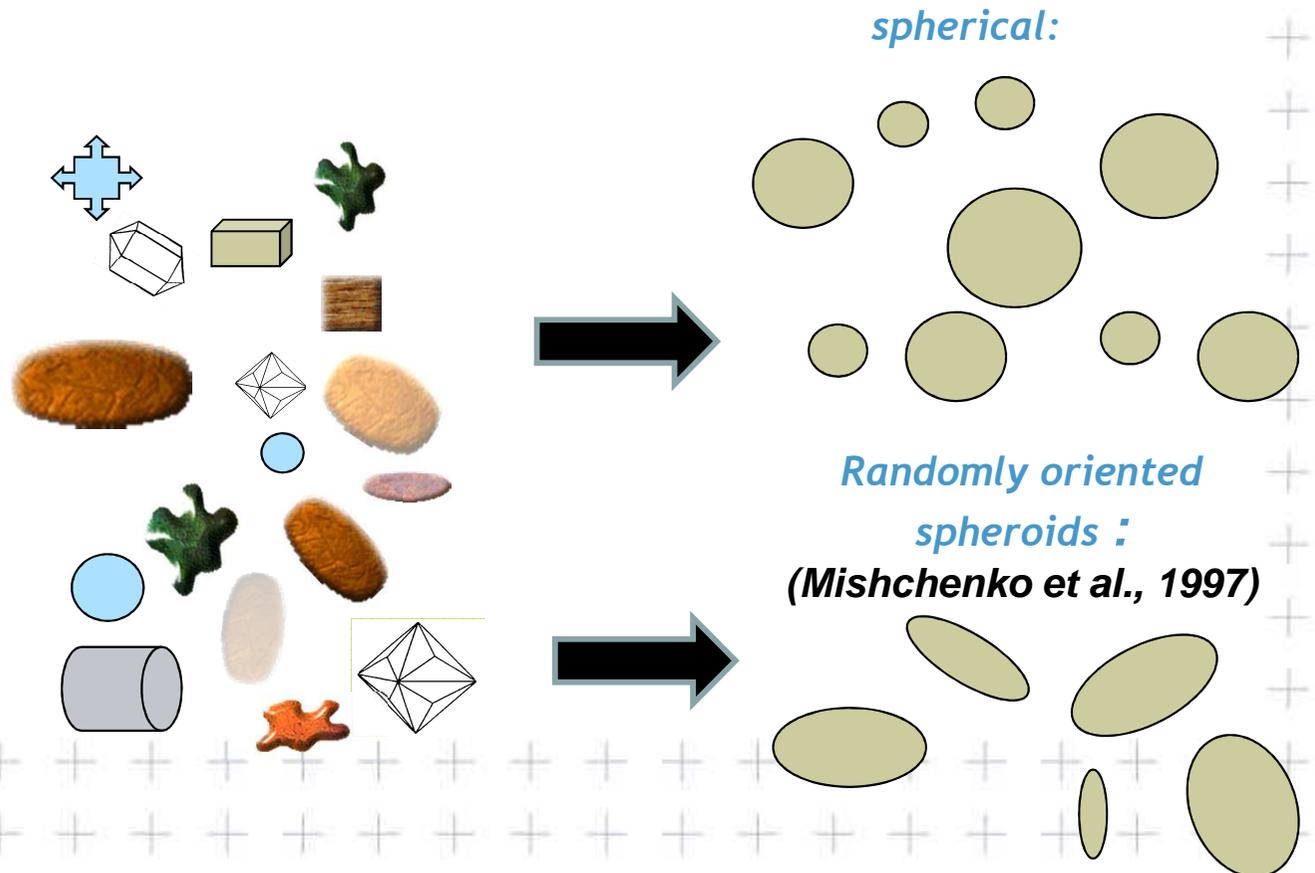




Busek and Posfai, 1996

Dust Particle Images

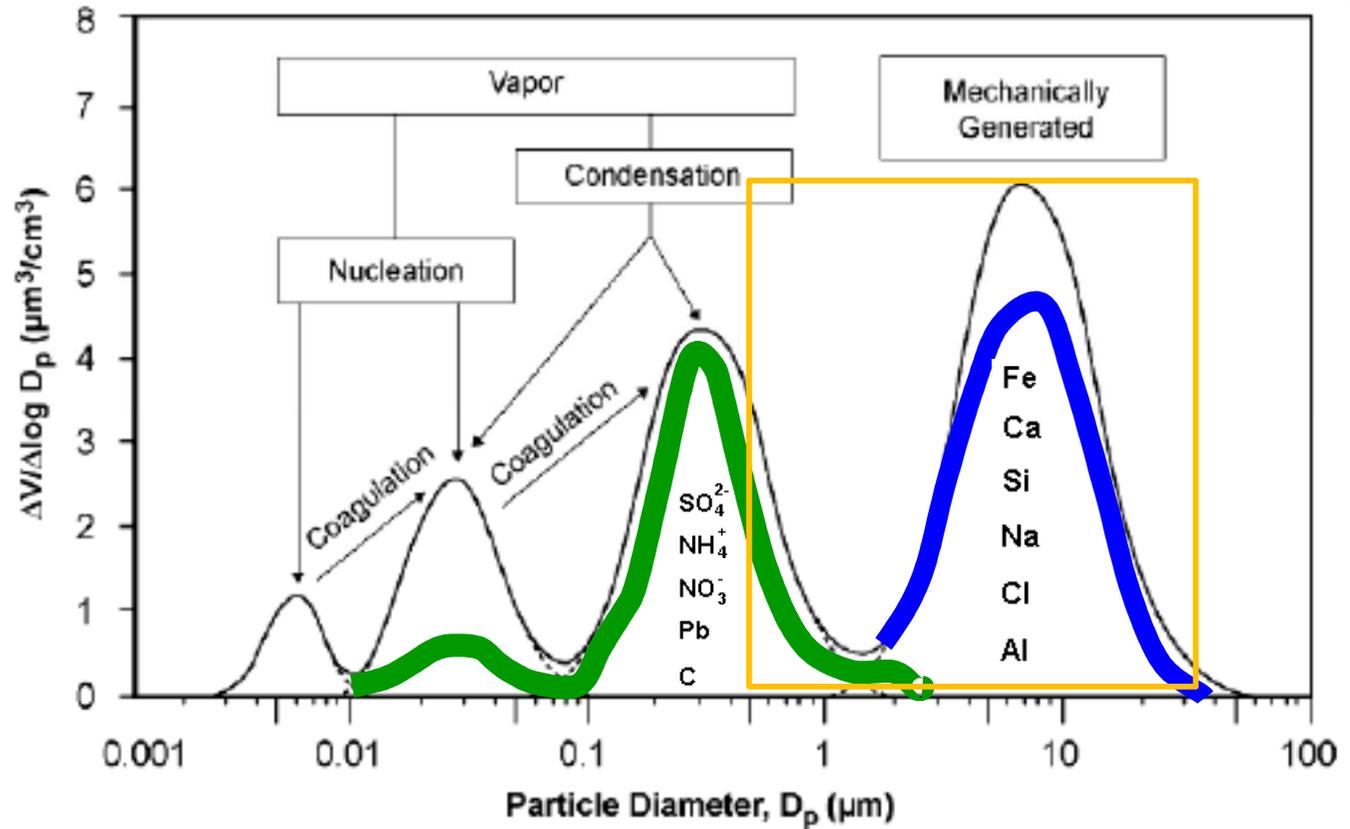
aerodynamic diameter, measured in microns or micrometers (μm), a unit equal to one millionth of a meter.



PM₁₀ (diameter <10 microm)

PM_{2.5}

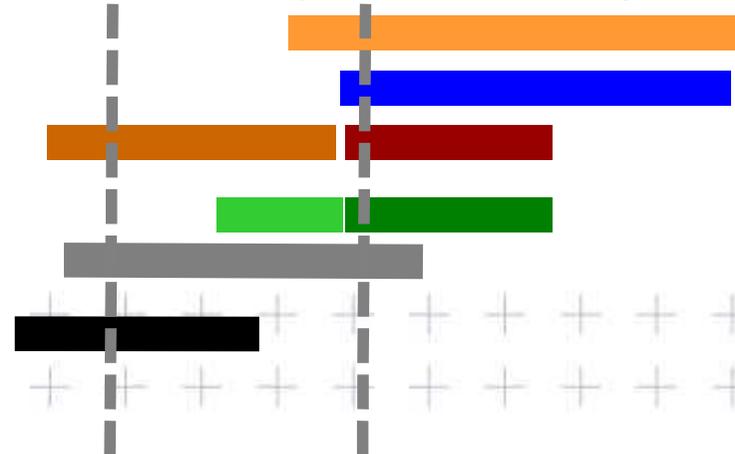
PM_{2.5-10}



ultrafine
<0.1 μm

accumulation
0.1 - 1 μm

Coarse
1 - 10 μm



Mineral dust :

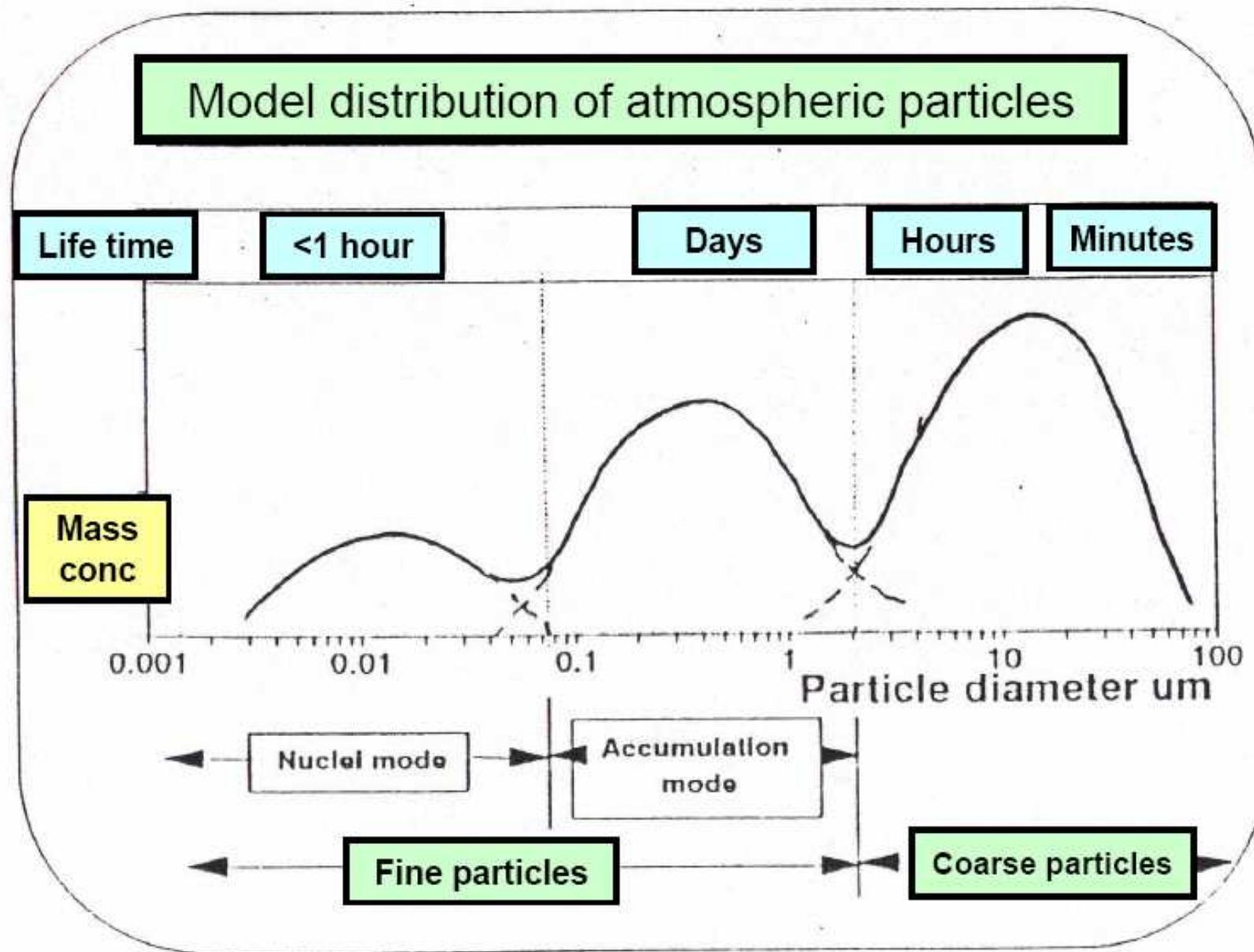
Marine salt:

Sulfate:

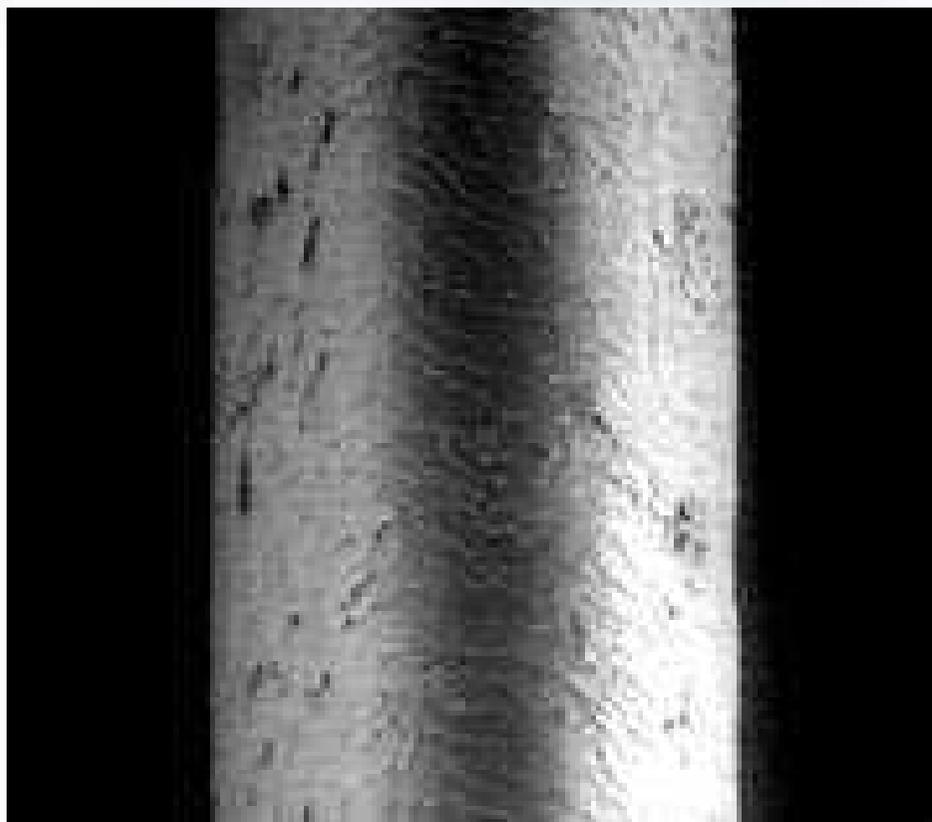
Nitrate:

Organic aerosol:

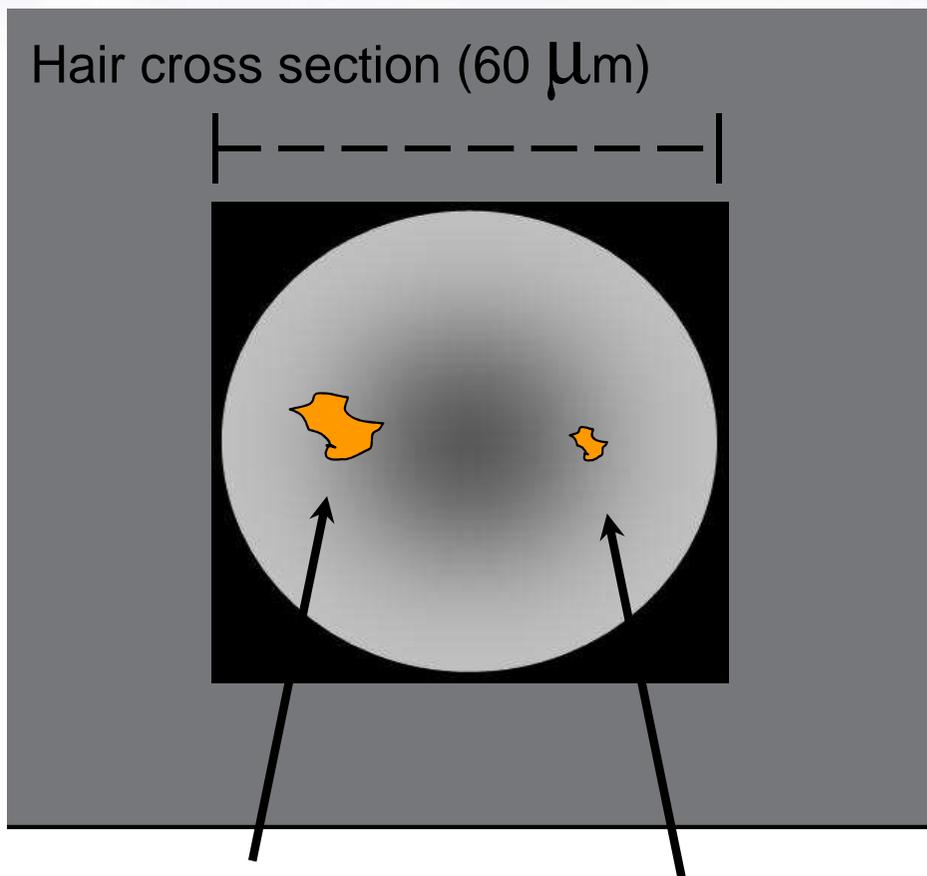
black carbon:



Particles in the Atmosphere: atmospheric residence time Model

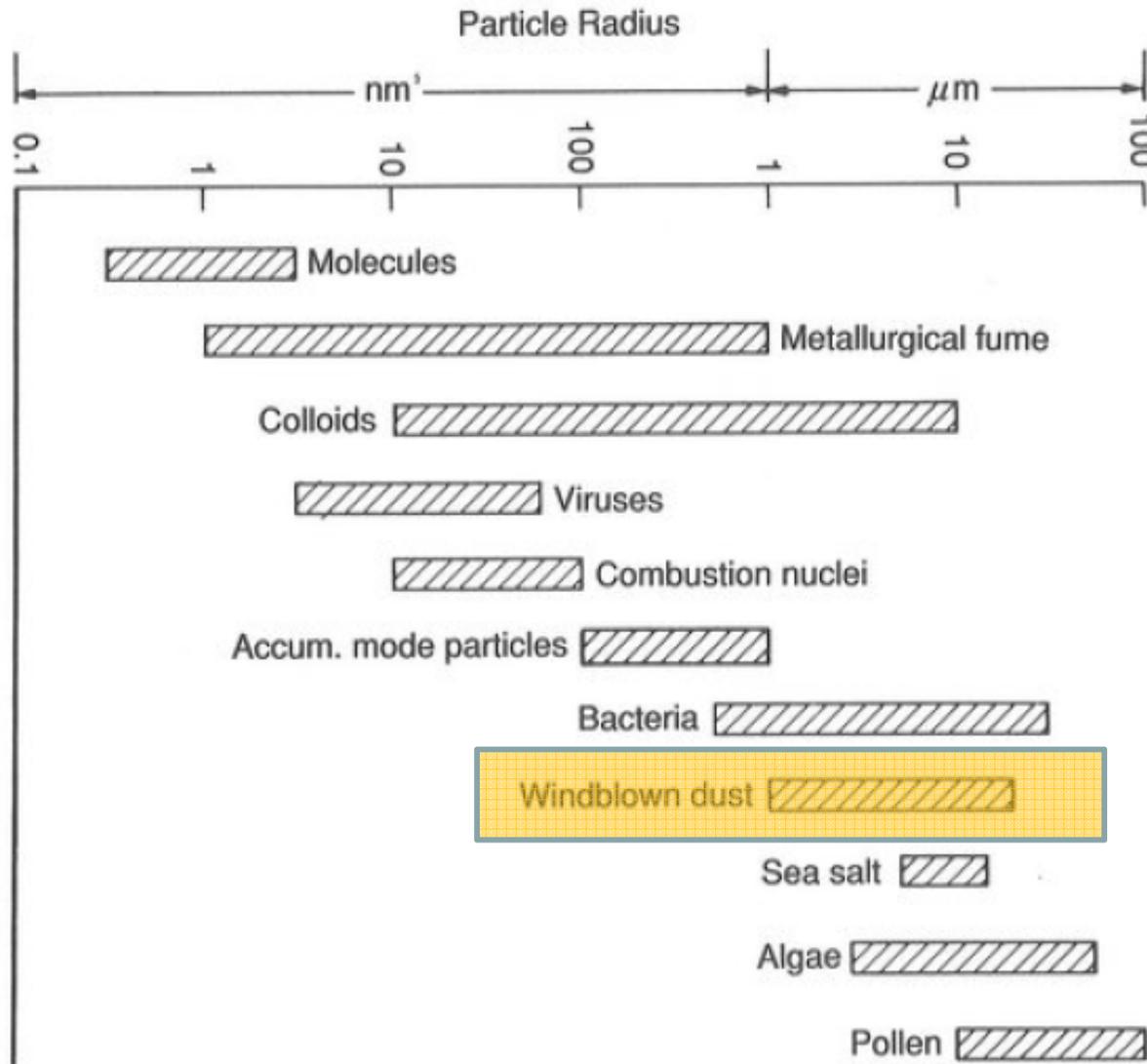


Human Hair
(60 μm diameter)



PM₁₀
(10 μm)

PM_{2.5}
(2.5 μm)



Size of different atmospheric aerosols, from (Graedel and Crutzen, 1994)

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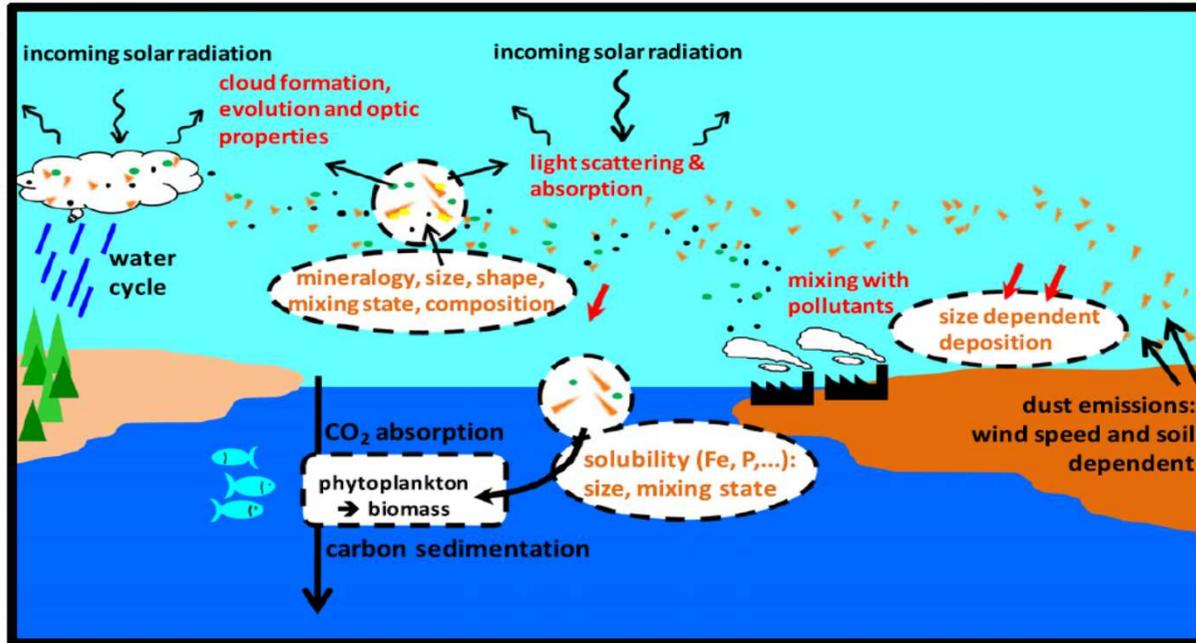
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There is a wide variety of aerosol properties that are relevant to climate forcing and human health:

List of recommended parameters by the Aerosol SAG of the Global Atmosphere Watch (GAW) program of WMO

- Multiwavelength optical depth
- Mass in two size fractions
- Major chemical components in two size fractions
- Scattering and hemispheric backscattering coefficient at various wavelengths
- Absorption coefficient
- Aerosol number concentration
- Cloud condensation nuclei (at various supersaturations)
- Aerosol size distribution
- Detailed size fractionated chemical composition
- Dependence on relative humidity
- Vertical distribution of aerosol properties (e.g. LIDAR)

What property of aerosol dust we want to measure ?



number size distribution
mass concentration
chemical composition
mixing state
mineralogy
optical properties

↓
in-situ techniques

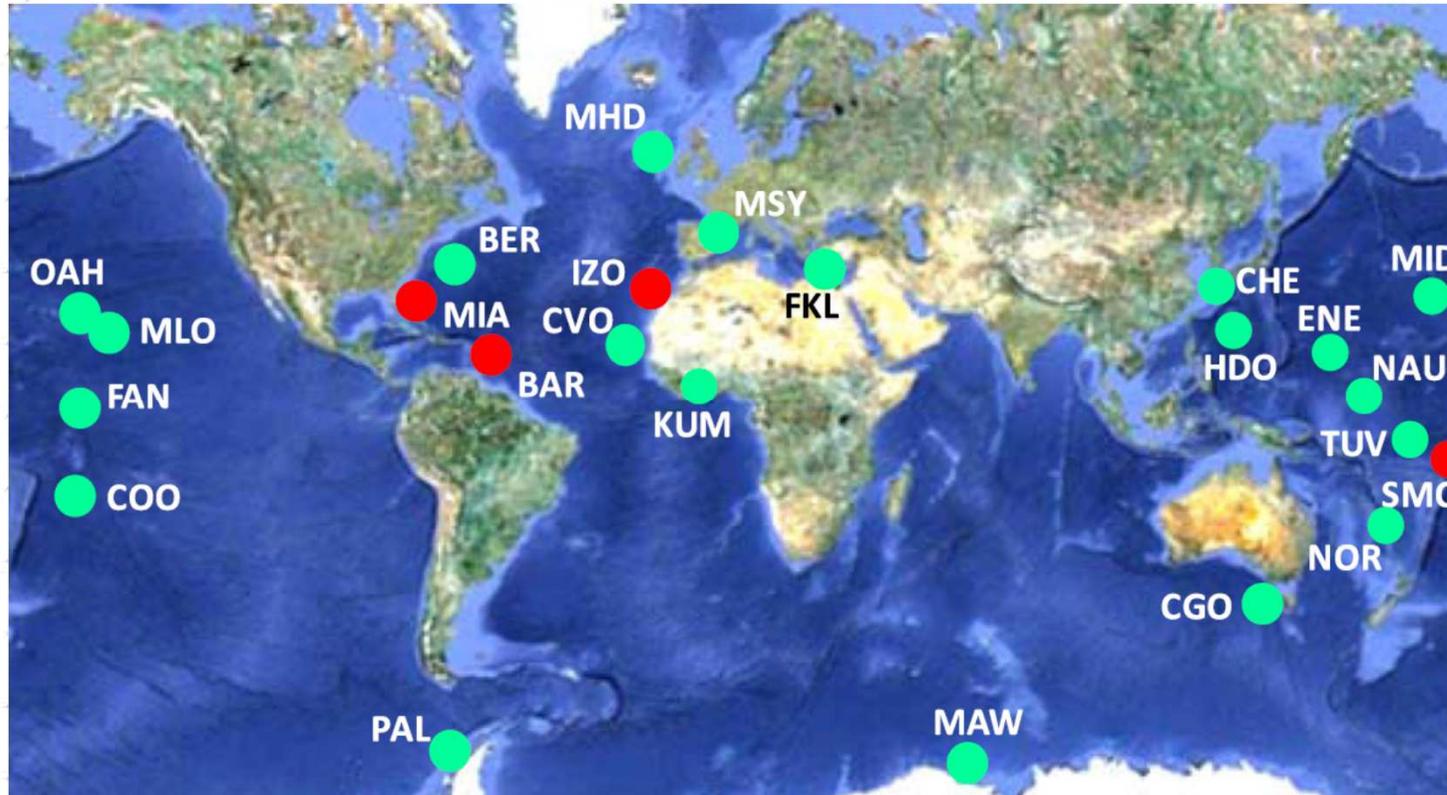
Review Article

Aeolian Research Aeolian Research 6 (2012) 55–74

A review of methods for long term in situ characterization of aerosol dust

Sergio Rodríguez^{a,*}, Andrés Alastuey^b, Xavier Querol^b

Long term monitoring dust background-observatories:



- at least 4 years
- Active during the last 20 years

Review Article

Aeolian Research Aeolian Research 6 (2012) 55–74

A review of methods for long term in situ characterization of aerosol dust

Sergio Rodríguez^{a,*}, Andrés Alastuey^b, Xavier Querol^b

property of aerosol dust:

number size distribution

mass concentration

chemical composition

mixing state

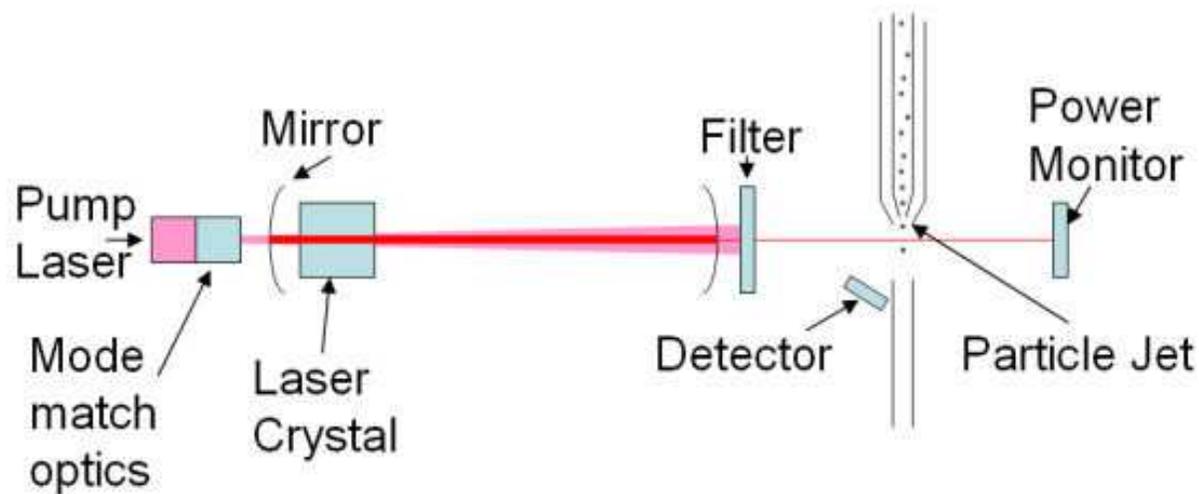
mineralogy

optical properties

property of aerosol dust: **number size distribution**

1. Optical Particle Counter OPC: 0.5 - 20 μm

Optical Particle Counters (OPC) use a high-intensity light source (a laser), a controlled air flow (viewing volume), and highly sensitive light gathering detectors (a photodetector).



Relationship between pulse height (signal in detector) and the size of the particle depends on unknown particle parameters

Particle counters count pulses of scattered light from particles, or in some cases, they count the shadows cast by backlit particles. The amount of light a particle scatters, can vary with several different factors, including the following:

- The shape of the particle:
- The albedo (reflectivity) of the particle:

OPC are very useful instruments, but sources of uncertainties should be known:

Particle size (?) \rightarrow diameter of the calibration polystyrene spheres (PLS)

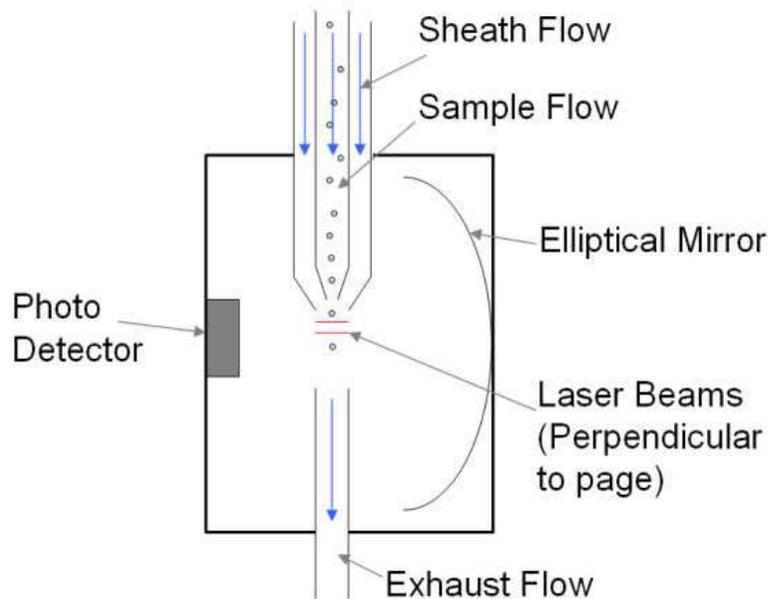
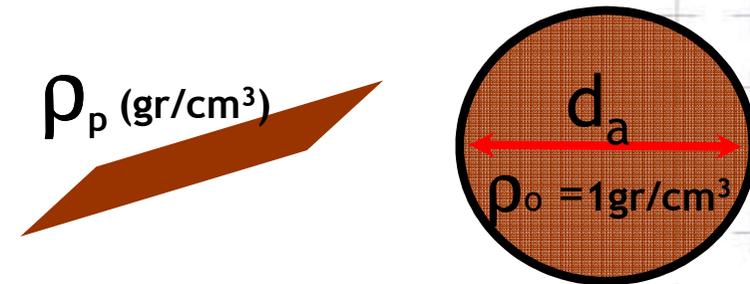
property of aerosol dust: **number size distribution**

1. Optical Particle Counter OPC

2. Aerodynamic Particle Sizer: 0.7 - 20 μm

The Aerodynamic Particle Sizer (APS) uses the principle of inertia to size particles.

Particles within the airflow are accelerated, but by different amounts depending on particle surface area and mass, thus particles exiting the jet have a velocity related to their aerodynamic diameter. Aerodynamic diameter is defined assuming spherical particles and unity density.



The APS measures particle velocity by passing the particles through two laser beams separated by about 200microns. An elliptical mirror collects scattered light onto a photodetector.

A particle passing through both beams produces two pulses of scattered light, the time delay between the pulses being related to the velocity and hence aerodynamic diameter of the particle. The APS also records the height of the peaks allowing a secondary calculation of particle size based on optical scattering.

property of aerosol dust: **number size distribution**

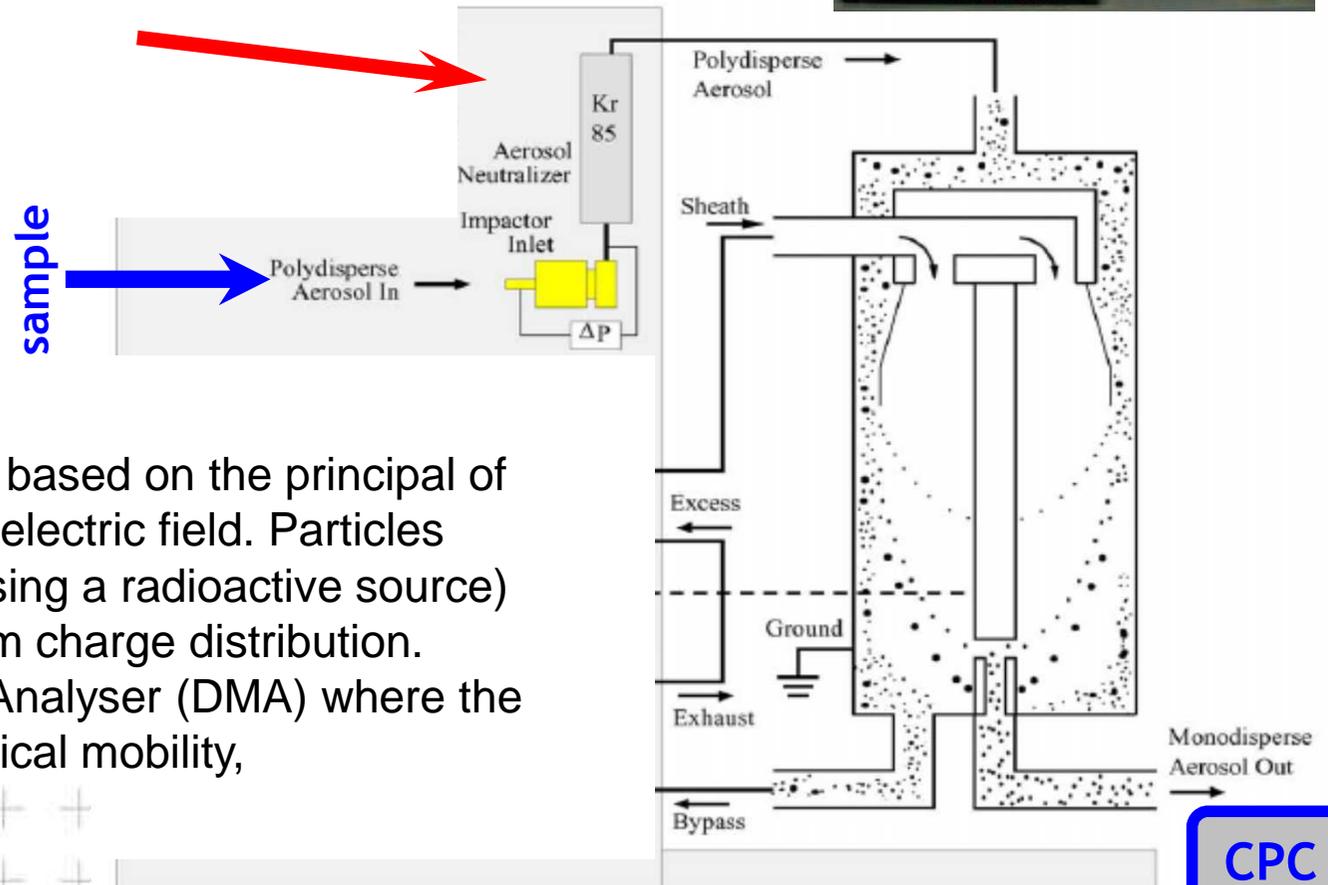
1. Optical Particle Counter OPC: 0.5 - 20 μm

2. Aerodynamic Particle Sizer: 0.5 - 20 μm

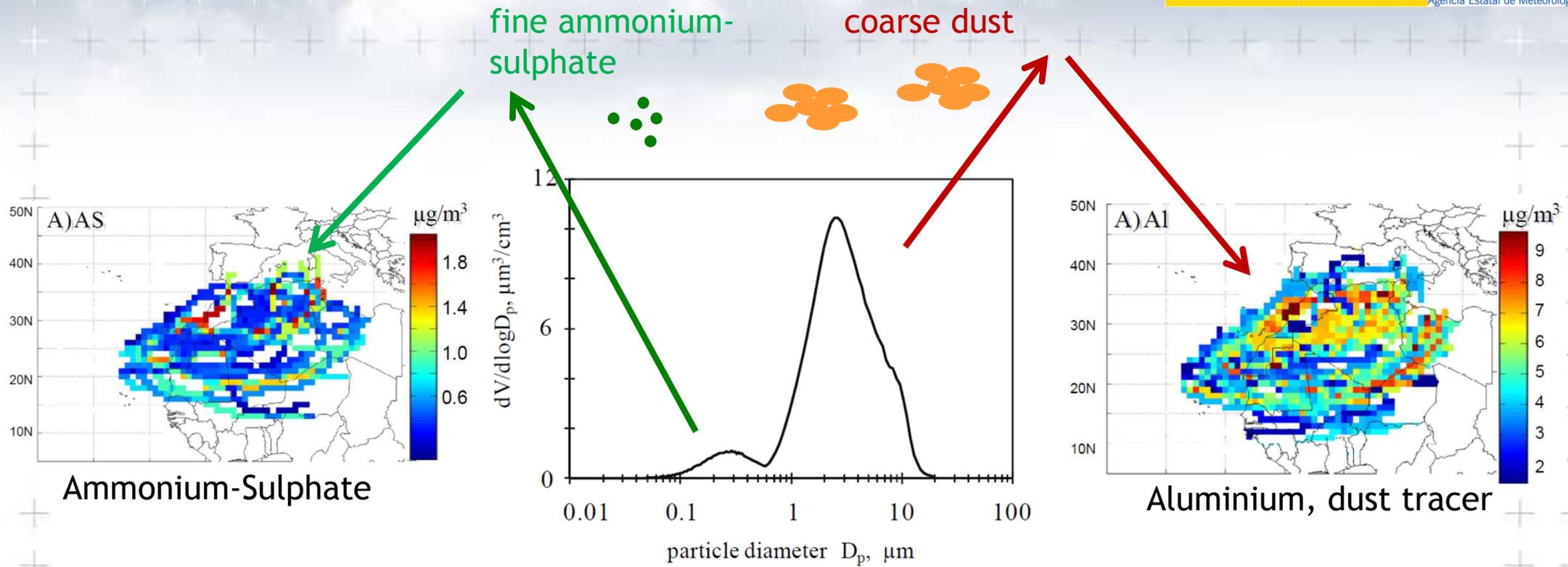
3. Scanning Mobility Particle Sizer: 3 nm - 1 μm



1. Neutralizer: known charge distribution



The Scanning Mobility Particle Sizer is based on the principle of the mobility of a charged particle in an electric field. Particles entering the system are neutralized (using a radioactive source) such that they have a Fuchs equilibrium charge distribution. They then enter a Differential Mobility Analyser (DMA) where the aerosol is classified according to electrical mobility,



Scanning Mobility Particle Sizer

Aerodynamic Particle Sizer

Disadvantage of particle sizers (OPC, APS SMPS): cannot differentiate dust from other particles

property of aerosol dust:

number size distribution

mass concentration

chemical composition

mixing state

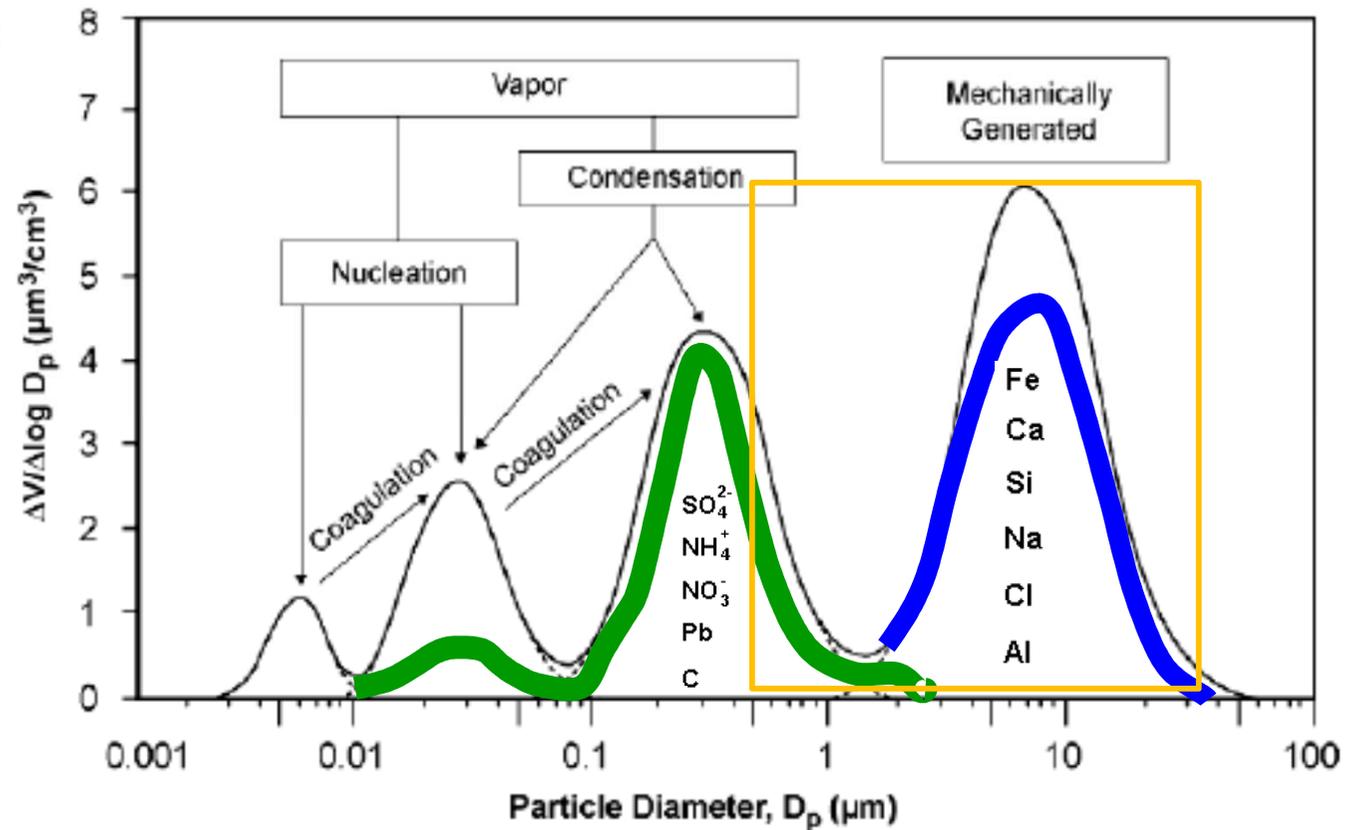
mineralogy

optical properties

PM₁₀ (diameter <10 microm)

PM_{2.5}

PM_{2.5-10}



ultrafine
<0.1 μm

accumulation
0.1 - 1 μm

Coarse
1 - 10 μm

Mineral dust :

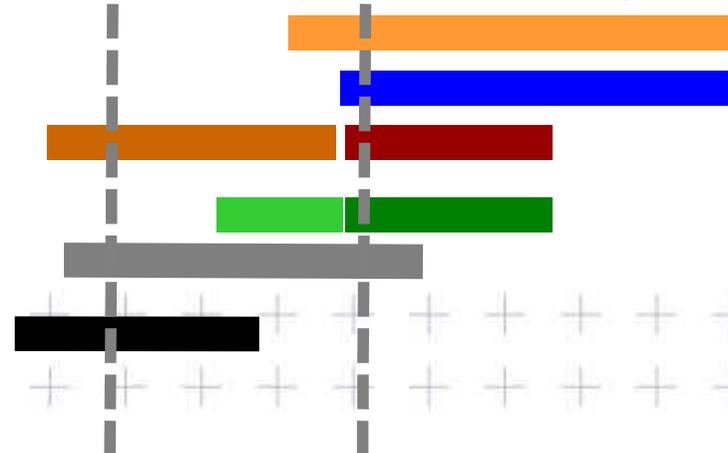
Marine salt:

Sulfate:

Nitrate:

Organic aerosol:

black carbon:



property of aerosol dust: **mass concentration**

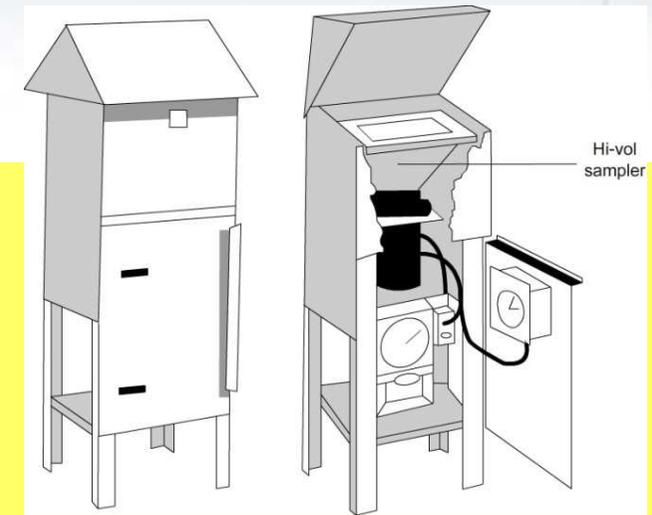
bulk aerosol mass concentration

1. Reference method: gravimetric method
2. Automated analyzers

PM₁₀ and PM_{2.5} measurements in air quality networks

1. Reference method: gravimetric method

$$PM = \frac{(W2 - W1)}{\text{Volume}} \mu\text{g}/\text{m}^3$$

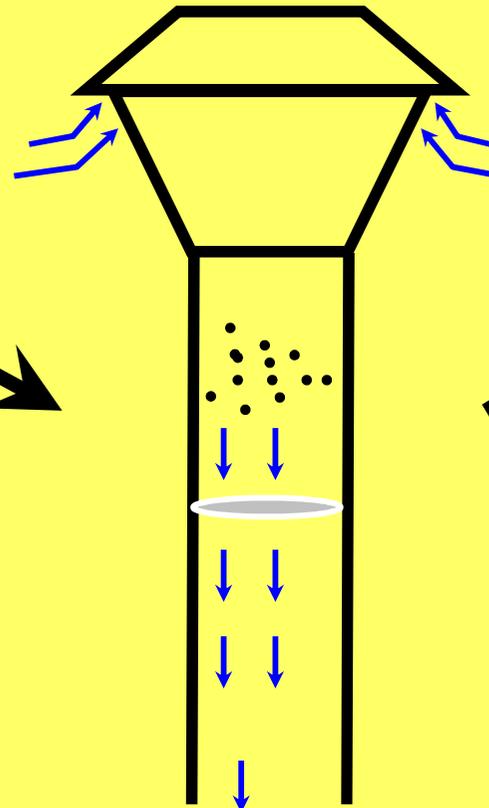
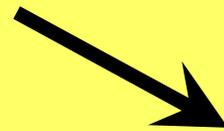


Blank filter

Conditioning

RH (50±5%) y T(20±1°C) 24-h

- Filter weight (W1)



Pump



Sampled filter

Conditioning

RH (50±5%) y T(20±1°C) 24-h

- Filter weight (W2)



Common Gravimetric Ambient Aerosol Sampling Techniques

- High volume methods: TSP, PM_{10} , $PM_{2.5}$
- Low volume methods: (PM_{10} , $PM_{2.5}$, PM_{Coarse})

Micro-Balance room



- Filters conditioning 48-h, HR=50±5 % and T=20±1°C
- balance, LVS resolution ≥ 5 digits (0.00001g)
- balance, HVS resolution ≥ 6 digits (0.000001g)

This sample filter is equilibrated at some set of thermodynamic conditions for a period of time before and after sampling. Through the use of a laboratory gravimetric balance, the difference in pre- and post-sample weights yields the PM mass collected. Knowing the volume of air passed through the filter allows the determination of the PM mass concentration.



PM₁₀ and PM_{2.5} measurements in air quality networks

1. Reference method: gravimetric method

Low Volume Sampler

LVS: **2.3 m³/h**



High Volume Sampler

HVS: **68 m³/h**

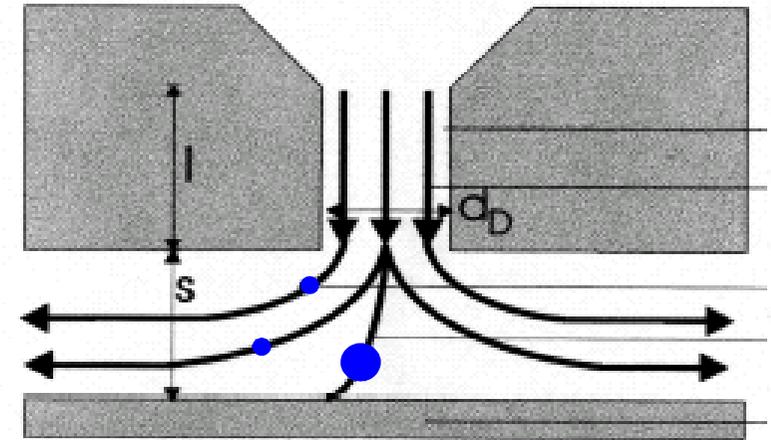
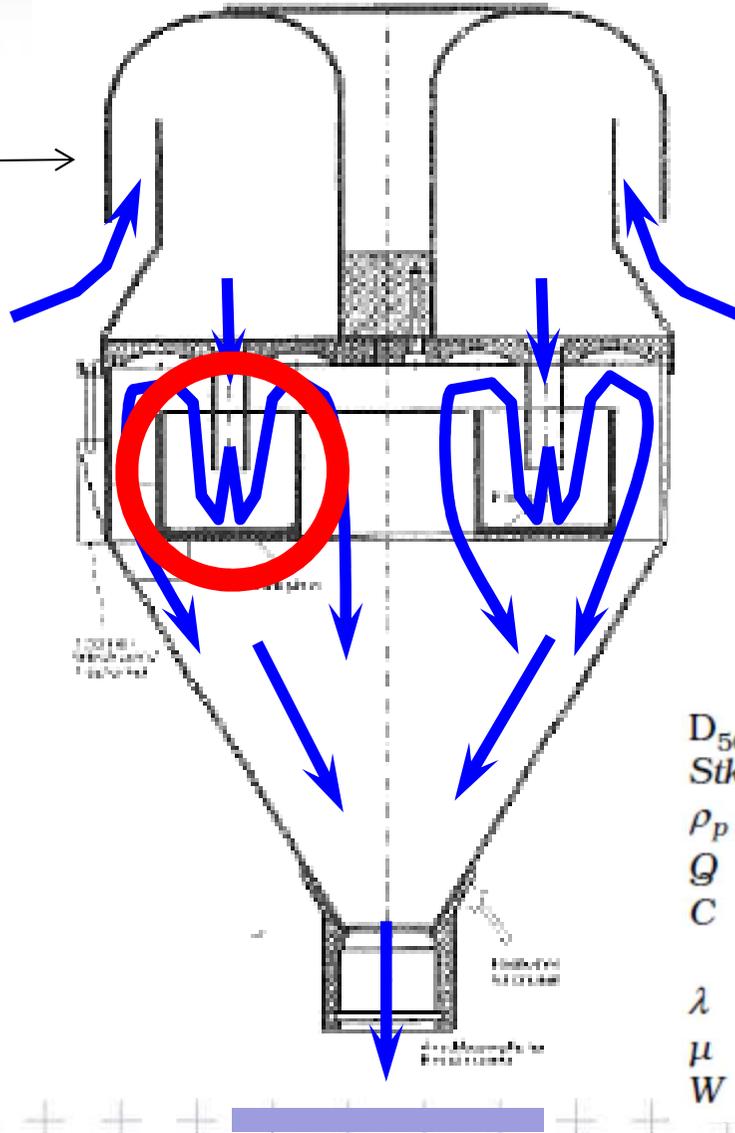


HVS: **30 m³/h**



TSP, PM₁₀, PM_{2.5}, PM₁:

aerodynamic diameter (as the APS)



$$D_{50} = \sqrt{\frac{9\pi Stk \mu W^3}{4\rho_p C Q}}$$

- D_{50} = particle cut-point diameter centimeter
- Stk = Stokes number = 0.23
- ρ_p = particle density (g/cm³)
- Q = volumetric flow rate (cm³/s)
- C = Cunningham slip correction
 $= 1 + 2.492 \lambda/D_{50} + 0.84 \lambda/D_{50} \exp(-0.435 D_{50}/\lambda)$
- λ = gas mean free path
- μ = gas viscosity (dyne•s/cm²)
- W = nozzle diameter (cm)

The Stokes number is a dimensionless parameter that characterizes impaction.

Filter

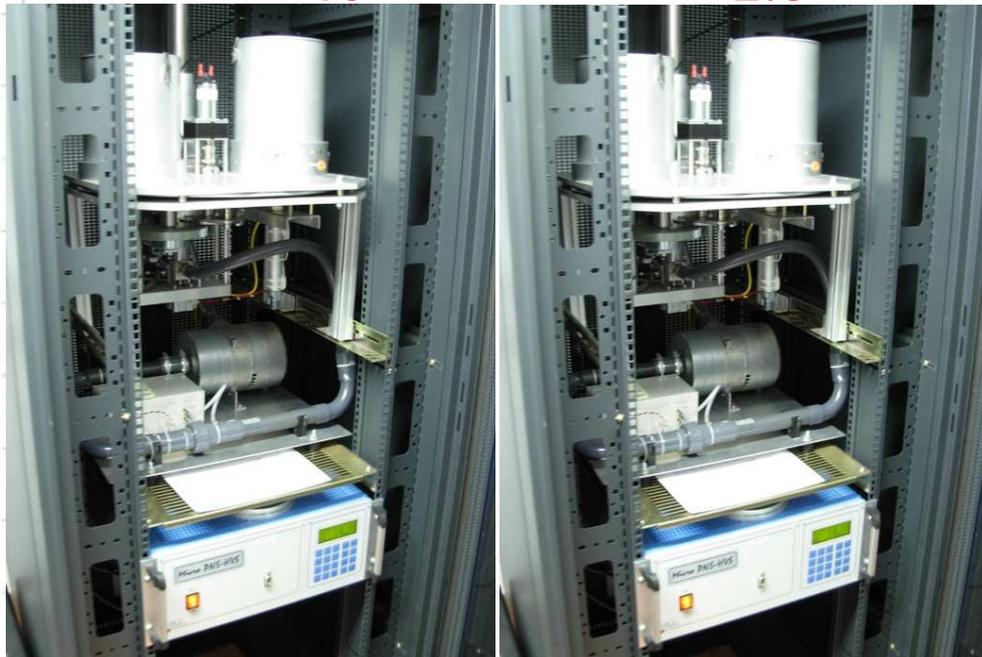
Complete PM gravimetric method set-up at Izana Atmospheric Research Center

Weight filters conditioned room



PM₁₀

PM_{2.5}



Common Gravimetric Ambient Aerosol Sampling Techniques

- Advantages: Recognized reference method, low capital cost
- Disadvantages: Limited time resolution (typically 24-hr), long turnaround times, labor intensive, and gravimetric lab maintenance/cost

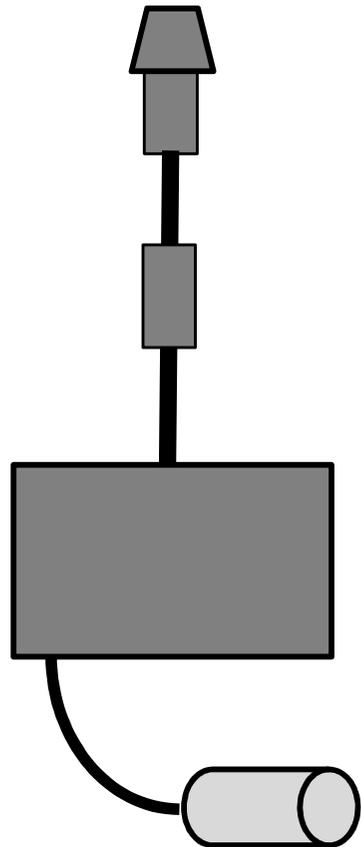
Common Continuous Ambient Aerosol Sampling Techniques

$$(Dm / Dt) / (DV / Dt) = \text{mg/m}^3$$

- Tapered Element Oscillating Microbalance
- Beta (Electron) Attenuation

PM₁₀ and PM_{2.5} measurements in air quality networks

2. Automated analyzers



1. Impactor PM₁₀ / PM_{2.5}

2. RH reductor / heater

3. Sensor (Beta radiation attenuation or Tapered Oscillating microbalance-TEOM-) → instead of weighting filters

4. Pump / Flow meter

Continuous measurements of PM (PM₁₀, PM_{2.5}, PM₁ or TSP)

Mass concentration

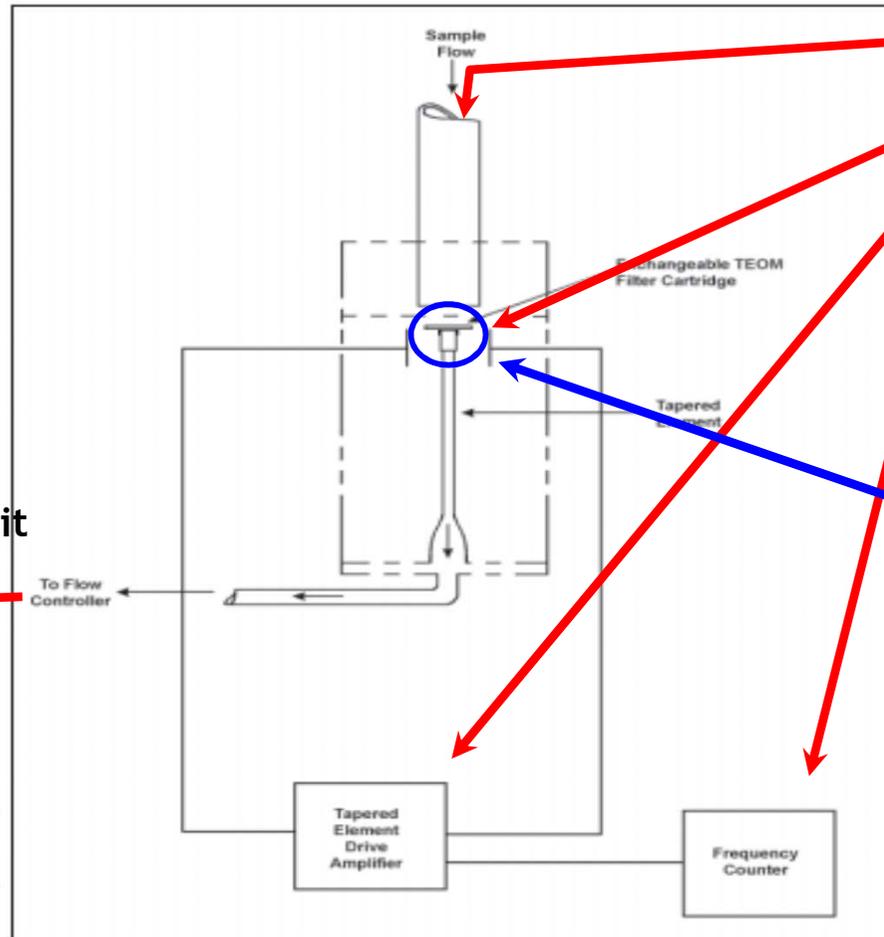
Automatic continuous measurements

TEOM : Tapped Element Oscillating Microbalance

1. TEOM mod. 1400a

mass=function (frequency)

sensor



Sampling flow rate (16.67 l/m)

Sample accumulated in the filter

Micro-oscillation of constant amplitude
GENERATOR

Frequency sensor

An increase in the amount of sample (dust) accumulated in the filter → decrease in the oscillation frequency

Mass concentration

Automatic continuous measurements

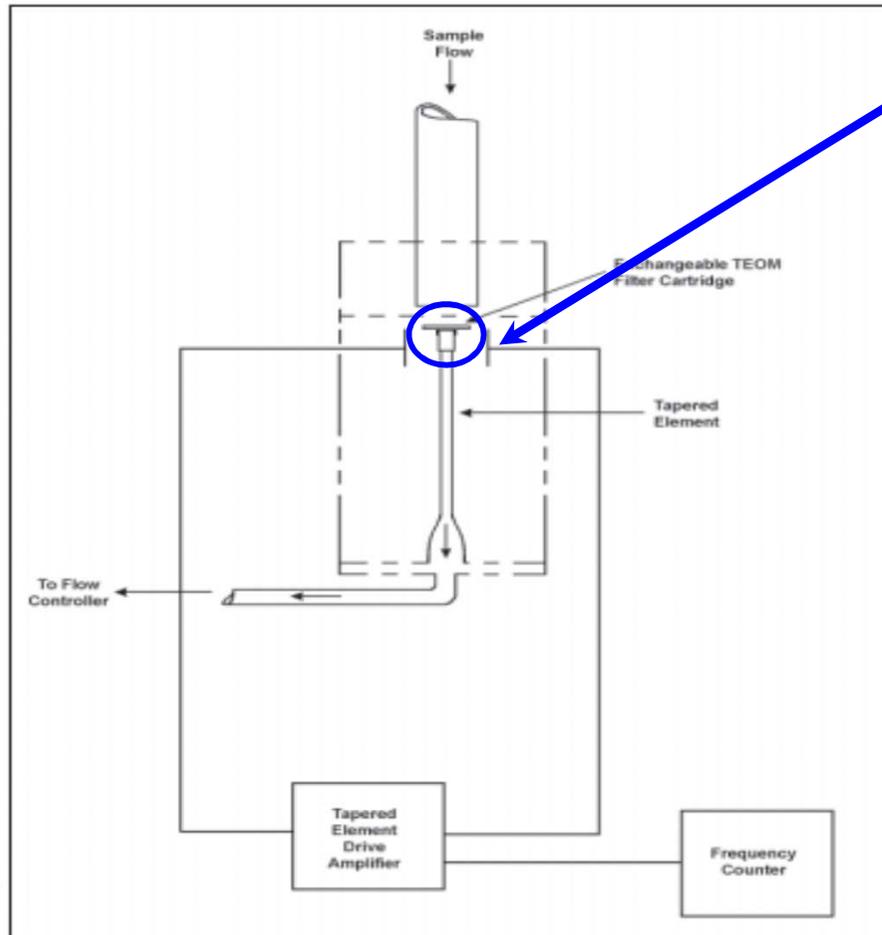
TEOM :Tapped Element Oscillating Microbalance

1. TEOM mod.1400a

sensor

mass=function (frequency)

more dust → lower oscillation frequency



In a spring-mass system the frequency follows the equation:

$$f = (K / M)^{0.5}$$

where:

- f = frequency (radians/sec)
- K = spring rate
- M = mass

K and M are in consistent units. The relationship between mass and change in frequency can be expressed as:

$$dm = K_0 \left(\frac{1}{f_1^2} - \frac{1}{f_0^2} \right) \quad (2)$$

where:

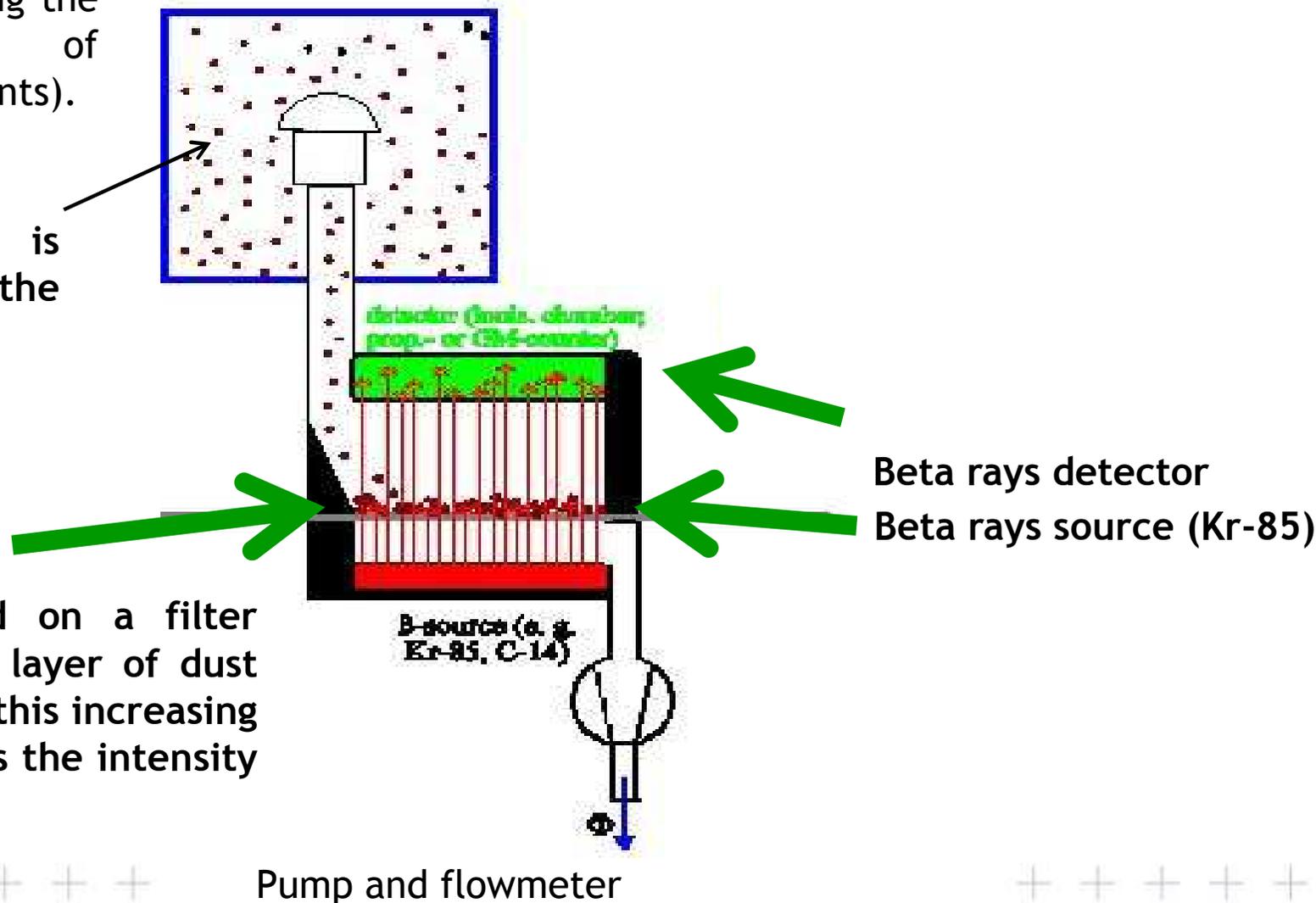
- dm = change in mass
- K₀ = spring constant (including mass conversions)
- f₀ = initial frequency (Hz)
- f₁ = final frequency (Hz)

Krypton-85 or Carbon-14 is used as source of beta radiation (emitted by electrons during the nuclear decay of radioactive elements).

Ambient air is drawn through the sample system

Dust is deposited on a filter continuously. The layer of dust is building up and this increasing dust mass weakens the intensity of the beta beam.

**Beta Attenuation:
 β -Ray Absorption in Matter**

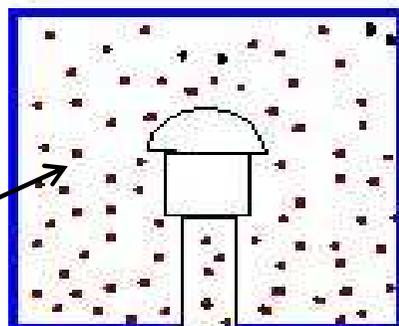


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Ambient air is drawn through the sample system

Dust is deposited on a continuously. The layer of is building up and this incre dust mass weakens the inte of the beta beam.

**Beta Attenuation:
β-Ray Absorption in Matter**



$$I = I_0 e^{-\mu_\beta \cdot x}$$

μ_β is the mass absorption coefficient for beta radiation
 x mass thickness of the sample

$x = f(\text{atomic number to atomic mass ratio } (Z/A))$
 Z/A (C, Si, Al, Ca, Fe, Mg, K, Cl, Na, N, O and S) 0.47–0.50

Standard foil calibration

typical elements of aerosols; fixed Z/A ratio: error of about 10%

Pump and flowmeter

PM with Beta attenuation (2)

$$m = F_{cal} \ln \left(\frac{I_0}{I} \right)$$

- **m**: increasing particle mass [μg]
- **F_{cal}**: calibration factor
- **I₀** beta ray intensity at empty filter
- **I** beta ray intensity at loaded filter

The intensities I_0 and I are measured with the detector system. F_{cal} has to be measured directly during the calibration procedure. This is accomplished by replacing the filter with the element having a known mass (mass calibration kit)

The mass concentration is calculated from:

$$c = \frac{m}{Ft}$$

Where:

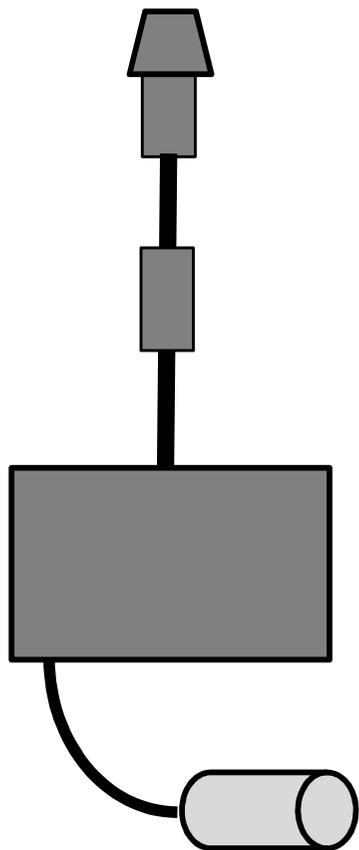
c: concentration [$\mu\text{g}/\text{m}^3$]

F: measured air flow [m^3/h]

t: time [h]

PM₁₀ and PM_{2.5} measurements in air quality networks

2. Automated analyzers



beta

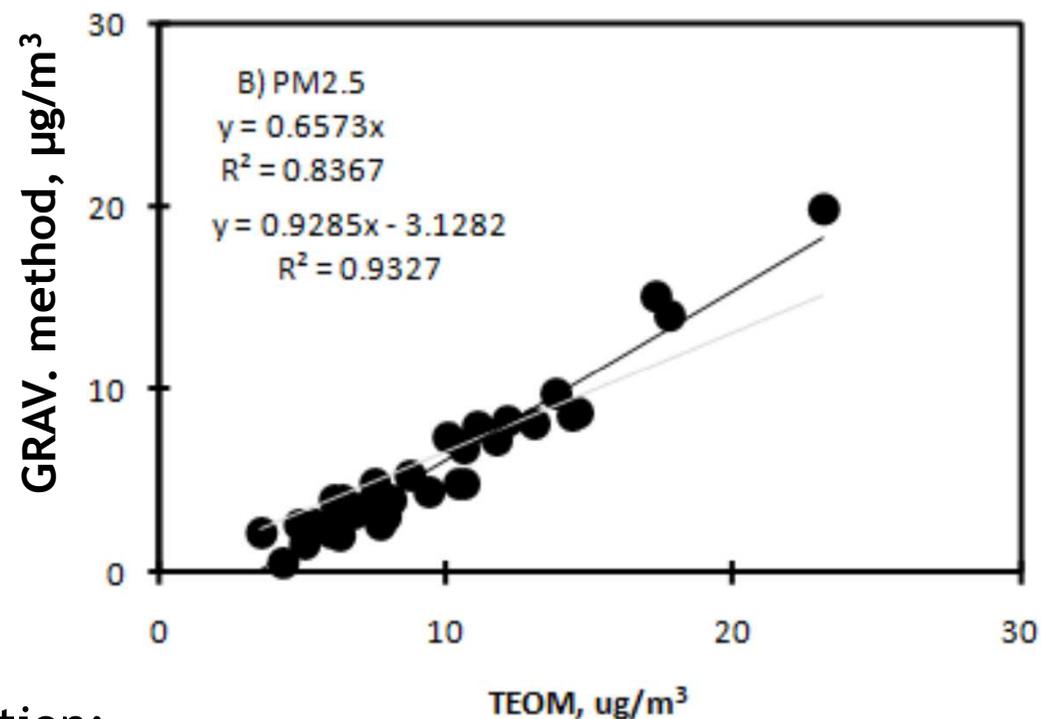
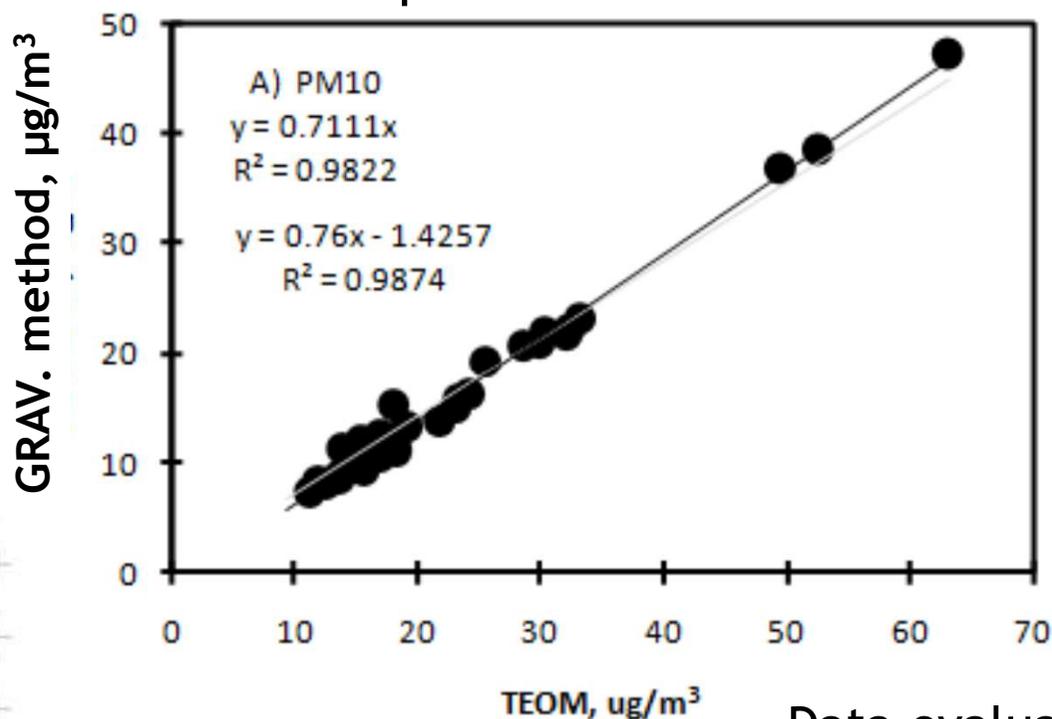


TEOM

Automatic versus the reference gravimetric method

Conversion of the 'automatic PM₁₀ and PM_{2.5}' data to GRAVIMETRIC EQUIVALENT data

Intercomparisons



Data evaluation:

Data from continuous analyzer are valid if they fit A or B:

A) $Y = a \cdot X; r^2 \geq 0.8$

B) $Y = a \cdot X + b; r^2 \geq 0.8; \text{abs}(b) < 5$

Y= Reference Method (gravimetric method),
X= Automatic analyzer

Common Continuous Ambient Aerosol Sampling Techniques

$$(Dm / Dt) / (DV / Dt) = \text{mg/m}^3$$

Advantages

- Continuous method
- Highly time resolved
- High resolution
- instantaneous turnaround
- Low operational cost

Disadvantages

Temperature dependency:

- Volatile losses
- Seasonal and regional dependencies

Affected by vibration

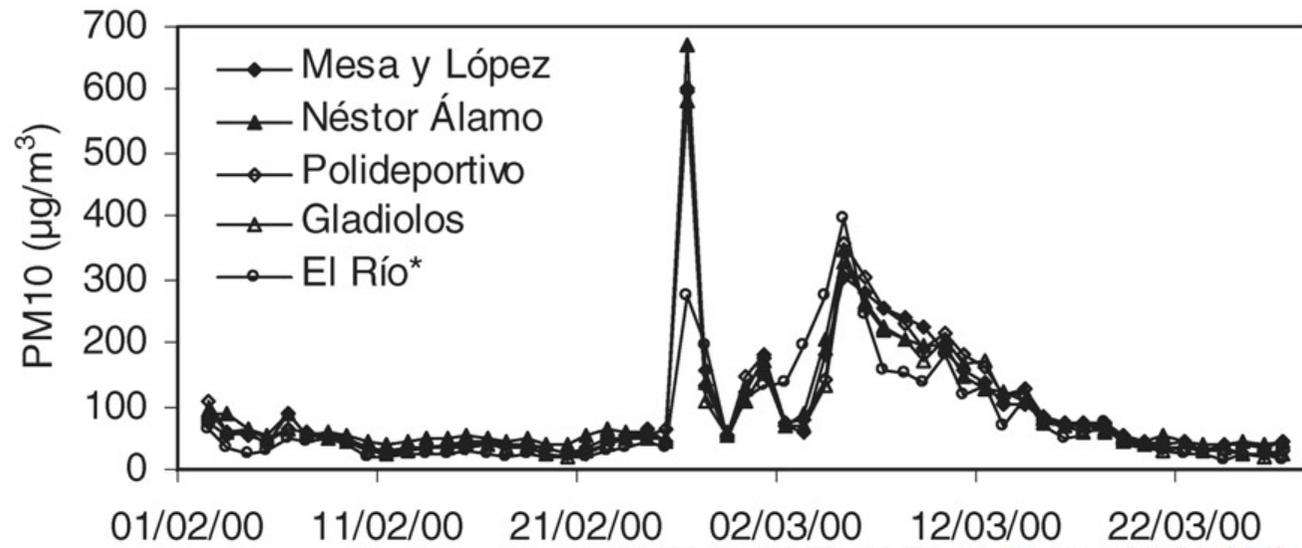
Manual filter changes necessary

Complex systems require some skill

X2 or X3 capita cost

Determination of Gravimetric Equivalent concentrations

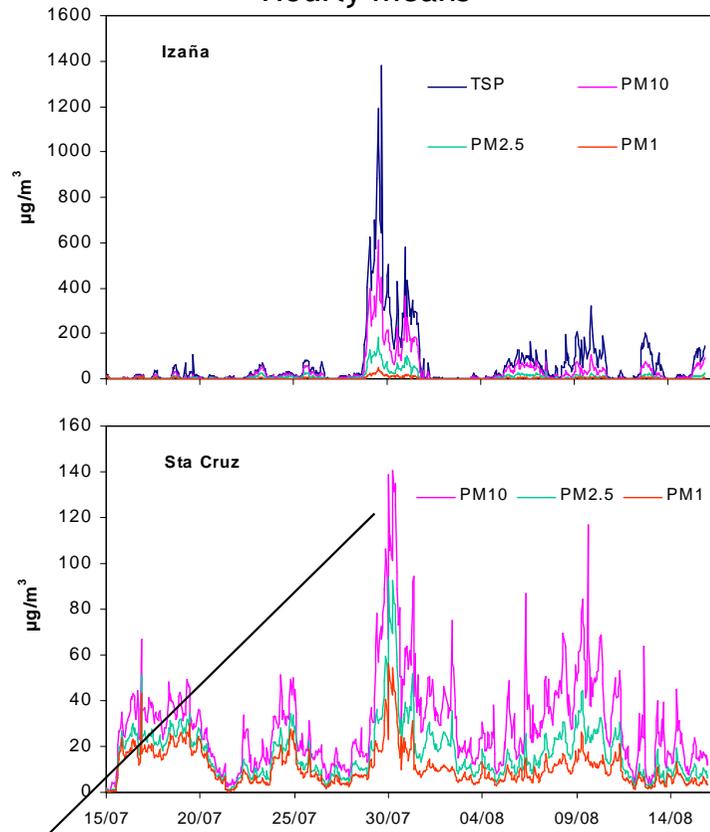
Air quality stations at Tenerife Island



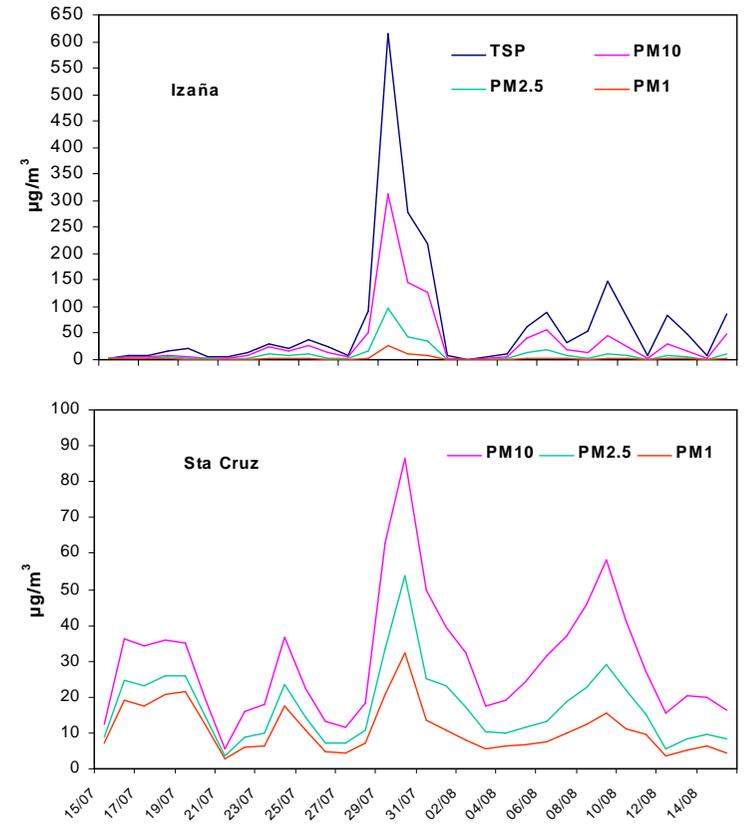
Viana et al., Atmospheric Environment , 2002

In-situ dust characterization

Hourly means



Daily means



property of aerosol dust: **mass concentration**

bulk aerosol mass concentration

bulk dust mass concentration

bulk dust mass concentration

method-1 : filter ash (J.M. Prospero)

step-1: samples collected on filters are extracted with de-ionized water and the extracts are analyzed for major soluble inorganic ions:

- Na⁺ by flame atomic absorption
- Cl⁻, NO₃⁻ and SO₄⁼ by suppressed ion chromatography
- NH₄⁺ by automated colorimetry

step-2: then, non sea salt sulfate is calculated using the SO₄⁼/Na⁺ ratio in bulk sea water (0.2517).

$$\text{sea salt} = \text{Na}^+ + \text{ss-SO}_4^= (0.2517 \cdot \text{Na}^+)$$

step-3: the extracted filters are then placed in a muffle furnace for 14-h (overnight) at 500° C. The ash residue weight.

$$\text{bulk dust} = \frac{\text{weight ash residue}}{\text{volume of sampled air}} \times 1.3$$

normalization: Al accounts for 8% of dust

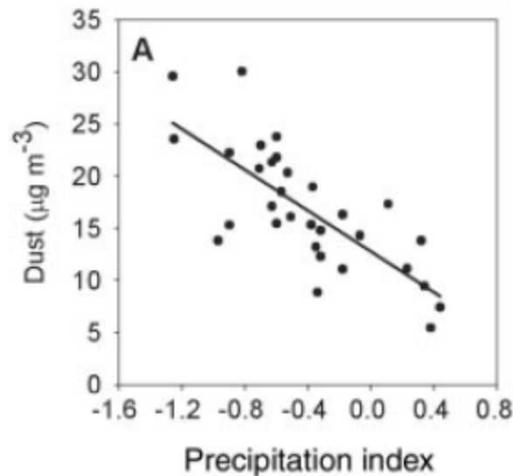
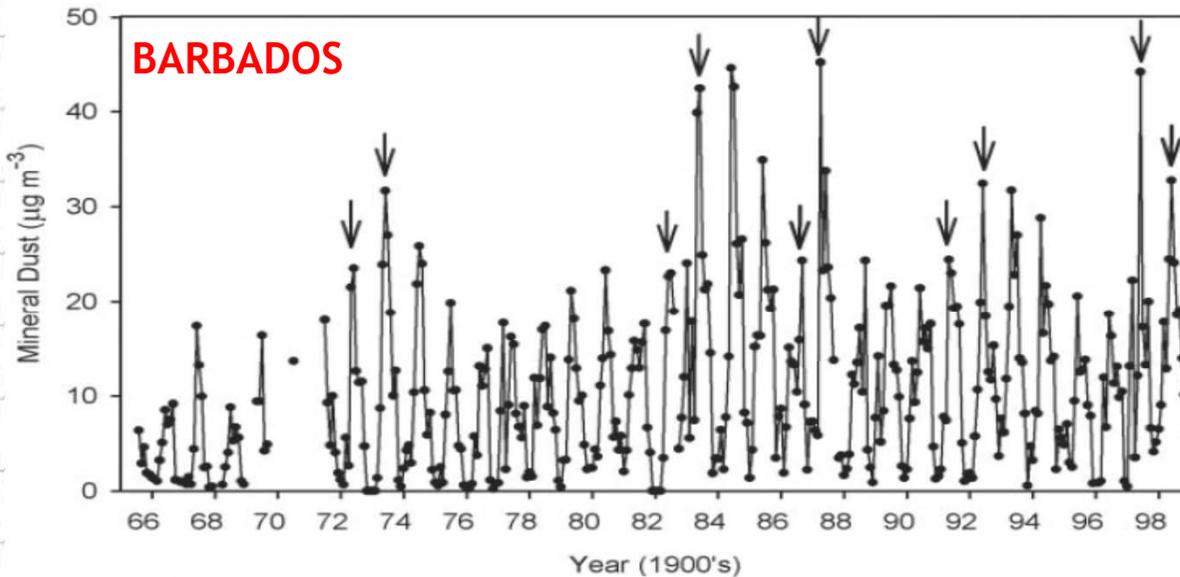
this technique may underestimate dust concentrations because of the loss of soluble minerals (carbonates, halides).

-standard error is considered: $\pm 0.1 \mu\text{g}/\text{m}^3$ for concentrations $< 1 \mu\text{g}/\text{m}^3$
10% for higher concentrations.



method-1 : filter ash (J.M. Prospero)

Longest records of dust: since 1964



(A) Dust plotted against the prior-year

African Droughts and Dust Transport to the Caribbean: Climate Change Implications

Joseph M. Prospero^{1*} and Peter J. Lamb²

7 NOVEMBER 2003 VOL 302 SCIENCE



Review Article *Aeolian Research* Aeolian Research 6 (2012) 55–74
 A review of methods for long term in situ characterization of aerosol dust
 Sergio Rodríguez^{a,*}, Andrés Alastuey^b, Xavier Querol^b

bulk dust mass concentration method-2: tracer analysis

In a filter with the dust sample, one or more dust tracer are analysed by chemical methods, and then total dust is calculated using the mean proportion of that element in dust:

$$\text{Al (8\% of soil)} \quad \longrightarrow \quad \text{dust} = \text{Al} \cdot (100/8) \quad \text{Eq-1}$$

$$\text{Si (33\% of soil)} \quad \longrightarrow \quad \text{dust} = \text{Si} \cdot (100/33) \quad \text{Eq-2}$$

ratio element / oxide

Na2O	0.47
MgO	0.43
Al2O3	1.89
SiO2	5.98
K2O	0.38
CaO	0.62
TiO2	0.09
Fe2O3	0.88

stoichiometry

$$\text{dust} = 0.47 \cdot \text{Na} + 0.43 \cdot \text{Mg} + 1.89 \cdot \text{Al} + 5.98 \cdot \text{Si} + 0.38 \cdot \text{K} + 0.62 \cdot \text{Ca} + 0.88 \cdot \text{Fe} + 0.09 \cdot \text{Ti} \quad \text{Eq-3}$$

Can be applied with elements with low enrichment factors!!!!



bulk dust mass concentration method-2: tracer analysis

Mean ratio to Al in soil

	Mason (1966)	Taylor (1964)
Na	0.34809	0.28676
Mg	0.25707	0.28311
Al	1.00000	1.00000
Si	3.40959	3.42041
K	0.31857	0.25395
Ca	0.44649	0.50425
Ti	0.05412	0.06926
Fe	0.61501	0.68408
P	0.01292	0.01276
Mn	0.01169	0.01154
Sr	0.00456	0.00461
S	0.00316	0.00320
V	0.00166	0.00164
Cl	0.00160	0.00158
Cr	0.00123	0.00122
Ni	0.00092	0.00091
Zn	0.00086	0.00085
Cu	0.00068	0.00067

$$EF\text{- Enrichment Factor (X)} = \frac{(X / Al)_{\text{sample}}}{(X / Al)_{\text{soil}}}$$

EF (X) = 1 **➡** element X is due to soil emissions

EF (X) >> 1 **➡** there is an additional contribution to element X from other sources. Then, do not include X in Eq-3.

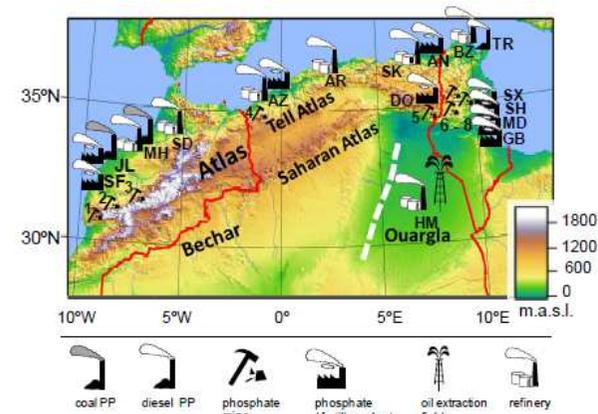
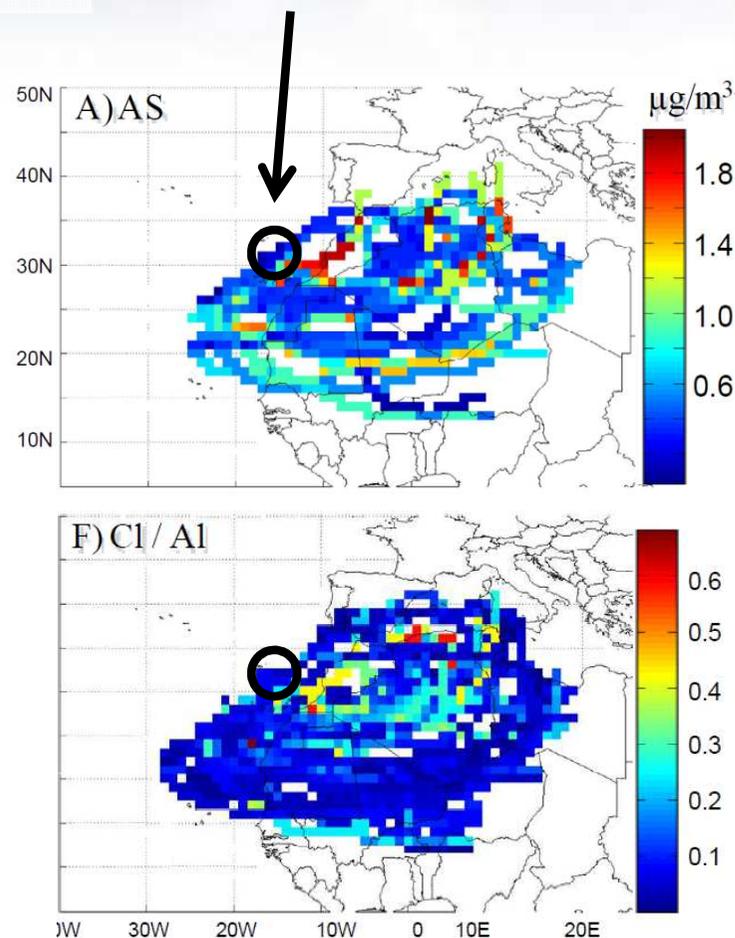
bulk dust mass concentration method-2: tracer analysis

Izaña - Tenerife

	EF
Al	1.00
Ca	0.97
K	0.72
Na	0.31
Mg	0.78
Fe	0.85
S	35.71
Cl	77.70
P	0.70
Ni	0.56



Izaña: measurement site



S and Cl⁻ emitted
by industry mixed
with dust

Transport of desert dust mixed with North African industrial pollutants in the subtropical Saharan Air Layer

S. Rodríguez¹, A. Alastuey², S. Alonso-Pérez^{1,2}, X. Querol², E. Cuevas¹, J. Abreu-Afonso¹, M. Viana², N. Pérez², M. Pandolfi², and J. de la Rosa³

Atmos. Chem. Phys., 11, 6663–6685, 2011
www.atmos-chem-phys.net/11/6663/2011/
doi:10.5194/acp-11-6663-2011
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property of aerosol dust:

number size distribution

mass concentration

chemical composition

mixing state

mineralogy

optical properties

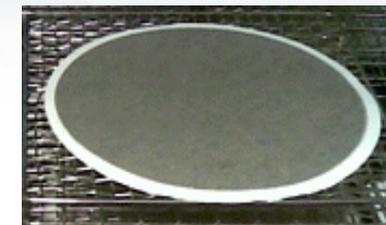
bulk chemical composition

PM samples: $\left\{ \begin{array}{l} \text{fine + coarse (TSP, PM}_{10}\text{)} \\ \text{fine (PM}_{2.5}\text{, PM}_1\text{)} \end{array} \right.$

Saharan dust



Urban particles



PM ($\mu\text{g}/\text{m}^3$) = **dust** + **trace elements** + **ions** (SO_4^- , NO_3^- , NH_4^+ , Na^+ , Cl^-) + OC + EC

Elemental Composition:

Major elements (Al, Si, Ca, K, Na, Mg) + trace elements (P, Li, Be, Sc, Ti, V, Cr, Mn, Co, Ni, Cu, Zn, Ga, Ge, As, Se, Rb, Sr, Y, Zr, Nb, Mo, Cd, Sn, Sb, Cs, Ba, La, Ce, Pr, Nd, Sm, Eu, Gd, Tb, Dy, Ho, Er, Tm, Yb, Lu, Hf, Ta, W, Tl, Pb, Bi, Th, U)

Ions: SO_4^- , NO_3^- , NH_4^+ , Na^+ , Cl^-

Ion Chromatography, ICP-AES, ICP-MS, selective electrodes and colorimetry

Thermal/optical reflectance (TOR) and/or thermal/optical transmission (TOT)

Inductively coupled plasma Atomic Emission Spectroscopy ICP-AES

Destructive techniques

destructive techniques

Inductively coupled plasma Mass spectroscopy ICP-MS

Destructive techniques

XRF, PIXE, INAA : none destructive techniques

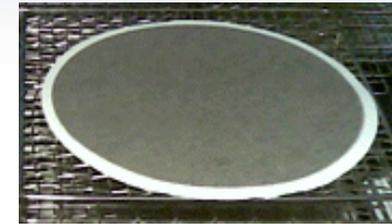
bulk chemical composition

PM samples: $\left\{ \begin{array}{l} \text{fine + coarse (TSP, PM}_{10}\text{)} \\ \text{fine (PM}_{2.5}\text{, PM}_1\text{)} \end{array} \right.$

Saharan dust



Urban particles

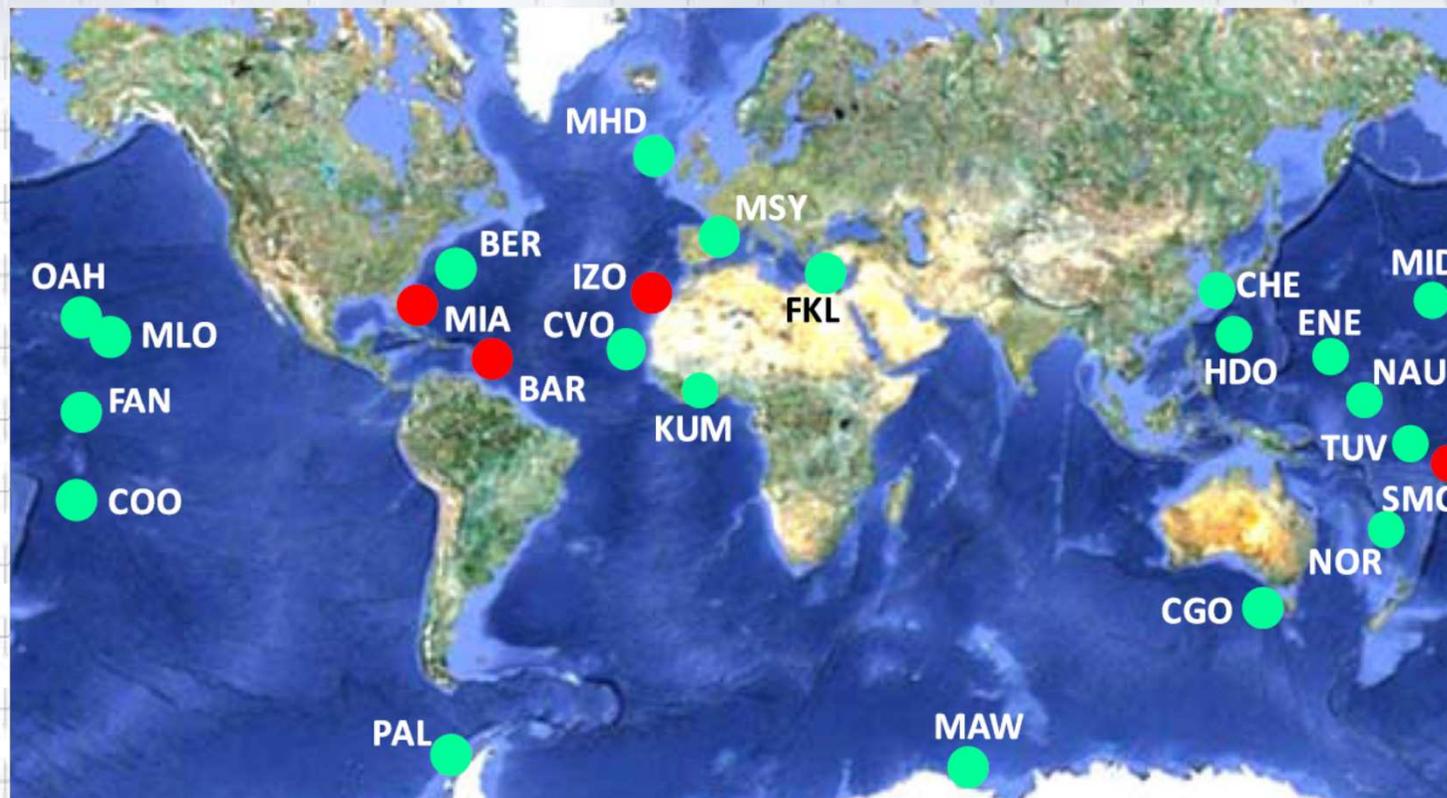


PM ($\mu\text{g}/\text{m}^3$) = **dust** + **ions** (SO_4^- , NO_3^- , NH_4^+ , Na^+ , Cl^-) + OC + EC + **trace elements**

bulk chemical composition is the most reliable technique for quantifying the concentration of dust and other species (if present, e.g. pollutants, sea salt).

This is considered a reference method for the quantification of dust.

Other analytical techniques are available. Scanning Electron Microscopy (SEM) and Transmission Electron Microscopy (TEM) coupled with Energy Dispersive X-ray analysis (EDX) allows individual particle characterization for size, morphology, chemical and mineral composition.



- at least 4 years
- Active during the last 20 years

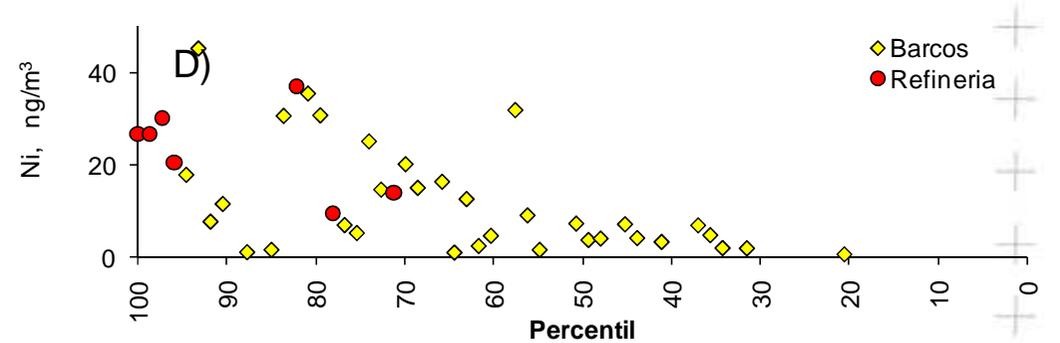
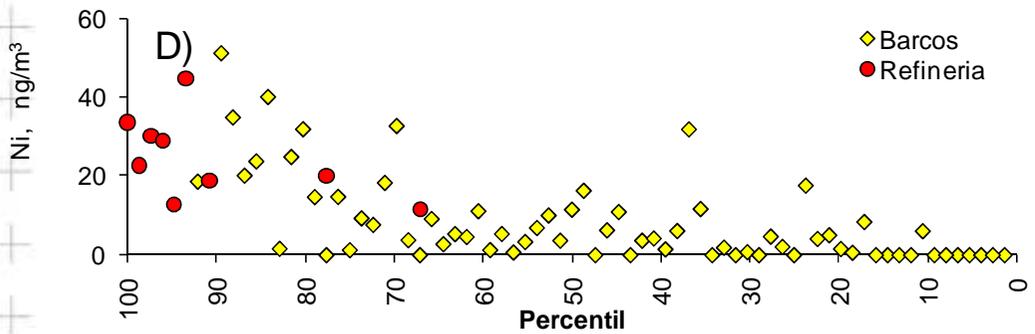
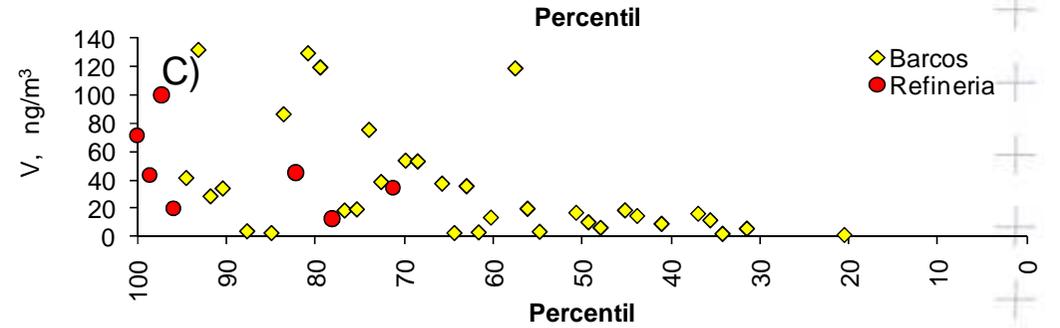
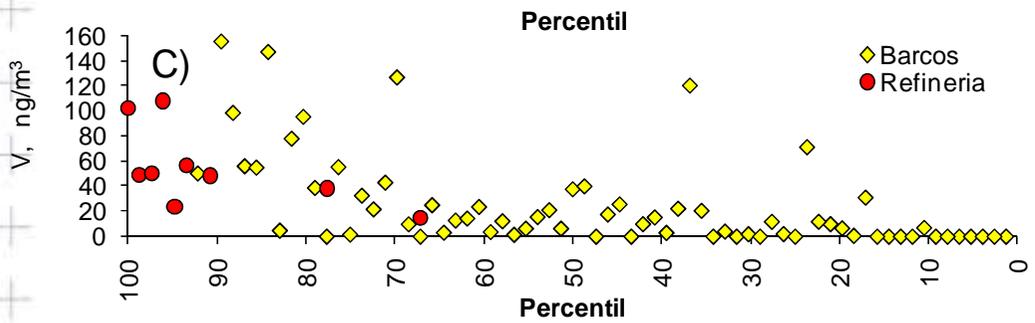
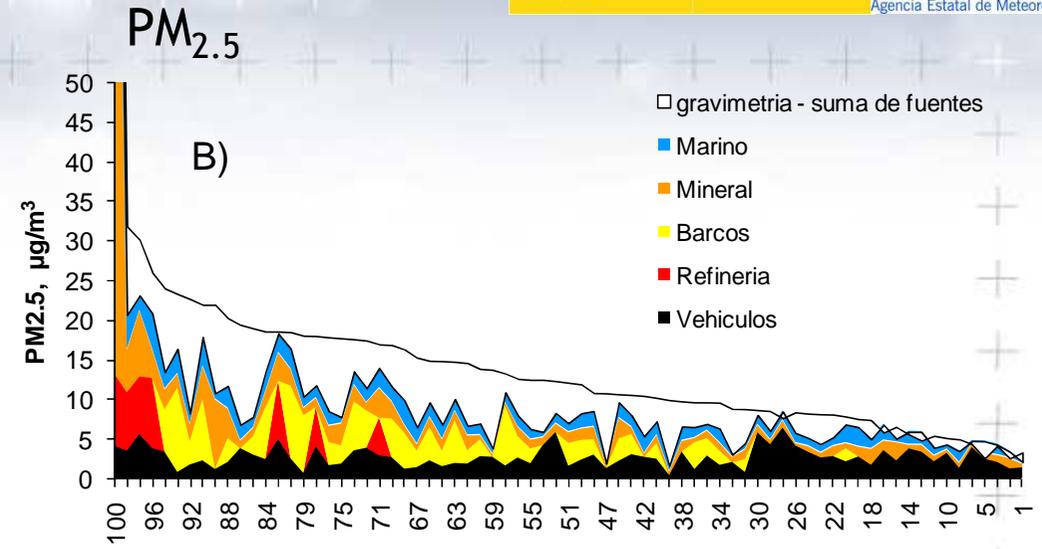
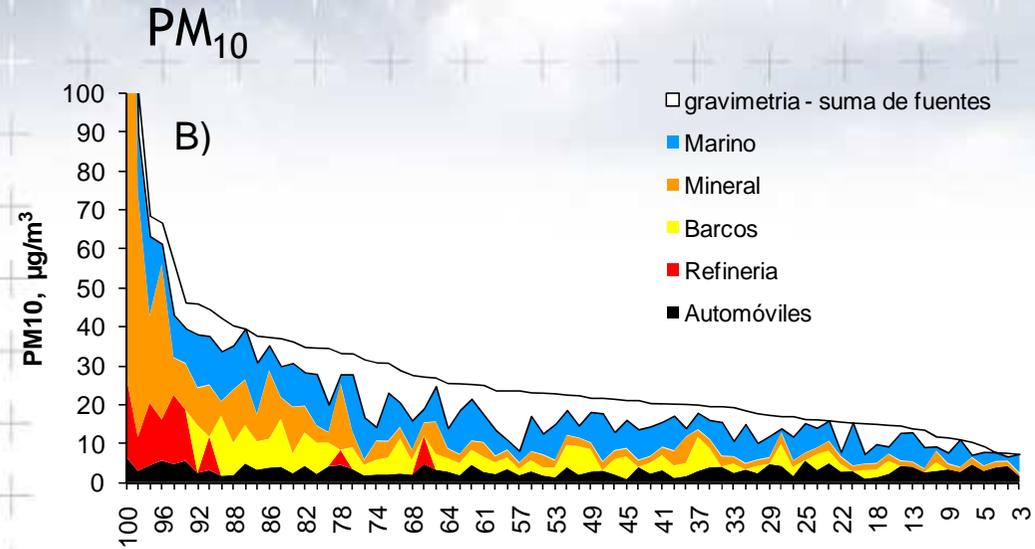
Review Article

Aeolian Research Aeolian Research 6 (2012) 55–74

A review of methods for long term in situ characterization of aerosol dust

Sergio Rodríguez^{a,*}, Andrés Alastuey^b, Xavier Querol^b

bulk chemical composition reference method for the quantification of dust.



Santa Cruz de Tenerife source apportionment study by receptor modeling
Rodríguez et al., 2009

property of aerosol dust:

number size distribution

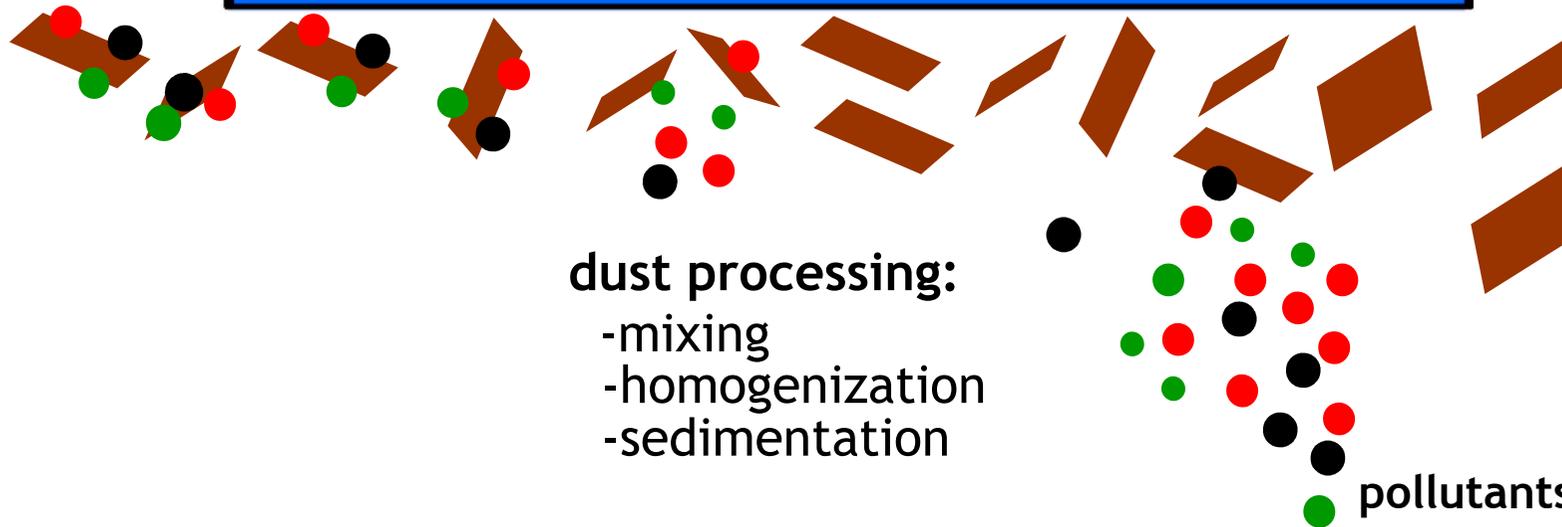
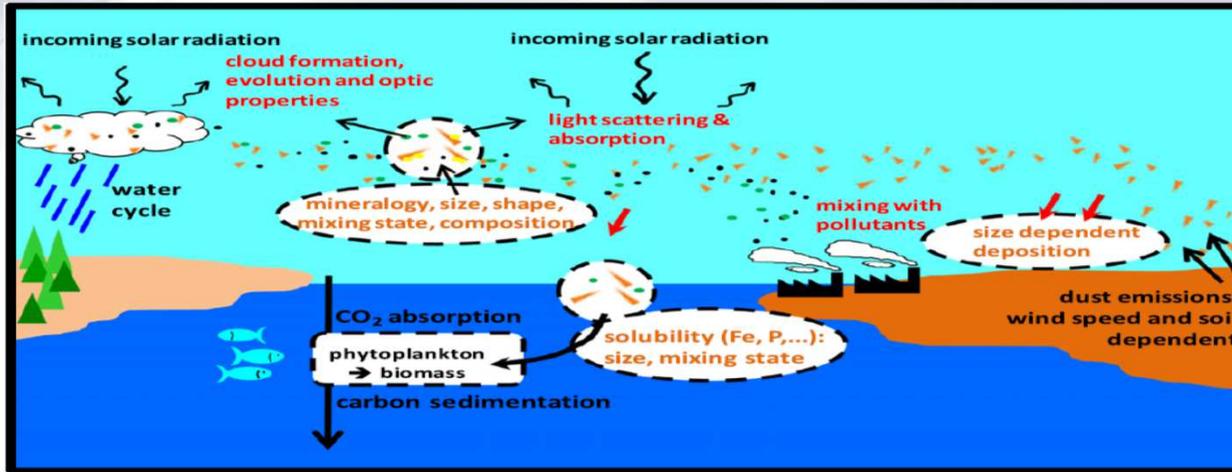
mass concentration

chemical composition

mixing state

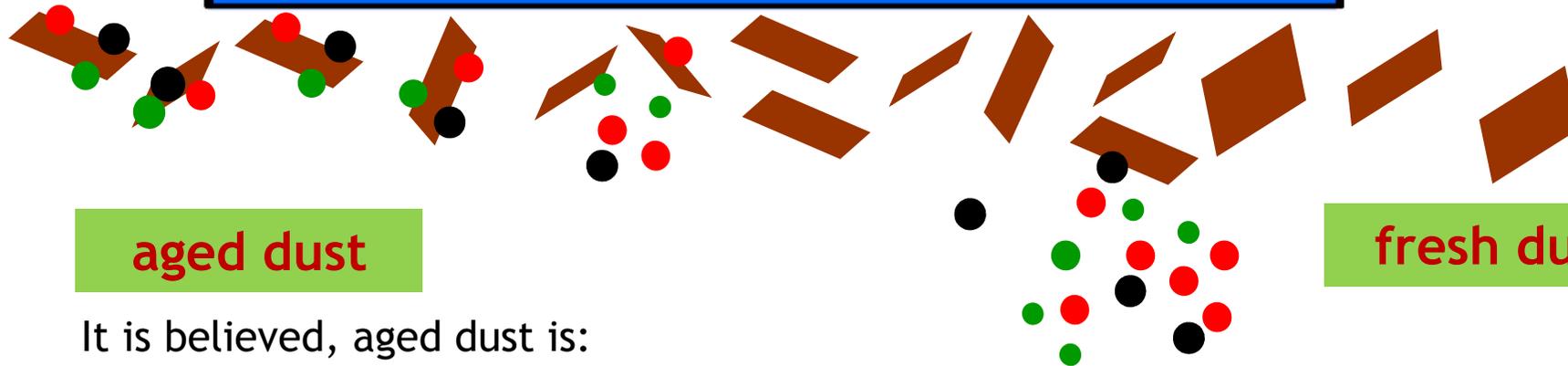
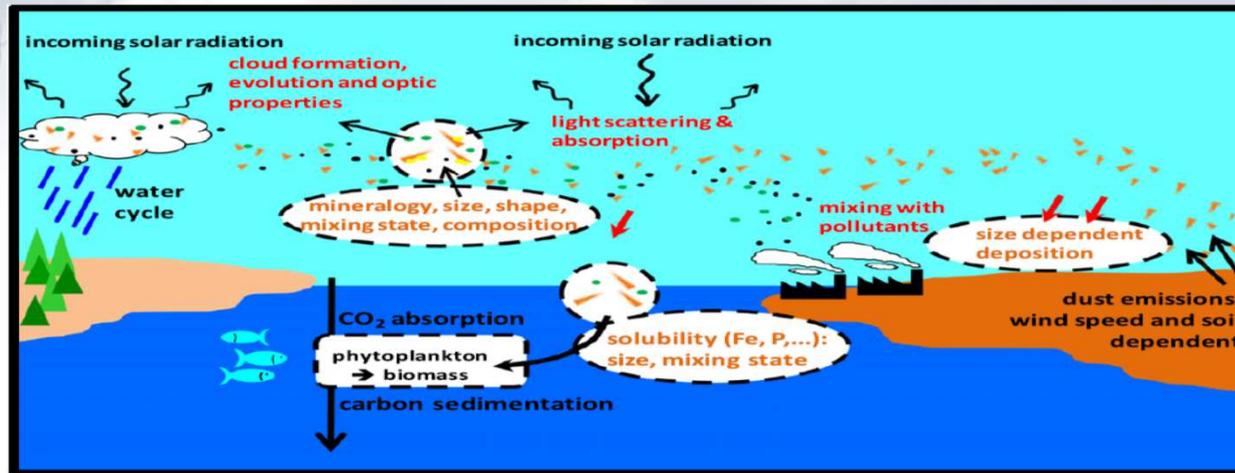
mineralogy

optical properties



internal mixing:
all particles the same composition (same mixing)

external mixing:
each particle different chemical composition

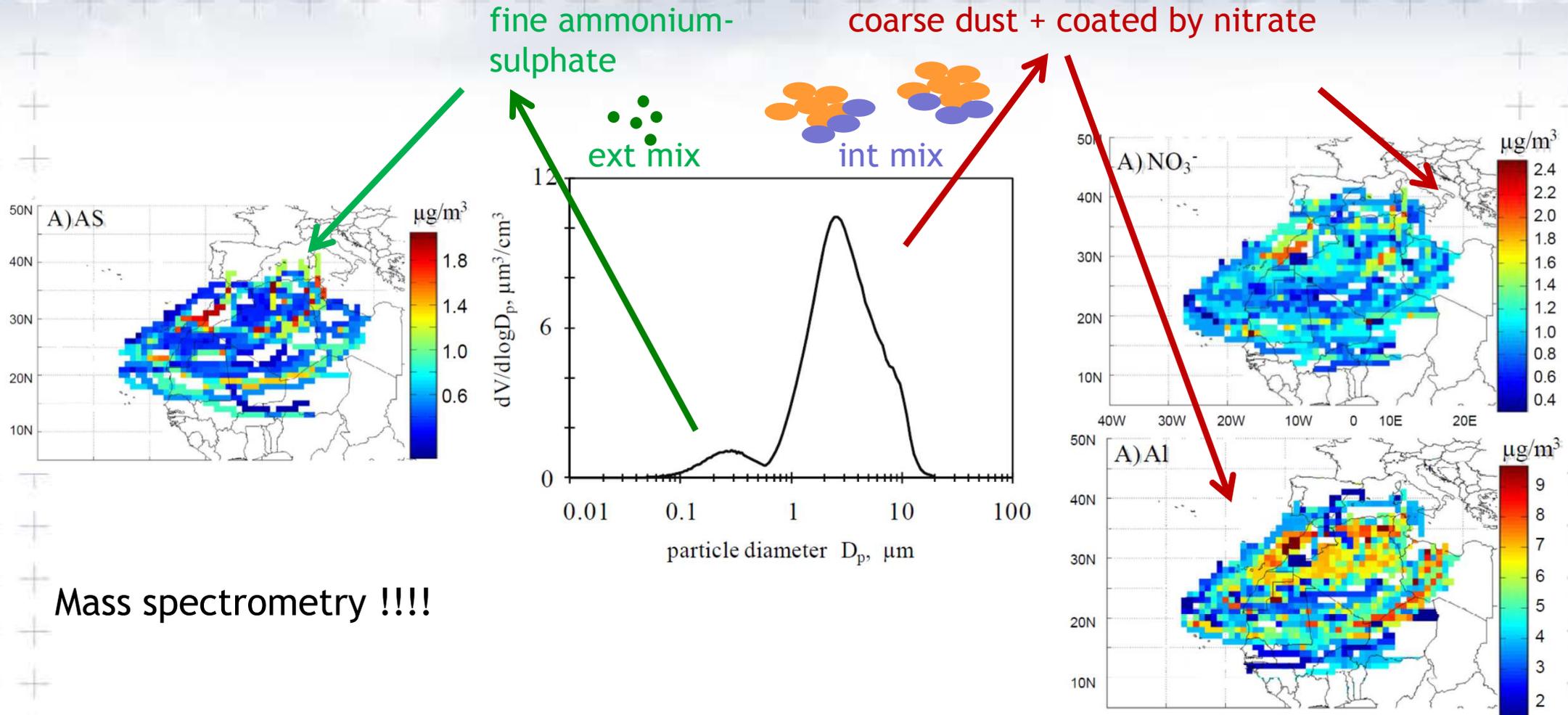


aged dust

It is believed, aged dust is:

- more soluble
- different refractive index

fresh dust



Mass spectrometry !!!!

Transport of desert dust mixed with North African industrial pollutants in the subtropical Saharan Air Layer

S. Rodríguez¹, A. Alastuey², S. Alonso-Pérez^{1,2}, X. Querol², E. Cuevas¹, J. Abreu-Afonso¹, M. Viana², N. Pérez², M. Pandolfi², and J. de la Rosa³

Atmos. Chem. Phys., 11, 6663–6685, 2011

www.atmos-chem-phys.net/11/6663/2011/

doi:10.5194/acp-11-6663-2011

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Atmospheric Environment 44 (2010) 3135–3146

Variation of the mixing state of Saharan dust particles with atmospheric transport

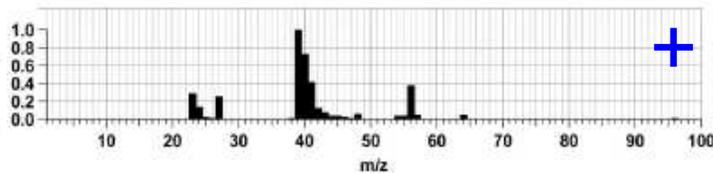
Manuel Dall'Osto^{a,b}, Roy M. Harrison^{a,*}, Eleanor J. Highwood^c, Colin O'Dowd^b, Darius Ceburnis^b, Xavier Querol^d, Eric P. Achterberg^e

Aerosol Time Of Flight Mass Spectrometer (ATOFMS)

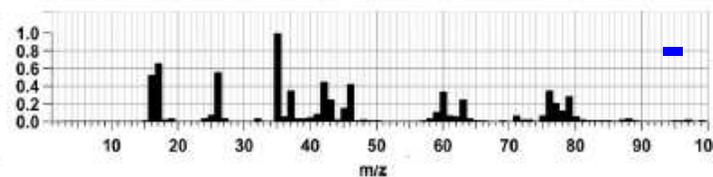
-aerodynamic size of particles (0.3 - 1 μm)

-chemical composition of individual particles

→ positive and negative ion mass spectrums of a single particle.



m/z: 27 (Al), 40 and 56 (Ca).....



m/z: -97 (HSO_4^-), -80 (SO_3^-), -62 (NO_3^-).....

The mass spectrum is qualitative in that the intensities of the mass spectral peaks are not directly proportional to the component mass but are dependent on the particle matrix.

Not for long term measurements

The ATOFMS can supply quantitative information on particle number as a function of composition, providing measurements of all the particle components (including OC, EC, sulfate, nitrate, dust and sea salt)

property of aerosol dust:

number size distribution

mass concentration

chemical composition

mixing state

mineralogy

optical properties

optical properties

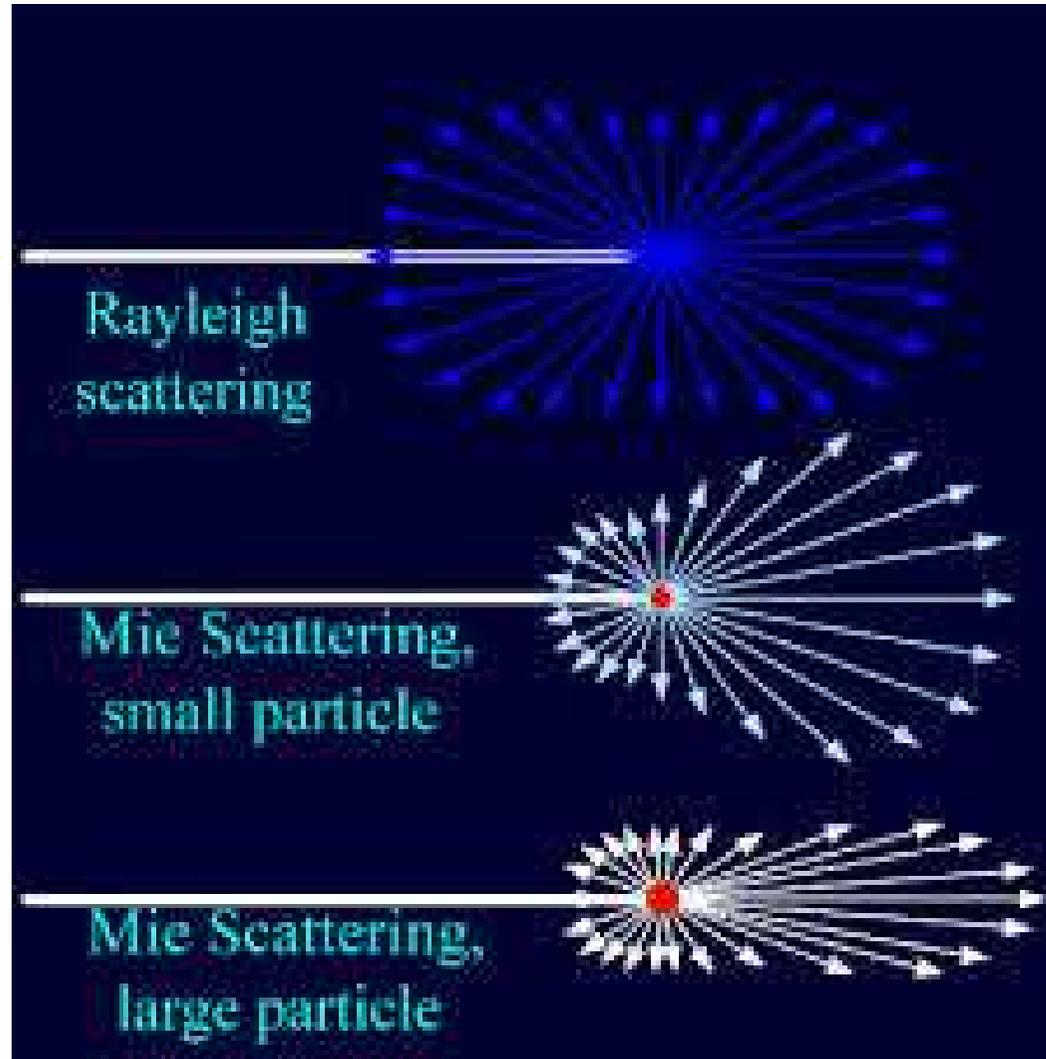
Redistribution of radiation depends on:

-particle size

forward scattering increase with particle size

-composition and mineralogy

mixing with pollutants
absorbing minerals (iron oxides)



$$\lambda > r$$

$$\lambda \sim r$$

$$\lambda \ll r$$

2 basic optical properties:

scattering and backscattering coefficient (several λ)

absorption coefficient (several λ)

$$I = I_0 \cdot e^{-\sigma_{\text{ext}} \cdot L}$$

$$\sigma_{\text{ext}} = \sigma_{\text{abs}} + \sigma_{\text{scat}}$$

Measured in Inverse Meters (m^{-1})
“How Much is Extinguished Per Meter?”

σ_{ext} aerosol extinction coefficient

σ_{abs} aerosol absorption coefficient : Absorption Photometer
(**MAAP, Aethalometer, PSAP**)

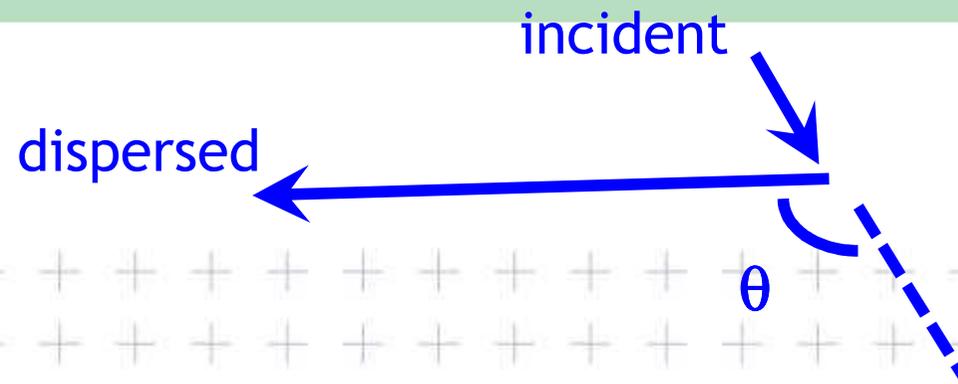
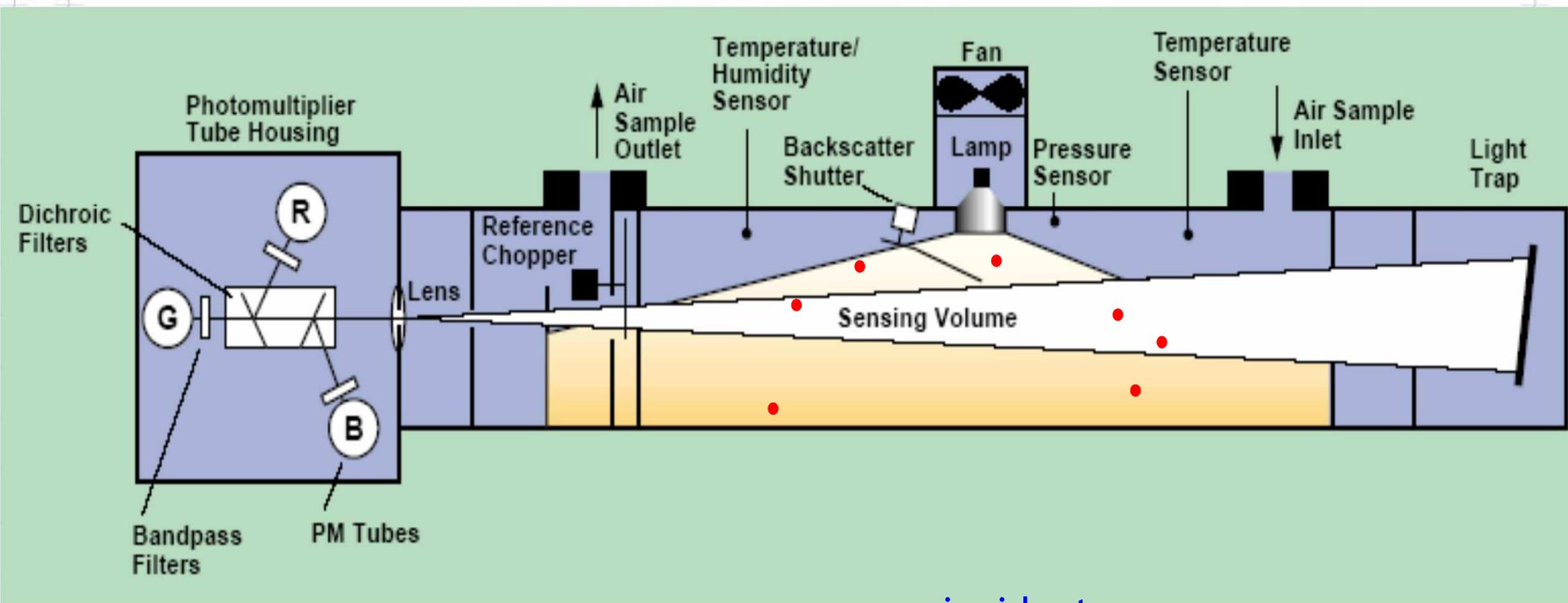
σ_{scat} aerosol scattering coefficient: **NEPHELOMETER**

2 basic optical properties:

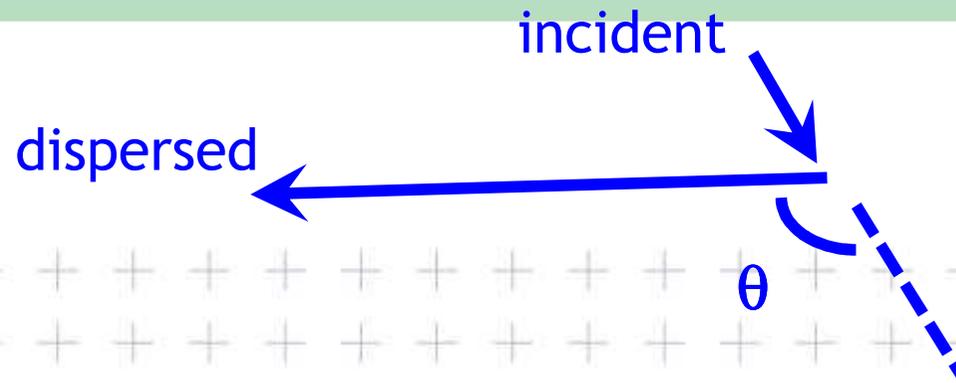
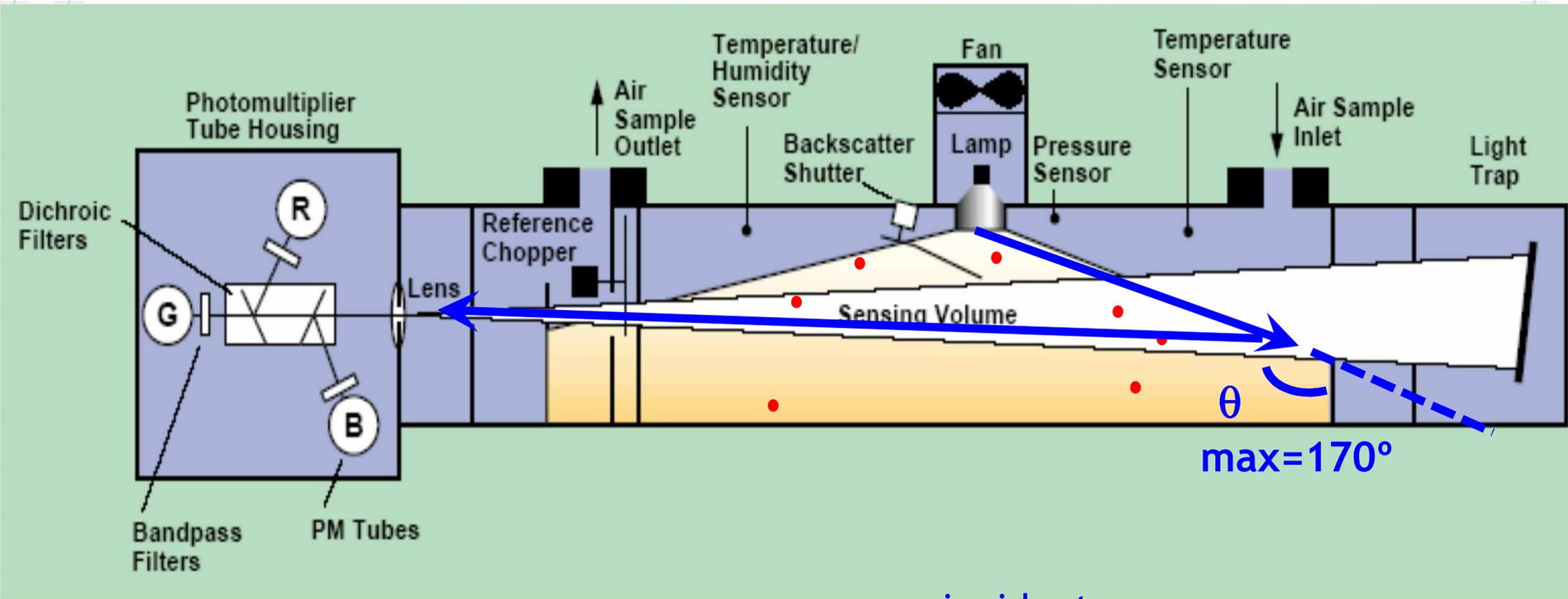
scattering (σ_{scat}) and backscattering coefficient (several λ)

absorption coefficient (several λ)

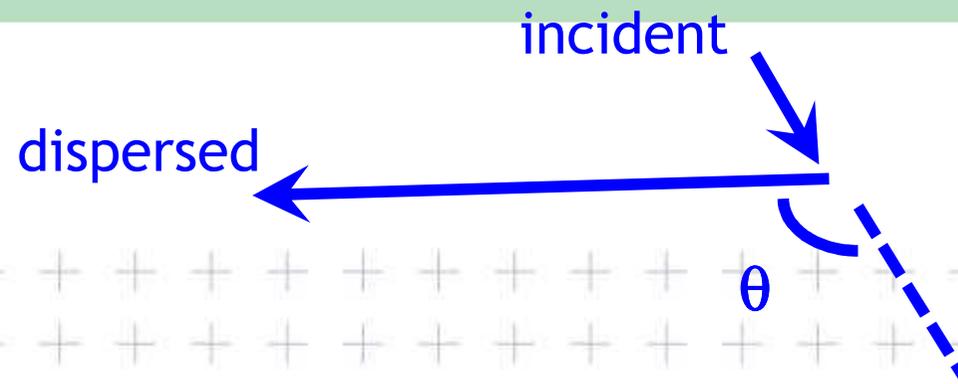
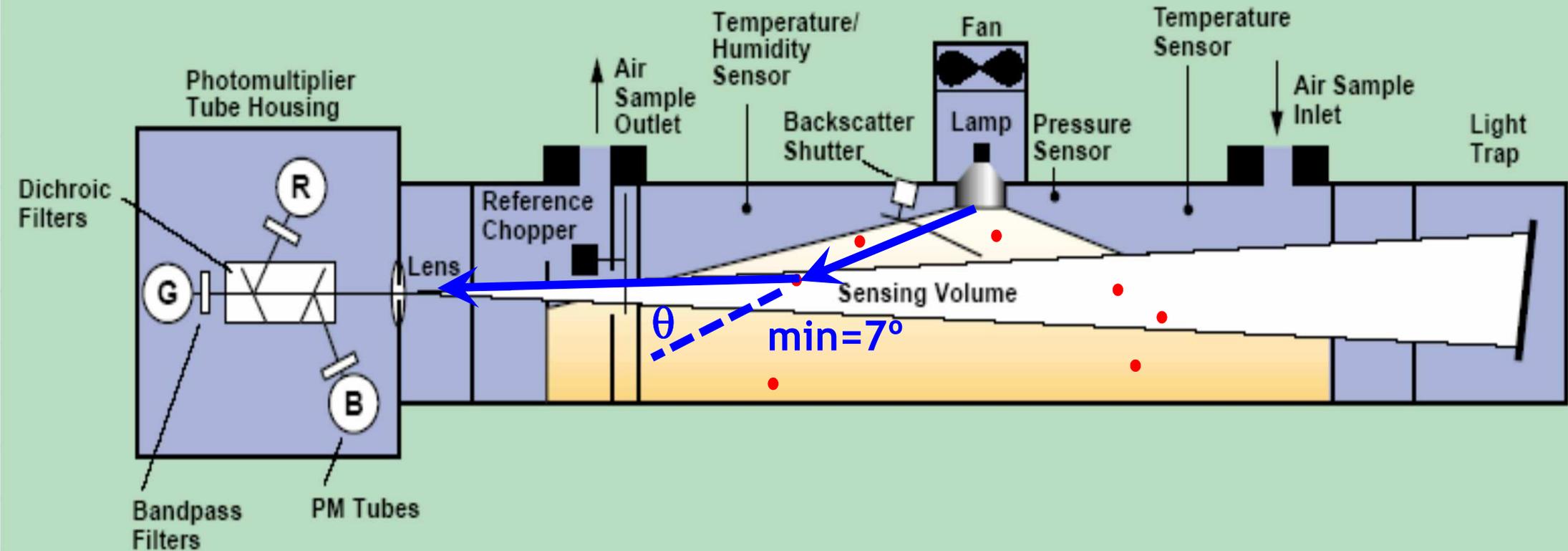
Integrating Nephelometer **coeficiente de scattering**



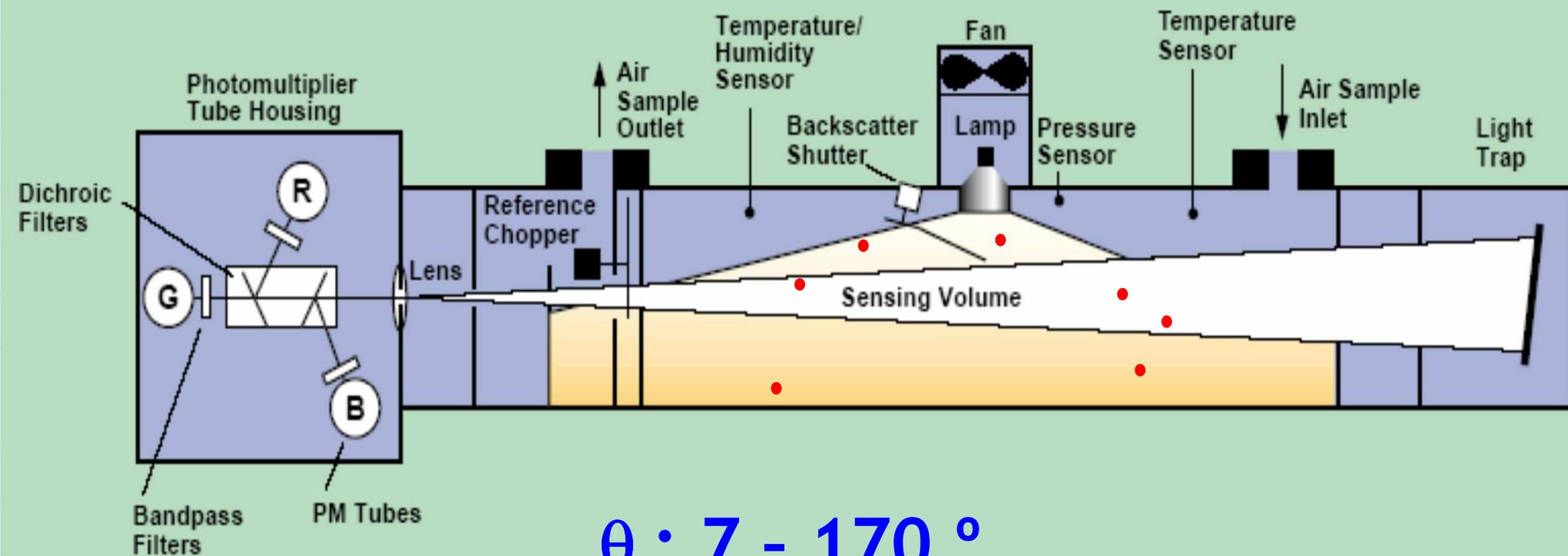
Integrating Nephelometer Scattering coefficient



Integrating Nephelometer Scattering coefficient



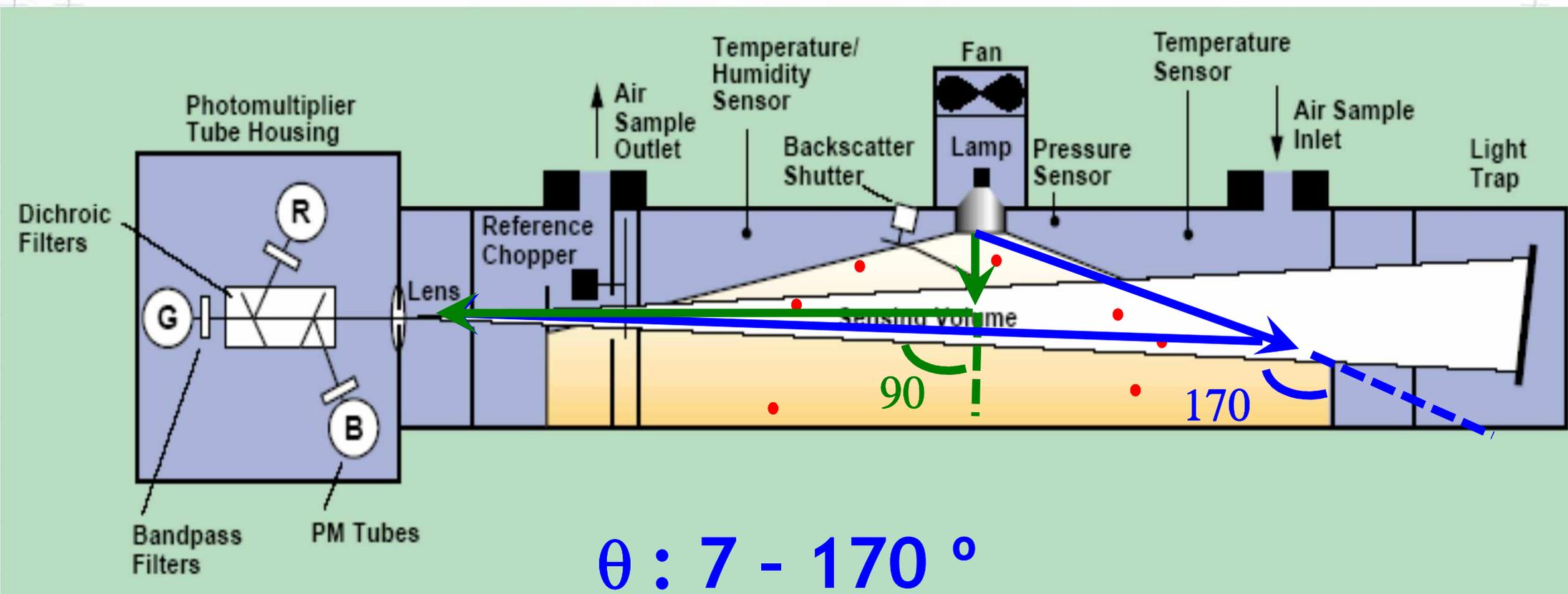
Integrating Nephelometer Scattering coefficient



Integrating nephelometer

Scattering Coefficient 7 - 170 °

Integrating Nephelometer Scattering coefficient

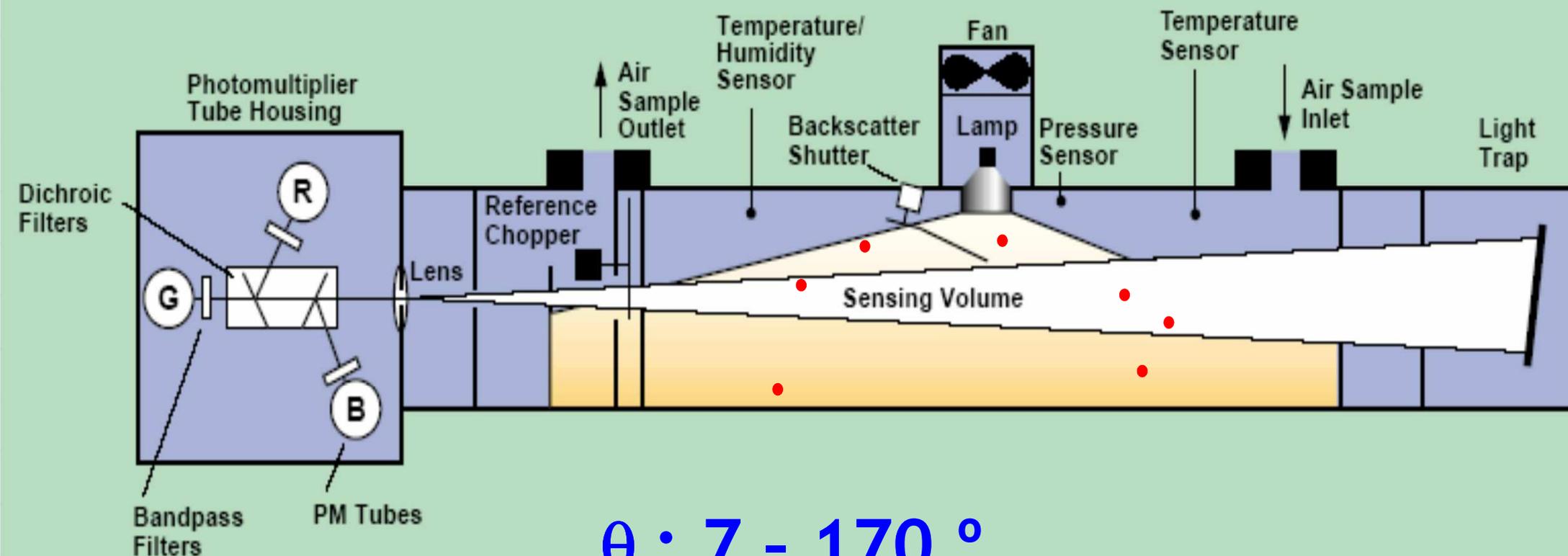


Integrating nephelometer

Total Scattering coefficient 7 - 170 °

Total Backscattering coefficient 90 - 170 °

Integrating Nephelometer Scattering coefficient



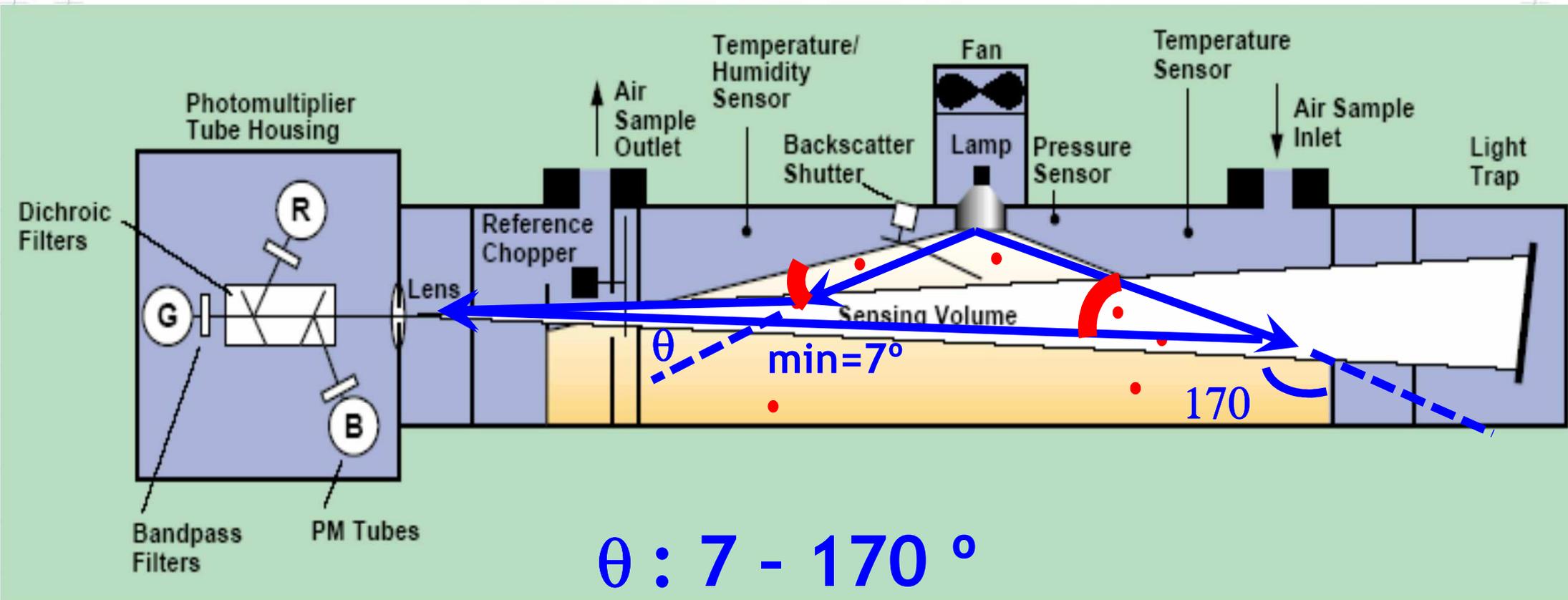
Integrating nephelometer

Total Scattering coefficient 7 - 170 °

Total Backscattering coefficient 90 - 170 °

$\lambda = 450, 550, 700 \text{ nm}$

Integrating Nephelometer: Truncation correction



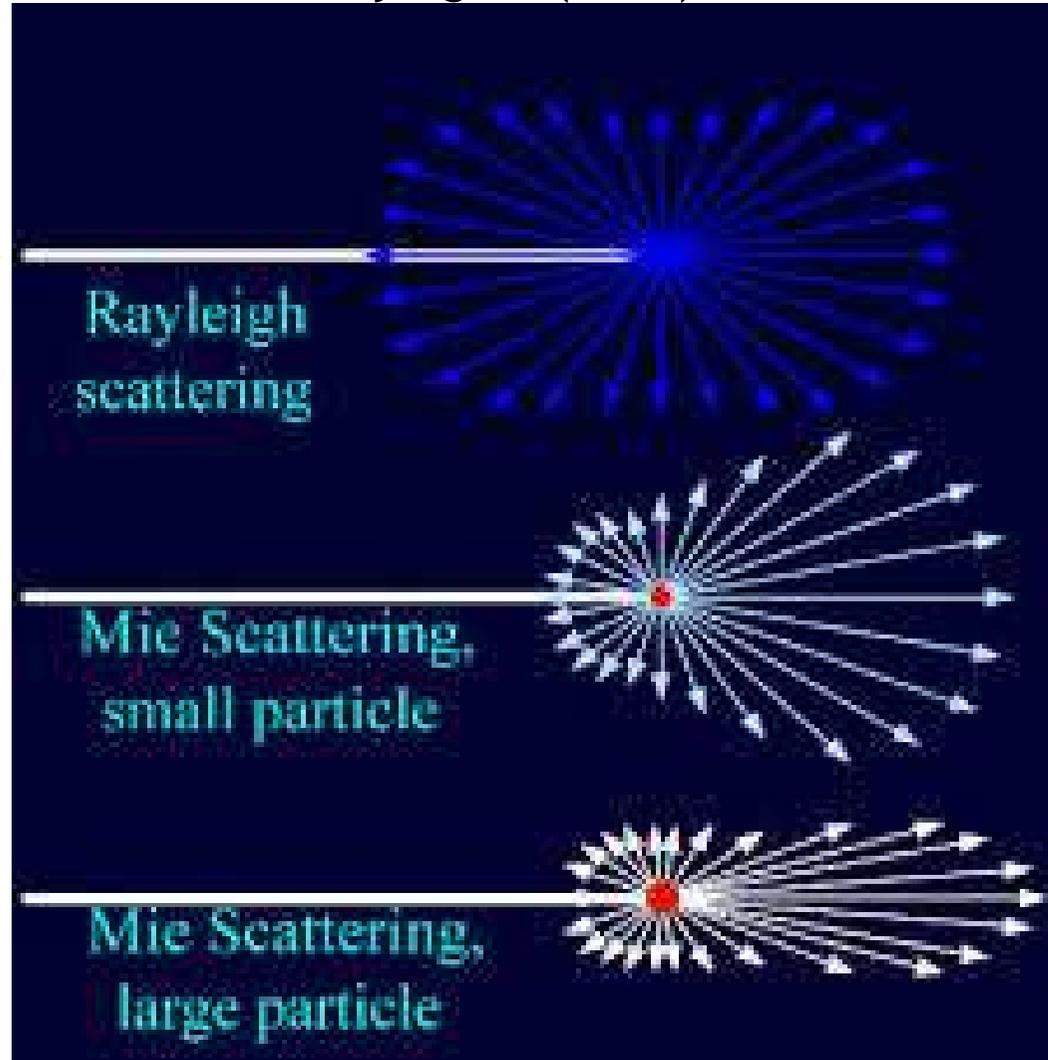
Total scattering coefficient $7 - 170^\circ$

Truncation error: light dispersed within the angles $0-7^\circ$ and $170-180^\circ$ is not measured

Forward scattering increase with particle size.
Coarse dust particles → TRUNCATION ERROR
TRUNCATION CORRECTION IS IMPORTANT FOR DUST
Correction scheme → Anderson y Ogren (1998).



If correction is not applied, the total scattering is underestimated by between 5-15% for submicron particles and by 40-60% for coarse particles

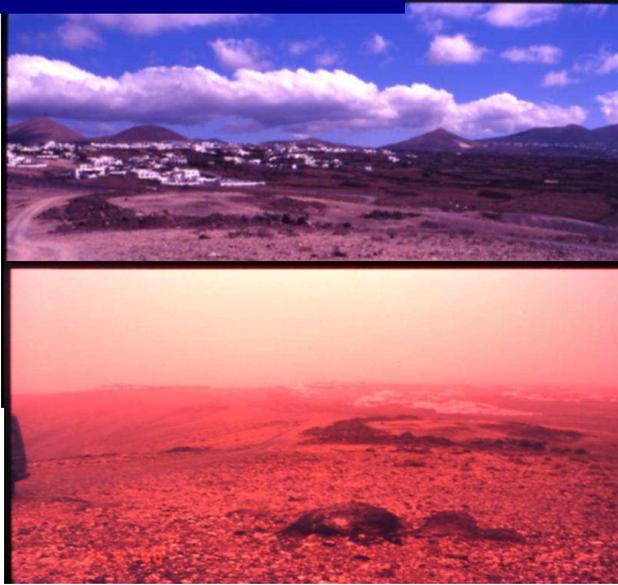


Back scattering

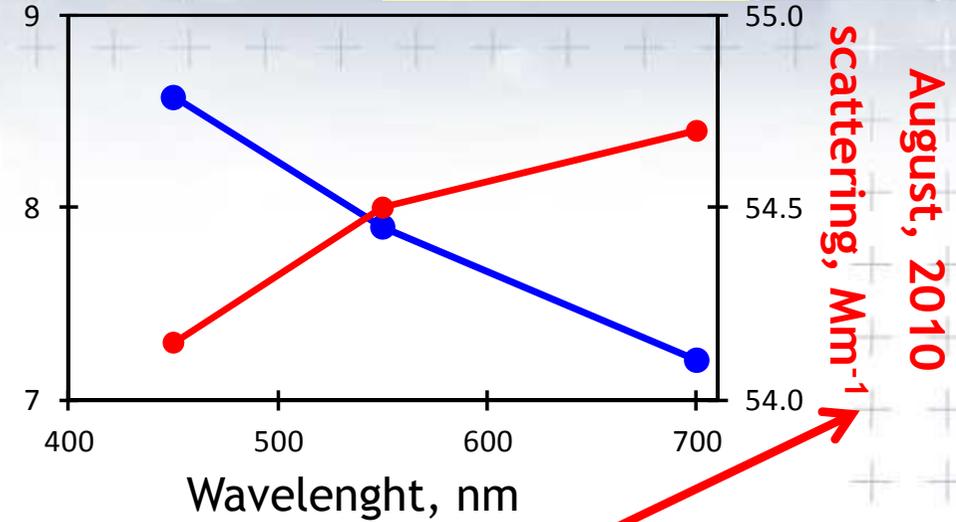


Forward scattering

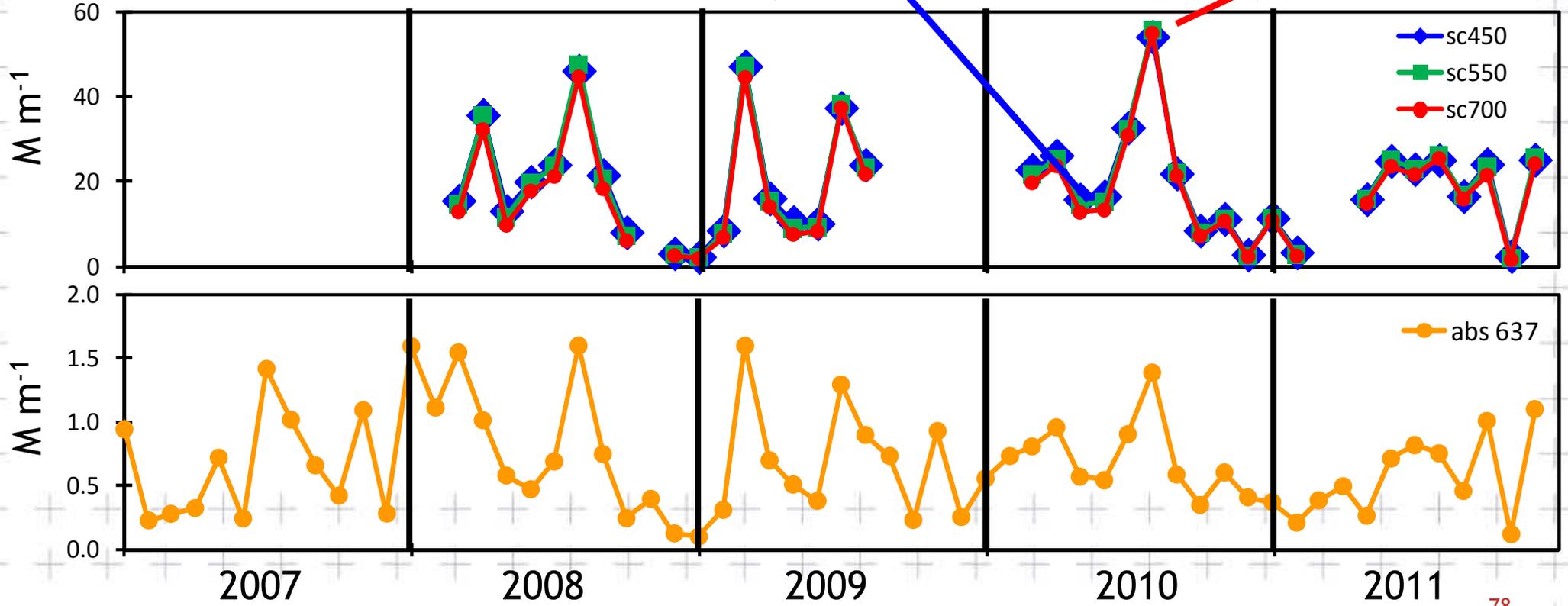
In-situ dust characterization



March, 2010
scattering, Mm^{-1}



Optical properties: scattering and absorption (note that $1 Mm^{-1} = 10^{-6} m^{-1}$)



optical properties

scattering (σ_{sp}) and backscattering coefficient (several λ)

2 basic optical properties:

absorption coefficient (several λ)

MAAP: Multi-Angle

Absorption Photometer

PSAP: Particle Soot

Absorption Photometer

3λ

Aethalometer

$5-7\lambda$



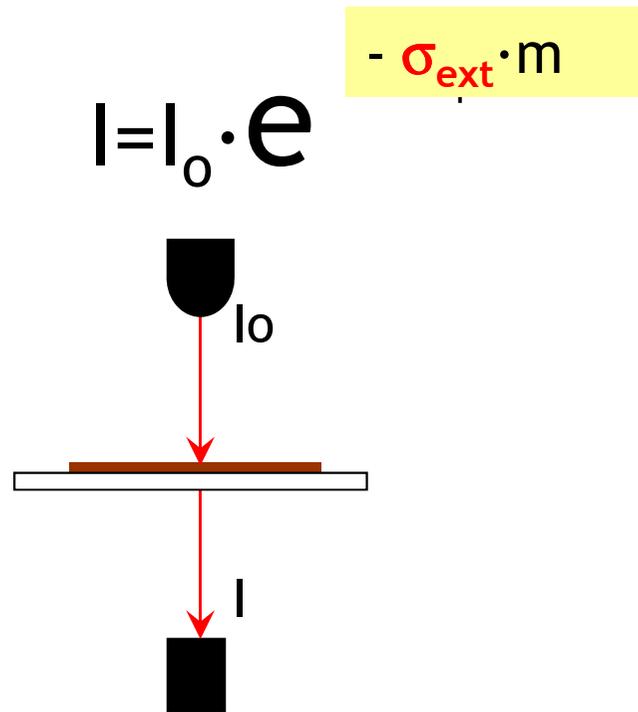
$\lambda=670\text{nm}$



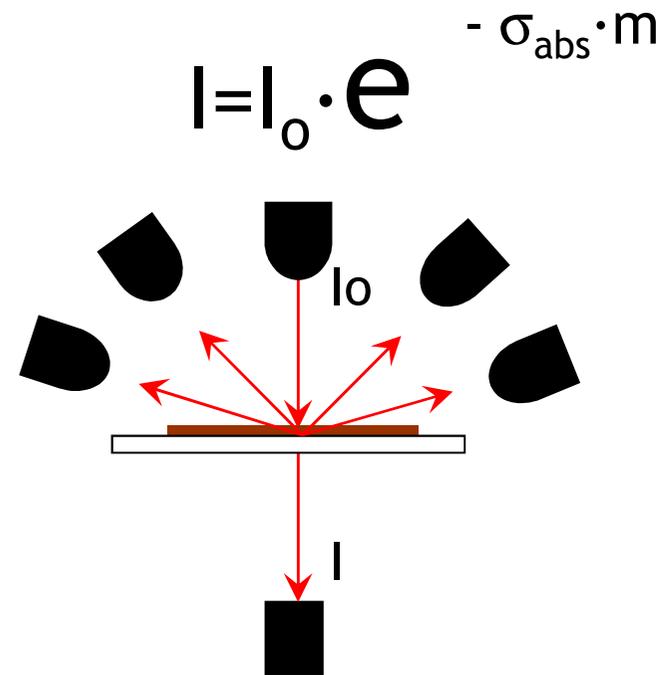
2 basic optical properties:

scattering (σ_{scar}) and backscattering coefficient (several λ)

absorption coefficient (several λ)



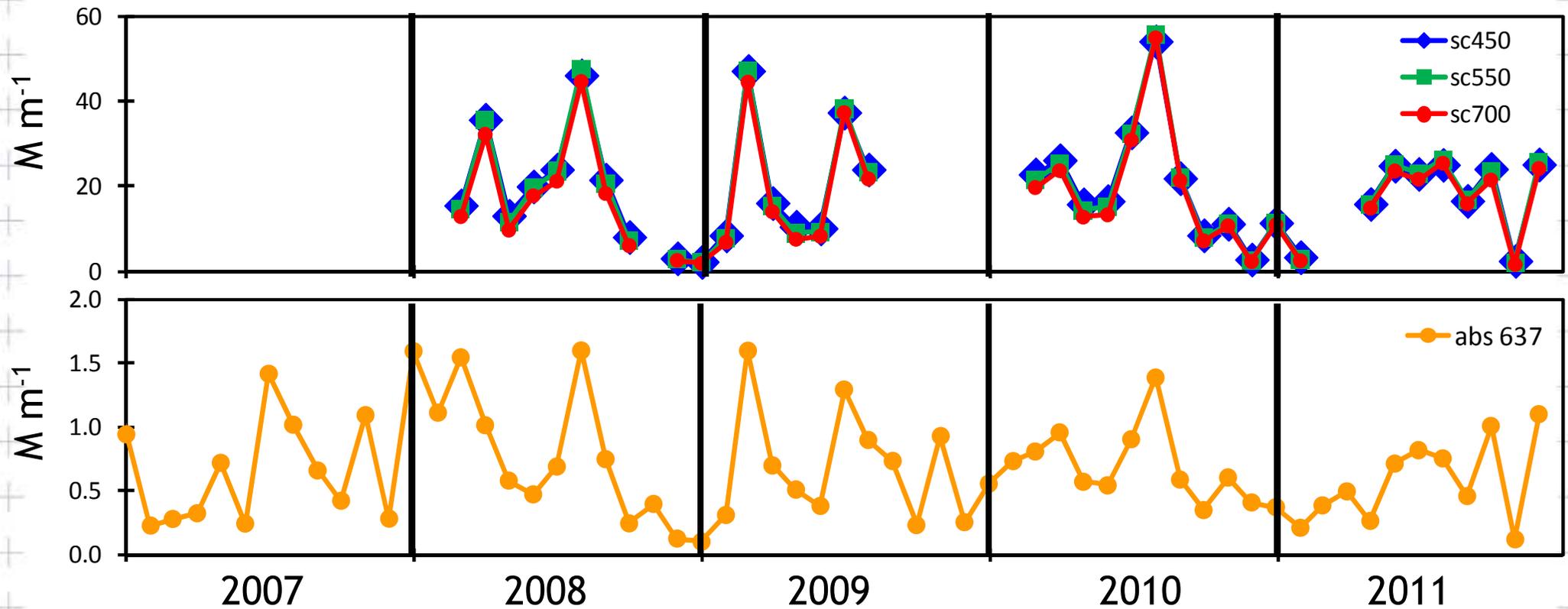
Aethalometer and PSAP



MAAP: MultiAngle Absorption Photometer

Abs. Coeff. (aethalometer and PSAP) > Abs. Coeff. (MAAP)

Long term monitoring of optical properties with simultaneous chemical and mineralogical characterization allows to understand potential changes in the optical properties due to changes in the dust and pollutants mixing or changes in the dust-sources



property of aerosol dust:

number size distribution

mass concentration

chemical composition

mixing state

mineralogy

optical properties

Dust is a mixing of different minerals:

type	common name	formula empirica
clay	Montmorillonite	$\text{Na}_{0.2}\text{Ca}_{0.1}\text{Al}_2\text{Si}_4\text{O}_{10}(\text{OH})_2(\text{H}_2\text{O})_{10}\cdot n(\text{H}_2\text{O})$
clay	Smectite	$(\text{Na}, \text{Ca})\text{Al}_4(\text{Si}, \text{Al})_8\text{O}_{20}(\text{OH})_4\cdot 2(\text{H}_2\text{O})$
clay	Chlorite	$\text{Na}_{0.5}(\text{Al}, \text{Mg})_8(\text{Si}, \text{Al})_8\text{O}_{18}(\text{OH})_{12}\cdot 5(\text{H}_2\text{O})$
Ca rich	calcite	CaCO_3
Ca rich	dolomite	$\text{CaMg}(\text{CO}_3)_2$
Ca rich	gypsum	$\text{CaSO}_4\cdot 2(\text{H}_2\text{O})$
Ca rich	anhydrite	CaSO_4
SiO2	quartz	SiO_2
Feldspars	mocricline	KAISi_3O_8
Plagioclase feldspar	Var oligoclase	$(\text{Na}, \text{Ca})\{\text{Si}, \text{Al}\}_4\text{O}_8$
Plagioclase feldspar	Var albite	$\text{NaAlSi}_3\text{O}_8$
Plagioclase feldspar	Var anorthite	$\text{CaAl}_2\text{Si}_2\text{O}_8$
Oxides	hematite	Fe_2O_3
Oxides	goethite	$\text{FeO}(\text{OH})$
Oxides	gibbsite	$\text{Al}(\text{OH})_3$
Oxides	rutile	TiO_2
Salt	halite	NaCl

close chemical composition,
but different mineralogy



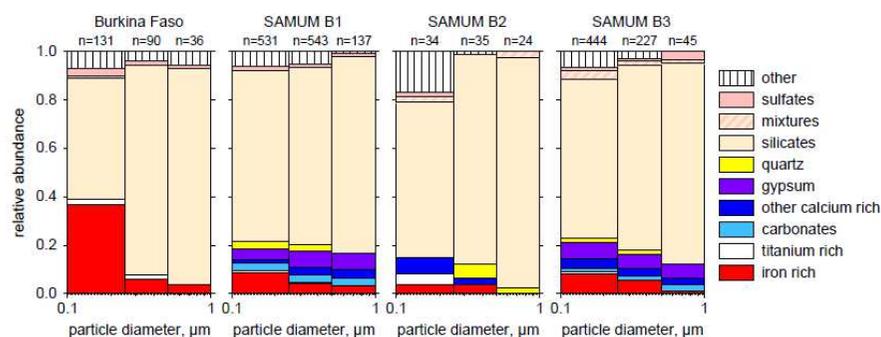
different optical properties

Dust is a mixing of different minerals:

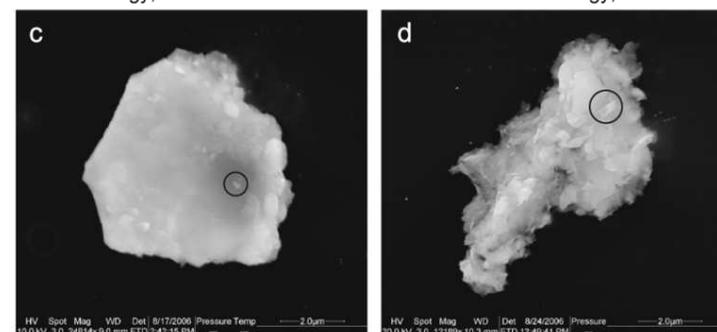
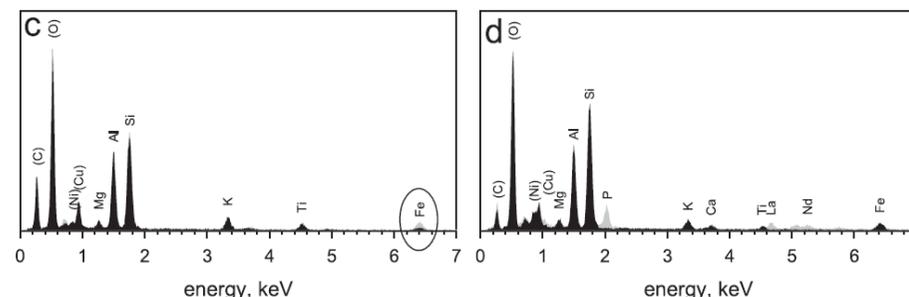
Techniques for the identification of different minerals:

- X-Ray diffraction
- Scanning Electron Microscopy (SEM) and Transmission Electron Microscopy (TEM) individual particle characterization for size, morphology, chemical and mineral composition.

not for long term monitoring



Component	Burkina Faso	SAMUM B1	SAMUM B3
"Quartz"	0.004	0.012	0.006
"Hematite"	0.027	0.016	0.011
"Calcite"	0.001	0.092	0.055
"Average Silicate"	0.949	0.842	0.867
"Sulphate"	0.019	0.037	0.061



Chemical composition and complex refractive index of Saharan Mineral Dust at Izaña, Tenerife (Spain) derived by electron microscopy

Konrad Kandler^{a,*}, Nathalie Benker^a, Ulrich Bundke^b, Emilio Cuevas^c, Martin Ebert^a, Peter Knippertz^d, Sergio Rodríguez^{c,e}, Lothar Schütz^d, Stephan Weinbruch^a

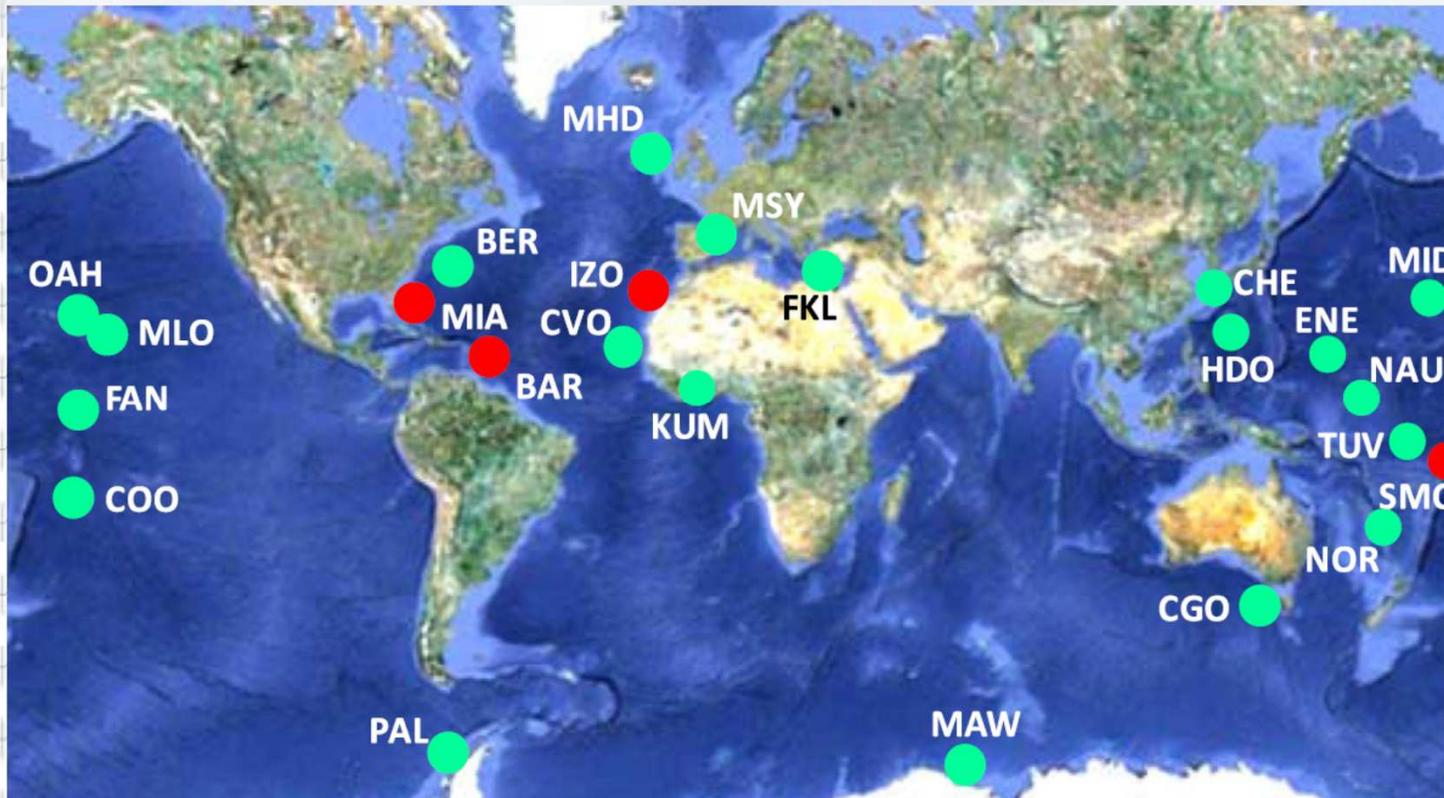
Atmospheric Environment 41 (2007) 8058–8074

Complex refractive indices of Saharan dust samples at visible and near UV wavelengths: a laboratory study

R. Wagner¹, T. Ajtai², K. Kandler³, K. Lieke³, C. Linke¹, T. Müller⁴, M. Schnaiter¹, and M. Vragel¹

Atmos. Chem. Phys., 12, 2491–2512, 2012

Long term monitoring dust background-observatories:



- at least 4 years
- Active during the last 20 years

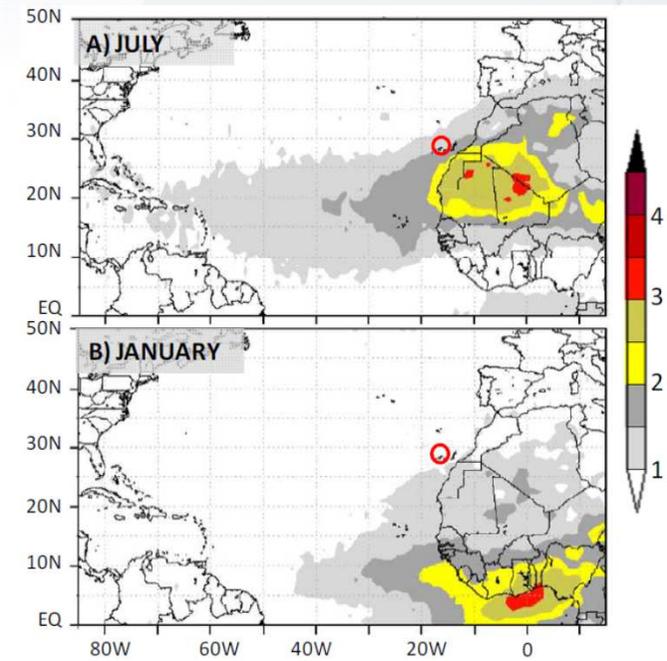
Review Article

Aeolian Research Aeolian Research 6 (2012) 55–74

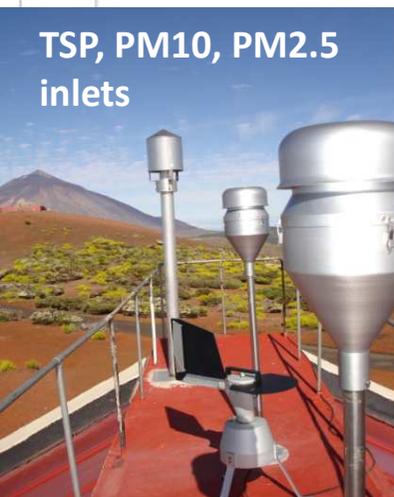
A review of methods for long term in situ characterization of aerosol dust

Sergio Rodríguez^{a,*}, Andrés Alastuey^b, Xavier Querol^b

Example of long term monitoring dust background-observatories, Izaña (Tenerife, The Canary Islands)



In-situ aerosols GAW program:



Chemical composition, TSP: 1987, PM_{2.5}: 2002, PM₁₀: 2005 ...

Ultrafine particles (CPC 3025A): 1997 - 2009

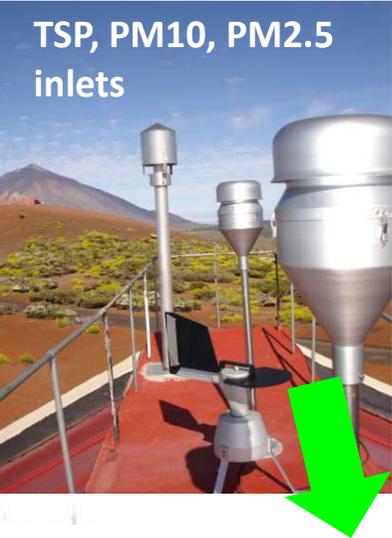
Size distribution of fine and ultrafine particles (SMPS): 2008 - ...

Size distribution of coarse particles (APS): 2006 - ...

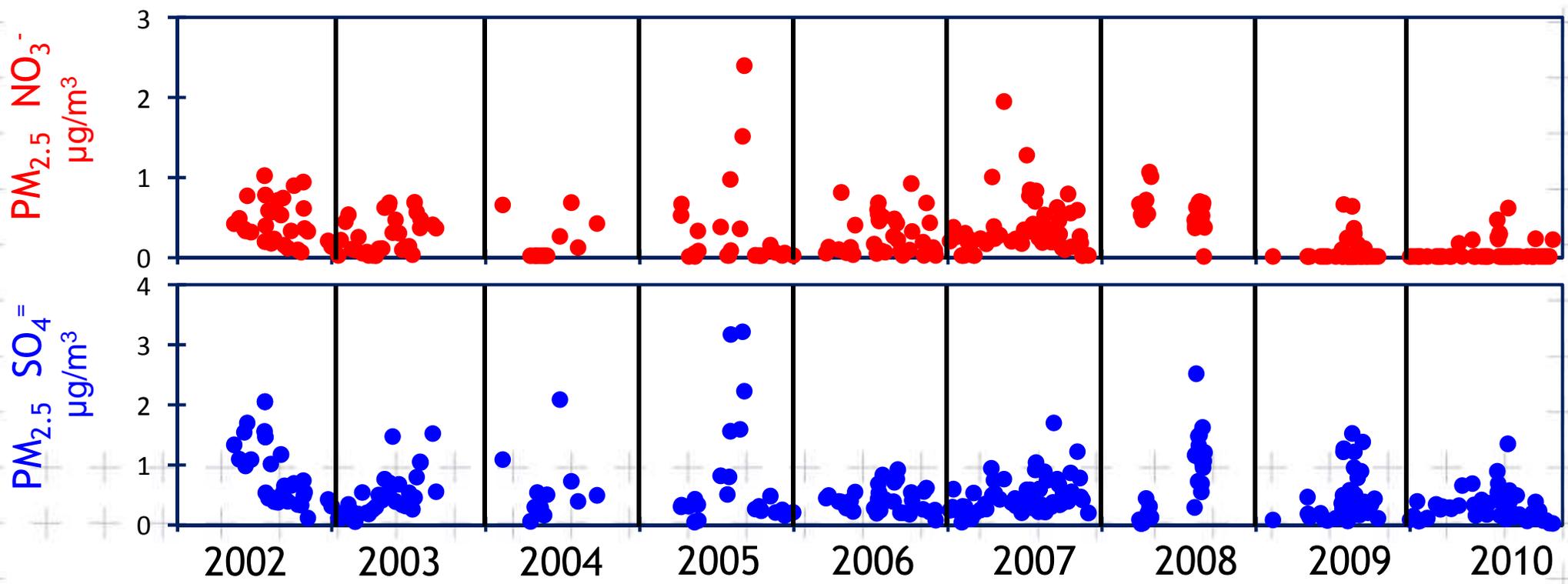
Scattering and backscattering (nephelometer): 2008 - ...

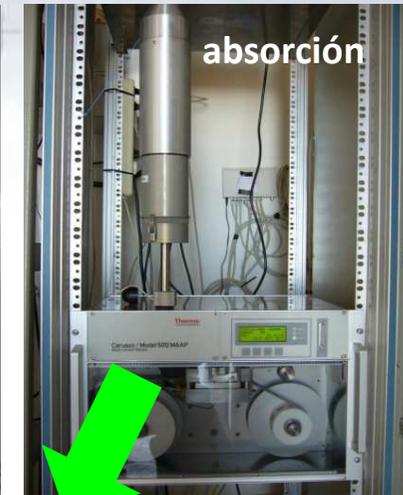
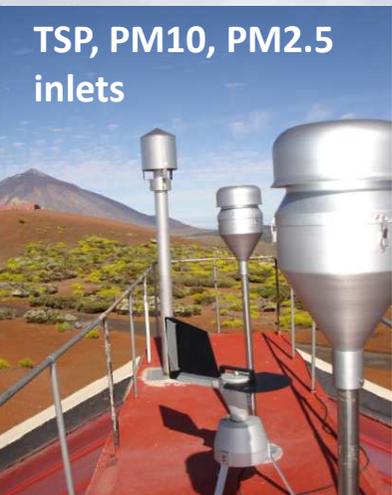
Absorption coefficient (1 λ): 2006 - ...

Absorption coefficient (7 λ): 2012 - ...

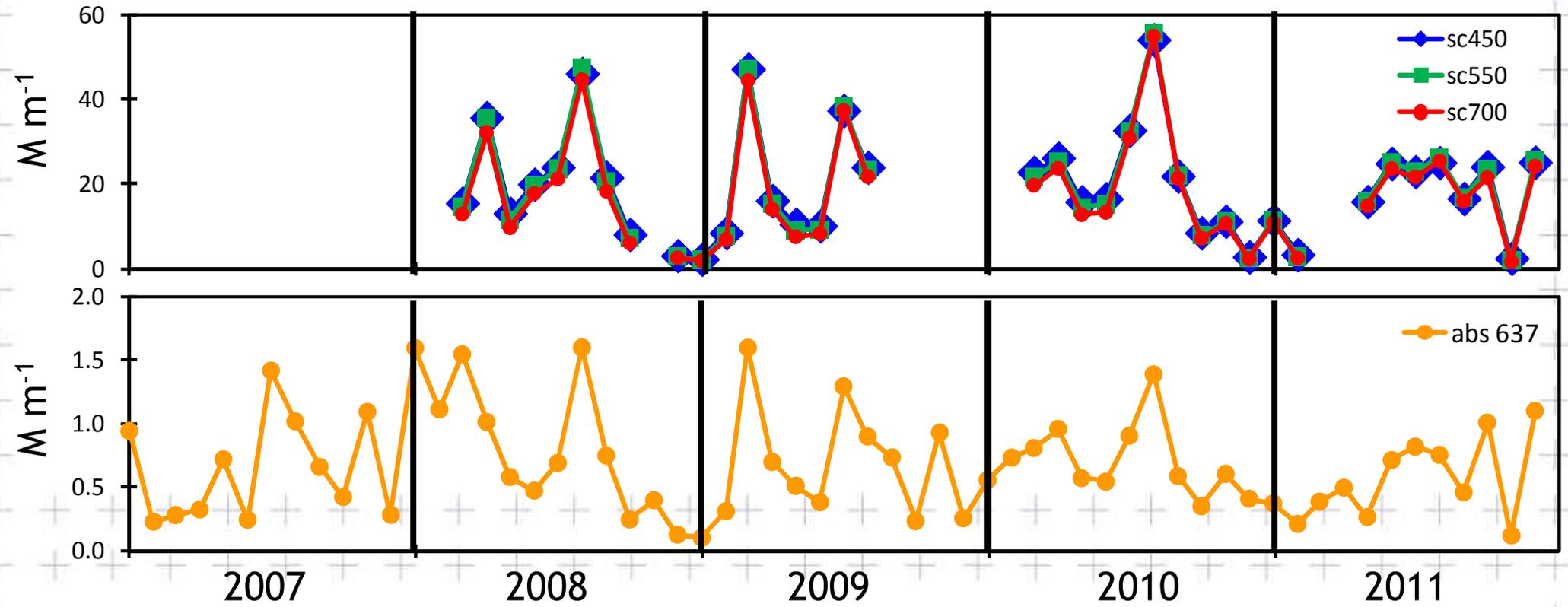


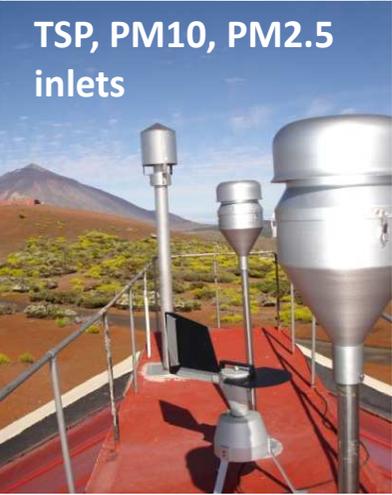
Chemical composition (TSP, PM₁₀, PM_{2.5}): elemental (ICP-AES+ICP-MS) , ions (SO₄⁼, NO₃⁻, NH₄⁺), OC, EC



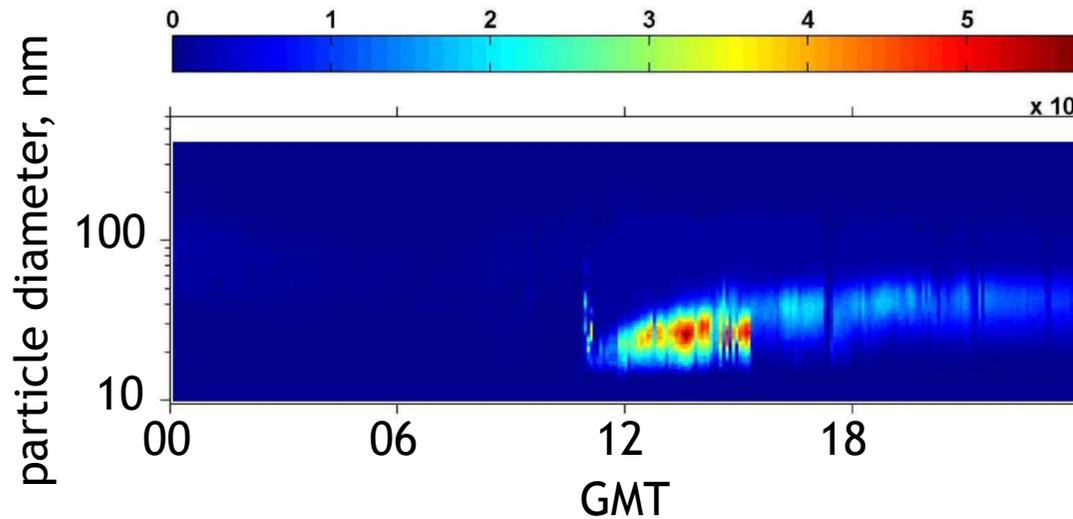


Optical properties: scattering and absorption





Size distribution: 10-500 nm (SMPS) + 0.5-20 μm (APS)

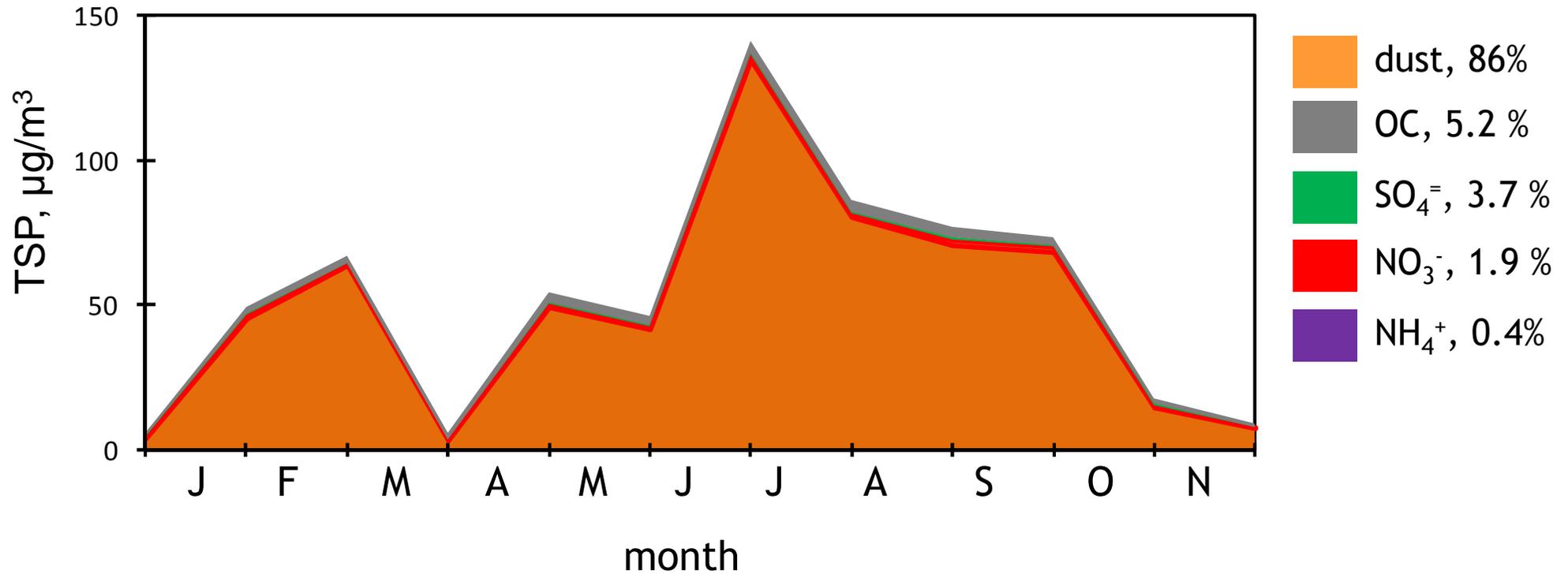


Example: new particle formation by nucleation

POLLUTANTS mixed with dust



Saharan dust is the most abundant aerosol we detect !!!!!



From ground observations...
to ground estimations...

Visibility

Index

- Aerosols and dust background
- In-situ dust characterization
- **In-situ dust estimations (Visibility)**
- Ground based remote sensing
- Recommended ground-based observations in Middle East
- Summary

WMO - visibility

The greatest distance that a black object of “suitable dimensions,” situated near the ground, can be seen and recognized when observed against a background of fog

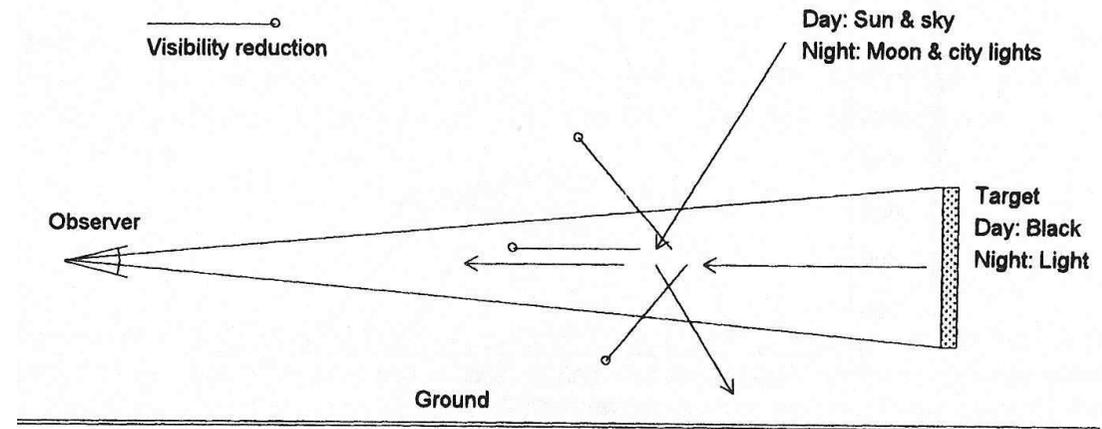


Fig. 11-1 Visibility reduction by scattering.

aerosols are the main cause of visibility reduction

- Operational surface synoptic weather station reports from Global Telecommunication System (GTS)
- Station reports include past & present weather, visibility (km), temperature (°C), dew point temperature, wind direction (°), and speed (knots)

62733	15.32	35.60	02040818	Dust, not at time of obs.	6	0	18	22	320	2	35.5			
62733	15.32	35.60	02041015	Dust, raised at time of obs.	7	0	99.	30	320	6	34.5			
62733	15.32	35.60	02041121		-9		-9	-9	-9	20	23	320	2	26.0
62733	15.32	35.60	02041212		-9		-9	-9	-9	20	34	340	3	37.5

Measurement of visibility - transmissometer & scatemeter

- A light source with one or two light detectors at fixed distances from the source
- Detectors are designed to receive light only from the source direction
- Often located along and parallel to a runway (runway visual range; RVR)

transmissometer

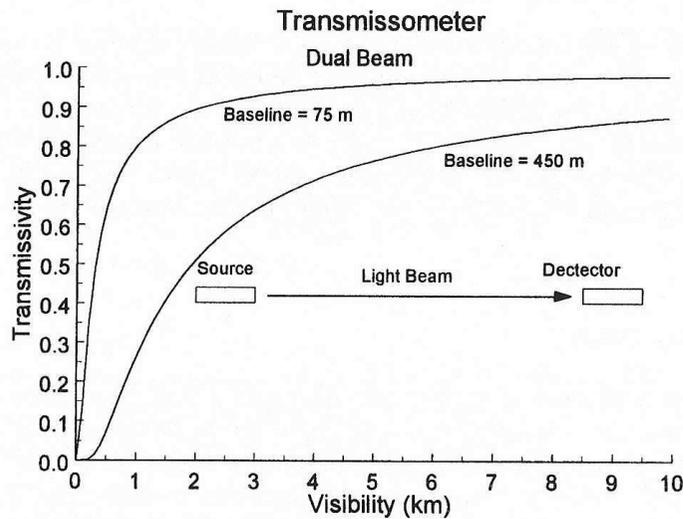


Fig. 11-3 Transfer function for a transmissometer.



scatemeter

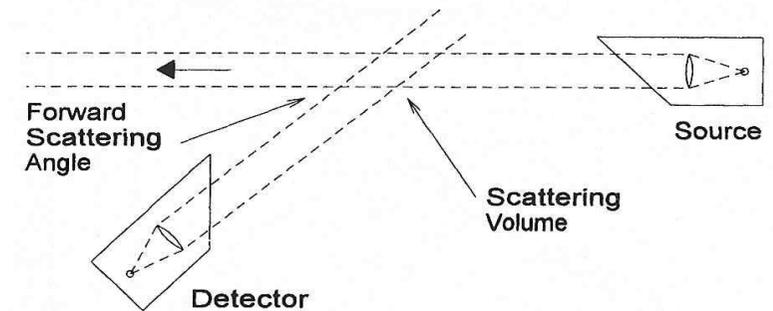
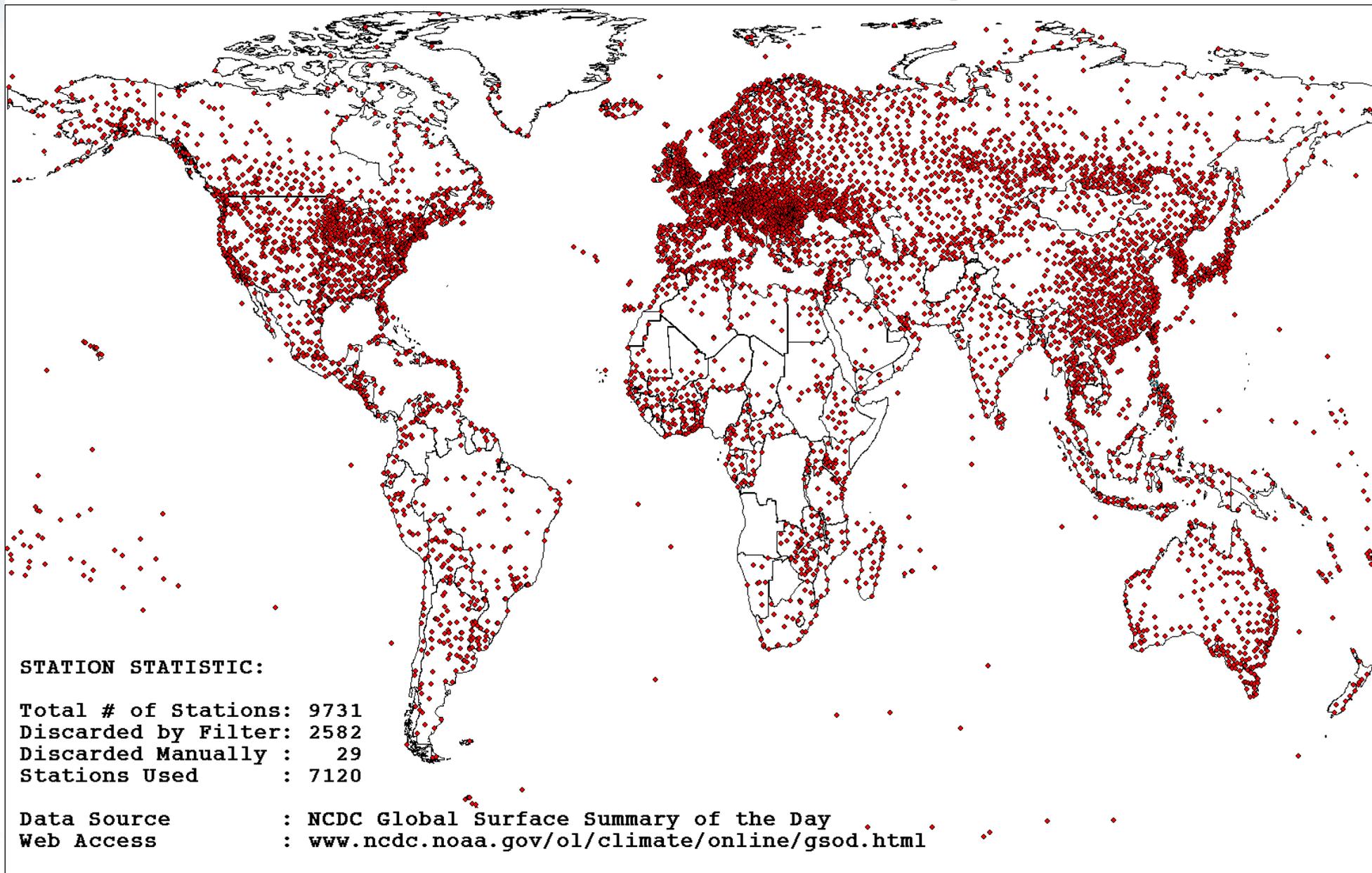


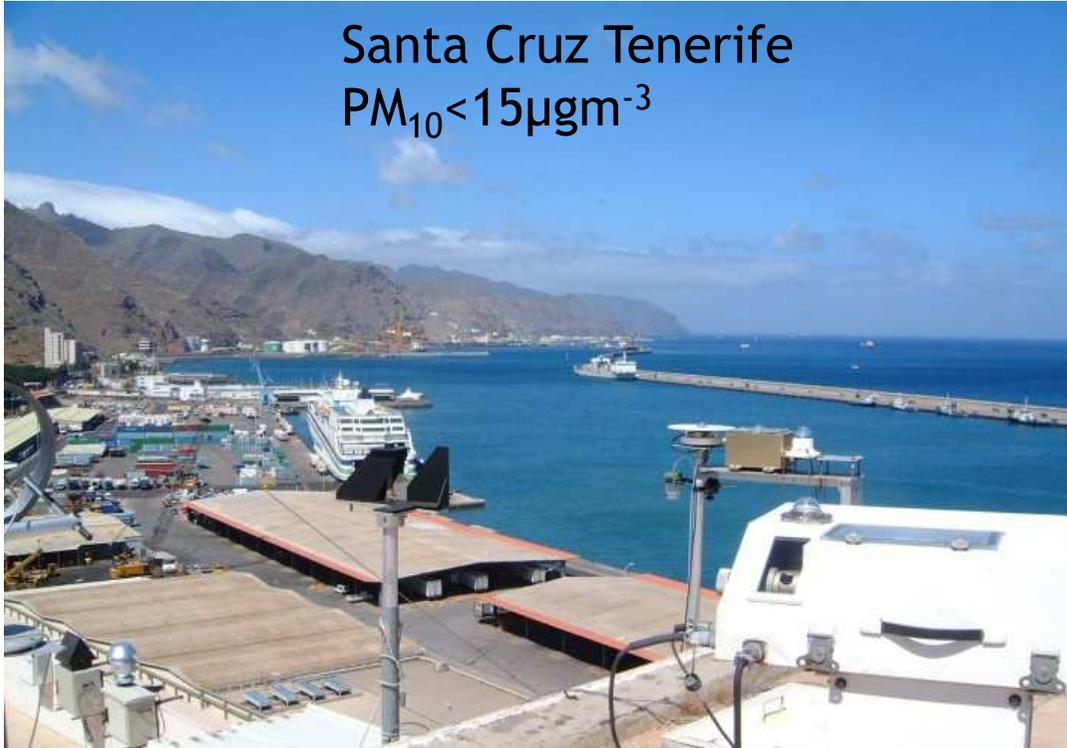
Fig. 11-4 A forward scatter visibility meter.

$$\text{Visual range (km)} = 3.912 / \sigma_{\text{ext}} \text{ (Mm}^{-1}\text{)}$$

WMO- World Wide Watch Global Surface Meteorological Network



Santa Cruz Tenerife
 $PM_{10} < 15 \mu g m^{-3}$



Santa Cruz Tenerife
 $PM_{10} > 60 \mu g m^{-3}$



Relation between horizontal visibility and TSP or PM10

Very few studies on the relation between horizontal visibility and TSP or PM10 levels of mineral dust mass concentration have been carried out in Africa.

$$C_{TSP} = 1339.84 VV^{-0.67}$$

Ben Mohamed et al. (1992)

$$C_{PM10} = 914.06 VV^{-0.73} + 19.03$$

D'Almeida's (1986)

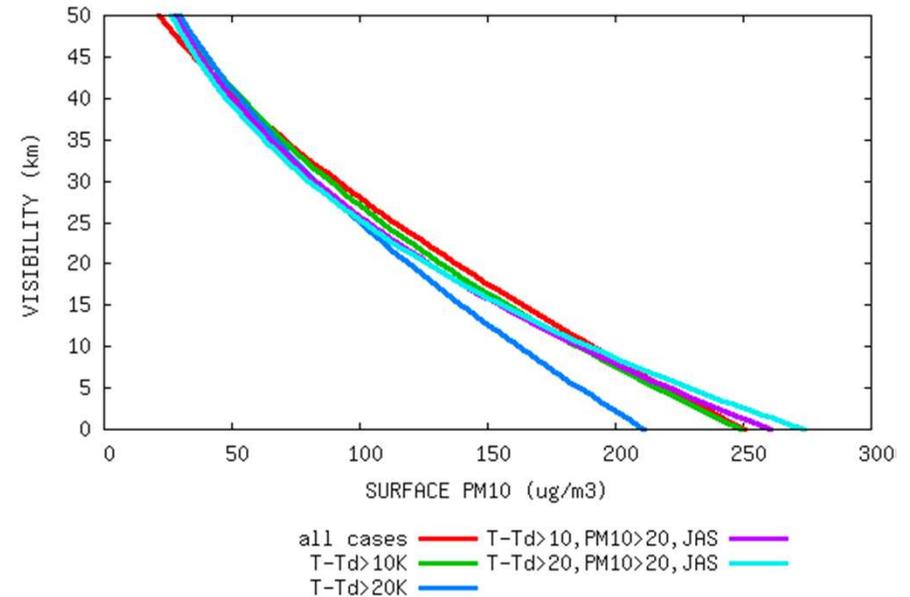
where C is the TSP concentration in $\mu\text{g m}^{-3}$ and VV is the horizontal visibility in km

E. Terradellas

$$\text{Vis} = 63023 - 1838(\text{PM}_{10})^{0.64}$$

PM₁₀ at Izaña

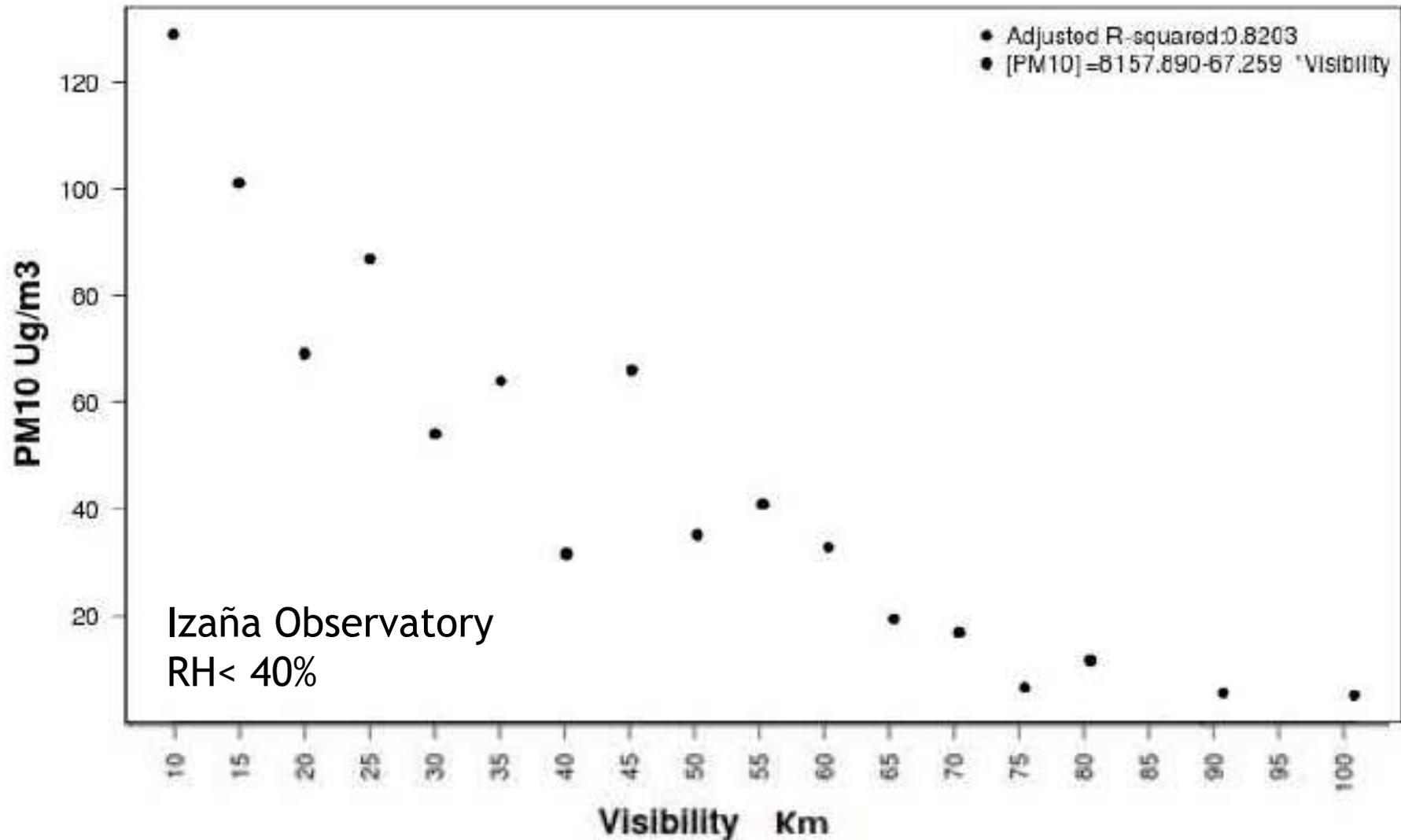
Visibility from SYNOP-Izaña



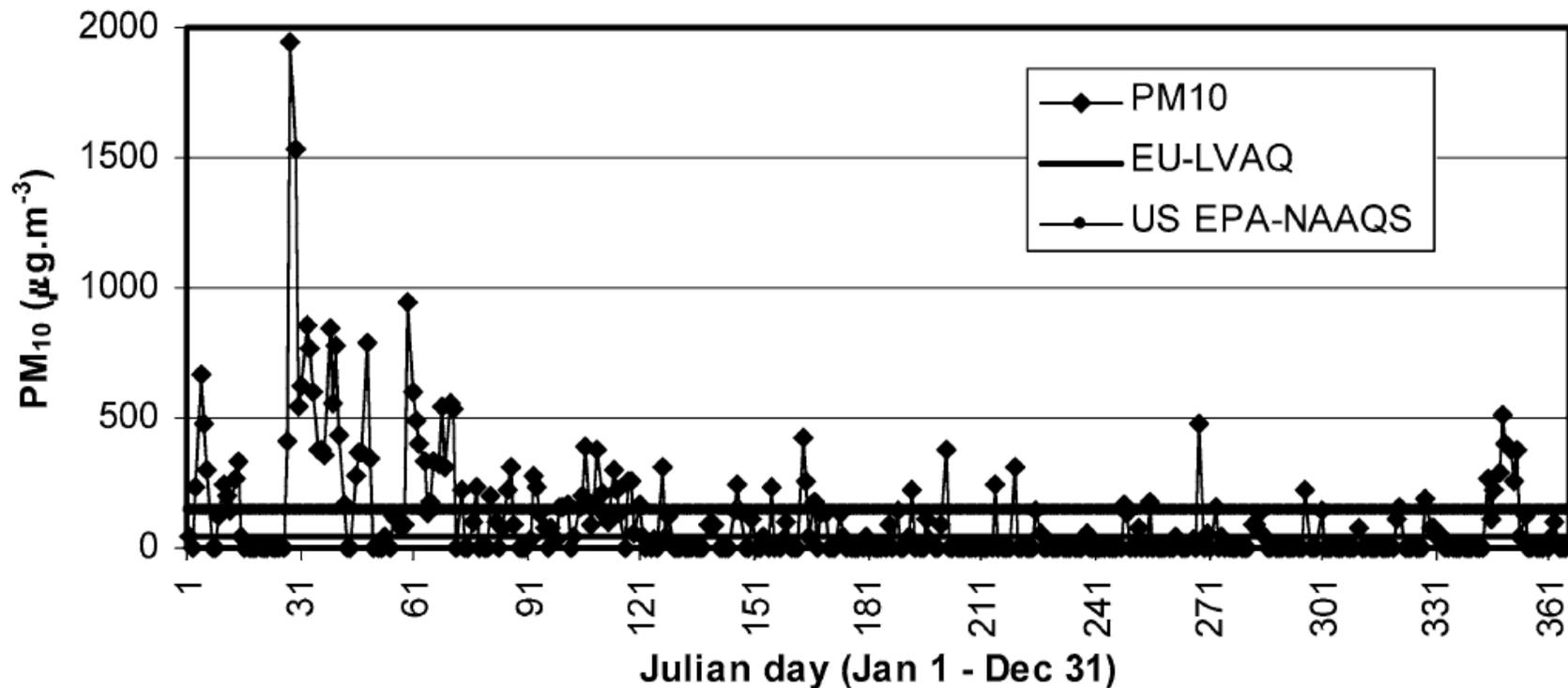
Identify surface station visibility reports that may be used in simple regression model for estimate ground PM10 or TSP

At above 60% RH, particles can experience hygroscopic growth because the water vapor can condense on the particles making them “Grow”

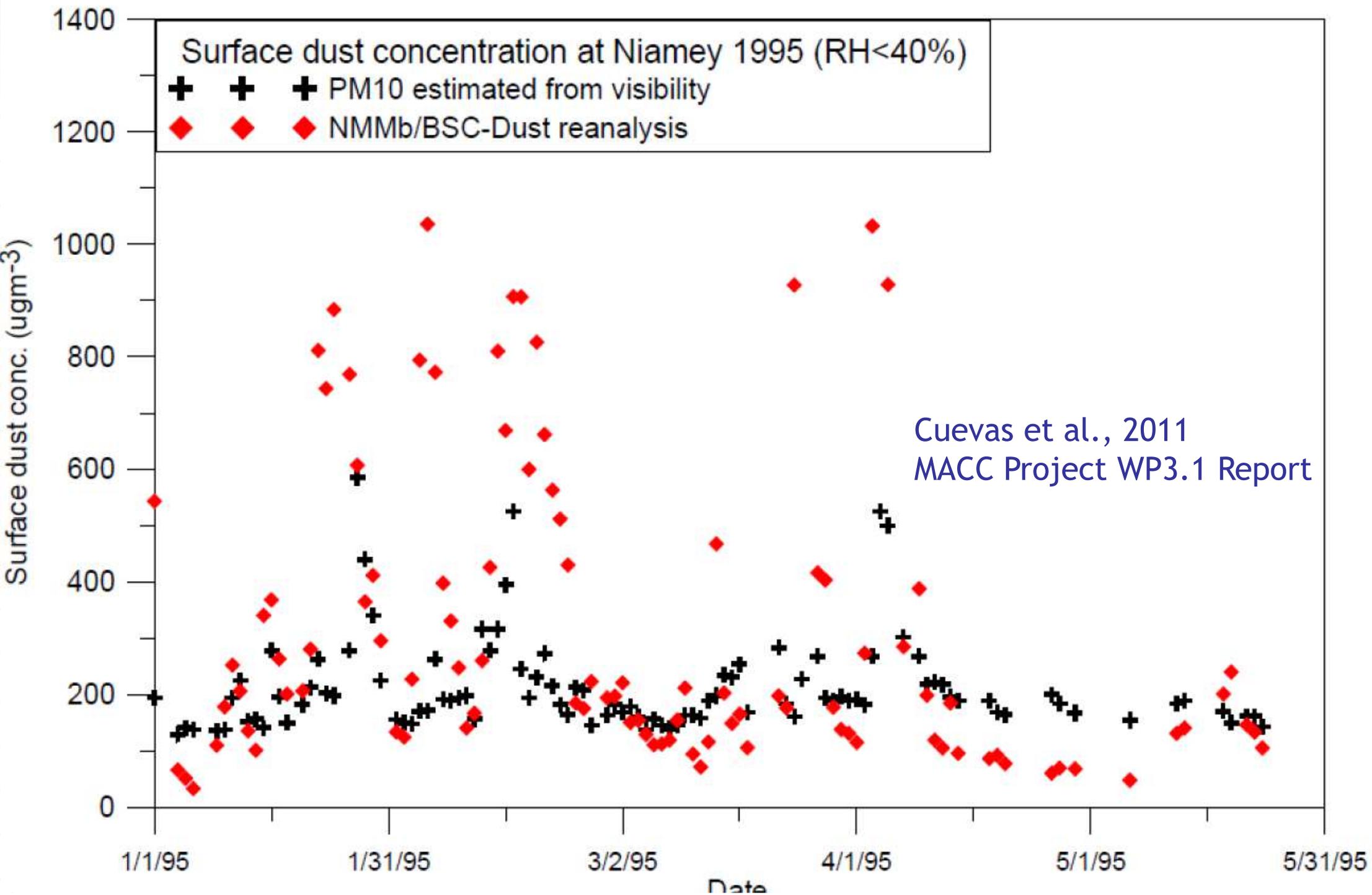
Scatterplot at Izaña



Variations of estimated daily mean concentrations of TSP and PM10 ($\mu\text{g}\cdot\text{m}^{-3}$) due to Saharan dust events at Nouakchott, Mauritania, in 2000

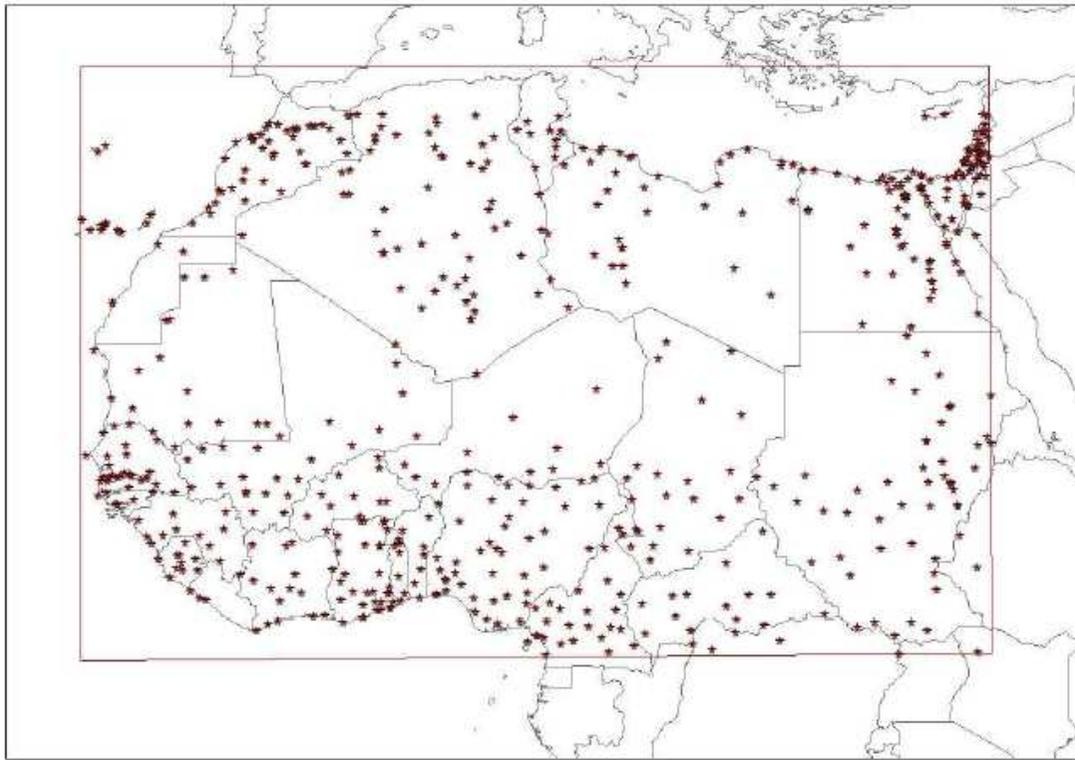


Ozer et al., (2006): Estimation of air quality degradation due to Saharan dust at Nouakchott, Mauritania, from horizontal visibility data

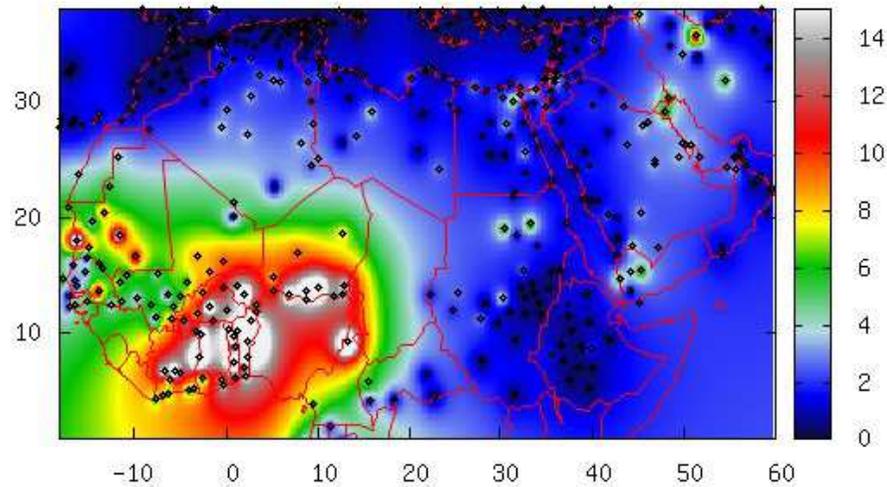


Synop and Metar reports from Met stations

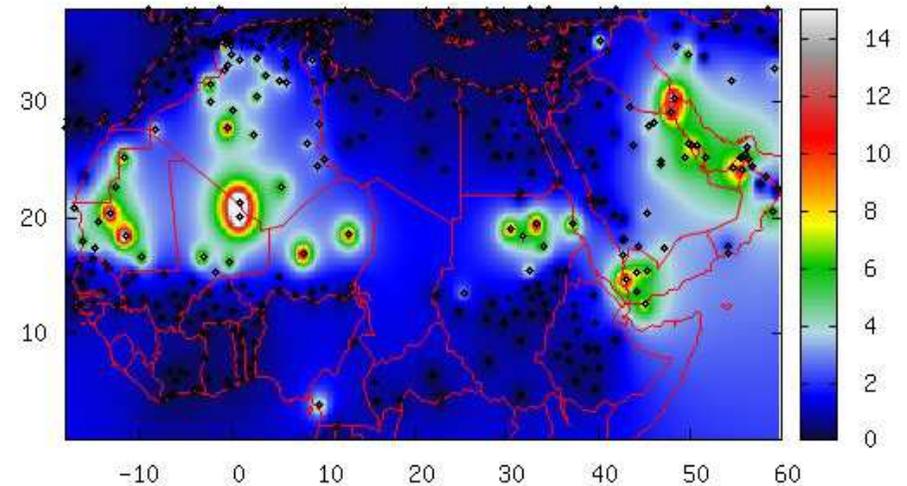
Provided by Enric Terradellas



WINTER



SUMMER



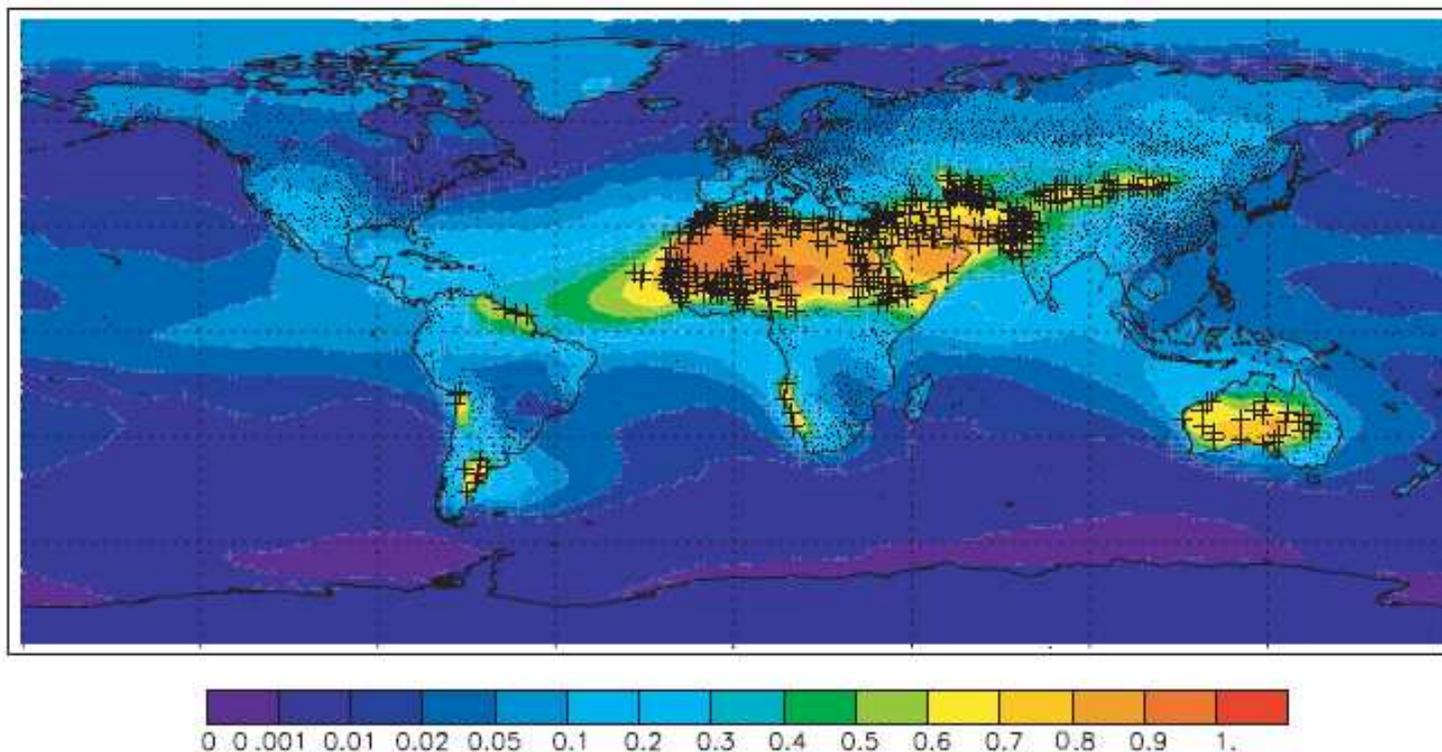


Fig. 4. Location of visibility stations with more than 30 years of data. Colored contours show fraction of surface extinction from desert dust. Pluses show stations in regions dominated by desert dust (>50%), while dots show other locations.

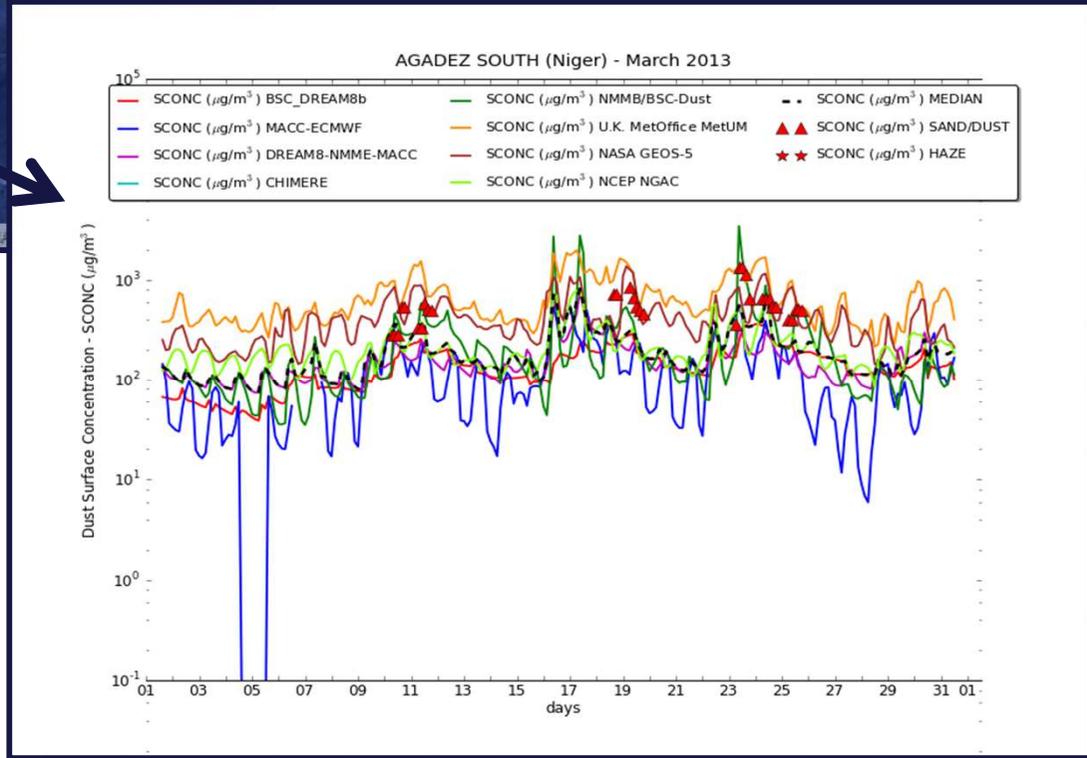
Mahowald et al. (2007) Atmos. Chem. Phys.; Global trends in visibility: implications for dust sources

SDS-WAS Regional Center NAMEE : Evaluation using VISIBILITY data



AGADEZ SOUTH, Niger
March 2013

PM10 = 1339.84 V^{-0.67}
Ben Mohamed et al. (1992)



Problems with station visibility estimates

1. Human observations are inherently subjective.
2. No all reductions of visibility are due to dust (fog, biomass burning...)
3. Coarse reporting bins
4. Judgment in distinguishing visibility beyond 10 km

Main advantages

1. Reports are abundant and widespread over land. There is information in remote areas (deserts)
2. There are *some* standards
3. Human detected visibility has been correlated well with surface extinction analyses (Husar et al., 2000)
4. Estimations of PM are possible

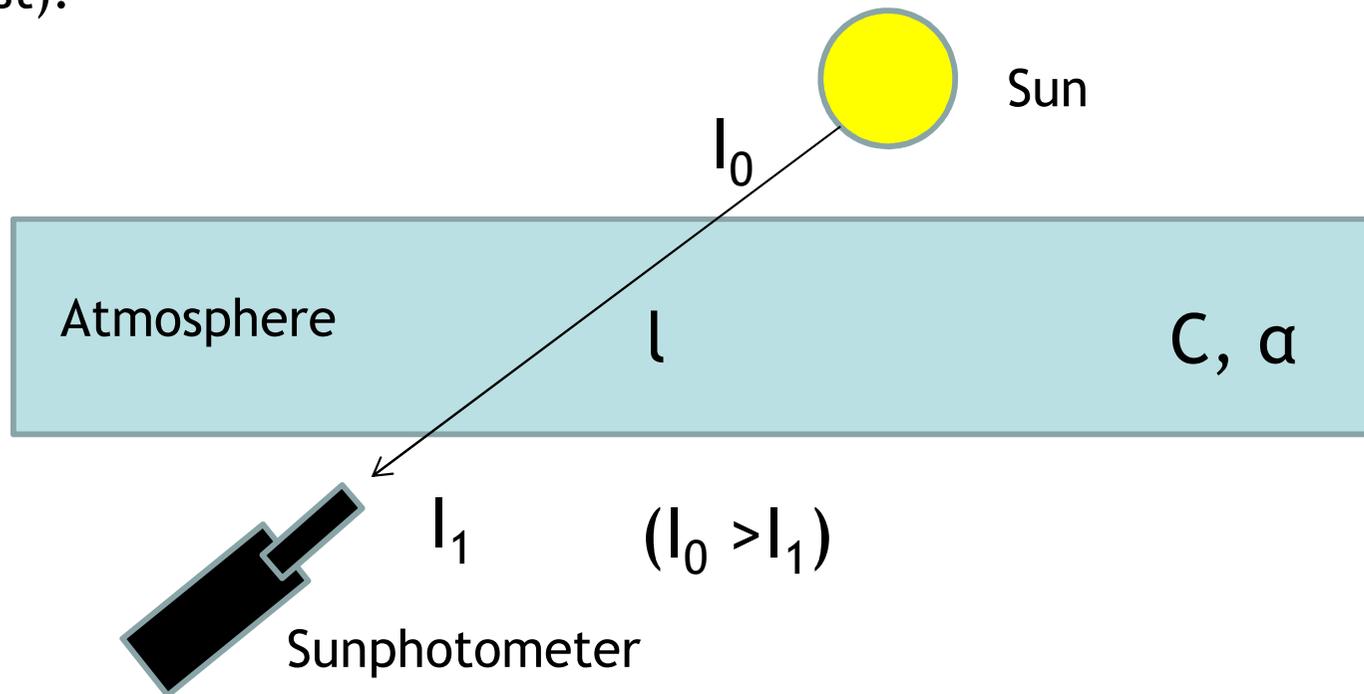
From ground observations...
to total atmospheric column observations

Sunphotometers

Index

- Aerosols and dust background
- In-situ dust characterization
- In-situ dust estimations (Visibility)
- **Ground based remote sensing**
- Recommended ground-based observations in Middle East
- Summary

Knowing the sunlight's energy at the top of the atmosphere, the thickness of the atmosphere, and the amount of sunlight transmitted to the earth's surface and can allows us to determine the amount of scattering, and thus, the amount of aerosols (dust).



Beer's Law

$$I = I_0 \cdot e^{-\sigma_{\text{ext}} \cdot L}$$

Transmissivity (T)

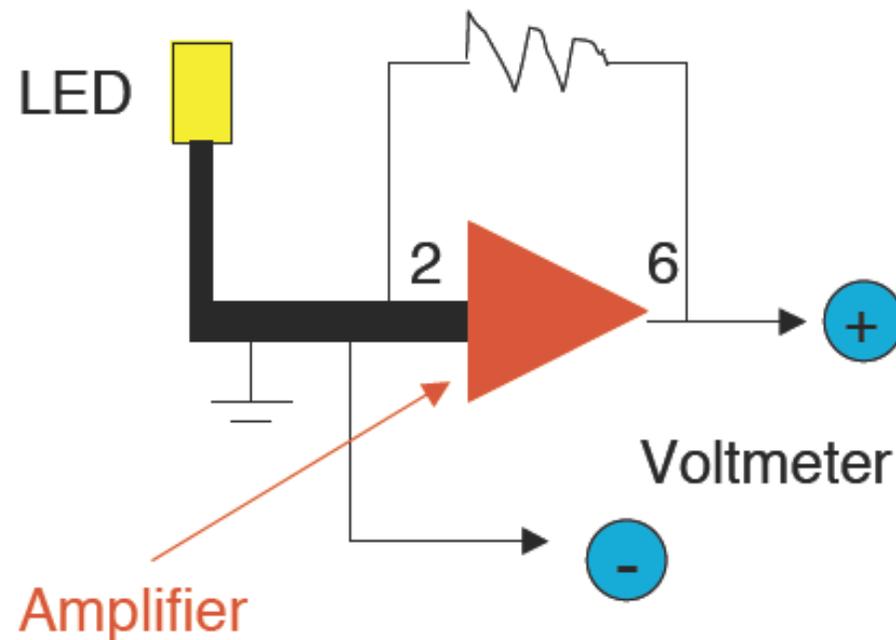
Extinction coefficient (σ_{ext}): ϵC

path length (L)

molar absorptivity of the absorber (ϵ)

concentration of absorbing species in the material (C)

Sun Photometers absorb *direct* sunlight energy with a LED light and convert the intensity into a quantified voltage to measure aerosols in the atmosphere.



The intensity of sunlight at the top of the earth's atmosphere is constant. While the sunlight travels through the atmosphere, aerosols can dissipate the energy by scattering (Rayleigh and Mie) and absorbing the light. More aerosols in the atmosphere cause more scattering and less energy transmitted to the surface.

ASSESSMENT OF OBSERVATIONS CONSISTENCY

Langley plot calibration (100 determination for each wavelength):

$$I = I_0 \cdot e^{-\sigma_{\text{ext}} \cdot L}$$

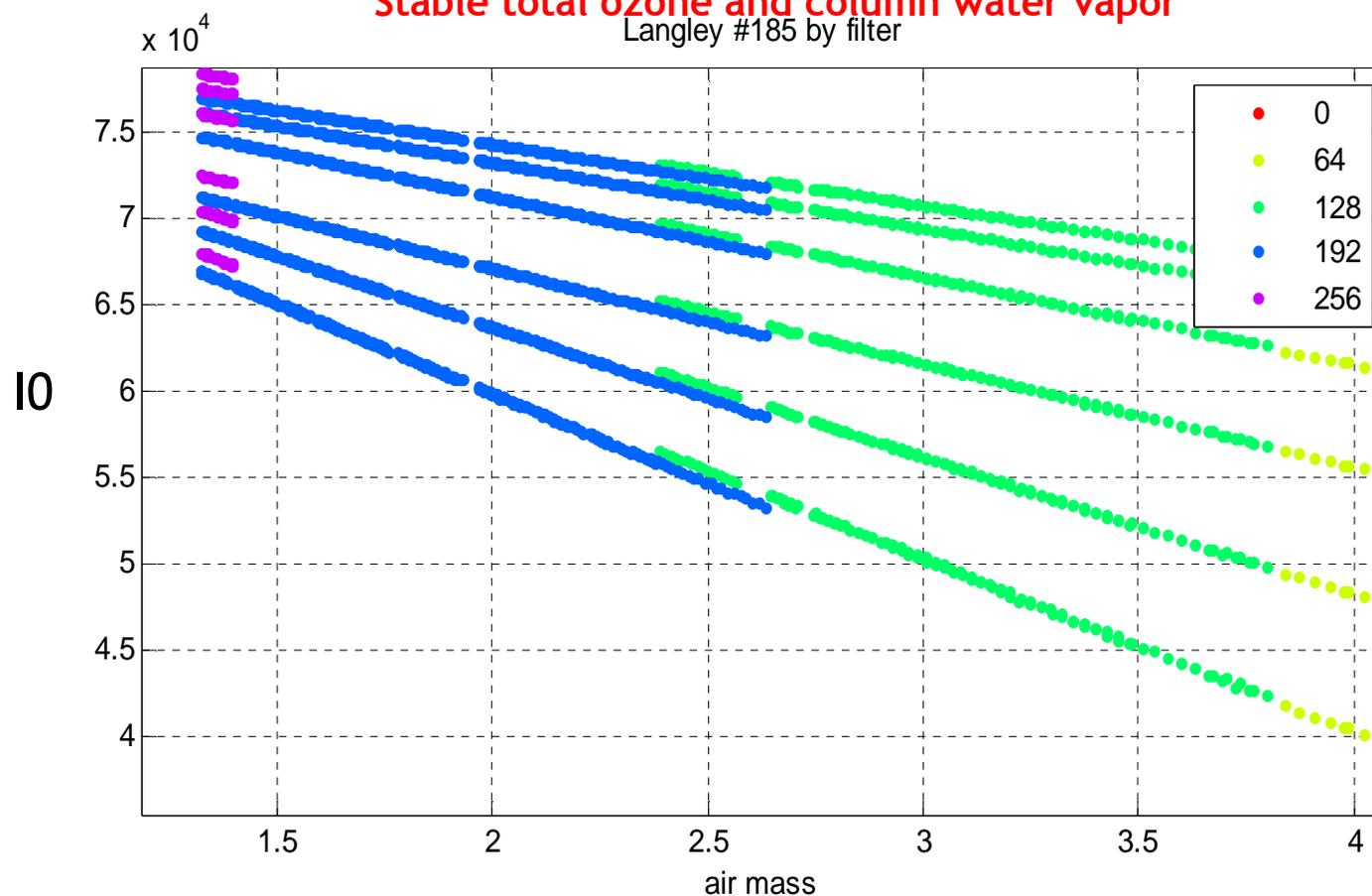
$$\ln I = \ln I_0 - \sigma_{\text{ext}} L$$

If σ_{ext} is constant during the observation We can determine I_0

Pristine conditions (very low and constant aerosol load)

No clouds

Stable total ozone and column water vapor



CONCEPTS:

Aerosol Extinction: A measure of attenuation of the light passing through the atmosphere due to scattering and absorption by aerosol particles.

Extinction coefficient is the fractional depletion of radiance per unit path length (also called attenuation). It has units of km^{-1} .

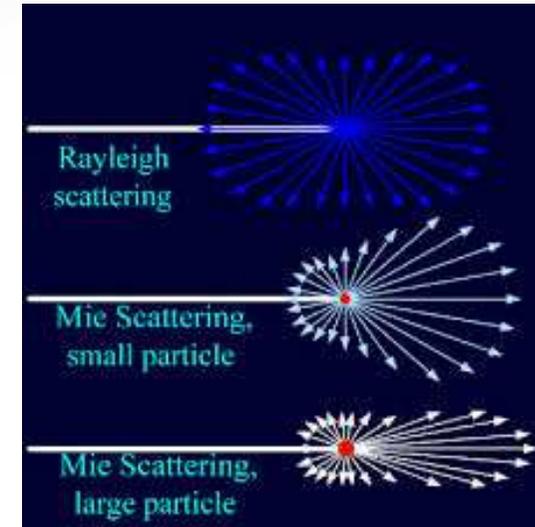
Aerosol Mass Load: The columnar aerosol mass concentration ($\mu\text{g}/\text{cm}^2$) is the total aerosol mass in a vertical column of atmosphere.

CONCEPTS:

Aerosol Asymmetry Factor A measure of the preferred scattering direction (forward or backward) for light encountering aerosol particles.

$$g = \frac{1}{2} \int_{-1}^{+1} \cos \Theta P(\cos \Theta) d \cos \Theta$$

$$P(\cos \Theta) = \frac{1 - g^2}{(1 + g^2 - 2g \cos \Theta)^{3/2}}$$



In general, $g=0$ indicates scattering directions evenly distributed between forward and backward directions, i.e. isotropic scattering (e.g. scattering from small particles)

$g < 0$ scattering in the backward direction (i.e. scattering angle > 90 deg.), often referred to as backscattering, is scattering at 180 deg.

$g > 0$ scattering in the forward direction (i.e. scattering angle < 90 deg.), often referred to as forward-scattering, is scattering at 0 deg. **For larger size or Mie particles, g is close to $+1$.**

CONCEPTS:

Aerosol Optical Depth (or Thickness)

"Aerosol Optical Depth" (AOD) is the degree to which aerosols prevent the transmission of light. The aerosol optical depth or optical thickness (τ) is defined as the integrated extinction coefficient over a vertical column of unit cross section.

$$AOD = \int_{z=0}^{z=toa} \sigma_{ext}(z) dz$$

Angstrom Exponent (α)

An exponent that expresses the spectral dependence of Aerosol Optical Depth (τ) with the wavelength of incident light (λ). The spectral dependence of aerosol optical thickness can be approximated (depending on size distribution) by:

$$AOD = \beta \lambda^{-\alpha}$$

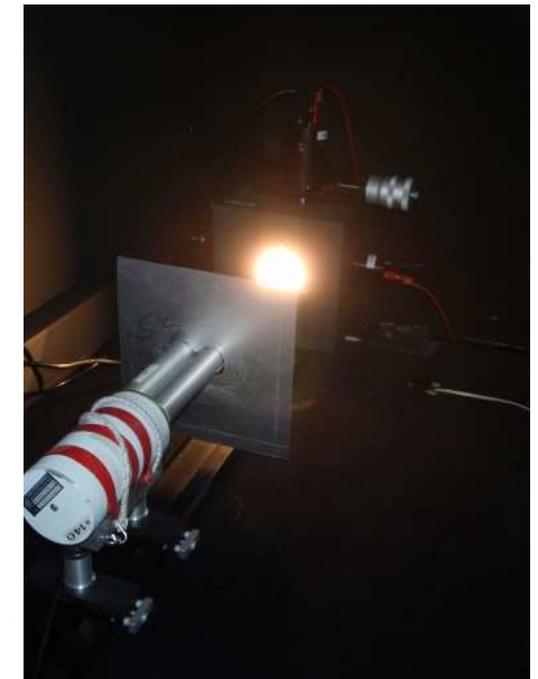
$\alpha \gg 0.9$ FINE particles

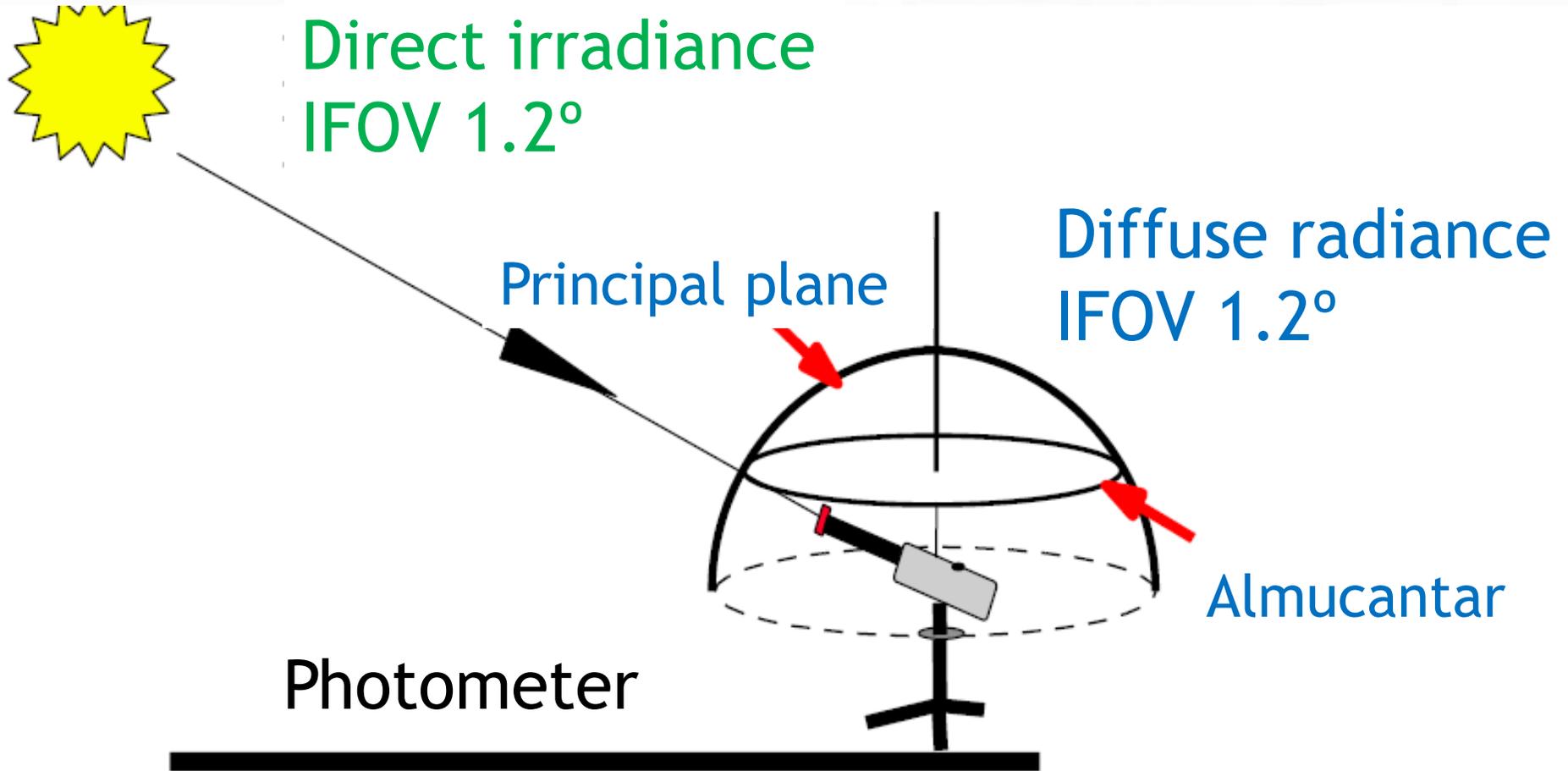
$\alpha \ll 0.7$ COARSE particles

where α is the Angstrom exponent (β = aerosol optical depth at 1 μm)

i.e. If $AOD > \sim 0.2$ and $\alpha < 0.7$ then we are observing dust (aprox.)

- The Cimel Electronique 318 spectral radiometer is a solar-powered, weather-hardy, robotically-pointed sun and sky spectral sun photometer.
- A sensor head points the sensor head at the sun according to a preprogrammed routine.
- The Cimel controller, batteries, and the optional Vitel satellite transmission equipment are usually deployed in a weatherproof plastic case.





Sun measurements
Sky measurements

AERONET Aerosol Robotic Network-Twenty Years of Observations and Research

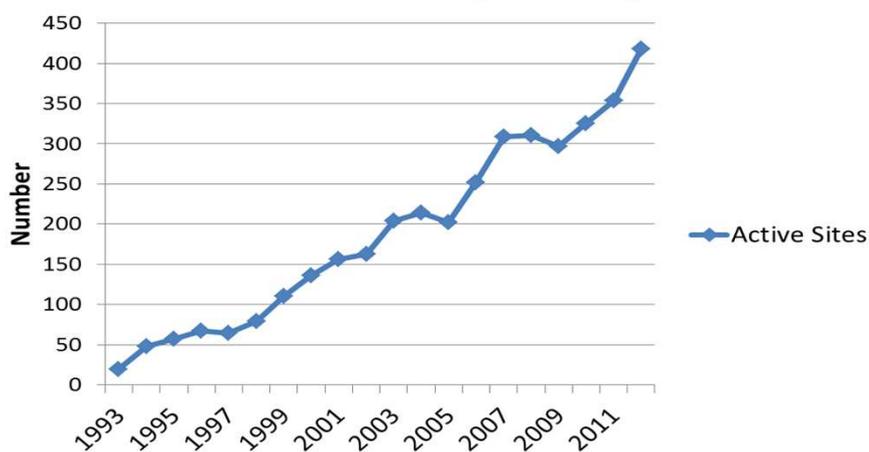
The **AERONET program** is a federation of ground-based remote sensing aerosol networks established by NASA and LOA-PHOTONS (CNRS) and has been expanded by collaborators from international agencies, institutes, universities, individual scientists and partners.



- >7000 citations
- >400 sites
- Over 80 countries
- <http://aeronet.gsfc.nasa.gov>



AERONET Growth (1993-2012)



AERONET provides a long-term, continuous public database of aerosol optical, microphysical, and radiative properties for aerosol research and characterization, validation of satellite measurements, and synergism with other databases.

AERONET Data Flows

<http://aeronet.gsfc.nasa.gov>

Holben et al.
RSE, 1998
Holben et al.
JGR, 2001

Eck et al.
JGR, 1999

Smirnov et al.
RSE, 2000

Dubovik and King
JGR, 2000
Dubovik et al.
JGR, 2000
GRL, 2002

Flux measurements

Direct - $\lambda=340, 380, 440, 500, 670, 870, 940, 1020$ nm

Diffuse - $\lambda=440, 670, 870, 1020$ nm (alm, pp, pol)

Calibration and processing information

Mauna-Loa and Izaña

CNRS-University of Lille and University of Valladolid

Aerosol optical depth and precipitable water computations

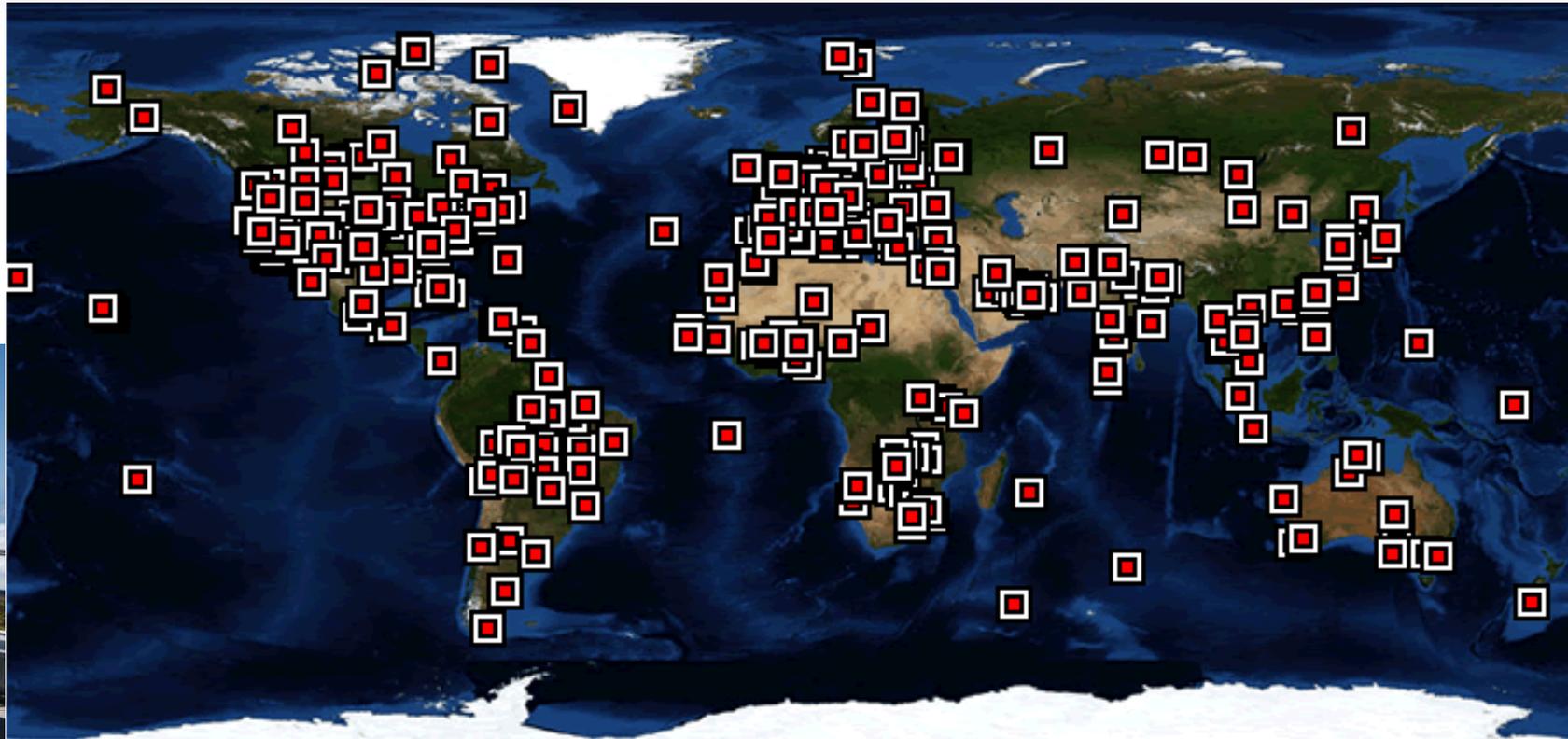
Cloud screening and quality control

Inversion products

Volume size distribution ($0.05 < \text{size} < 15 \mu\text{m}$),
refractive index, single scattering albedo
($\lambda=440, 670, 870, 1020$ nm)

AERONET (AErosol RObotic NETwork)-

<http://aeronet.gsfc.nasa.gov>



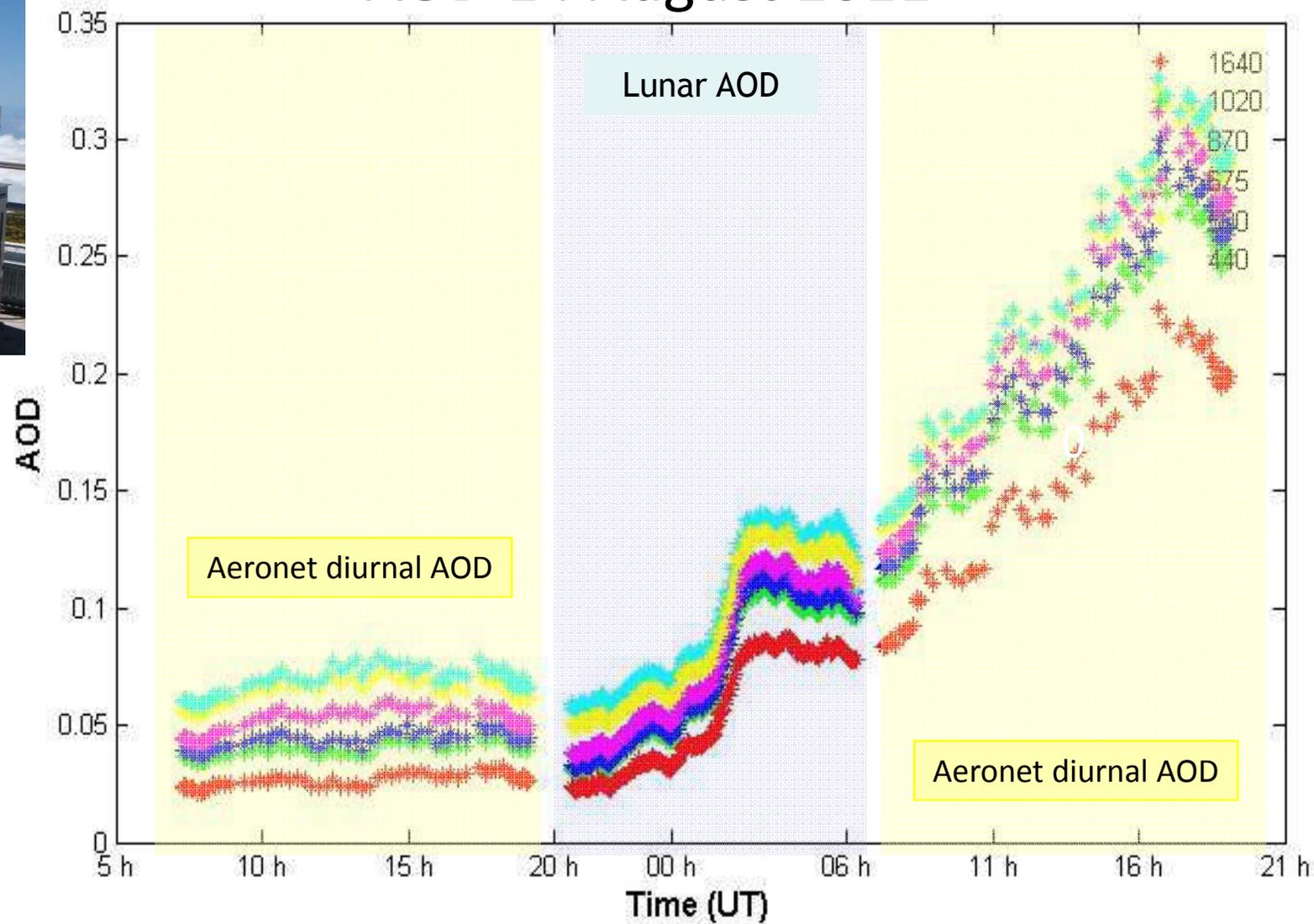
An internationally Federated Network

- Characterization of aerosol optical properties
- Validation of satellite aerosol retrieval
- Near real-time acquisition; long term measurements

AERONET provides:

- global Aerosol Optical Depth of Dust in near real-time
- robust optical properties of Dust: size distribution, ref. Index, etc. (e.g. Asian Dust has stronger and less spectral dependent absorption than Saharan Dust)
- climatological models that reproduce observed optical properties of aerosol (useful for satellite retrievals)

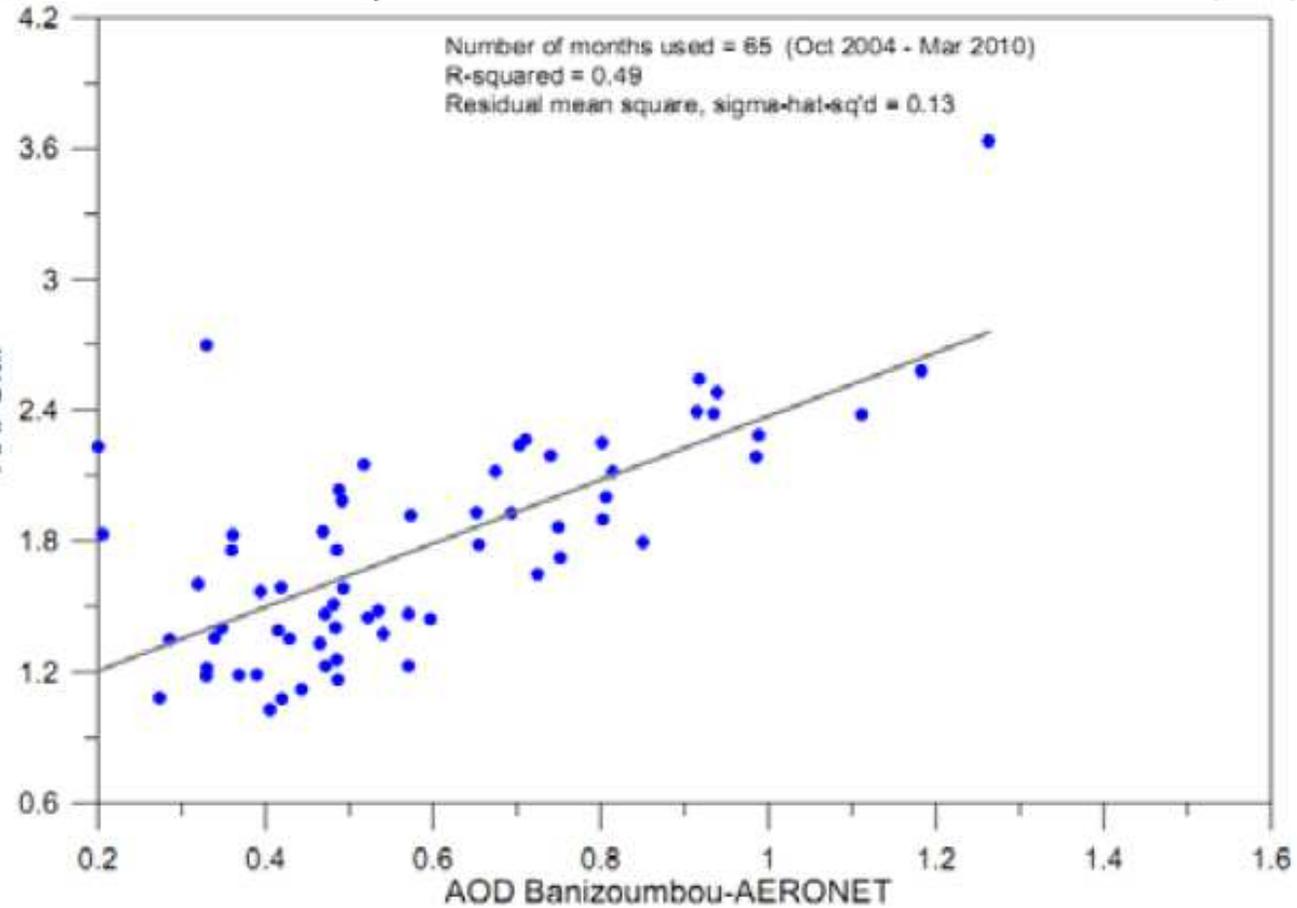
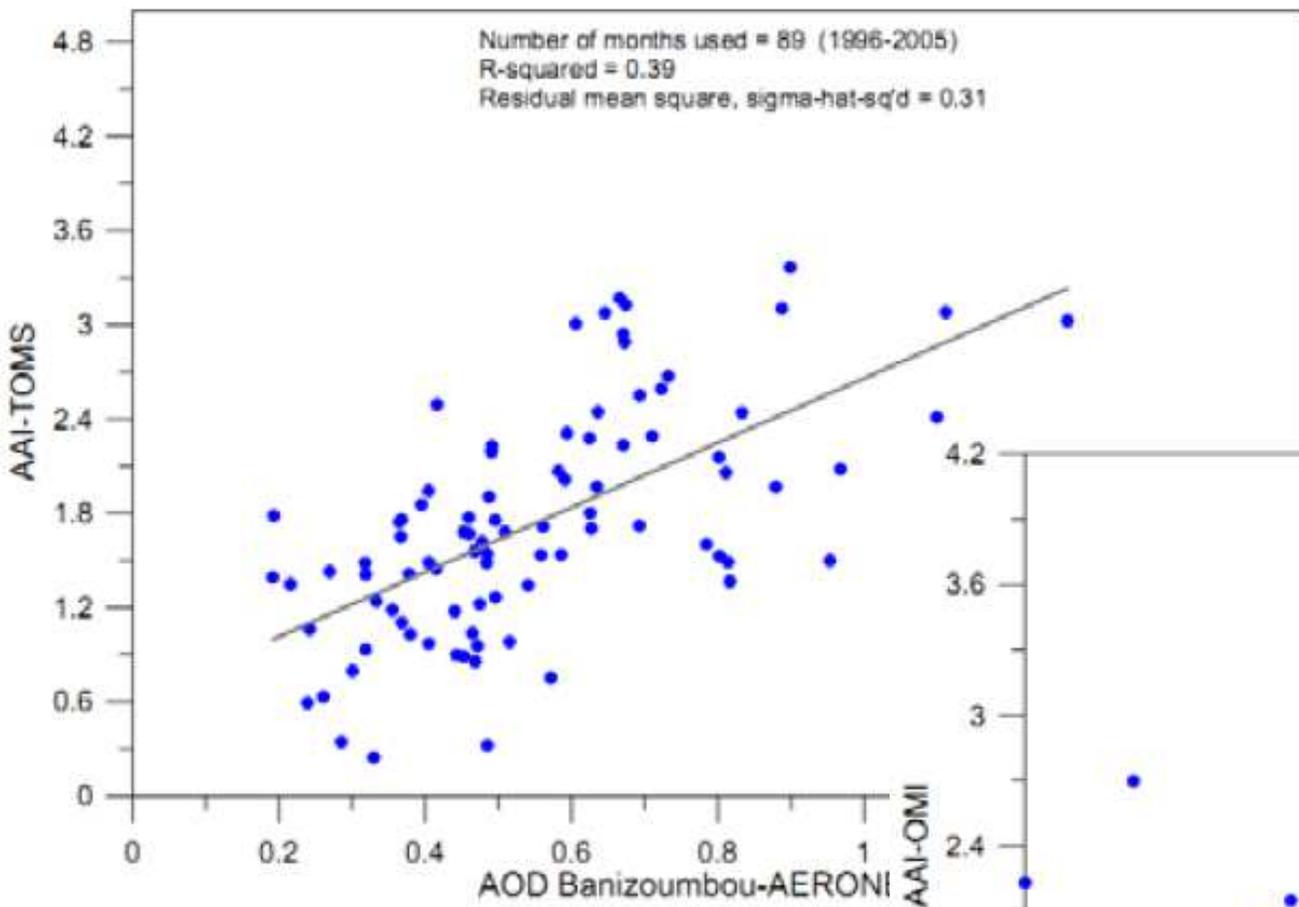
AOD 14 August 2011

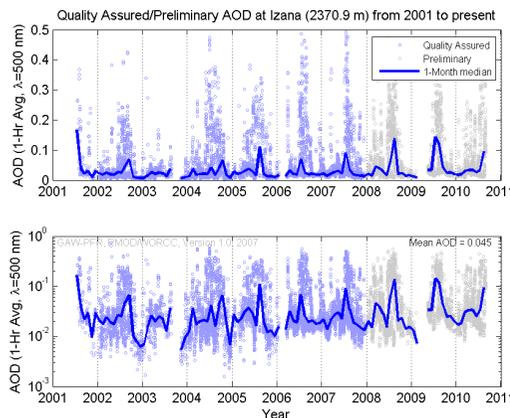
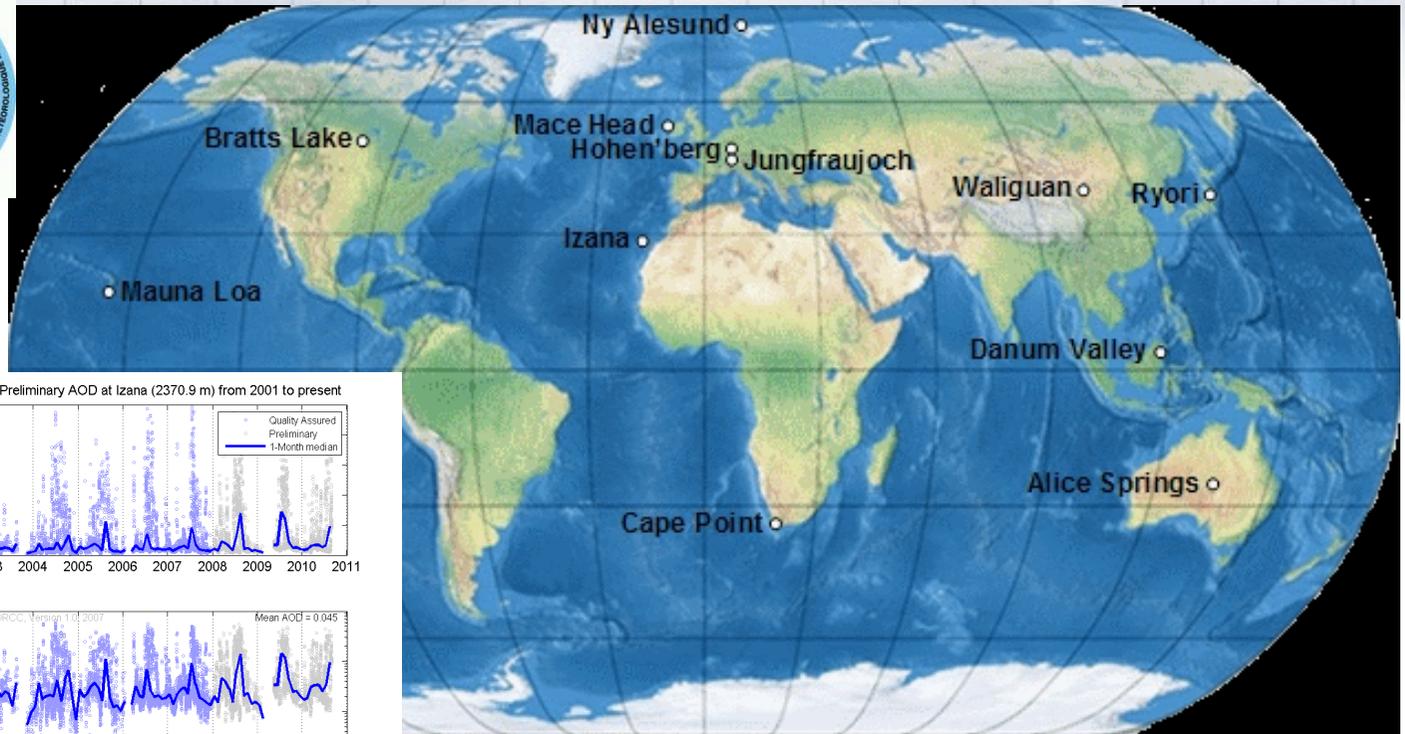


August 13

August 14







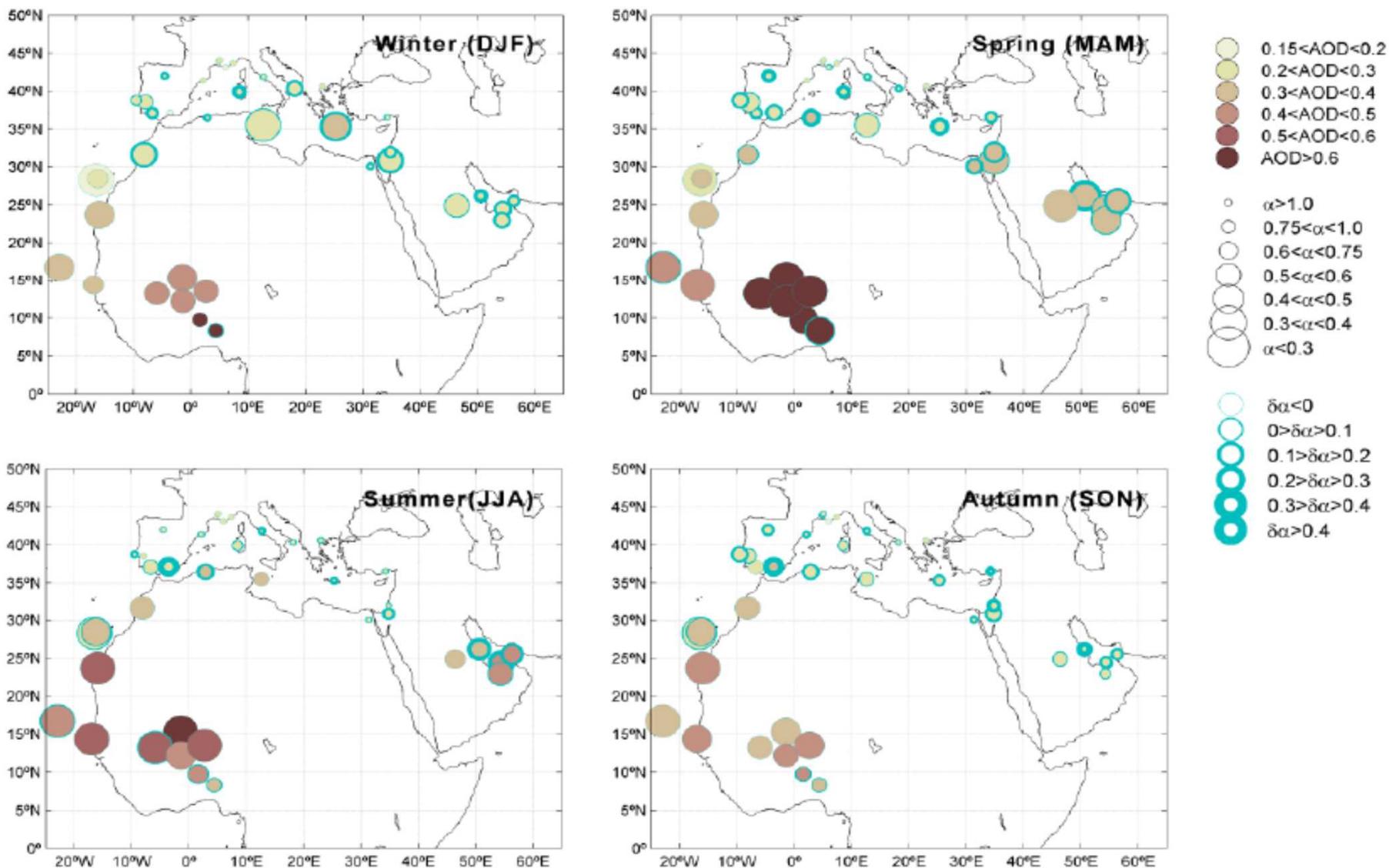
- Classic extinction measurements at the recommended 4 WMO wavelengths 368, 415, 500 and 862 nm using Precision Filter Radiometers (PFRs).
- Continuous sampling at a 1- minute frequency by automated systems.
- Data products: **AOD** and the **Angström coefficients alpha** and **beta** (no inversions).
- Hourly mean AOD archived at the World Data Center for Aerosols (WDCA). Data with a 1-minute resolution are available from WORCC upon request.

GAW-PFR provides:

- long-term high-accuracy AOD and Angström Coefficients
- GAW-PFR provides AOD Dust in near real-time

Aerosol characterization in Northern Africa, Northeastern Atlantic, Mediterranean Basin and Middle East from direct-sun AERONET observations

S.Basart, C. Pérez, E. Cuevas, J.M. Baldasano, and G.P. Gobbi
(Atmos. Chem. Phys. October-2009)



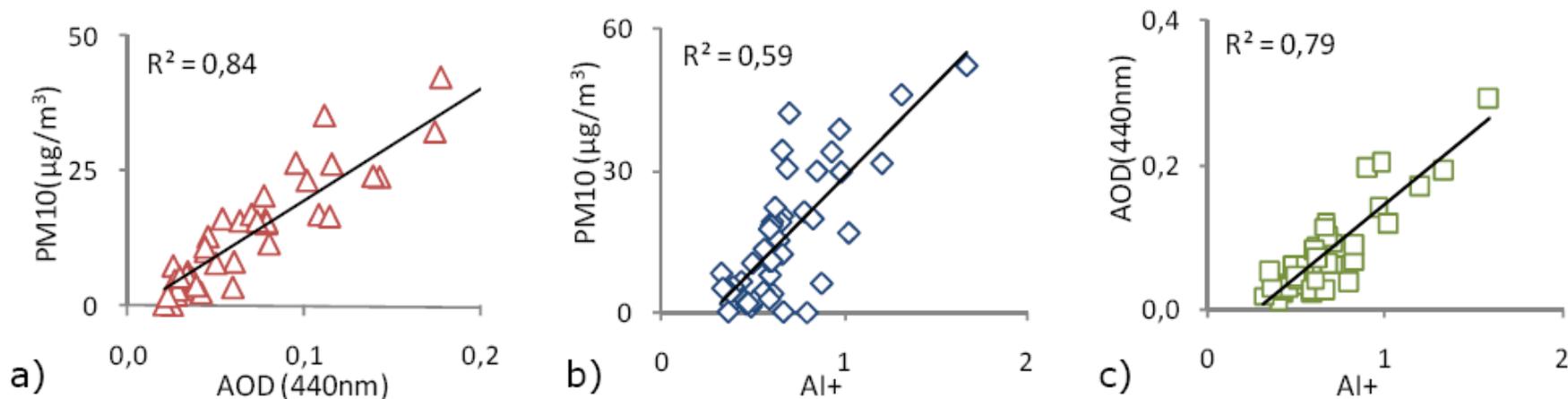


Figure 1. Scatterplot of monthly means of a) AOD vs PM₁₀; b) AI positive values vs PM₁₀; c) AI positive values vs AOD.

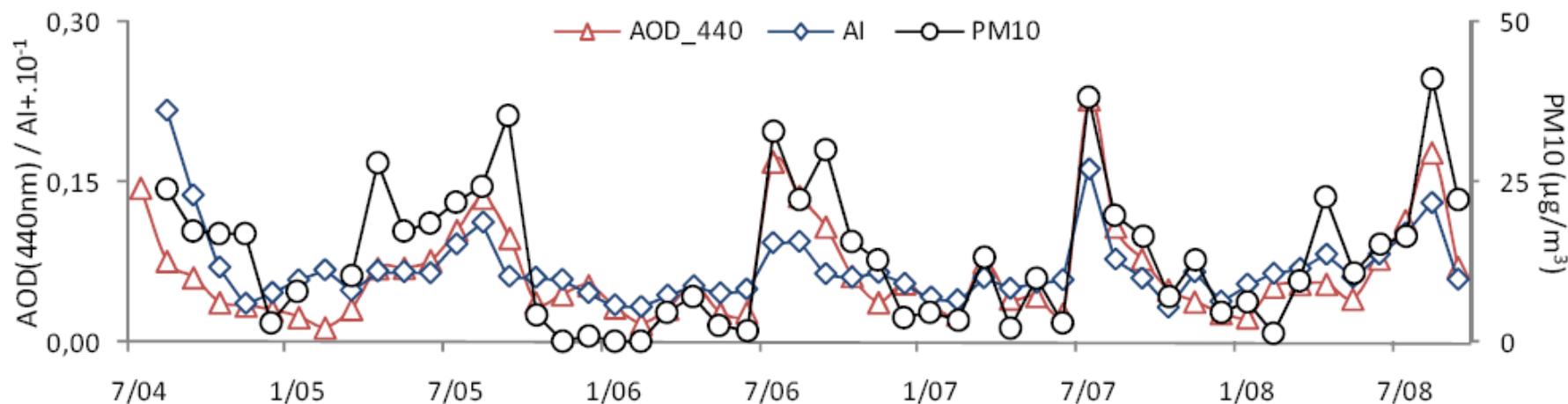


Figure 2. Monthly means of PM₁₀ (µg/m³), AOD and AI positive values.

Adam et al., 2010 (ACP-Interlaken): Detection of the Saharan dust air layer in the North Atlantic free troposphere with AERONET, OMI and in-situ data at Izaña Atmospheric Observatory

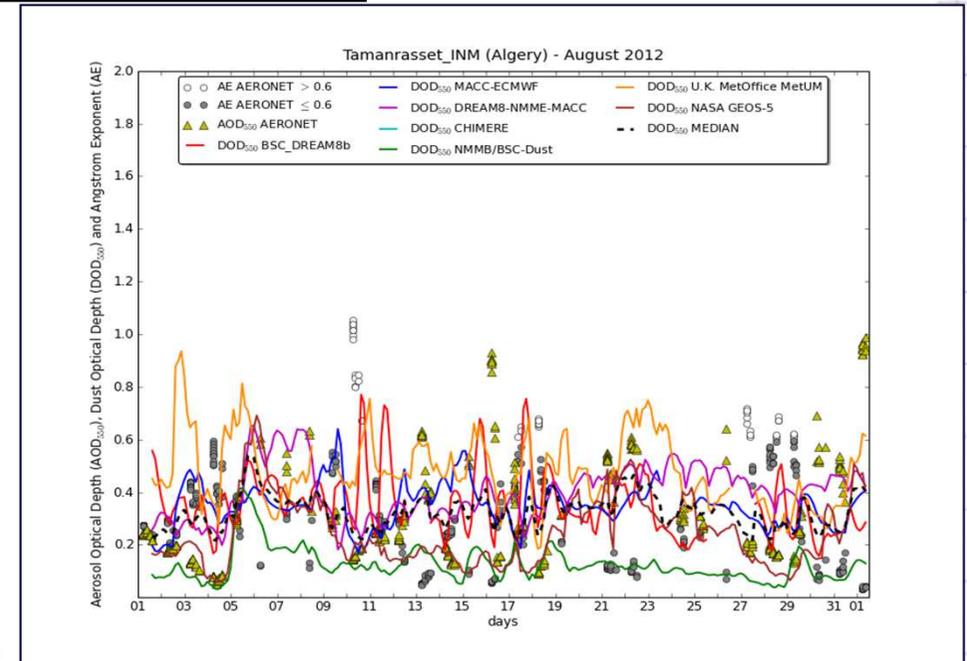
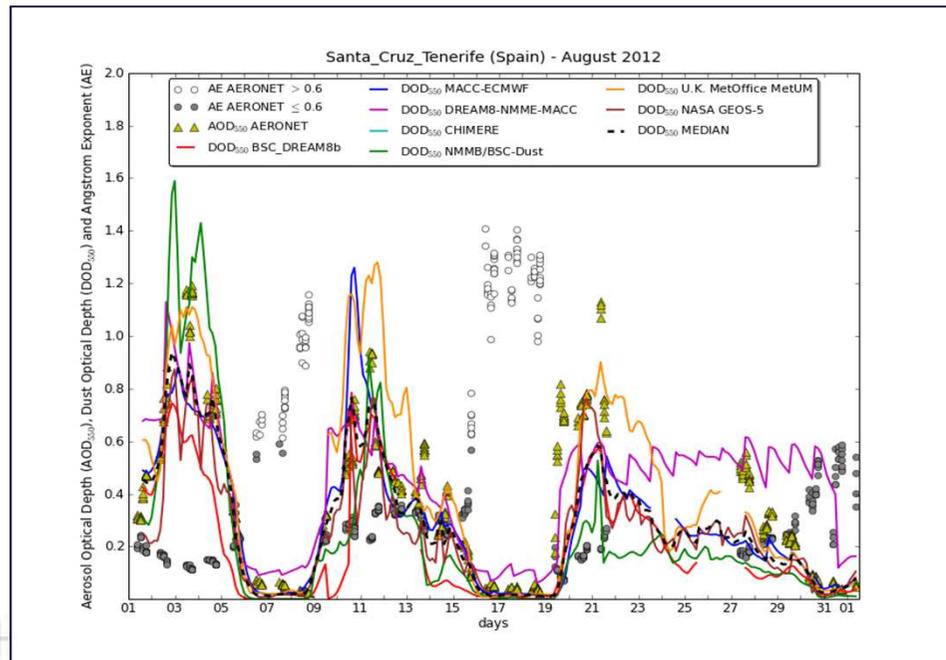
NRT evaluation using AERONET data



SANTA_CRUZ_TENERIFE



TAMANRASSET_INM



From total column observations...
to vertical resolved observations

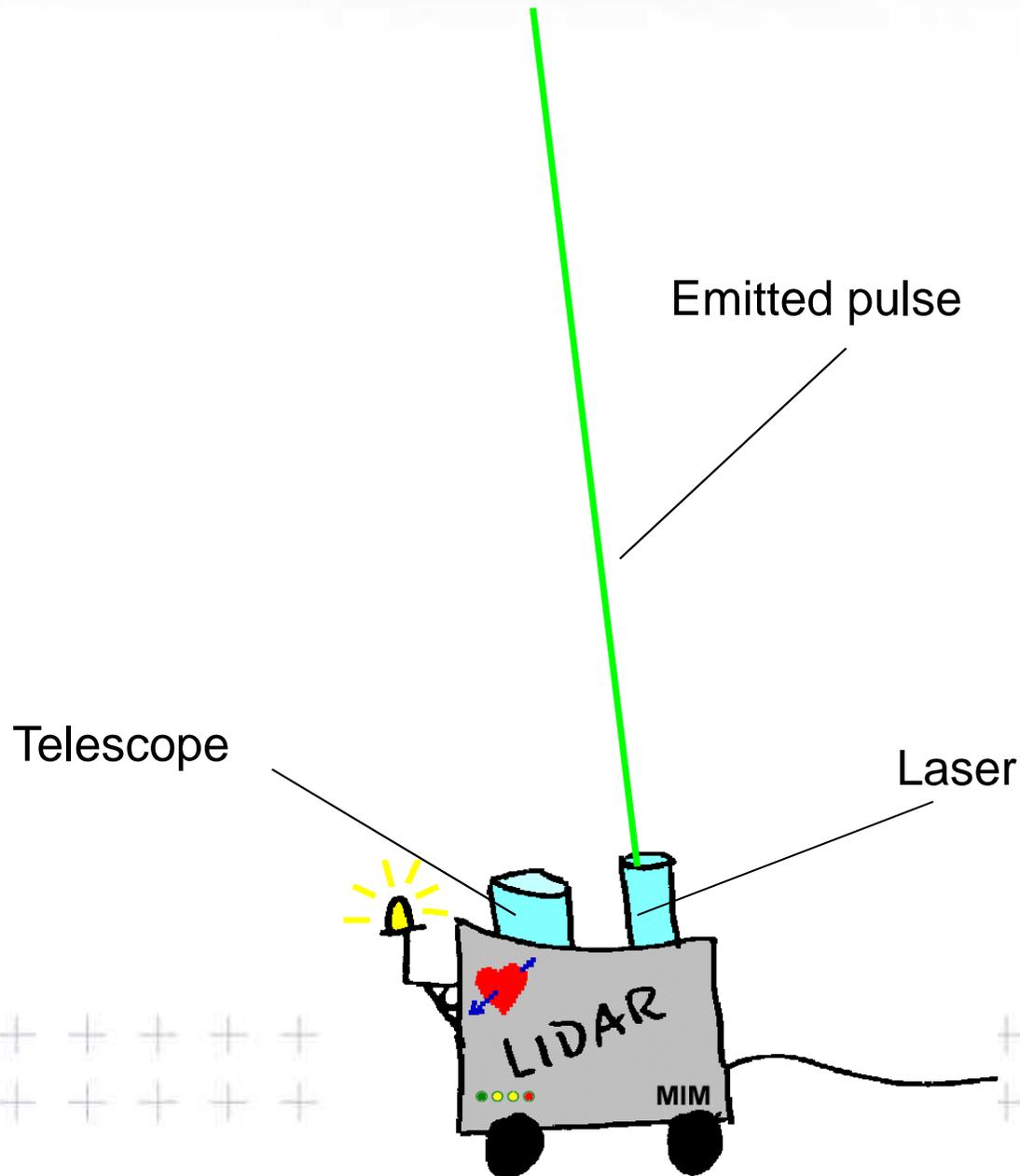
Lidars

Telescope

Laser

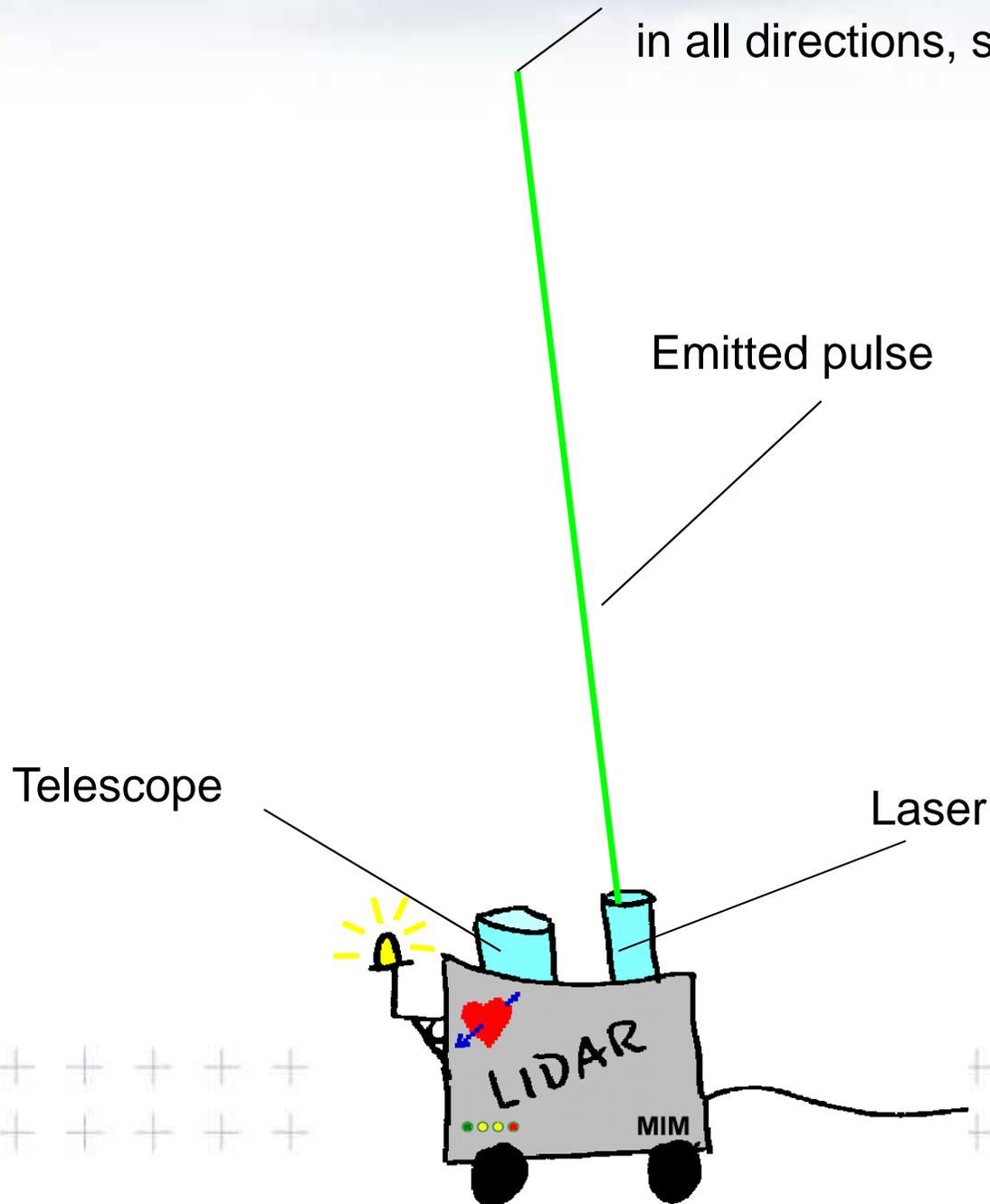


Taken from Matthias Wiegner's presentation (University of Munich, Meteorological Institute) to the SPIE 10 (International Symposium Remote Sensing)

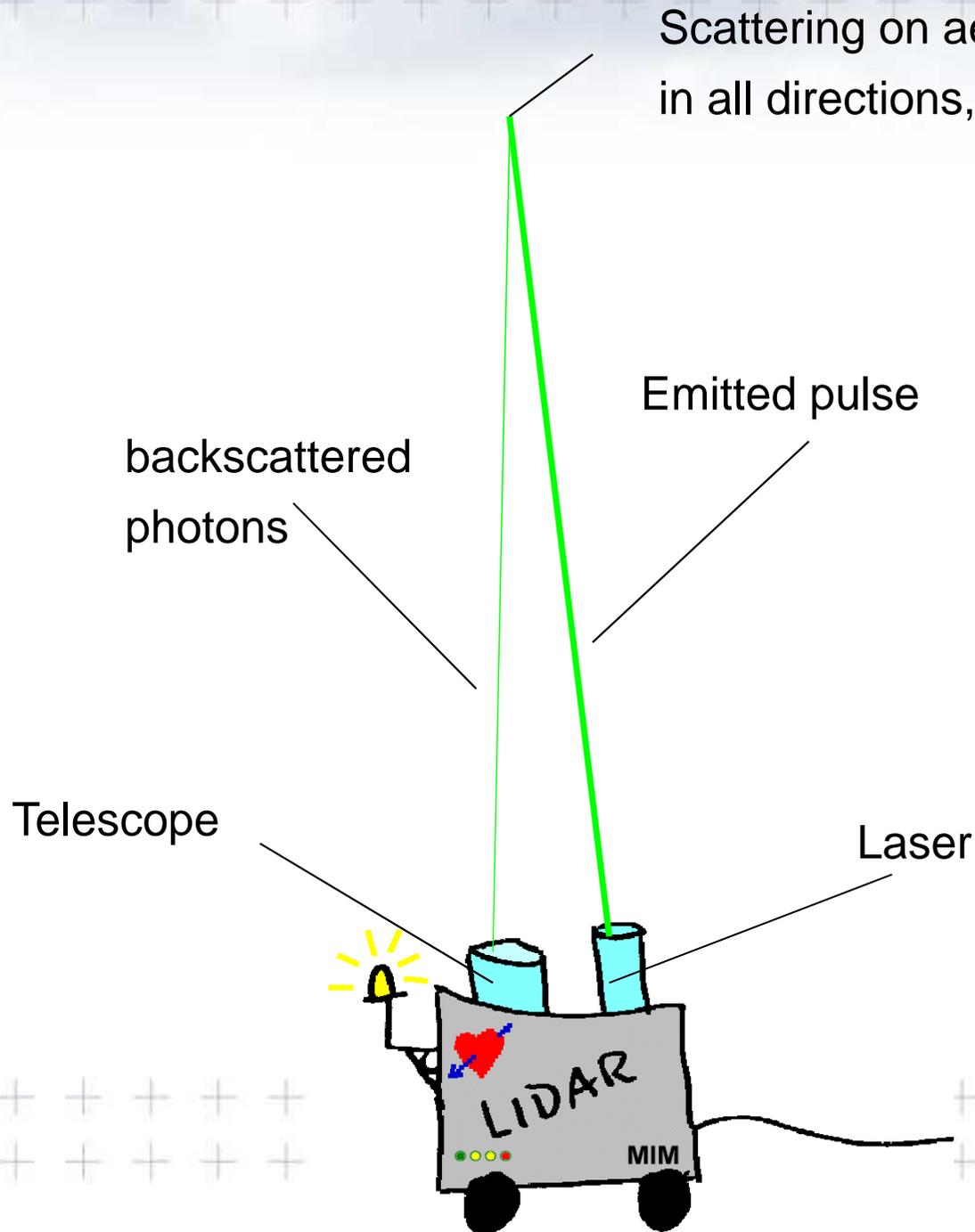


Taken from Matthias Wiegner's presentation (University of Munich, Meteorological Institute) to the SPIE 10 (International Symposium Remote Sensing)

Scattering on aerosols (and molecules) in all directions, some photons back



Taken from Matthias Wiegner's presentation (University of Munich, Meteorological Institute) to the SPIE 10 (International Symposium Remote Sensing)

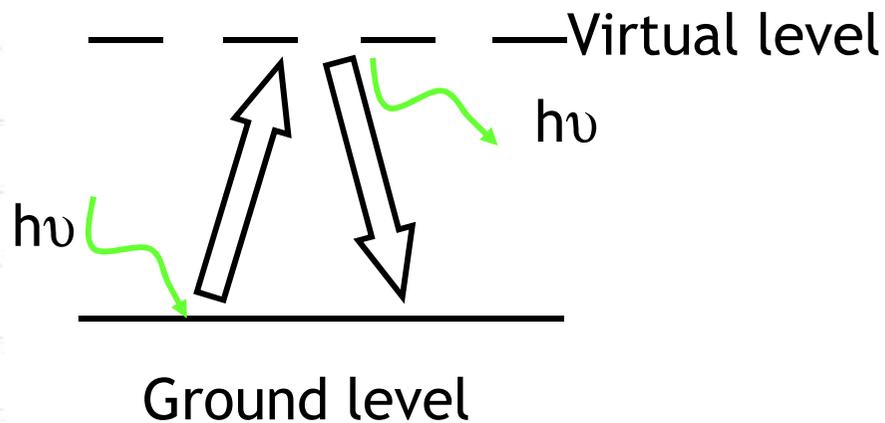


Scattering on aerosols (and molecules) in all directions, some photons back

Taken from Matthias Wiegner's presentation (University of Munich, Meteorological Institute) to the SPIE 10 (International Symposium Remote Sensing)

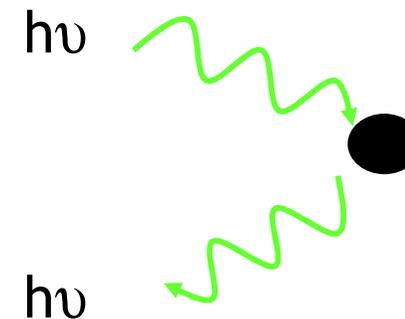
- Rayleigh Scattering

“Laser radiation elastically scattered from atoms or molecules is observed with no change of frequency”



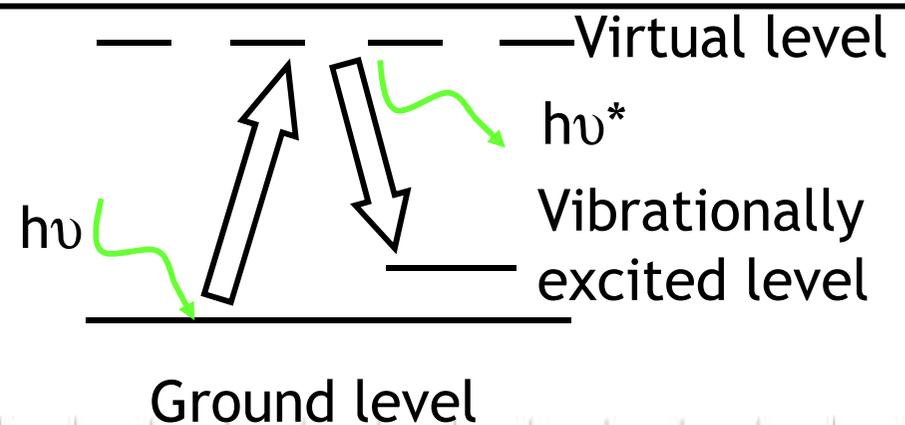
- Mie Scattering

“Laser radiation elastically scattered from small particulates or aerosols (of size comparable to wavelength of radiation) is observed with no change in frequency”



- Raman Scattering

“Laser radiation inelastically scattered from molecules is observed with a frequency shift characteristic of the molecule ($h\nu - h\nu^* = E$)”



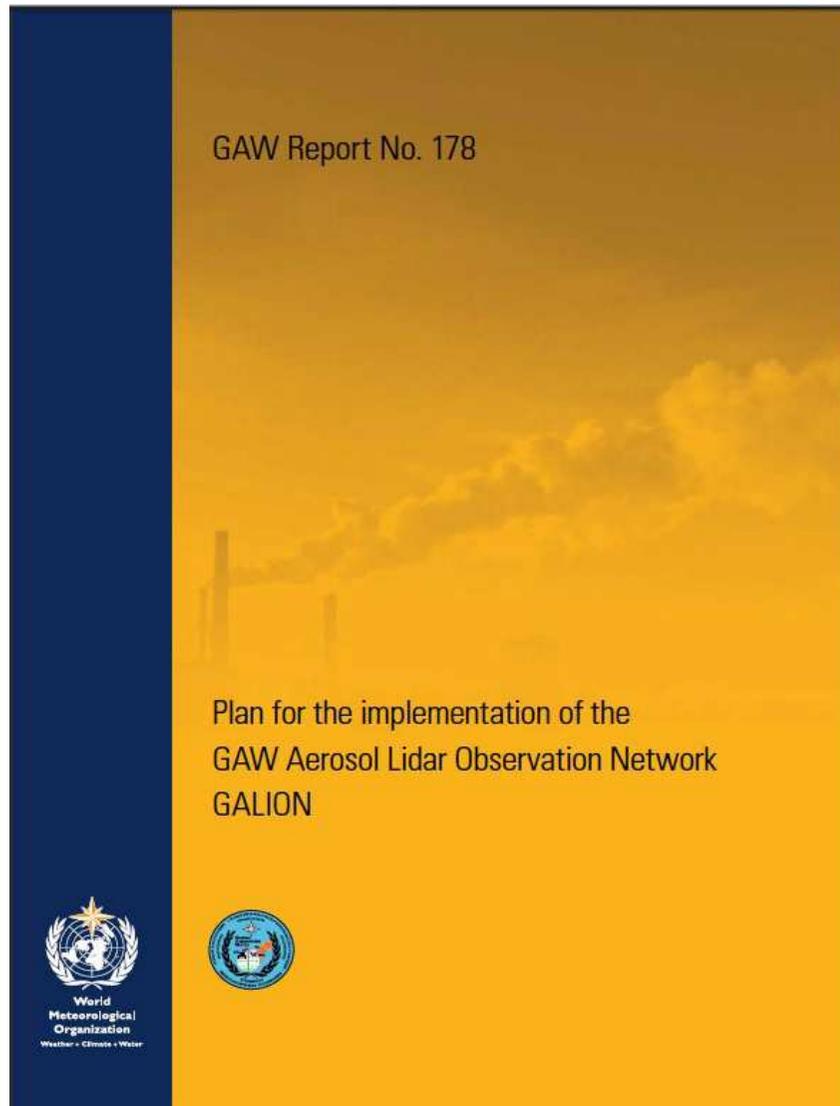


Lidar-Barcelona (UPC)
Raman Lidar
EARLINET-SPALINET

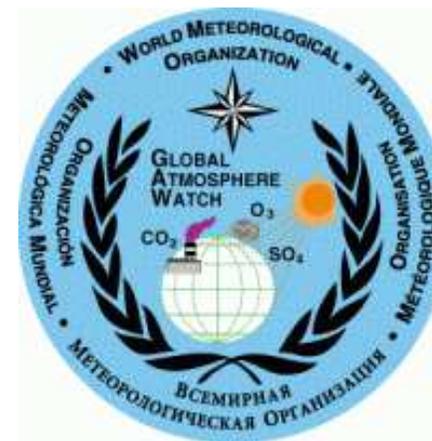


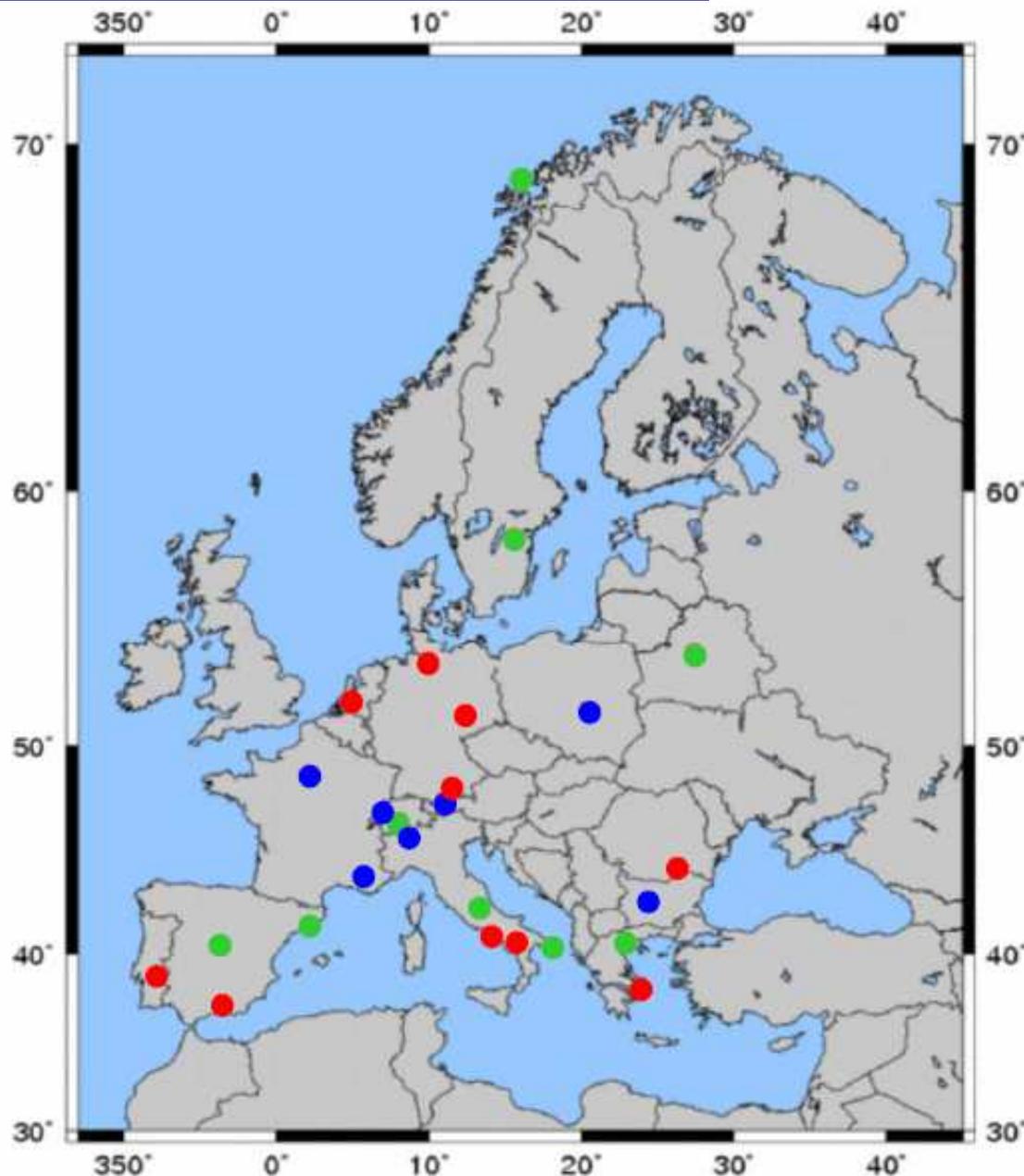
Lidar-Tenerife (INTA-AEMET); Elastic lidar
MPLNET

GAW Atmospheric Lidar Network (GALION)



<ftp://ftp.wmo.int/Documents/PublicWeb/arep/gaw/gaw178-galion-27-Oct.pdf>





EARLINET

EARLINET (European Aerosol Research Lidar NETwork) is a network of advanced lidar stations distributed over Europe with the main goal to provide a comprehensive, quantitative, and statistically significant data base for the aerosol distribution on a continental scale. EARLINET provides independent measurements of aerosol extinction and backscatter, and retrieval of aerosol microphysical properties.

10 EARLINET stations are equipped also with sunphotometers (they are part of AERONET).

26 lidar stations

- 10 multiwavelength Raman lidar stations

backscatter (355, 532 and 1064 nm)
+ extinction (355 and 532 nm) +
depol ratio (532 nm)

- 9 Raman lidar stations

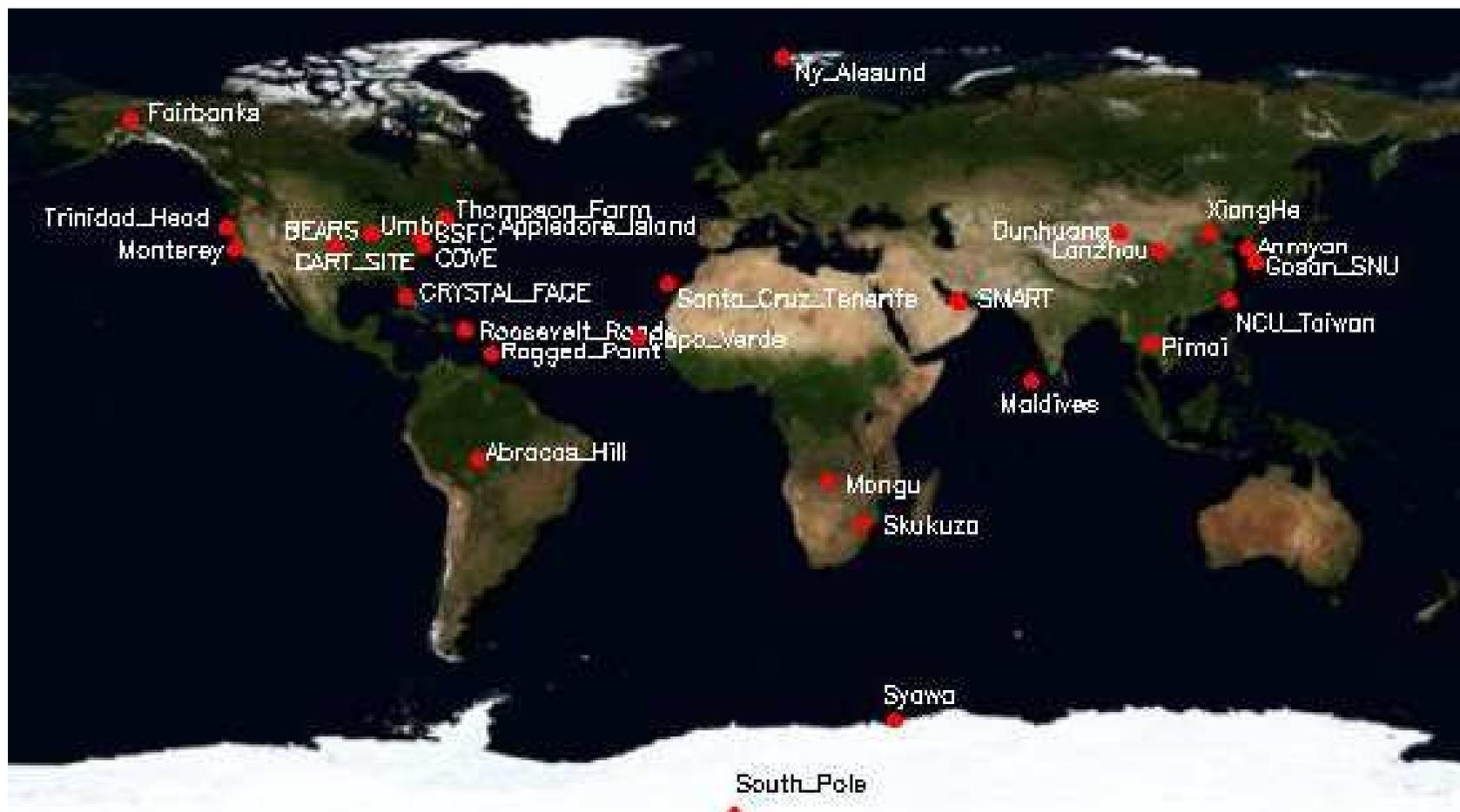
- 7 single backscatter lidar stations

Aerosol lidar (MPLNet)

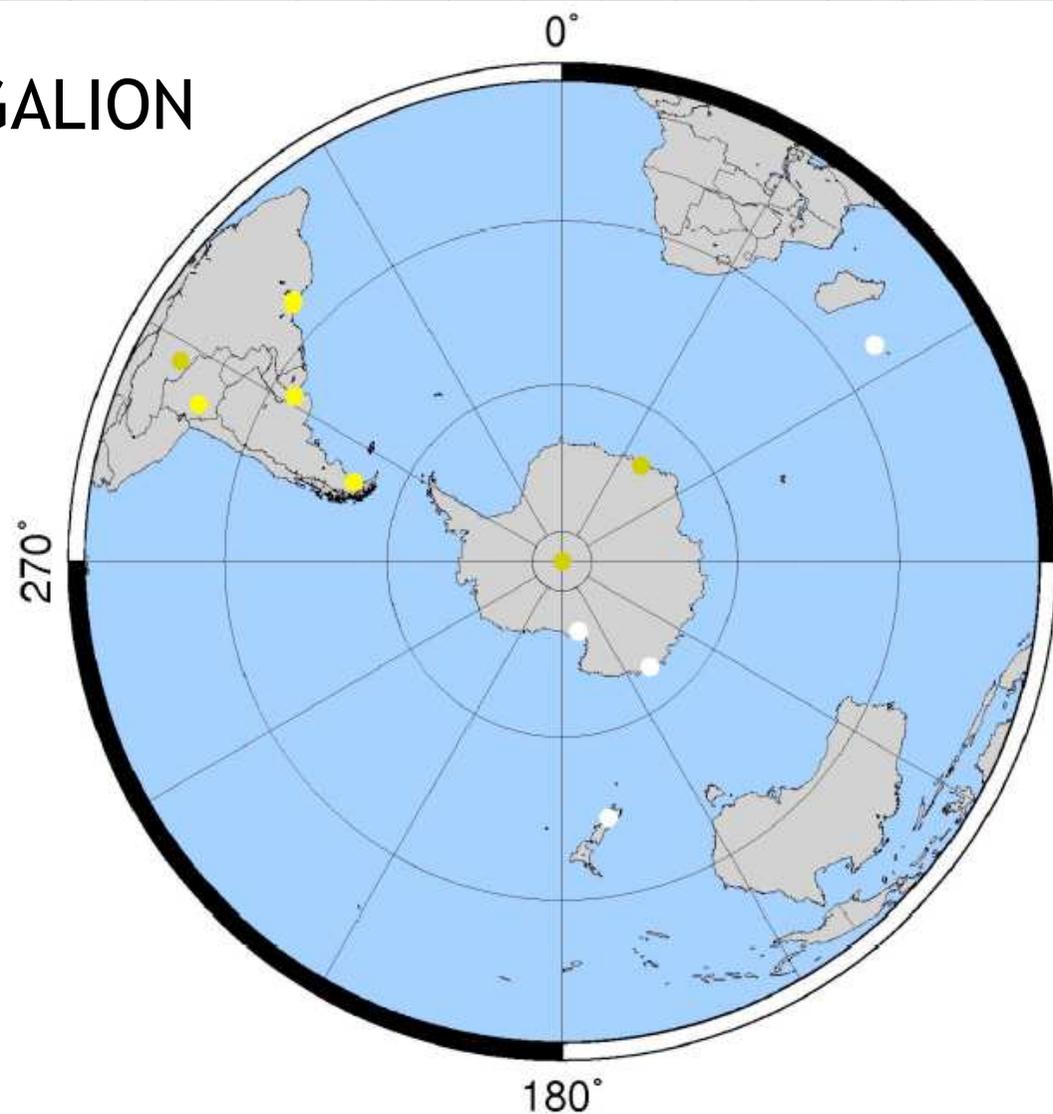
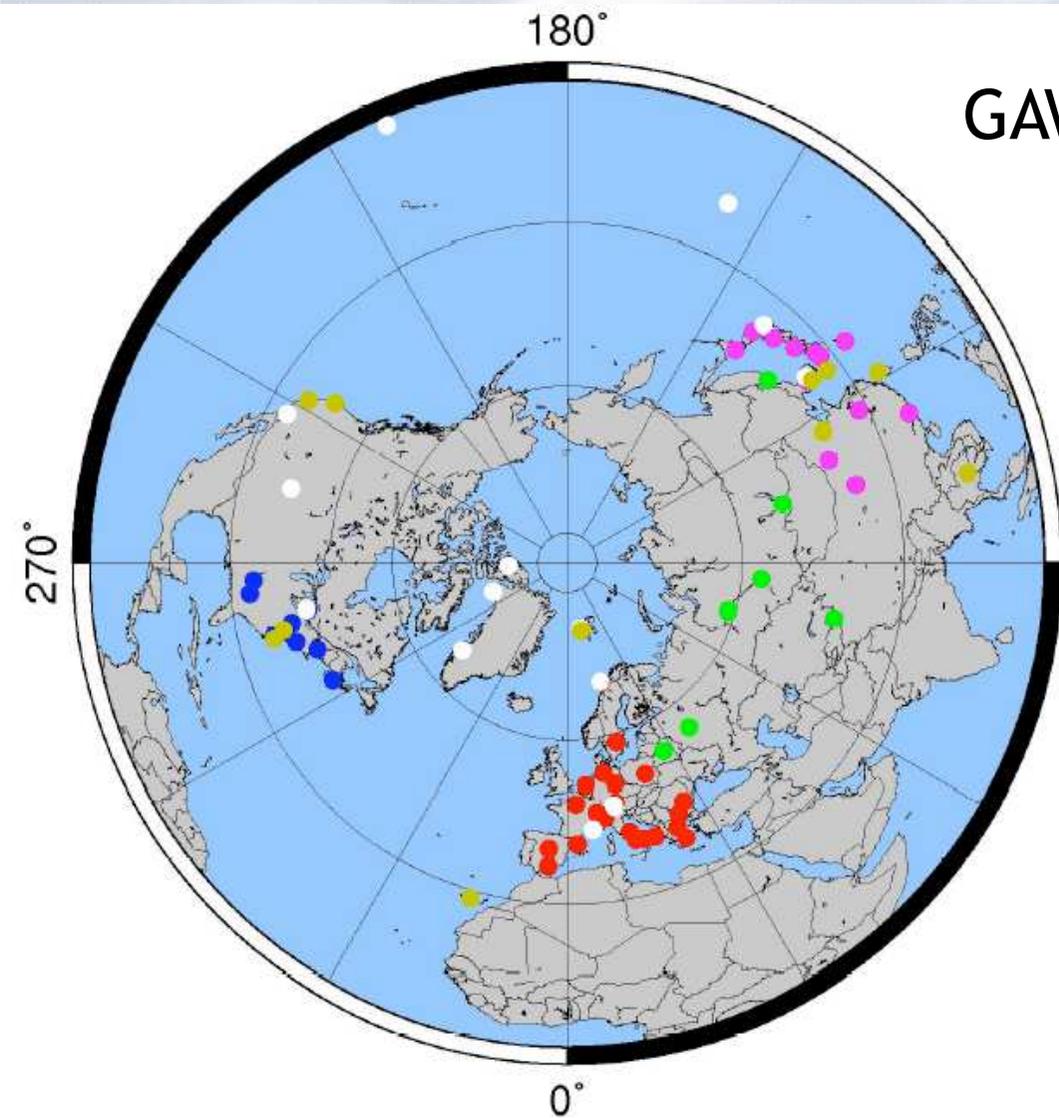
<http://mplnet.gsfc.nasa.gov/>



523 nm MPLNET
Automatized since July 2005

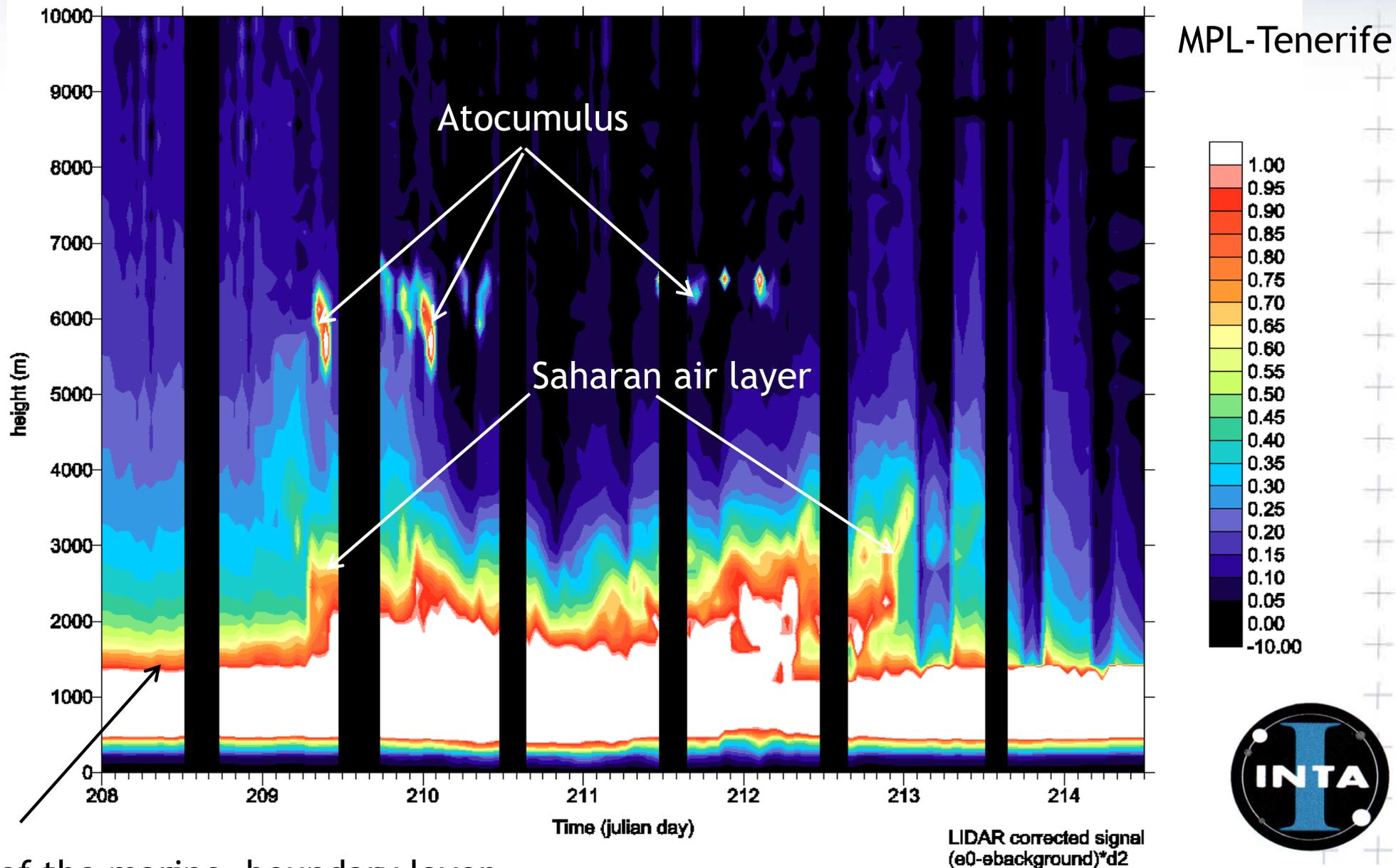


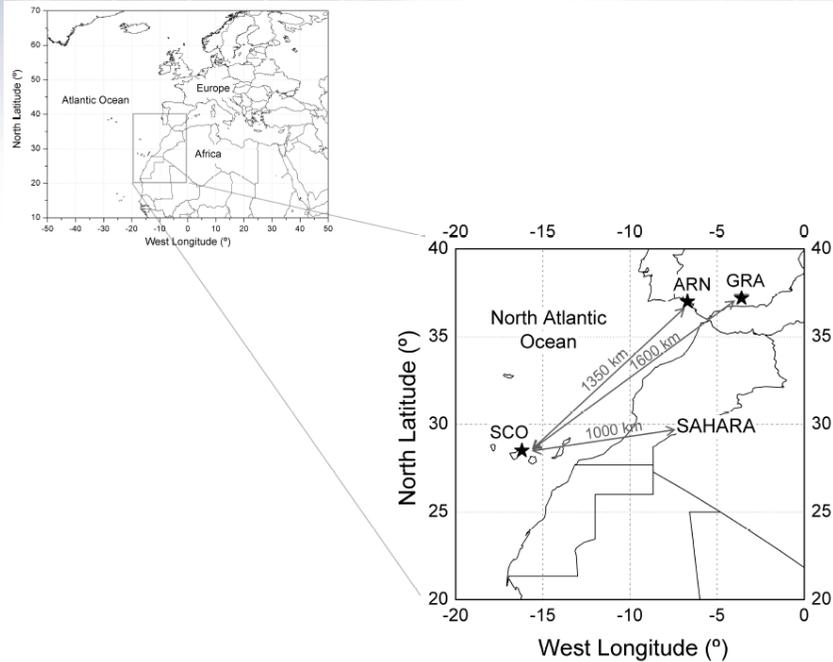
GAW-GALION



Distribution of stations as available through the cooperation between existing networks: **AD-NET** , **ALINE** , **CISLiNet** , **EARLINET** , **MPLNET** , **NDACC** , **REALM** .

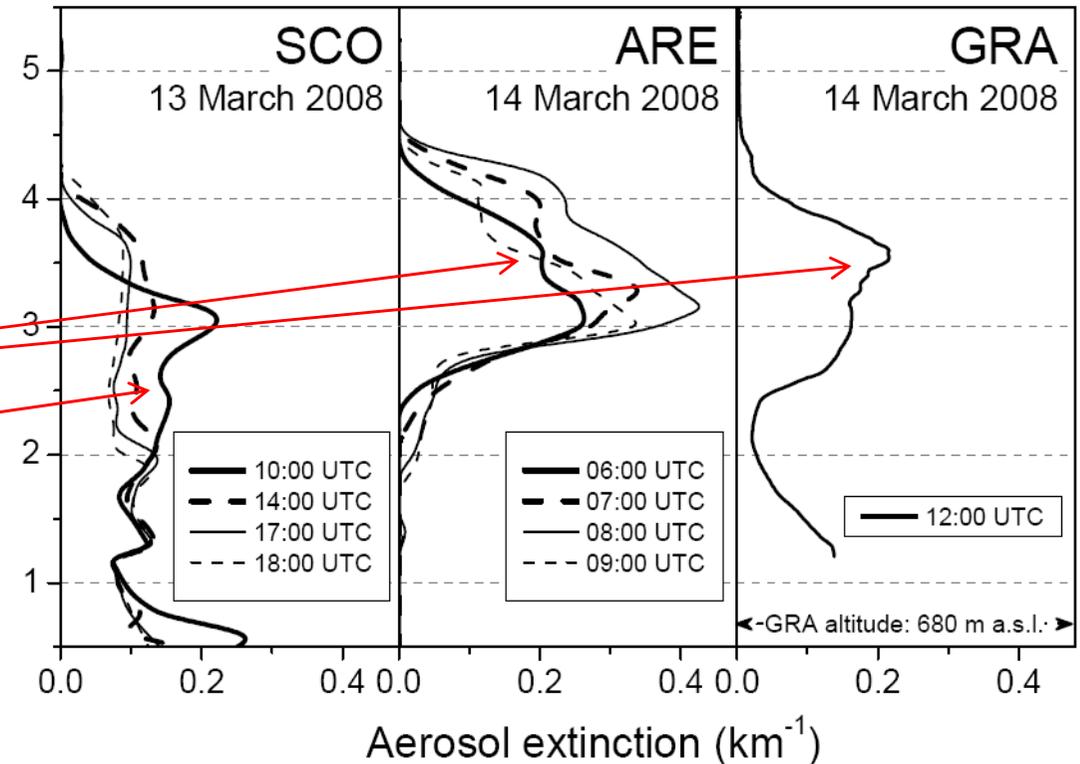
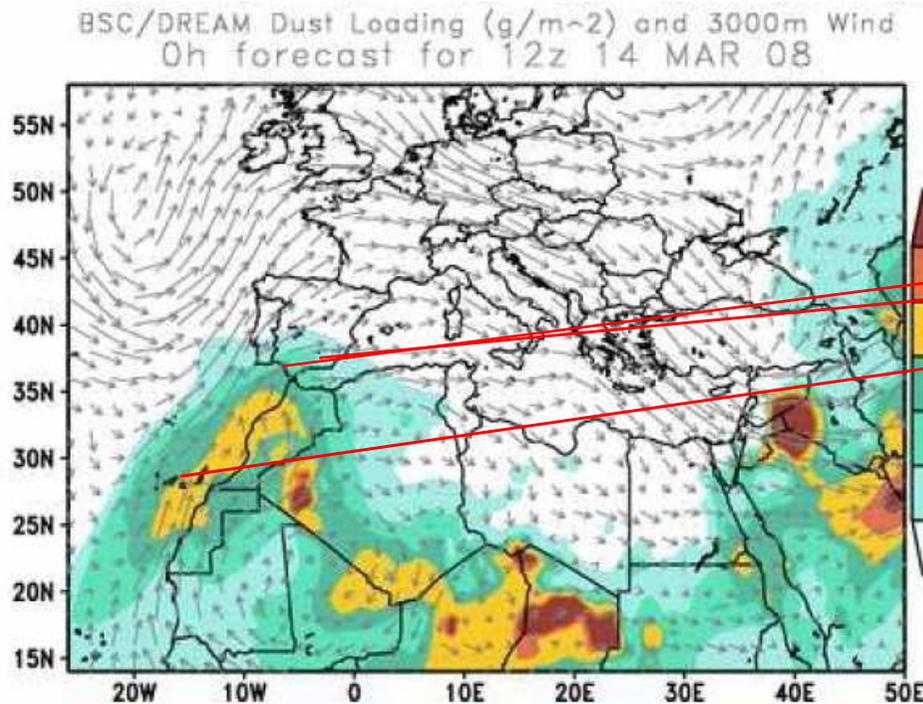
DUST EVENT 28 JULY - 2 AUGUST 2002



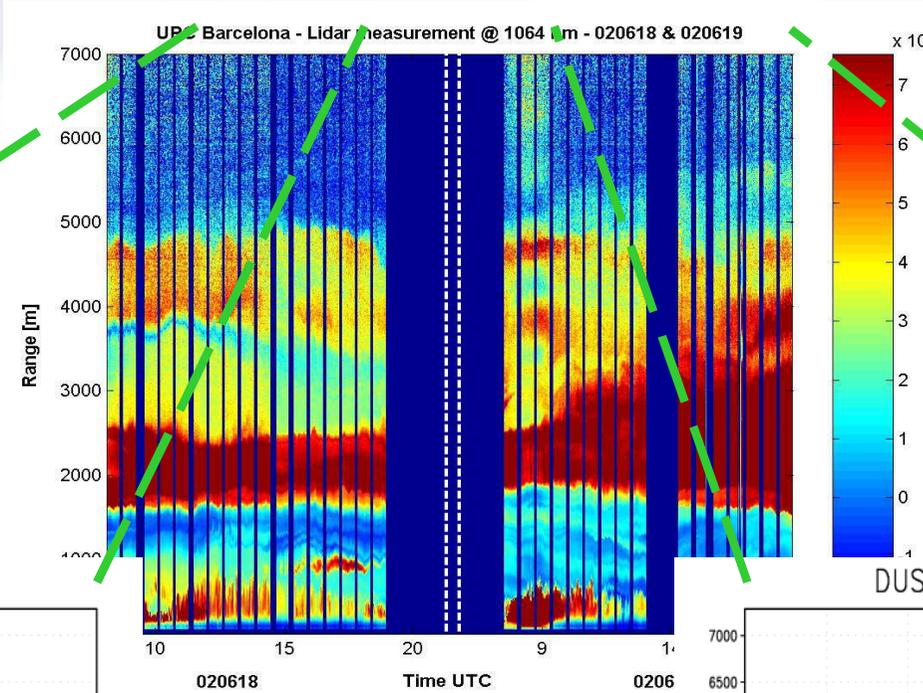


A case study of dust transport from Canary Islands to Iberian Peninsula

Córdoba-Jabonero et al., ACP Discuss., 2010



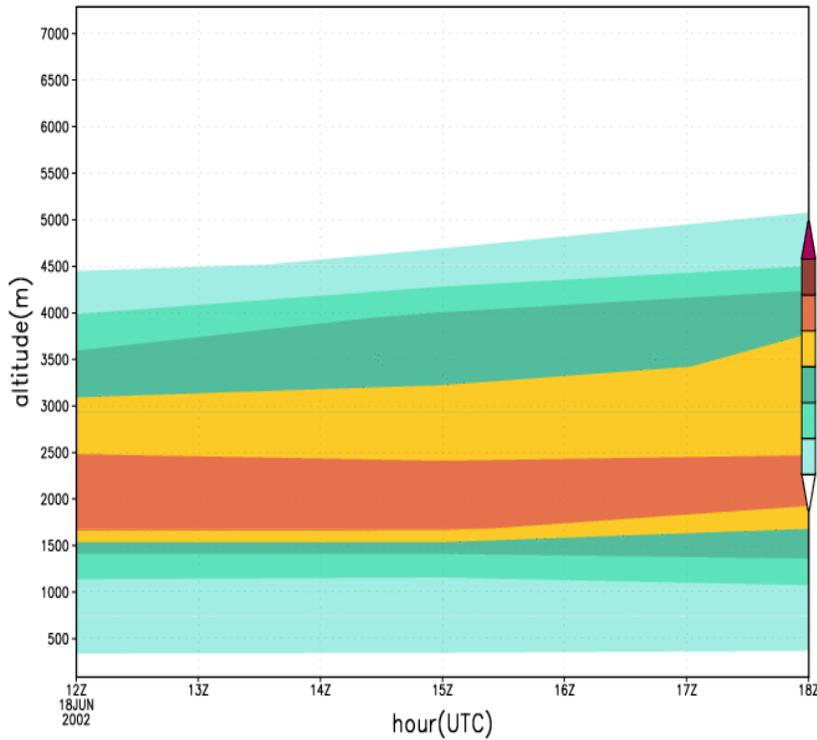
Barcelona lidar vs DREAM BSC



EARLINET: Lidar-UPC,
Barcelona

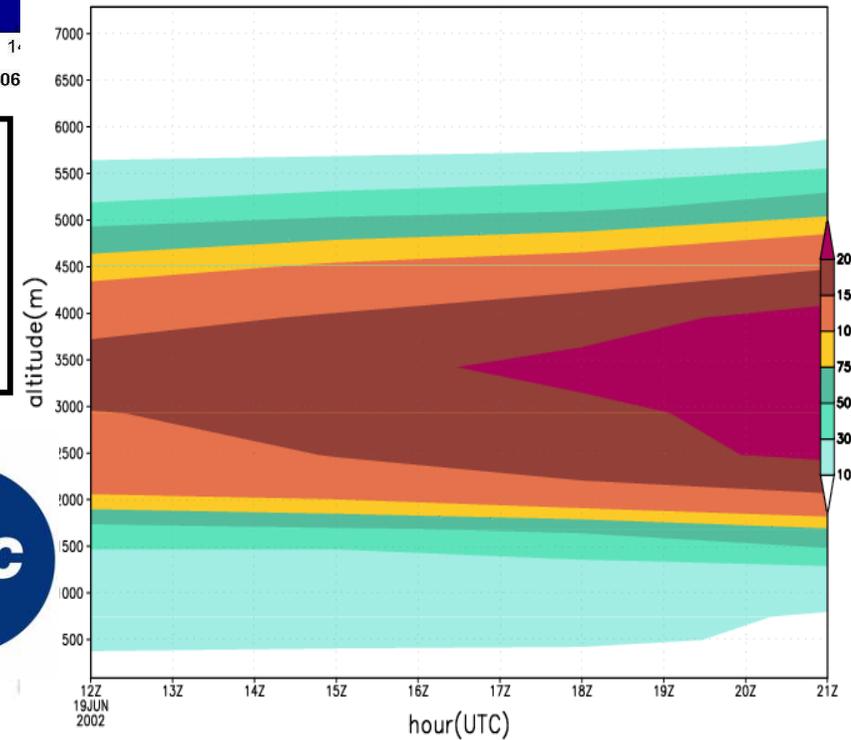
18-19 June 2002

DUST CONC. ($\mu\text{g}/\text{m}^3$) 18 JUN 2002

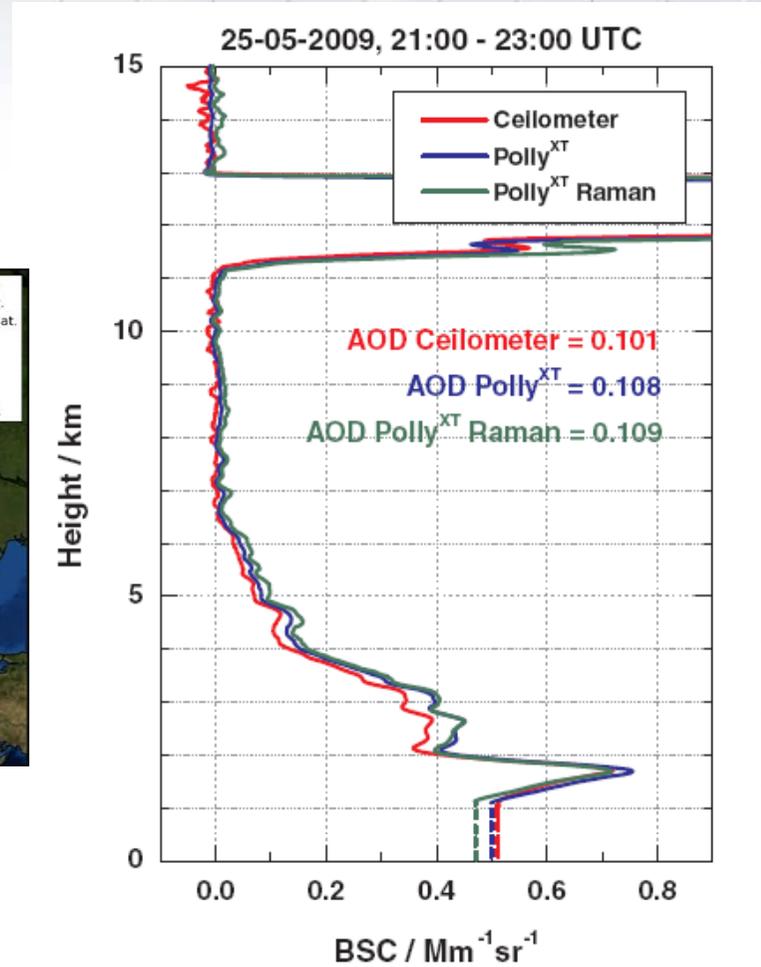
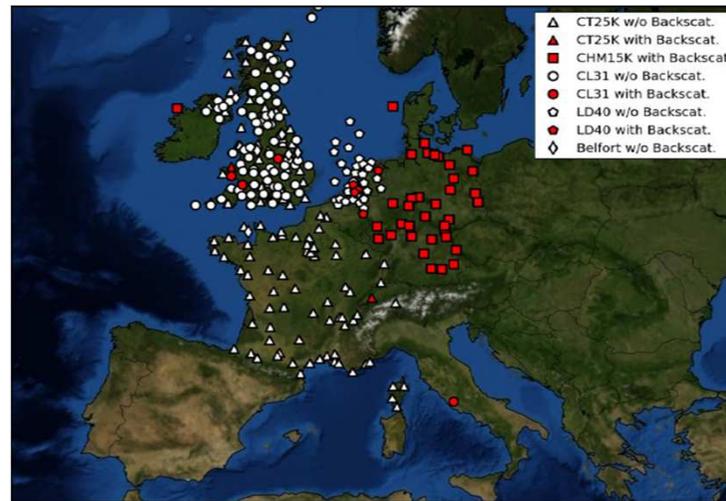


DUST CONC. ($\mu\text{g}/\text{m}^3$) 19 JUN 2002

Vertical dust
distribution
validation:
AIRLINET-DREAM

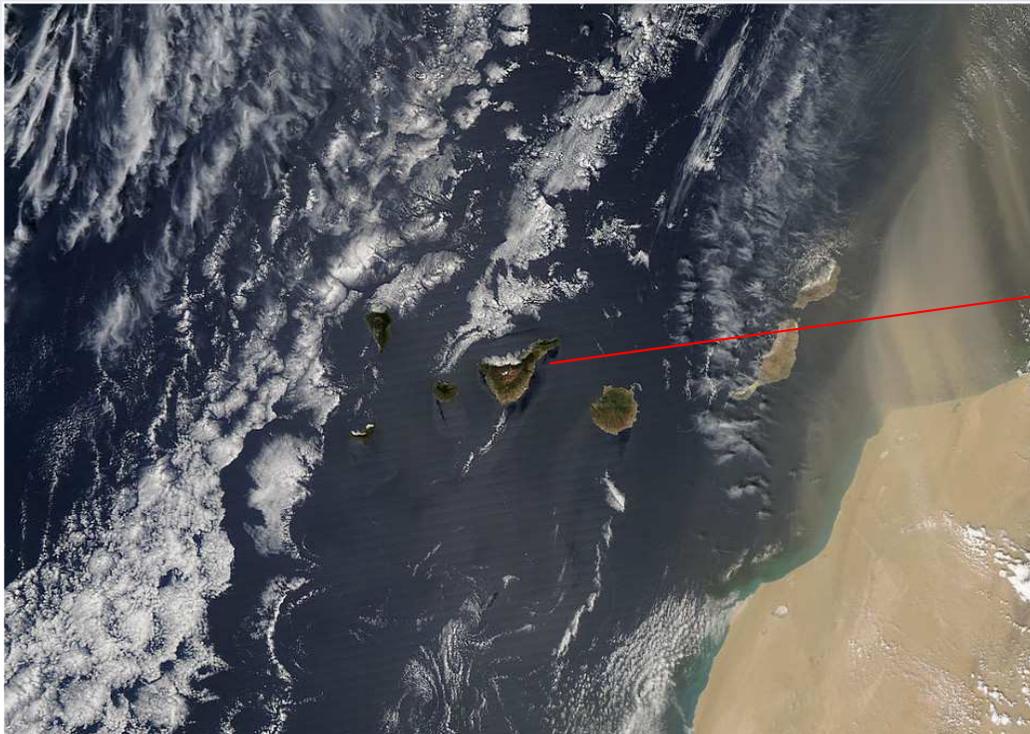


Met Services are replacing cloud-base ceilometer networks by aerosol backscatter profiling ceilometers (IR wavelength).
Objective: To monitor MLD (Mixing Layer Depth) based on several hundred profiling ceilometers (100km sampling)



Heese et al., Atmos. Mes. Tech. 2010, Ceilometer-lidar inter-comparison: backscatter coefficient retrieval and signal-to-noise ratio determination

Optimal for desertic areas !!

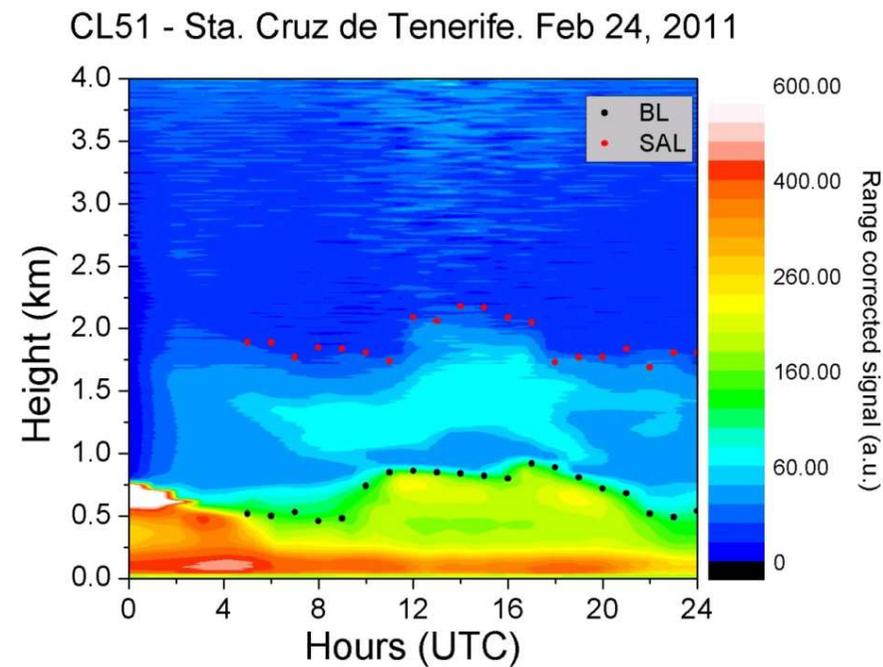
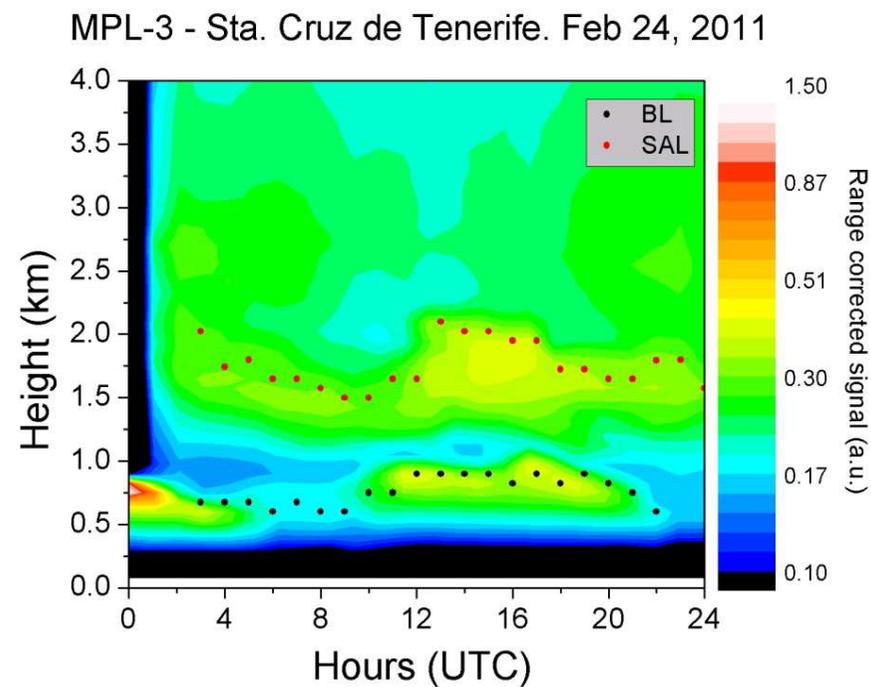
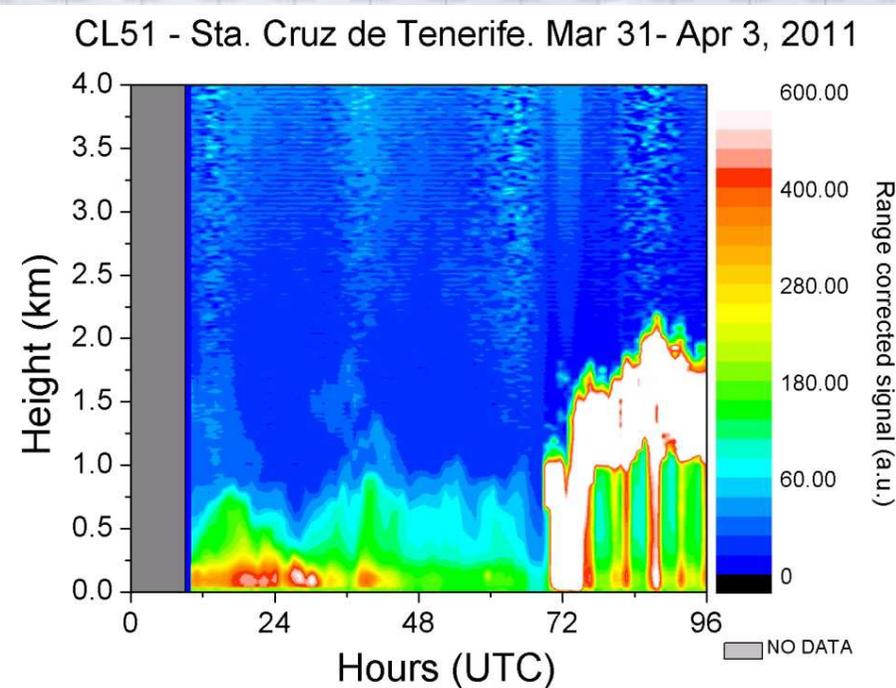
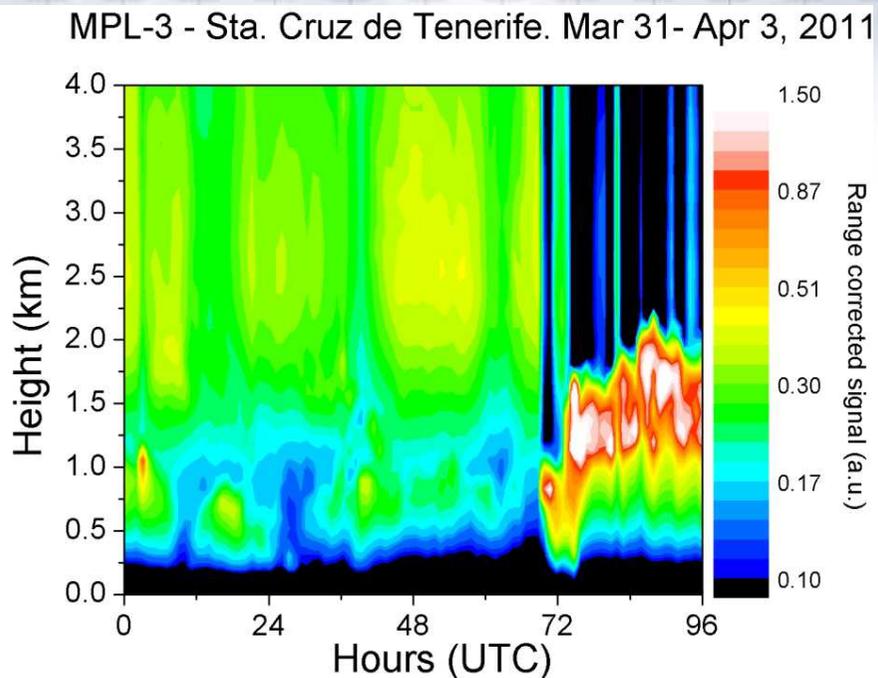


Viasala Ceilometer
CL-51

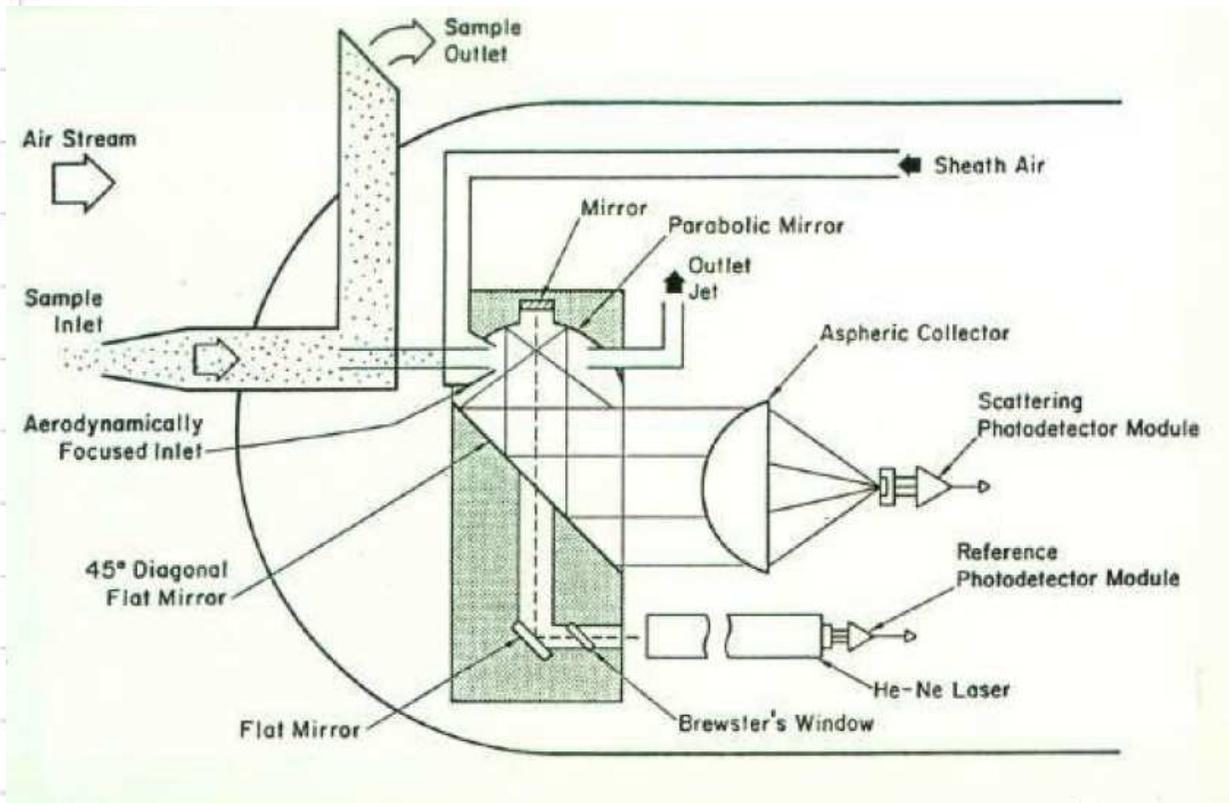
MicroPulse Lidar and Ceilometer inter-comparison during Saharan dust intrusions over the Canary Islands

Y. Hernández, S. Alonso-Pérez, E. Cuevas, C. Camino, R. Ramos, J. de Bustos, C. Marrero, C. Córdoba-Jabonero and M. Gil (2011)

Campaign performed from January to March 2011 in Tenerife island



INTA C212-200 N/S 301



PCASP-100X sonde

Optical counter

0.1 - 3.0 μm en 15 channels

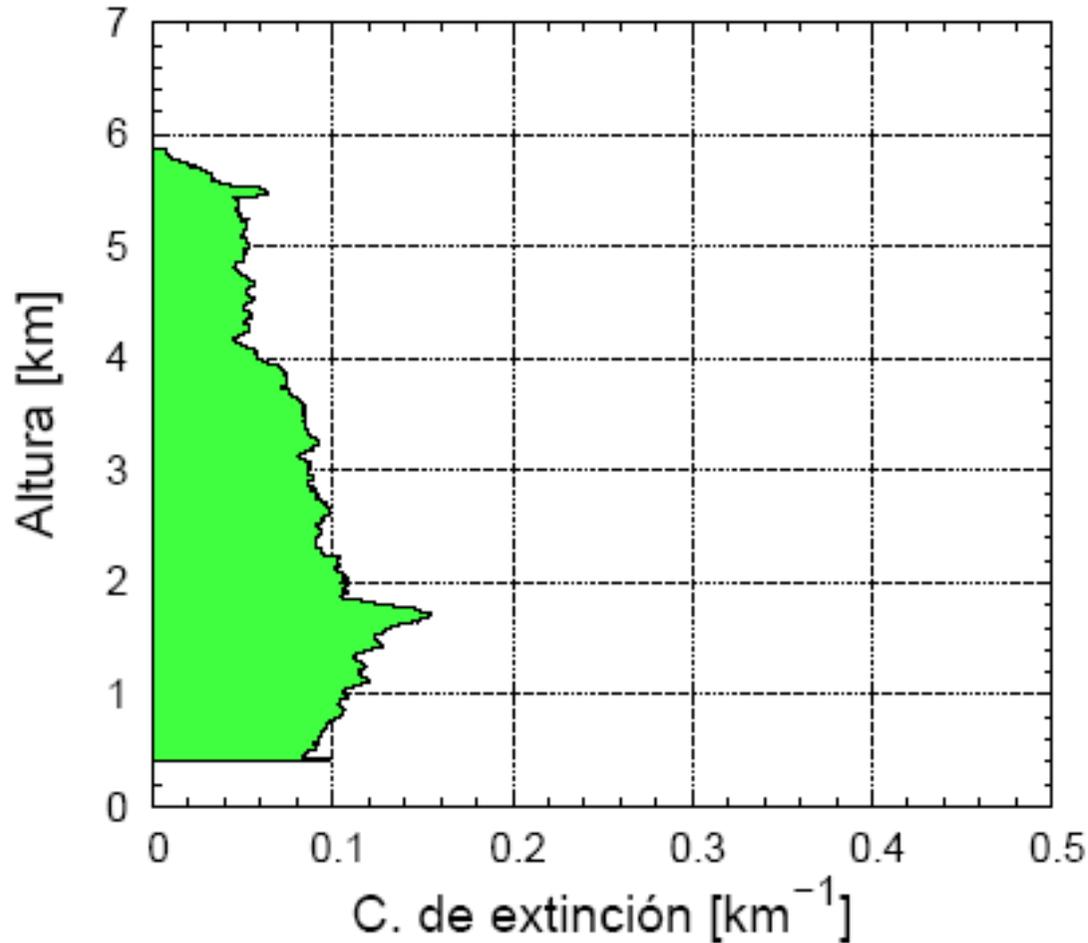
Up to 20,000 particles / s

A size distribution / s

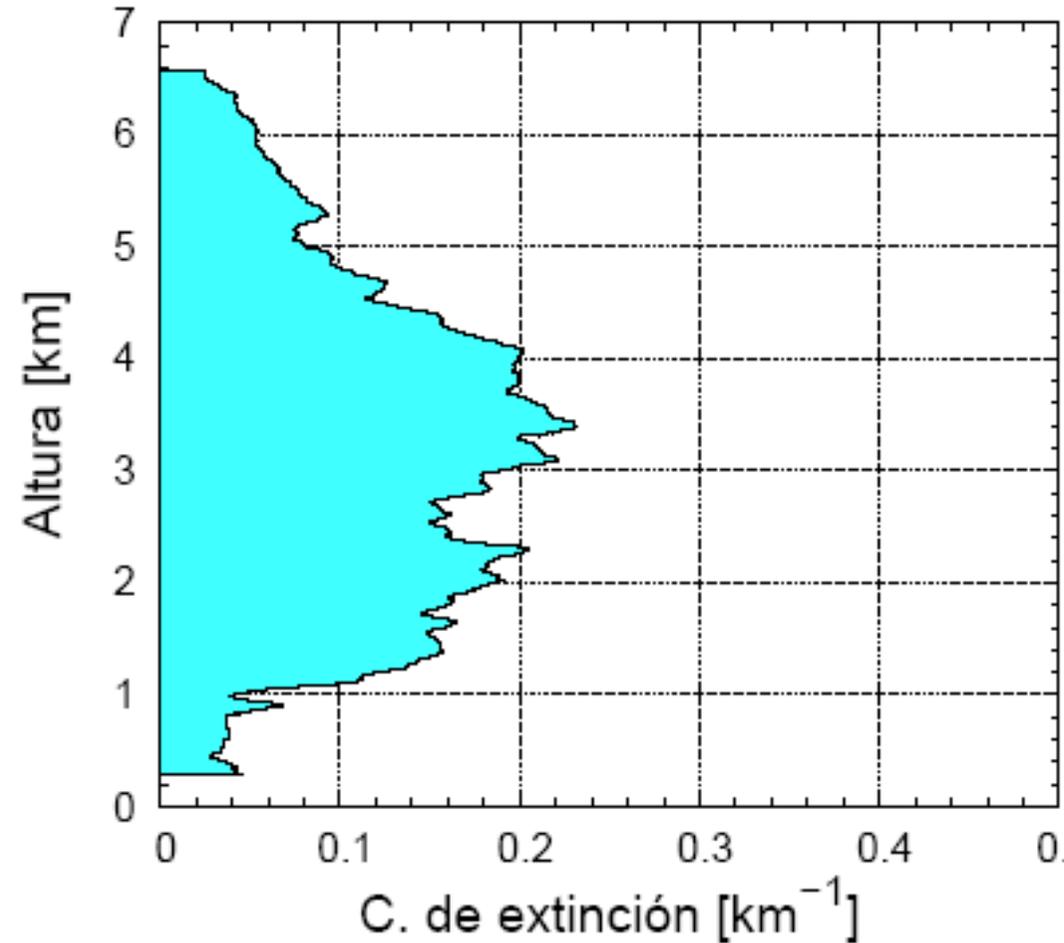




20 - julio - 05



22 - julio - 06

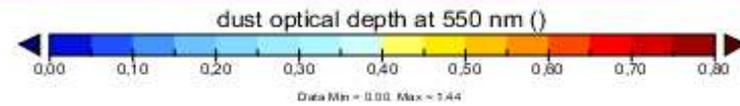
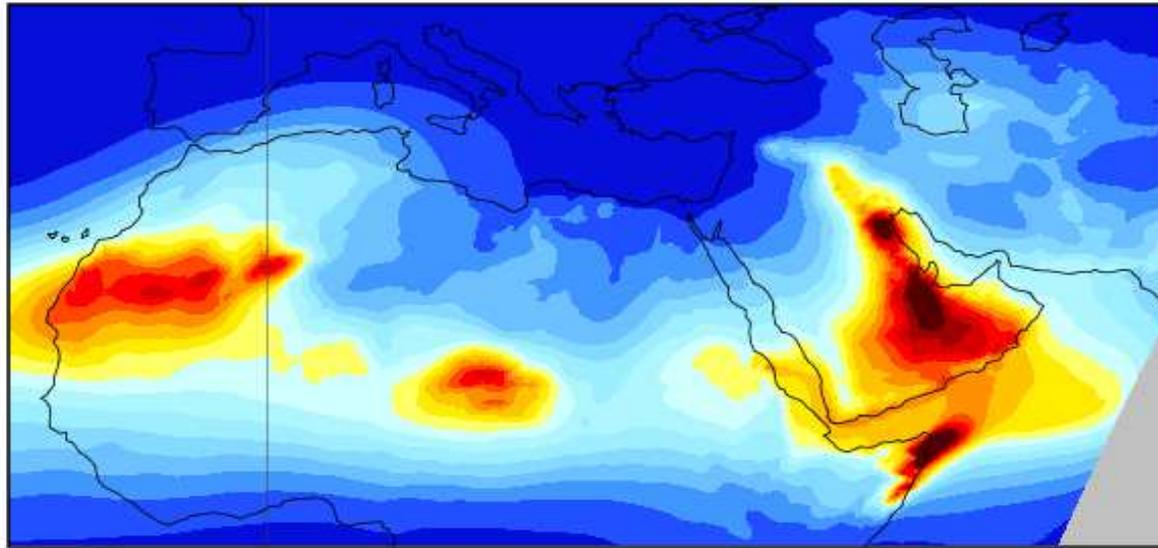


Andrey, 2011 (PhD)

Index

- Aerosols and dust background
- In-situ dust characterization
- In-situ dust estimations (Visibility)
- Ground based remote sensing
- Recommended ground-based observations in Middle East
- Summary

July 2003-2009 (DOD at 550nm: NMMB\BSC-Dust)



Impact of Dust on West Asia (Middle East)



Establishing a WMO Sand and Dust Storm Warning Advisory and Assessment System Regional Node for West Asia: Current Capabilities and Needs

Technical Report

Available at:

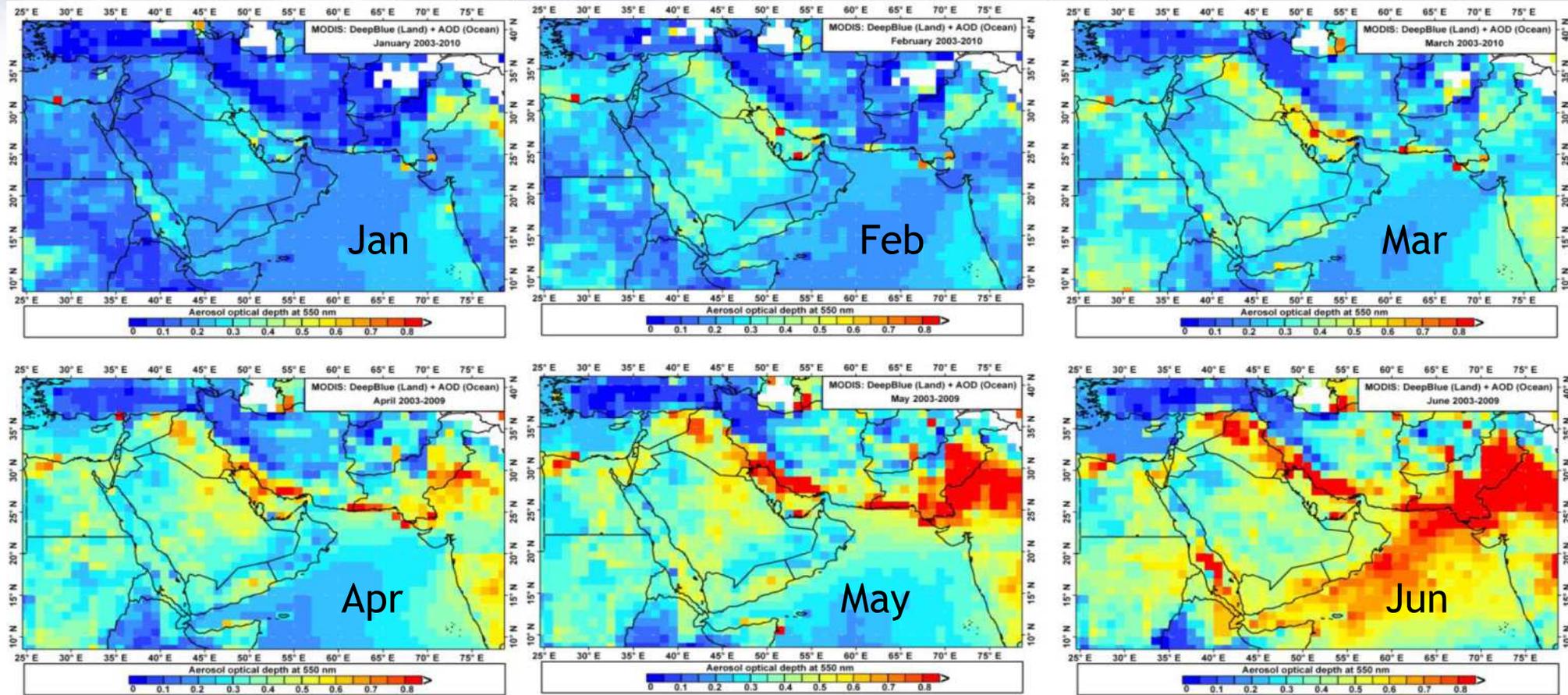
http://www.wmo.int/pages/prog/arep/wwrp/new/documents/1121_SDS_Technical_Report_en.pdf



WMO-No. 1121

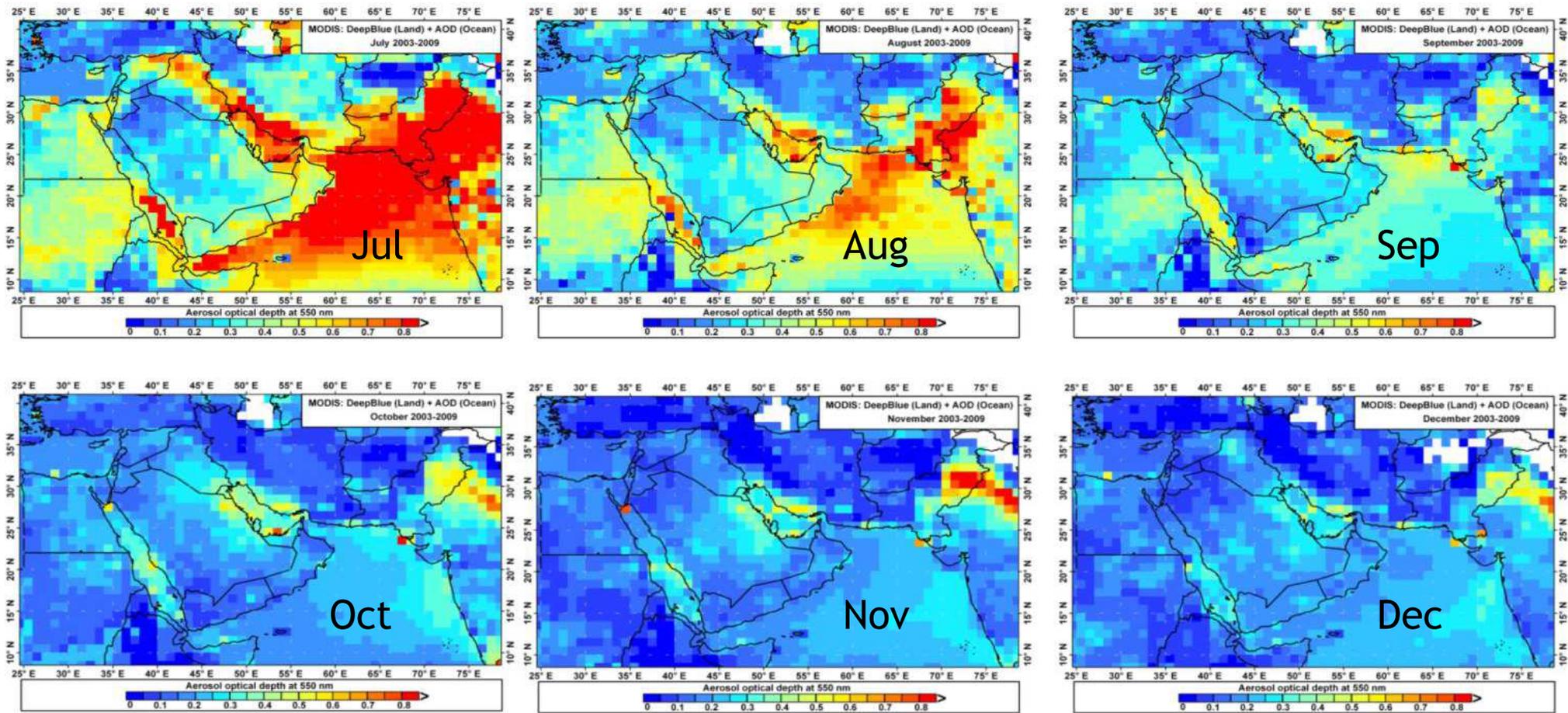


Recommended ground-based observations in Middle East



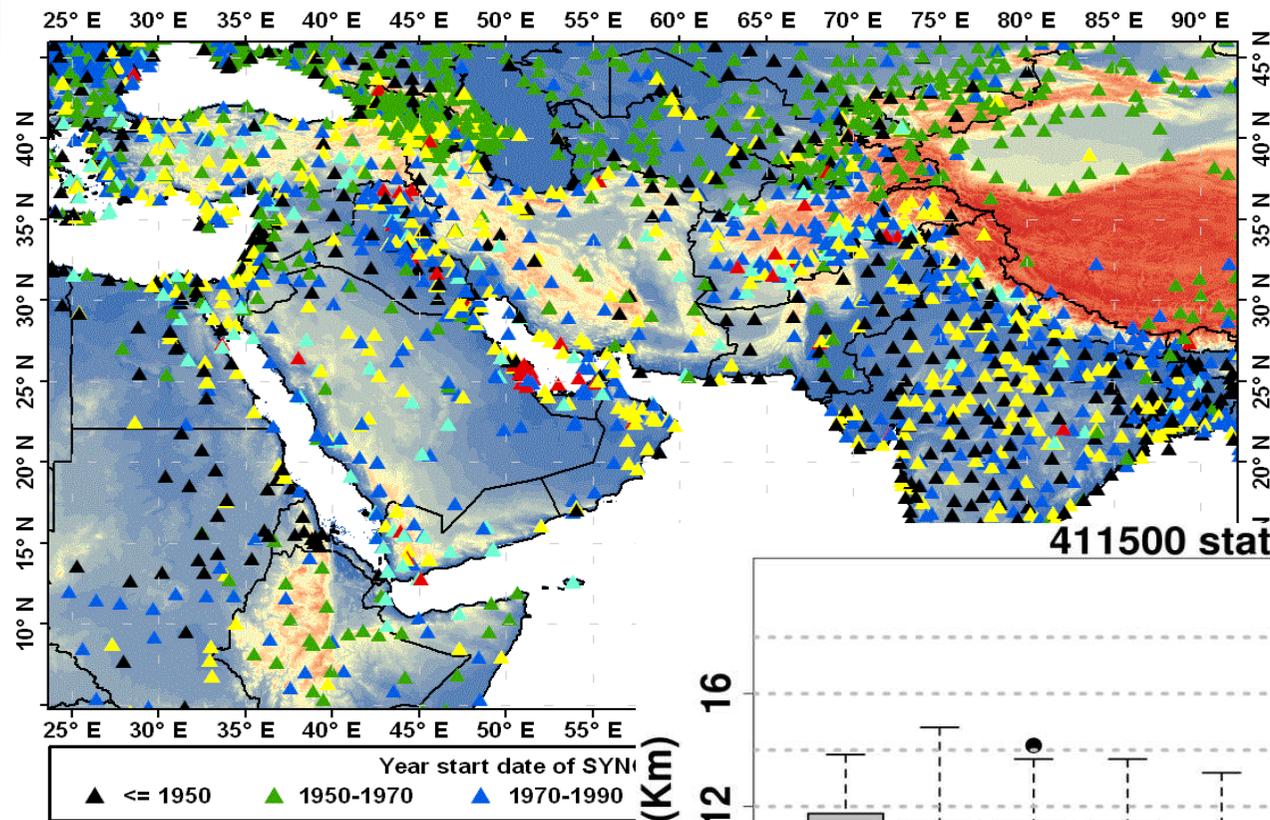
Monthly means AOD from MODIS DB for the period 2003-2010
January-June

Recommended ground-based observations in Middle East



Monthly means AOD from MODIS DB for the period 2003-2010
July-December

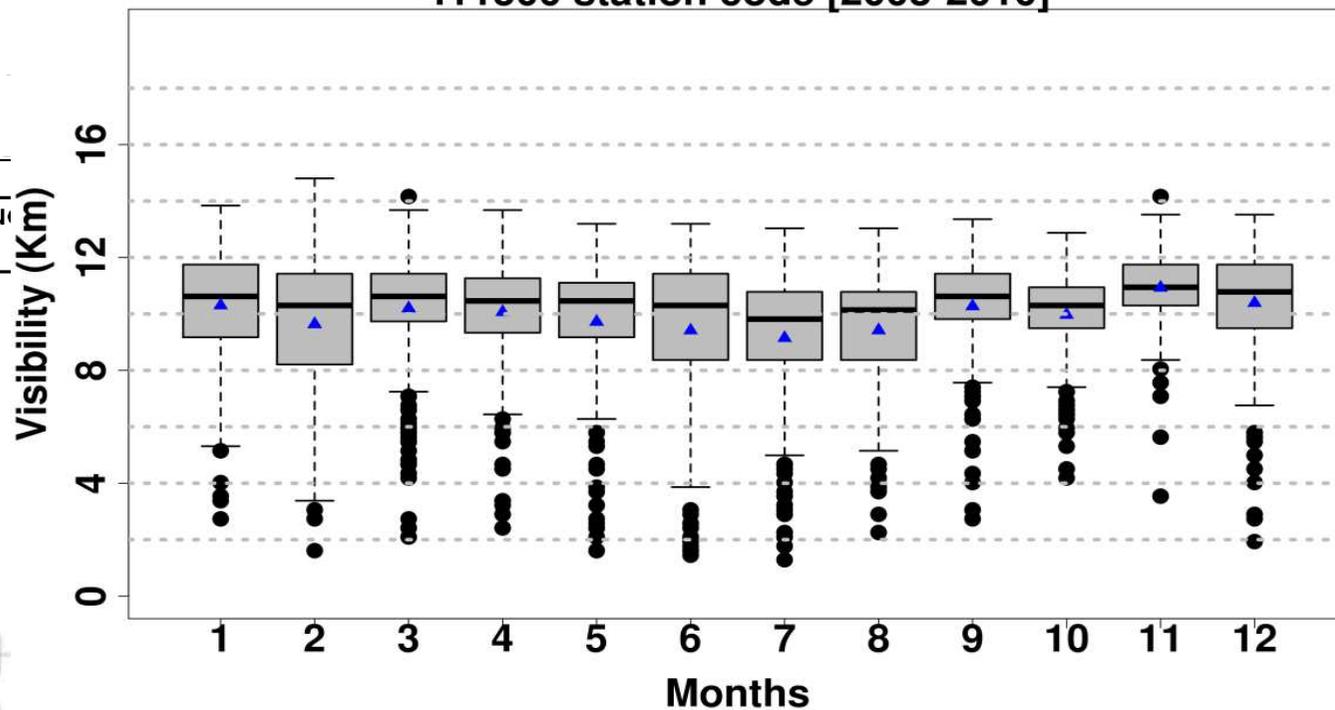
Recommended ground-based observations in Middle East

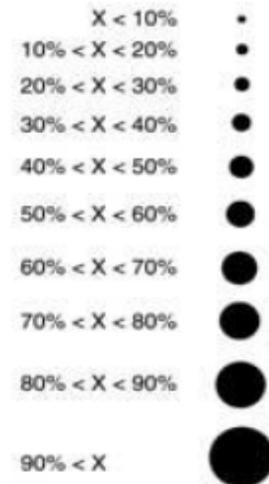
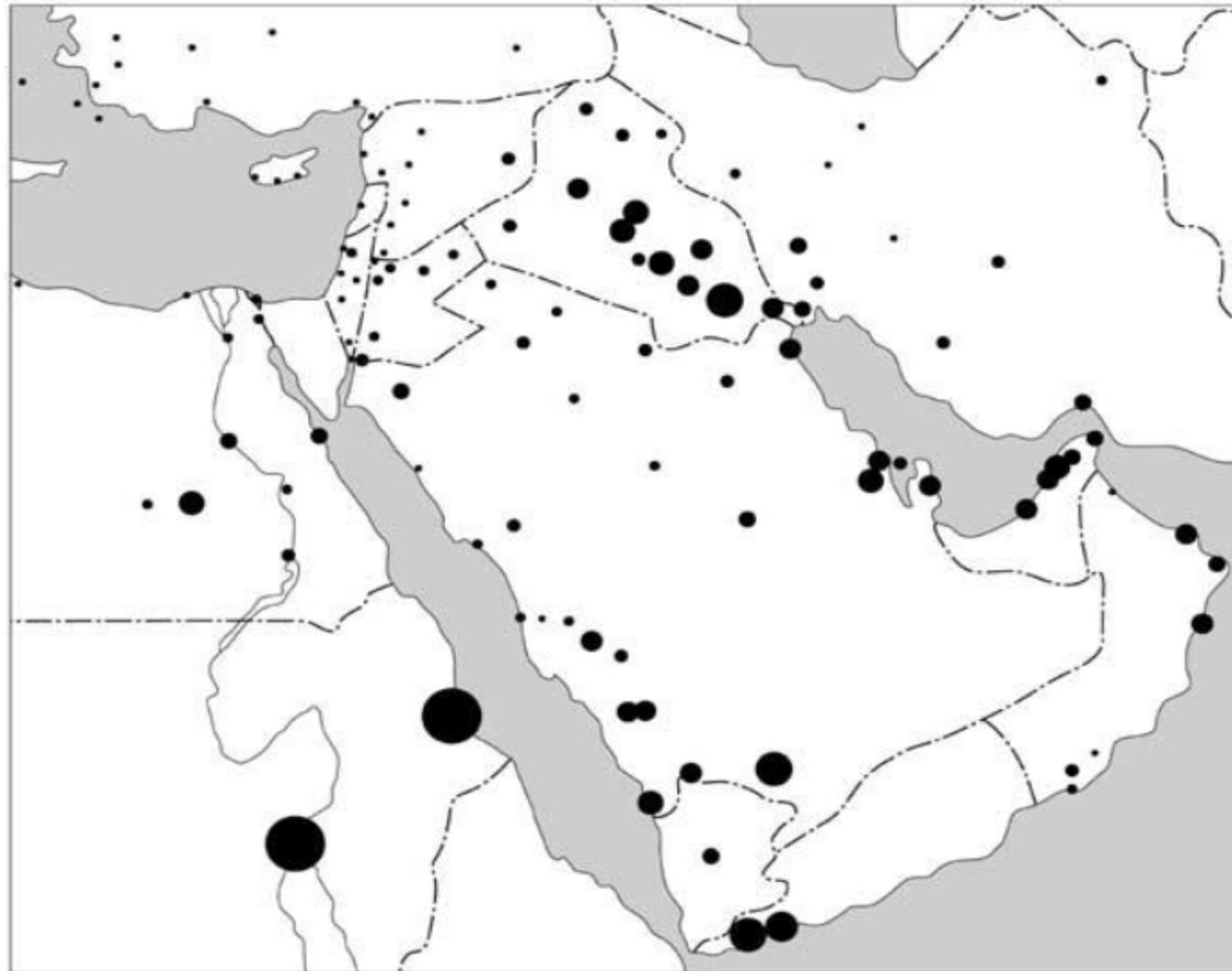


Transmissometer

SYNOP stations in West Asia. Each color indicates

Monthly climatology of horizontal visibility at Bahrain (2003-2010)





Spatial distribution of the maximum occurrence of visibility reduction in (%) of the time. Circles sizes' are proportional to percentage of visibility reduction (Kutiél and Furman (2003))

Visibility and sky conditions

-  < 50 m
-  > 50 m And < 2 km
-  > 2 km And < 4 km
-  > 4 km And < 6 km
-  > 6 km And < 8 km
-  > 8 km And < 10 km
-  > 10 km

Last Update:

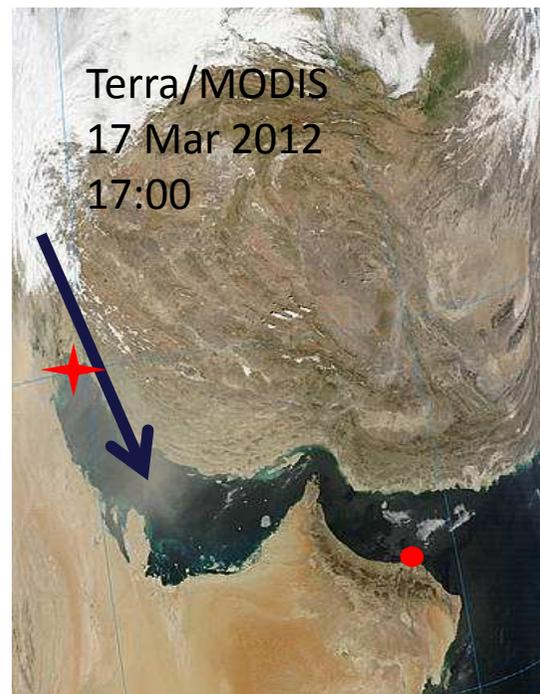
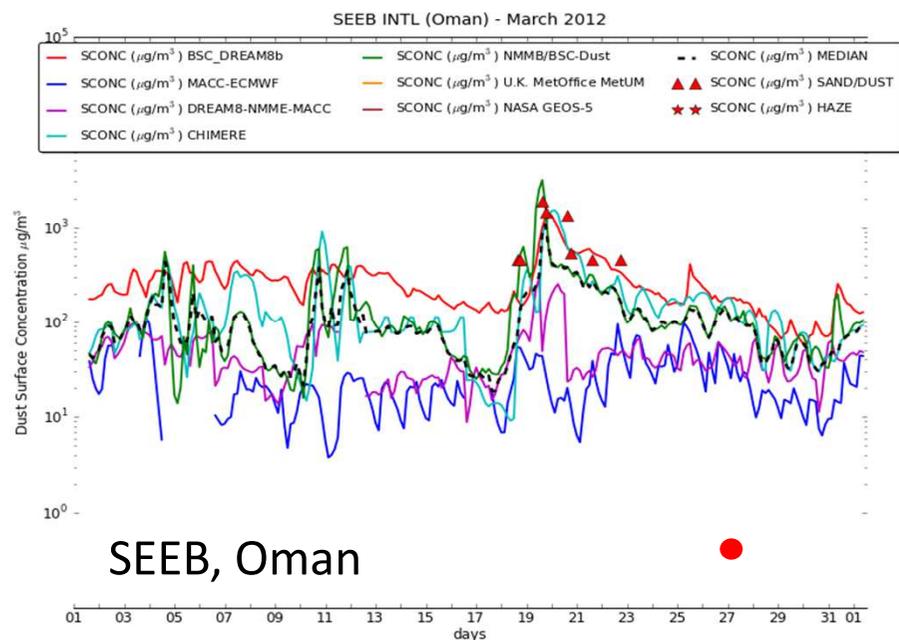
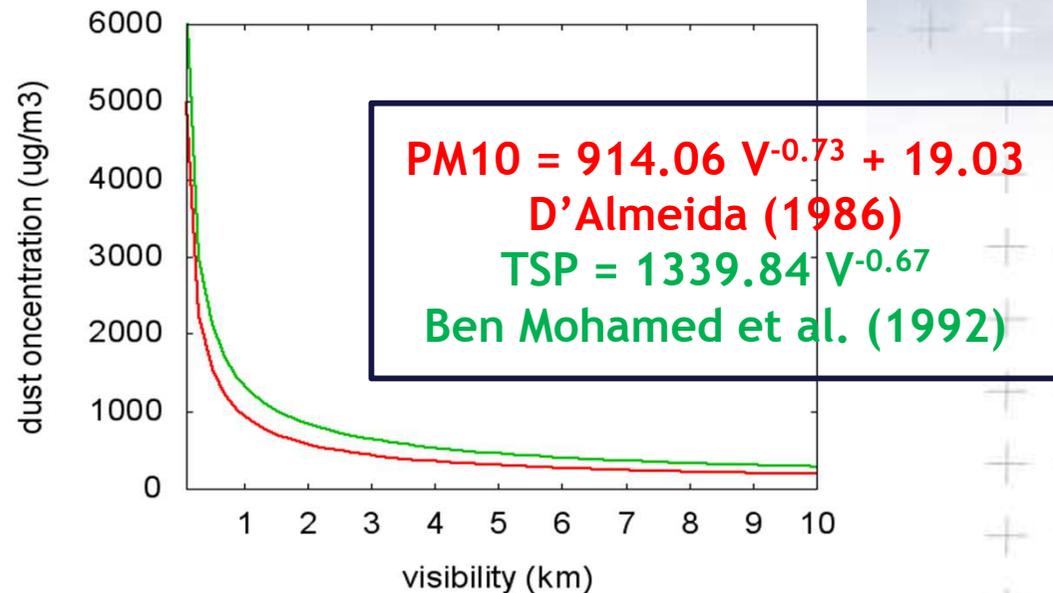
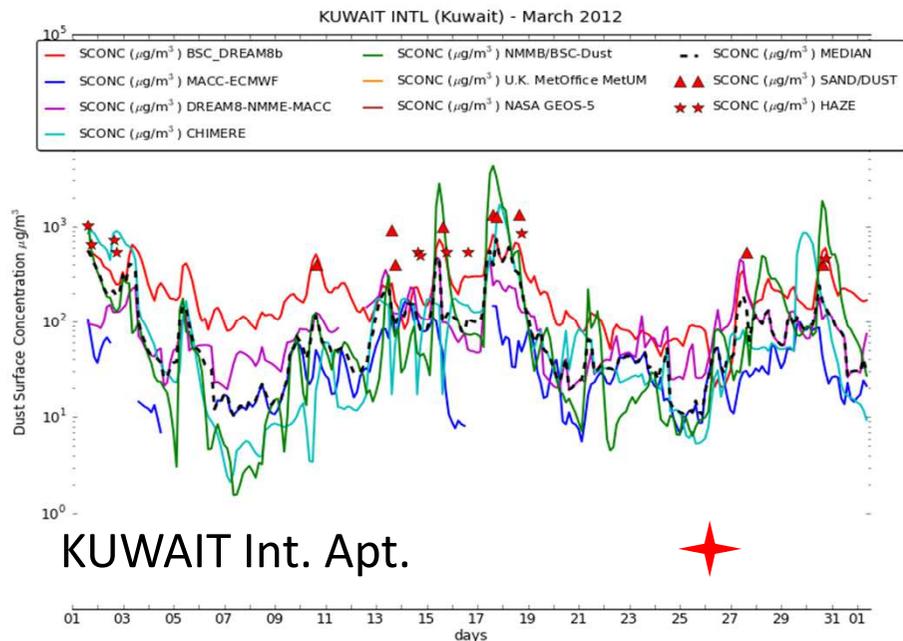
2012-09-08 04:26:55

CLICK ON A STATION FOR TIME OF OBSERVATION

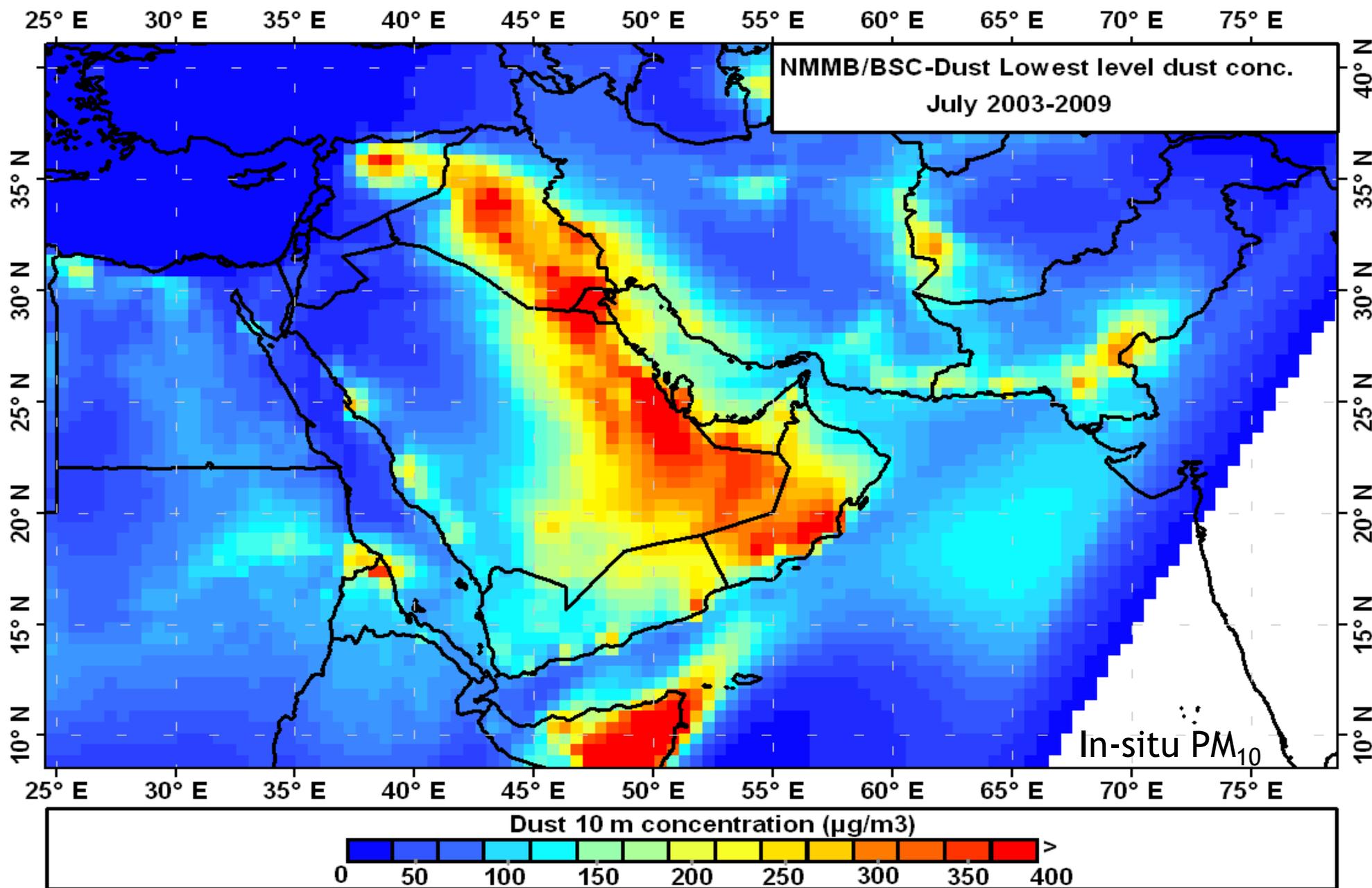


NRT Visibility semaphore at the SDS WAS NAMEE Regional center
(<http://sds-was.aemet.es/forecast-products/dust-observations/visibility>)

Recommended ground-based observations in Middle East



Visibility and sky conditions



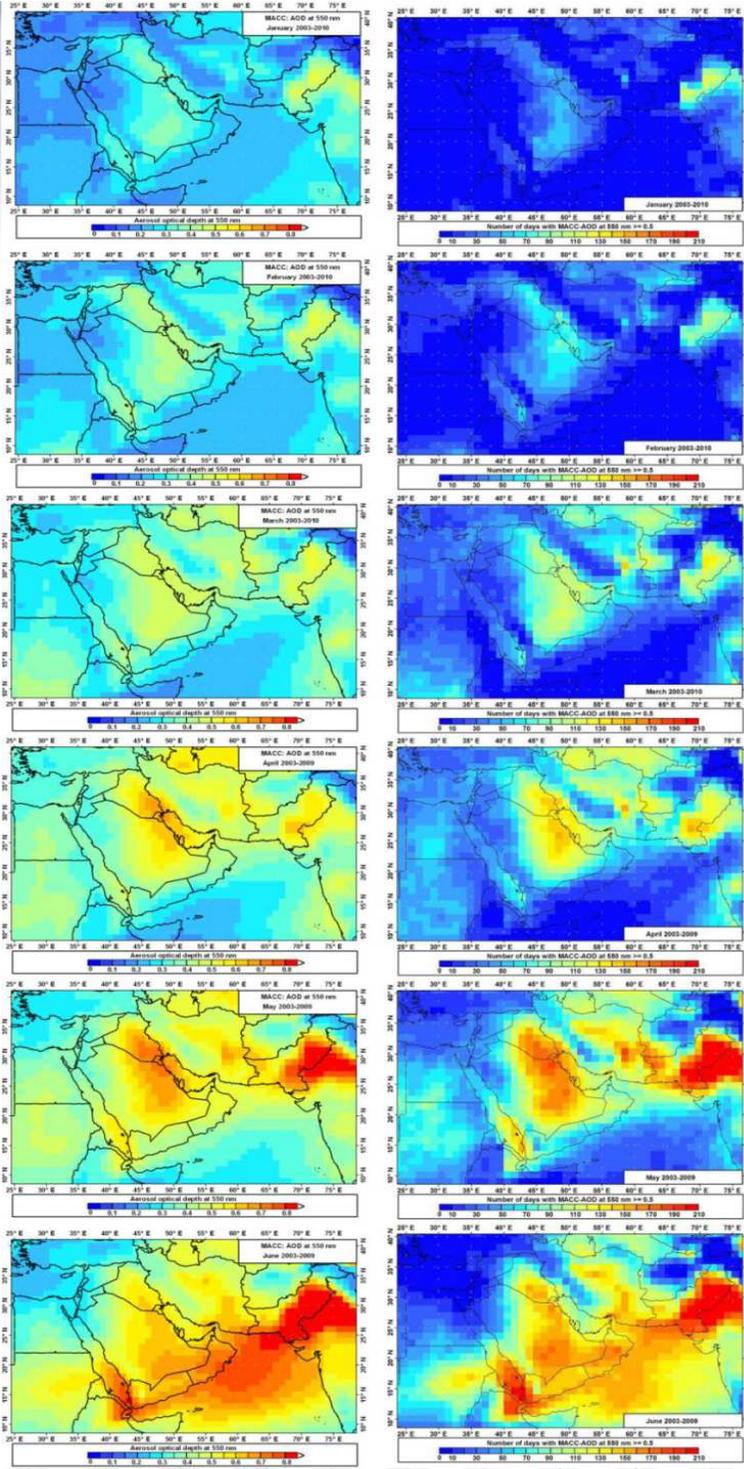
In-situ PM₁₀

Dust deposition gauges. This method measures dust deposition rate. Data are usually collected over monthly periods and results are expressed in g/m²/month (ie. the mass of dust deposited/m²/month). Recommended for spatial dust climatologies.

High volume samplers determine average dust concentrations. Data are usually collected over a 24 hour period and results are expressed in µg/m³/24hr. A selective inlet may be fitted to a high volume sampler to restrict the particle size being sampled (for example, to ensure only PM₁₀ particles are sampled).

Some few PM10 stations (around 10% of the stations) in the region must be setup in rural sites, far away from direct impact of anthropogenic sources located in populated cities and industrial centers, in order to obtain aerosol background measurements which should be affected, basically, by mineral dust from local resuspension or transported from other regions. It will help air quality managers to discriminate between industrial/vehicles pollution (regulated by air quality standards) and natural mineral dust pollution .

Quality assurance system: periodical (at least once per year) manual calibrations with gravimetric PM₁₀ high volume samplers should be performed by co-located continuous PM₁₀ analyzers (TEOM and Beta analyzers).

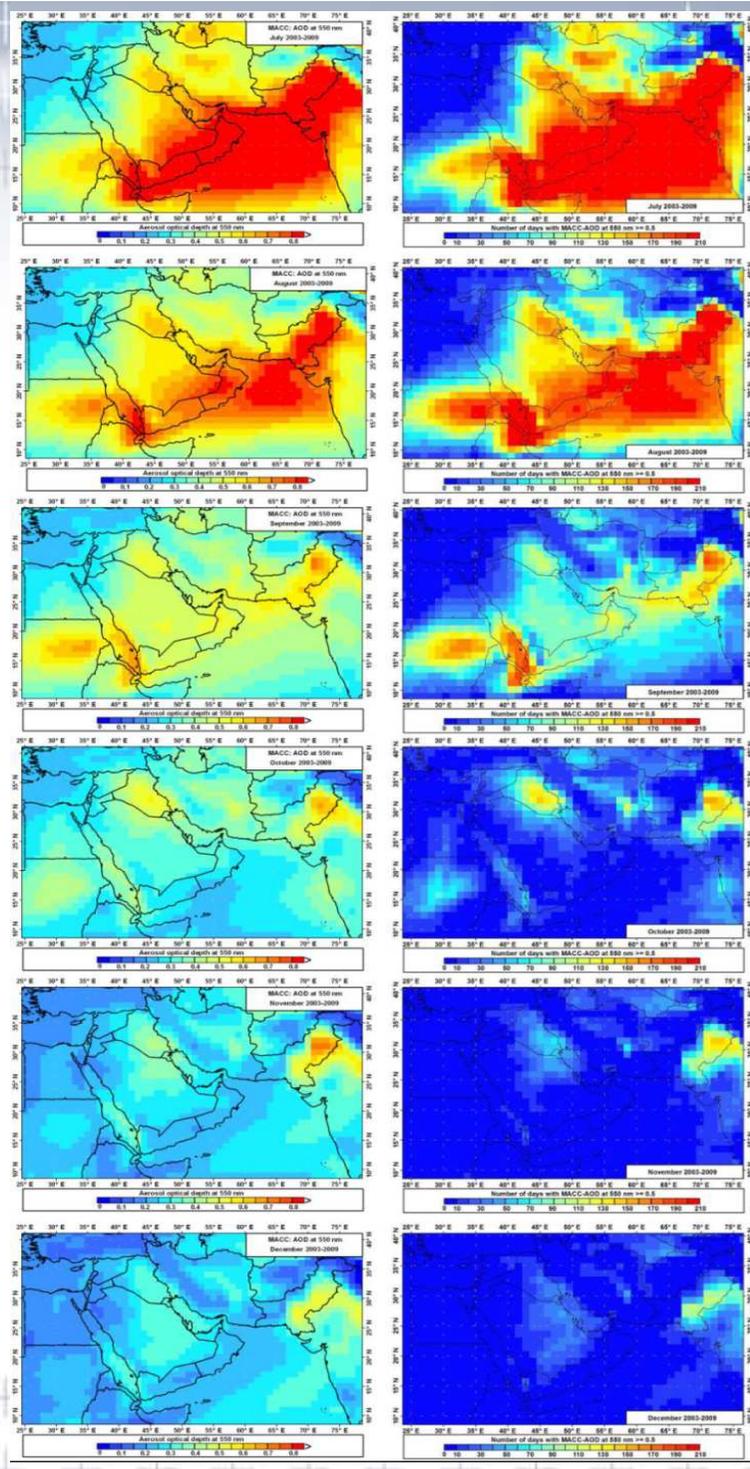


Total AOD from
MACC-ECMWF and
number of days with
AOD>0.5
2003-2010

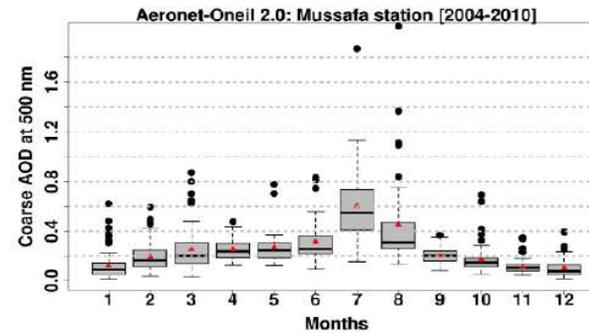
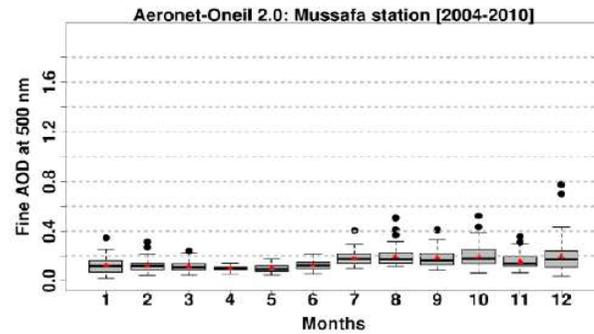
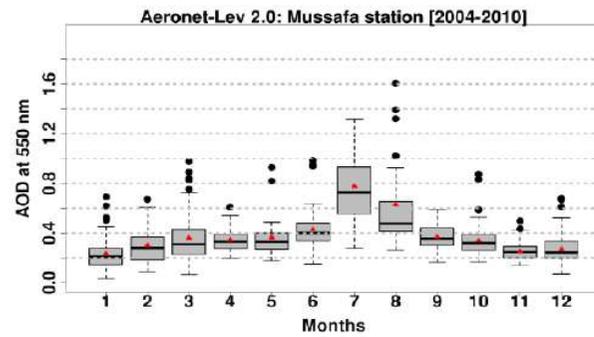
January-June



July-December

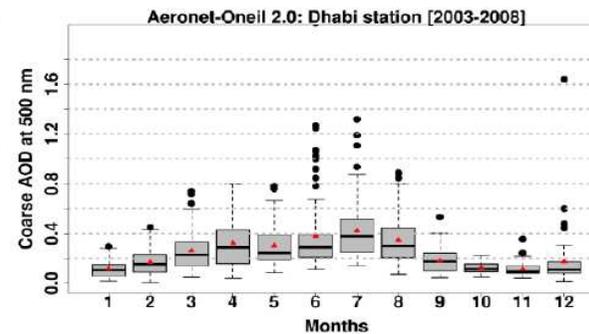
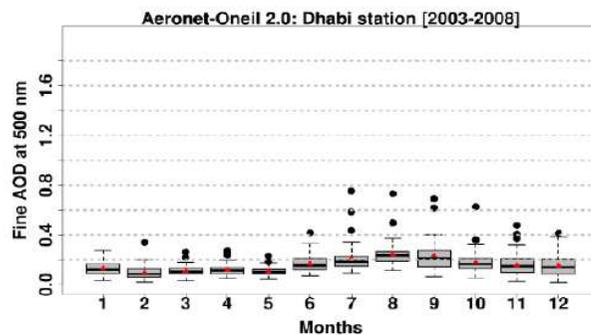
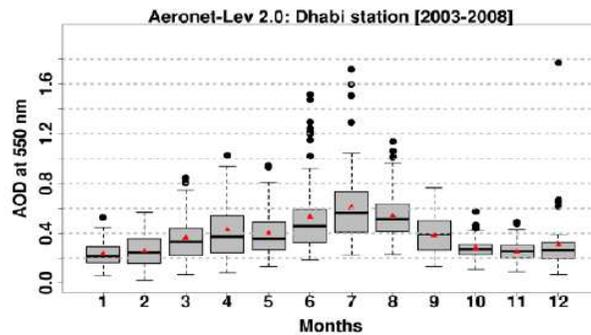


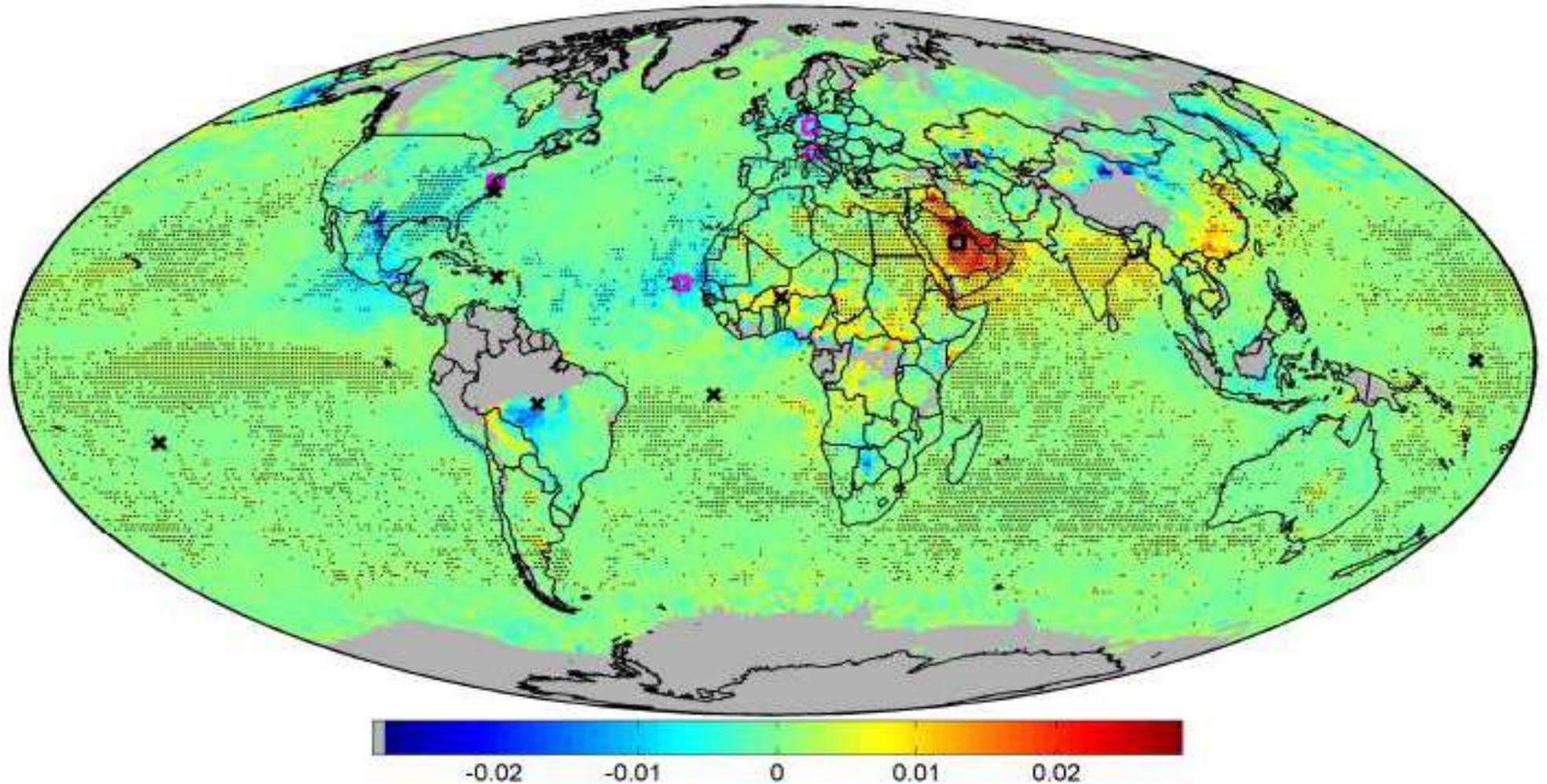
Mussafa (Emirates)
 24.3716° N
 54.4666° E
 10 m a.s.l.



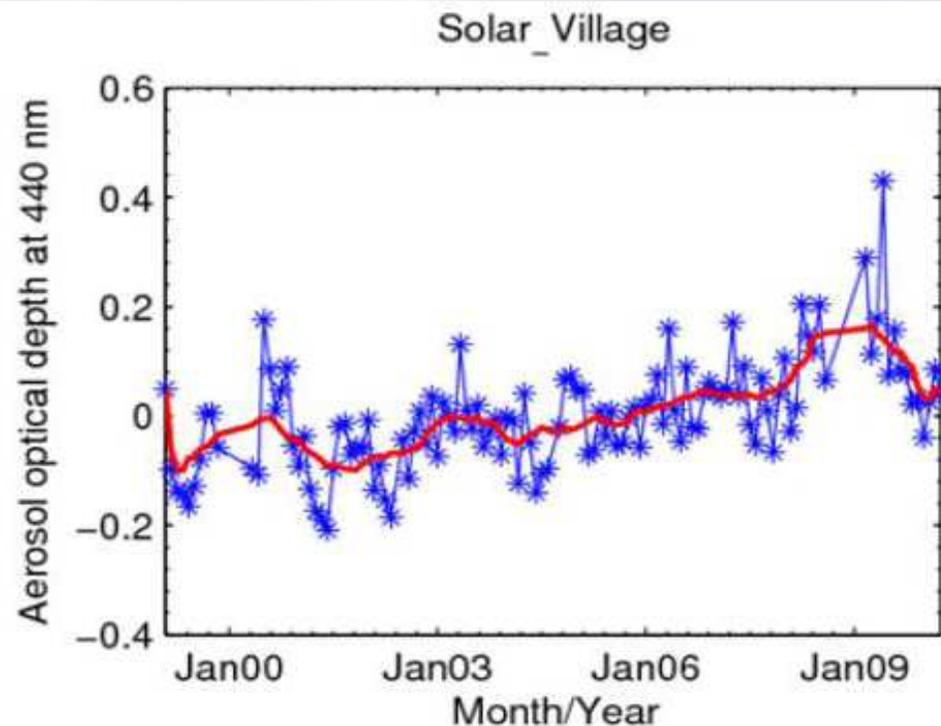
Seasonal variation of aerosols at Mussafa and Abu Dhabi, AERONET stations (Emirates)

Abu-Dhabi (Emirates)
 24.47611° N
 54.32889° E
 7m a.s.l.

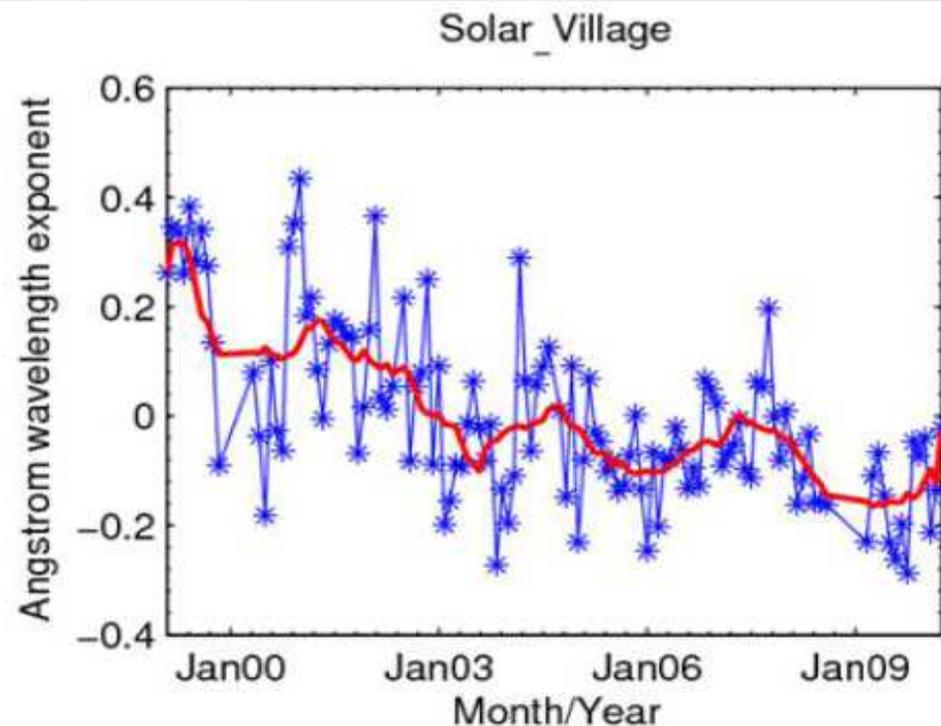




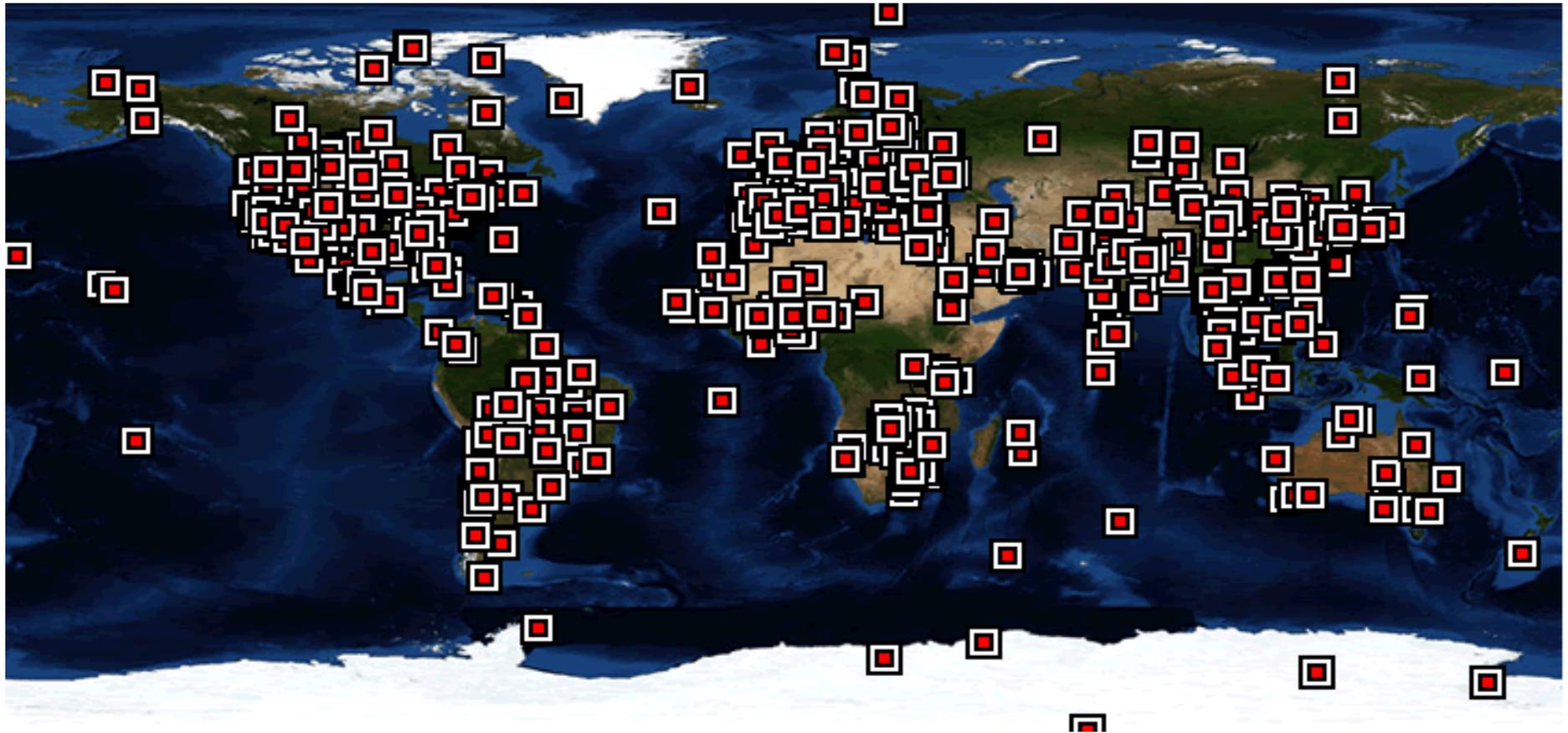
Linear trend based upon deseasonalized monthly anomaly of AOD at 550 nm with SeaWifs for the period 1998-2010. Units are AOD/yr. Dots indicate significance at 95% confidence level (Hsu et al., 2012)



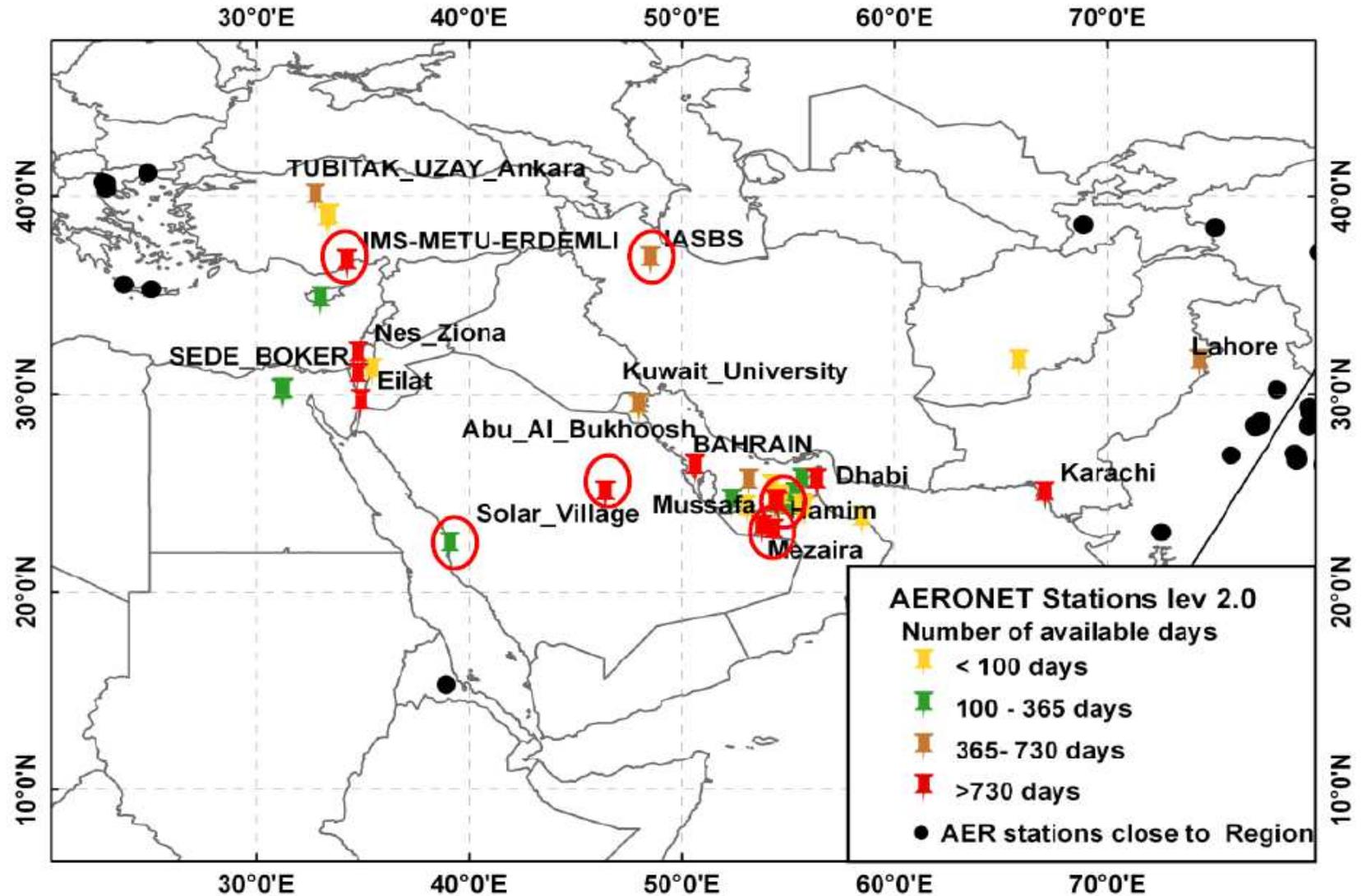
Monthly anomalies of AOD at Solar Village AERONET station (Saudi Arabia). (Xia, 2011)



Monthly anomalies of Angstrom Exponent at Solar Village AERONET station (Saudi Arabia). (Xia, 2011)

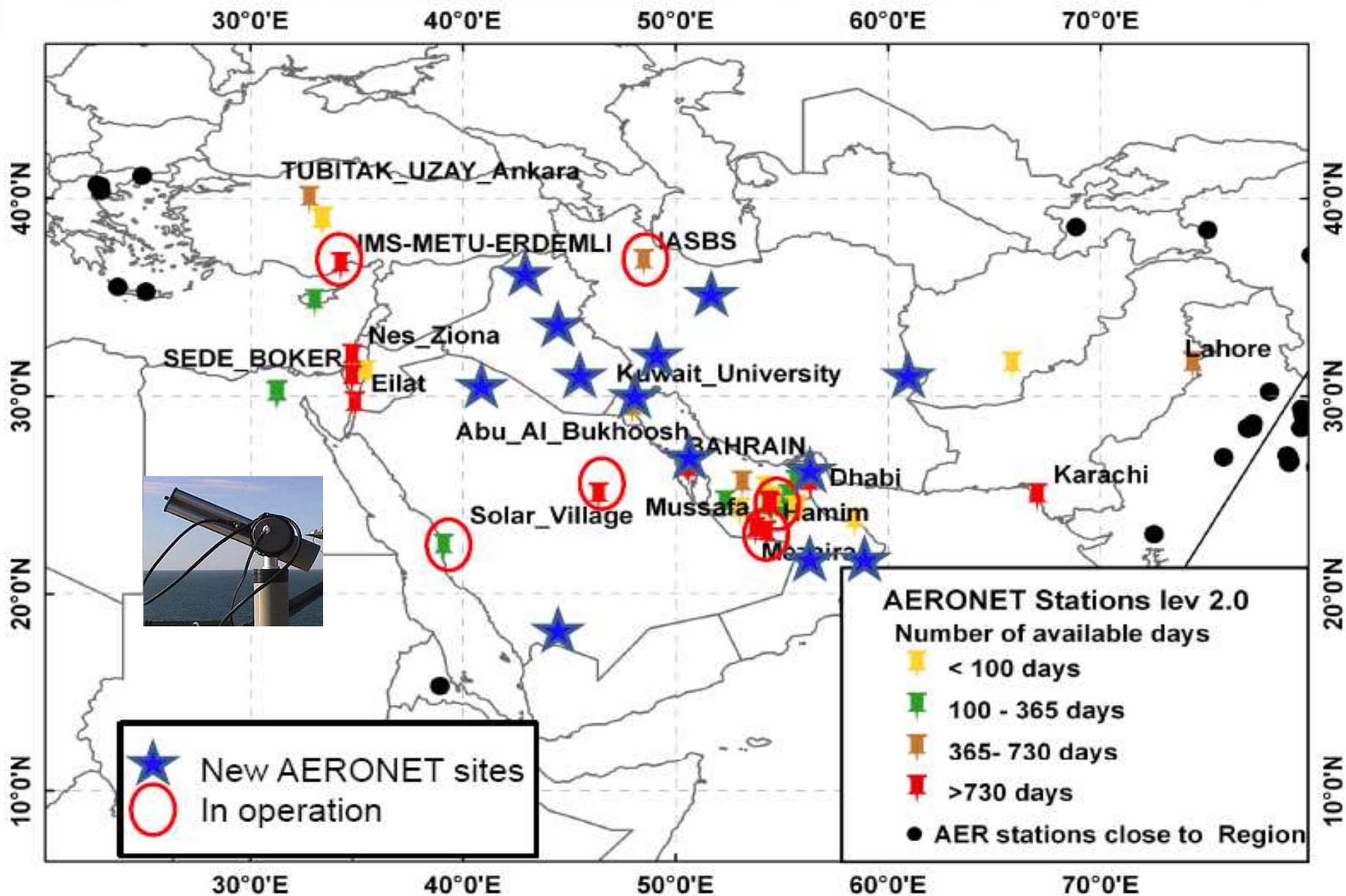


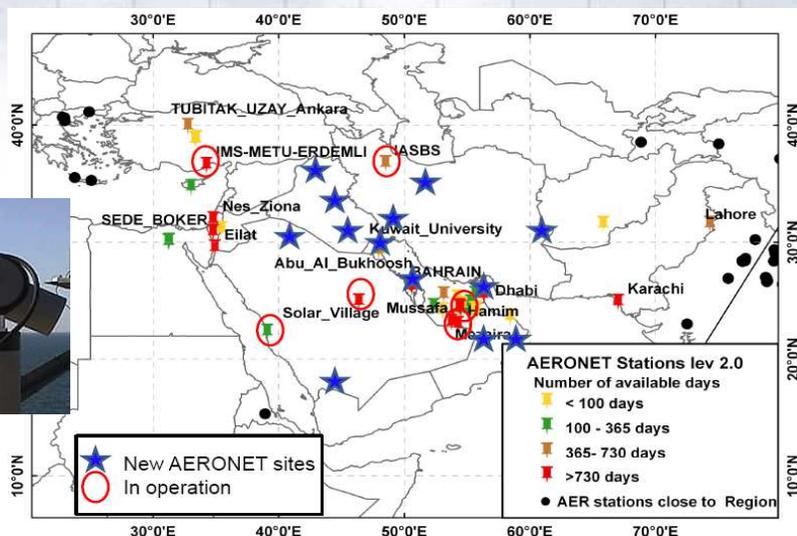
Map of AERONET stations (<http://gsfc.nasa.aeronet.gov>)



Map of AERONET stations in West Asia and Middle East. Stations encircled in red correspond to present operational active stations

Sunphotometer network (AERONET) in Middle East





Sunphotometer network (AERONET)

1) Re-start operations in the following former AERONET sites:

- Kuwait University, Khalidiyah Campus (Kuwait)
- Bahrain (re-start operations in the most convenient horizon-free site)
- Dhadnah (United Arab Emirates)

2) New possible AERONET sites chosen by geographical location of dust storms pathways and considering topology of the regional network:

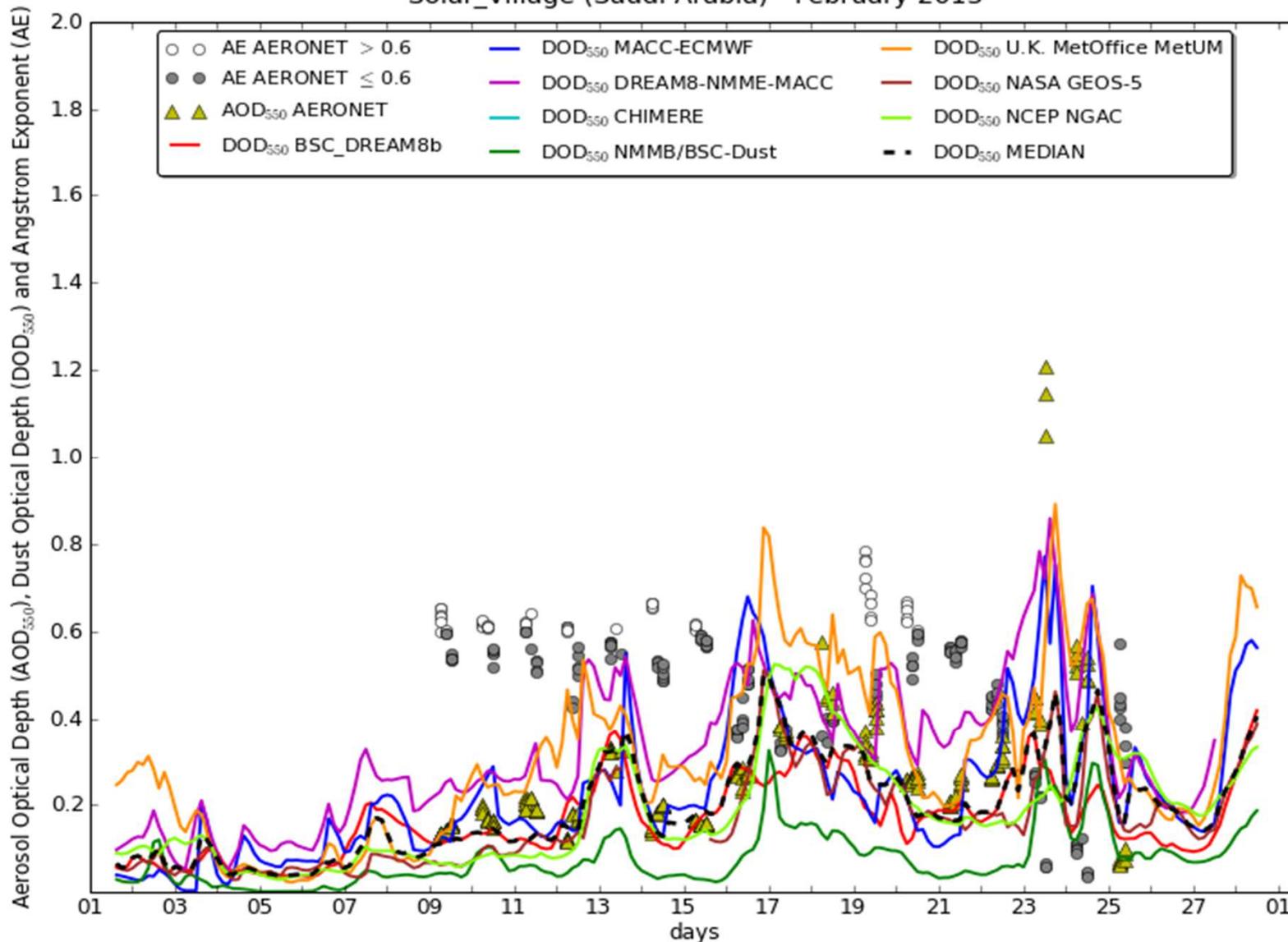
- Arar (North Saudi Arabia); Najran (Southwest Saudi Arabia)
- Somewhere in Empty Quarter (Saudi Arabia): not indicated in the map
- Mosul (North Iraq) ; Baghdad (Central Iraq); As Smawah (South Iraq)
- Faud, Dhahirah (Oman); Bani Bu Hassan (Oman)
- Ahvaz, Khuzestan (southwest Iran); Zabol (preferable) or Zahedan (Sistan basin, East/southeast Iran); Tehran (Iran)

- High accuracy dust climatologies and trends
- Model NRT and off-line validation
- Satellite NRT and off-line verification

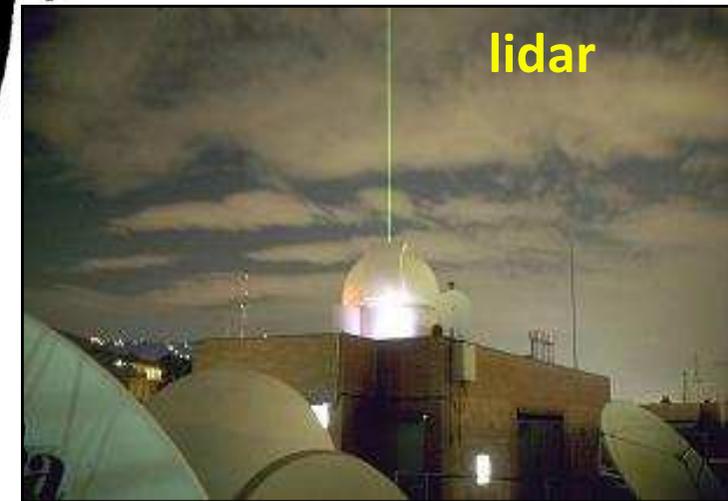
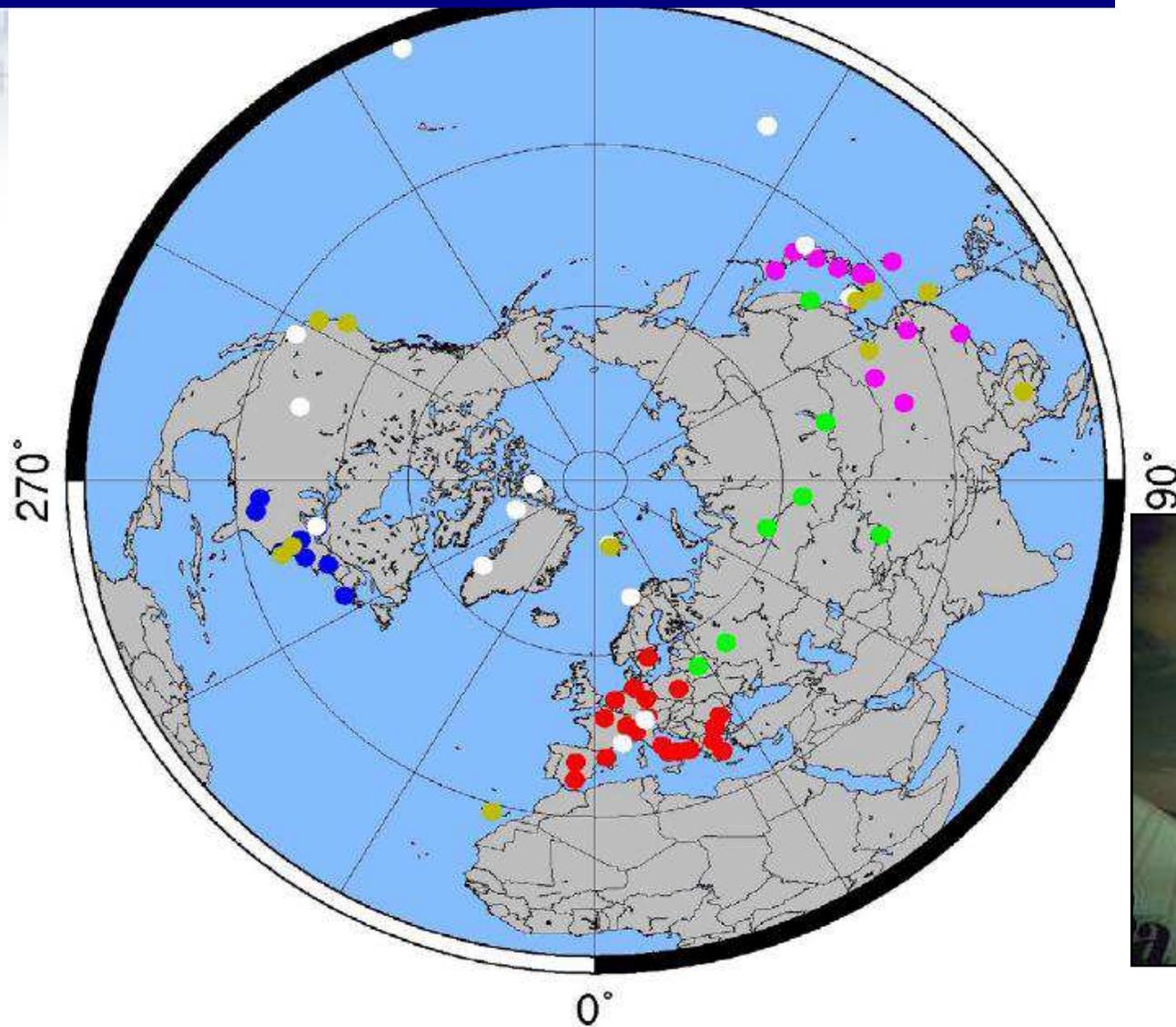
Regional network

http://sds-was.aemet.es/forecast-products/forecast-evaluation/Solar_Village

Solar_Village (Saudi Arabia) - February 2013



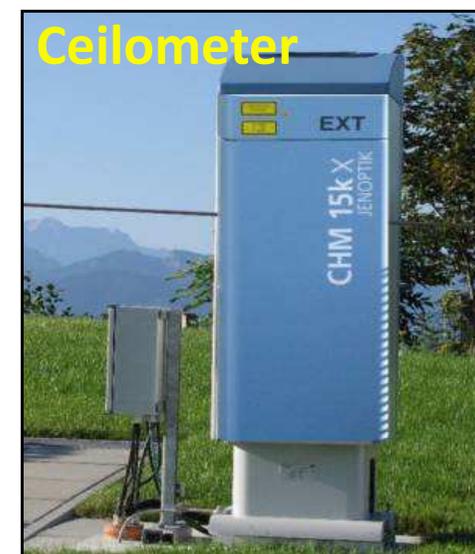
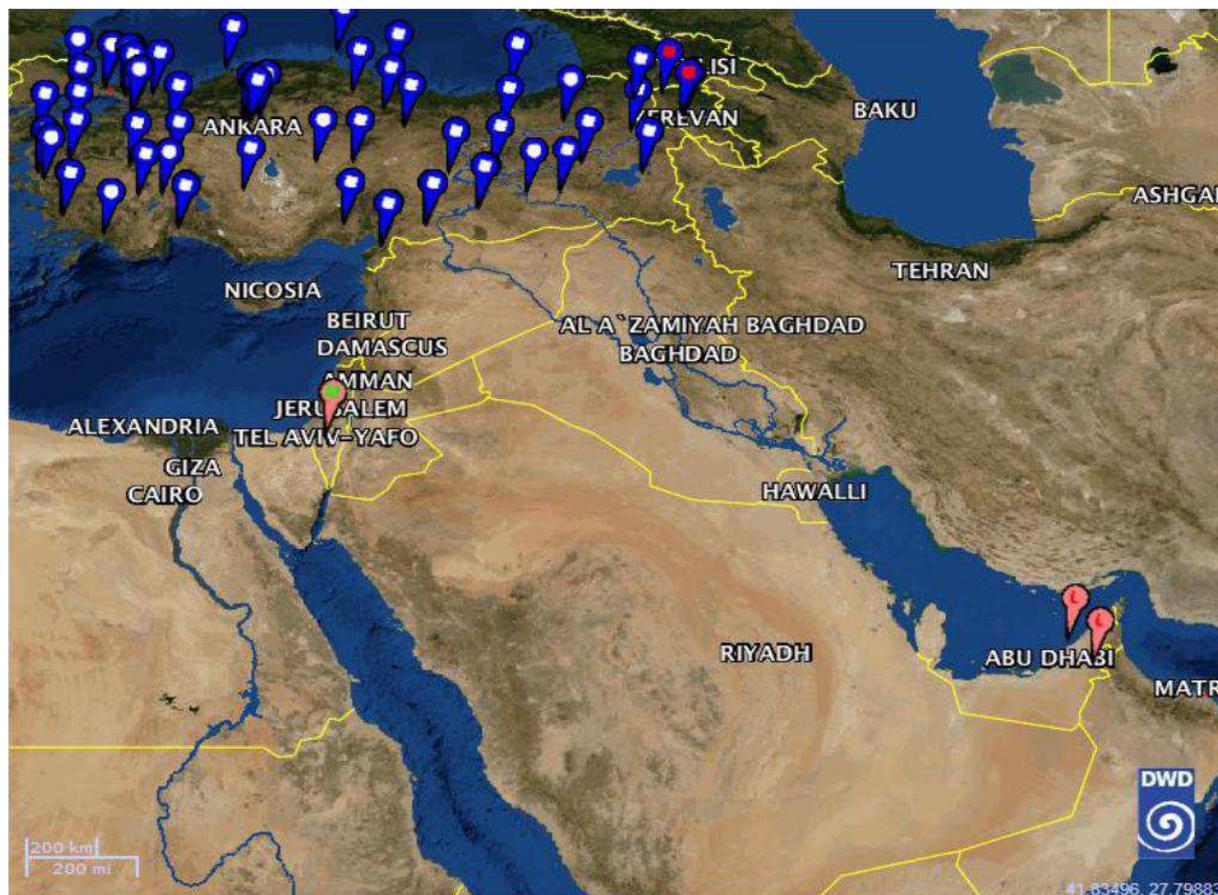
NRT validation



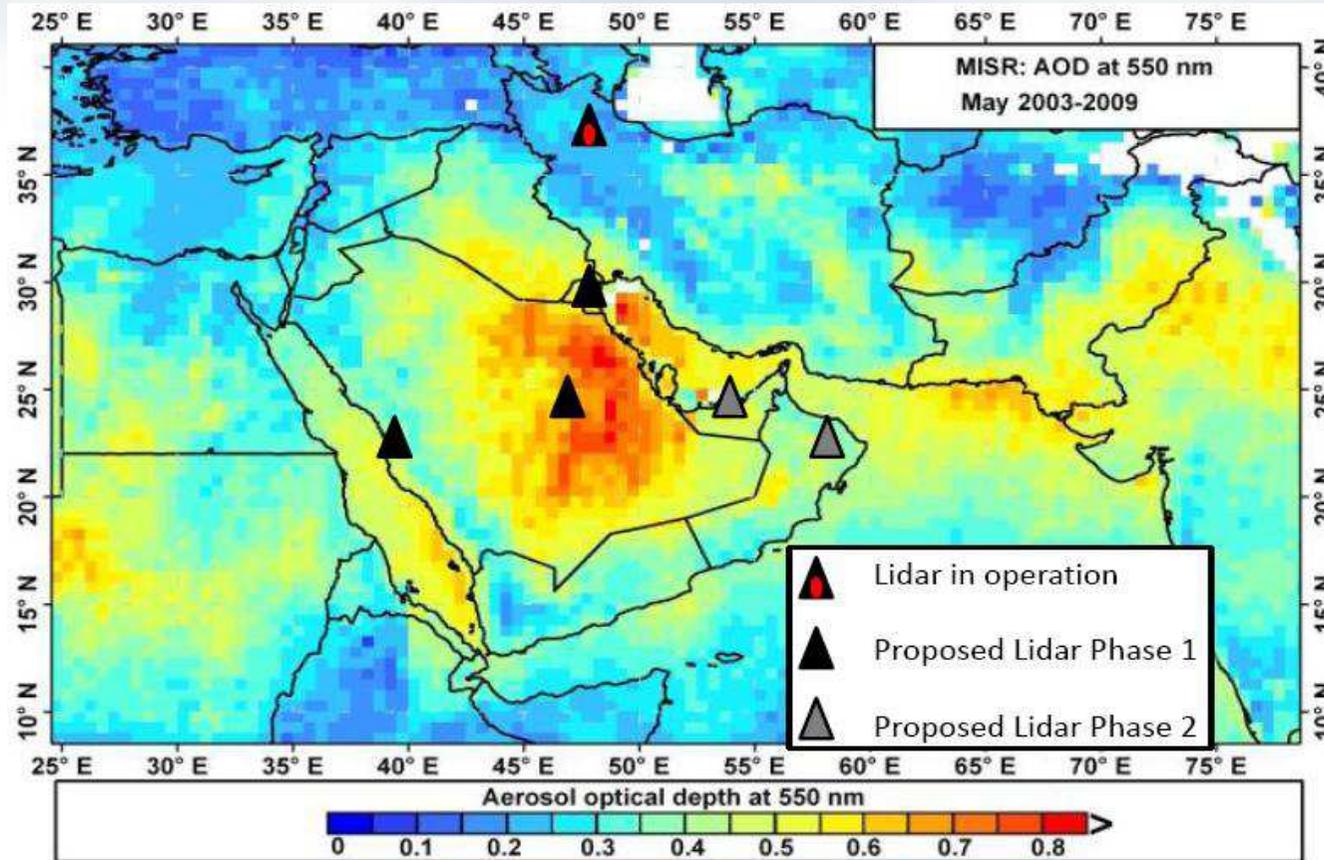
GALION network. The different networks are indicated by the colour of dots: AD-NET violet, ALINE yellow, CISLiNet green, EARLINET red, MPLNET brown, NDACC white, REALM blue.

West Asia Ceilometer network

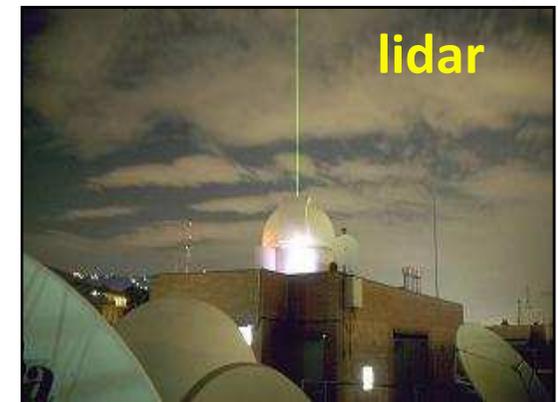
- Turkey a pilot experience to be integrated in GALION
- Replace old ceilometers by new generation ceilometers (NMSs)
- Contact Dr. Thomas Werner (German Weather Service) for new algorithms and data processing



West Asia Lidar network



To make operational two lidars at Zanzan (Iran) operated by ISBS



- Kuwait University where an AERONET station was in operation till August 2012?
- KAUST_Campus at the King Abdullah University of Science and Technology (KAUST; Saudi Arabia)
- Solar_Village at the Energy Research Institute of the King Abdulaziz City for Science and Technology (KACST; Saudi Arabia)
- Two additional Lidar stations could be set up in Emirates and Oman, respectively.
- Contact Gelsomina Pappalardo (AIRLINET and GALION Chair)

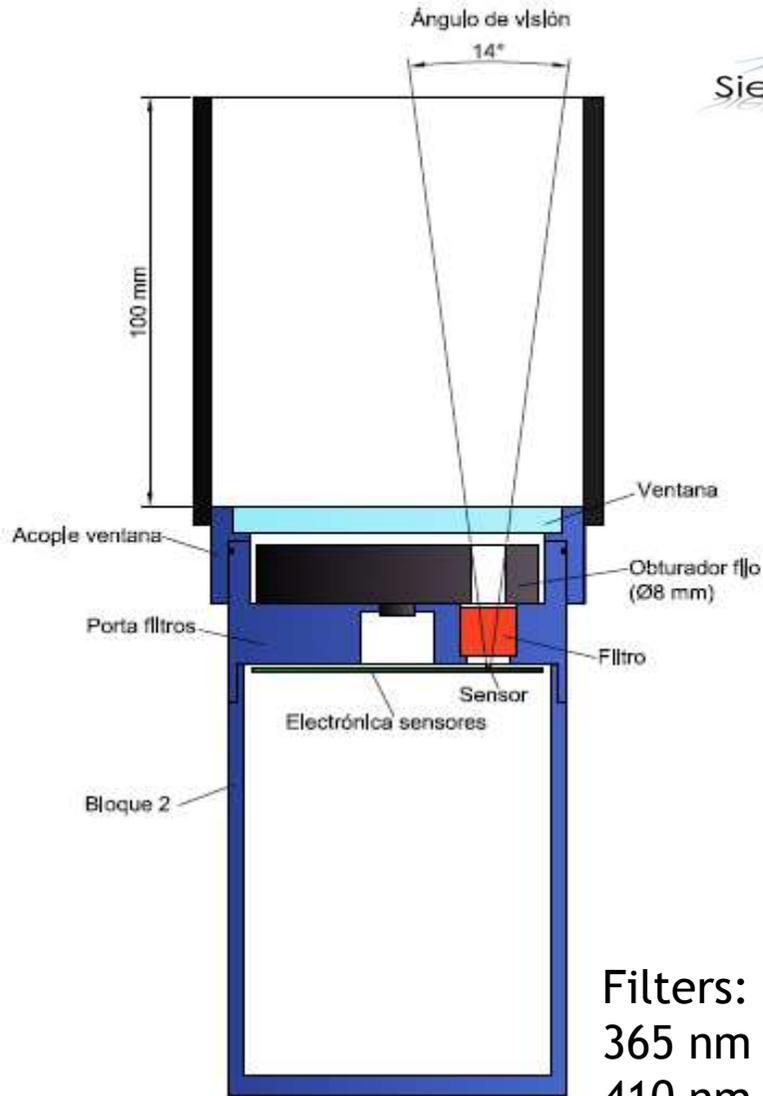
New developments for dust detection

- ❑ Total Sky cameras and webcams
- ❑ New inexpensive “Color Index” radiometers (NRT AOD and cloud attenuation estimation)



Dust storm seen by total-sky camera during UAE² in 2004

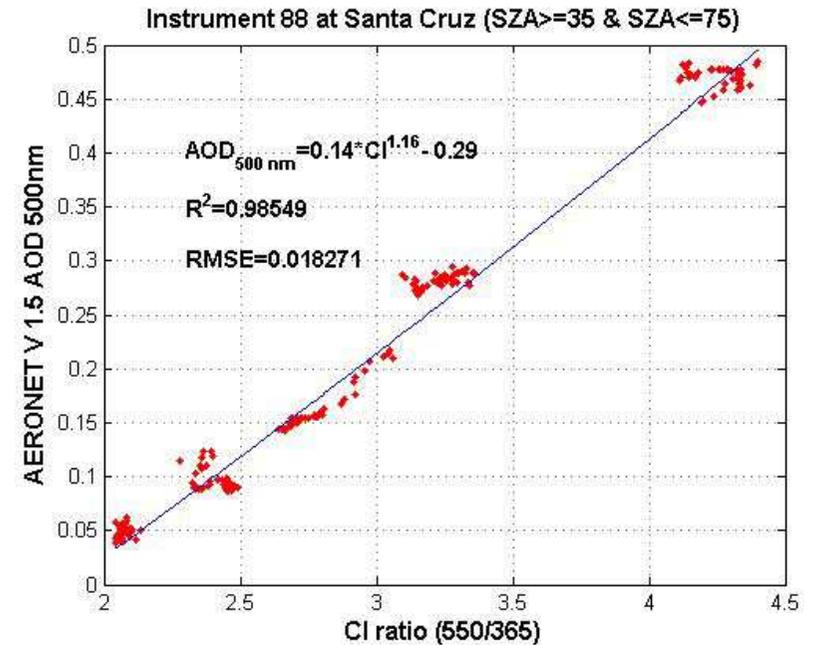
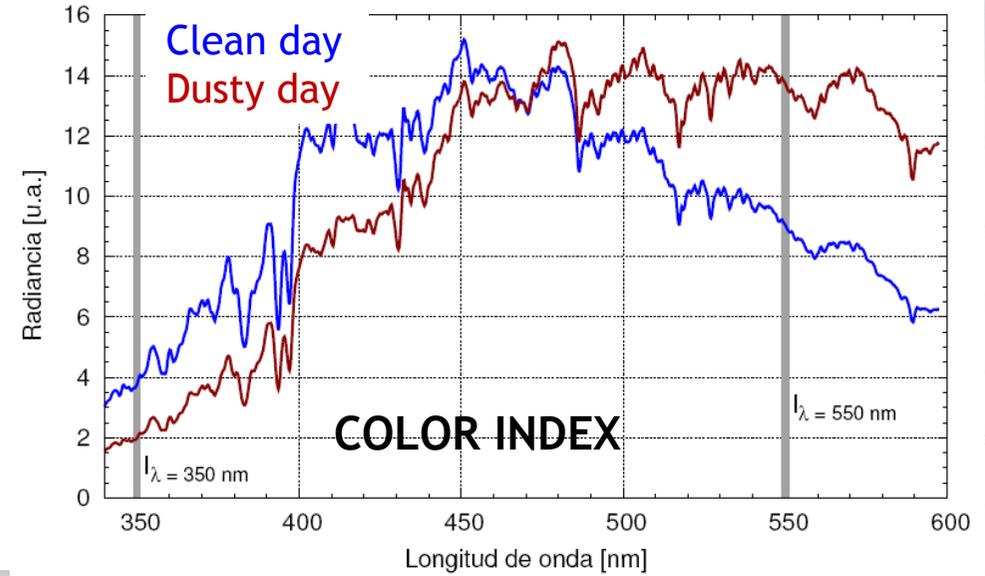
“Color-Index” Radiometer



www.sieltec.es
Sieltec Canarias S.L.



Filters:
365 nm
410 nm
470 nm
550 nm
870 nm



Index

- Aerosols and dust background
- In-situ dust characterization
- In-situ dust estimations (Visibility)
- Ground based remote sensing
- Recommended ground-based observations in Middle East
- **Summary**

In-situ measurements and surface remote sensing compared to satellite

Advantages

Disadvantages

In-situ
measurements

Ground-based
remote sensing

Satellite remote
sensing

In-situ measurements and surface remote sensing compared to satellite

Advantages

In-situ measurements

- very straightforward;
- unique dust physical and chemical information;
- universal applicability (no sky conditions dependent)
- Time high resolution (minutes)

Ground-based remote sensing

- high information on dust (transmitted light dominates over reflected);
- non-intrusive measurements;
- easy access to equipment;
- column dust information

Satellite remote sensing

- global coverage; (global dust)
- non-intrusive measurements

Disadvantages

- intrusive measurements;
- local coverage in some sites

- local coverage;
- indirect measurements;
- very limited capability in presence of clouds (Photom.)

- limited on information aerosol (No chemical composition, size distribution, low temporal resolution);
- no access to equipment

Ground-based dust observations are essential for:

- High-accuracy optical and chemical characterization of dust
- Dust model verification, validation and assimilation
- Dust satellite-based products validation

Ground based ‘supersites’ equipped with complete high-quality dust observation programs (in-situ and remote sensing) constitute unique platforms for satellite-based dust observations and dust models quality assurance

In-situ observations & long term monitoring:

mass concentration:

bulk aerosol mass (TSP, PM_{10} , $PM_{2.5}$)

bulk dust mass (total dust, $dust_{10}$, $dust_{2.5}$): 1 or more tracers methods

chemical composition

bulk aerosol = dust, pollutants (SO_4^- , NO_3^- , NH_4^+), sea-salt + trace metals, OC, EC

number size distribution

10nm - 500nm + 0.5 - 20 μm (no distinction between dust and other aerosols)

optical properties

scattering and absorption coefficients

complementary measurements: mixing state, mineralogy, isotopic characterization, etc...

Close cooperation and data exchange between Met. Services, Air quality networks/Agencies, and Universities is highly recommended !!

In-situ remote **remote sensing**:

Total column dust optical properties (sunphotometers):

Aerosol optical depth (AOD)

Angström Exponent

Dust vertical distribution

Corrected vertical backscattering

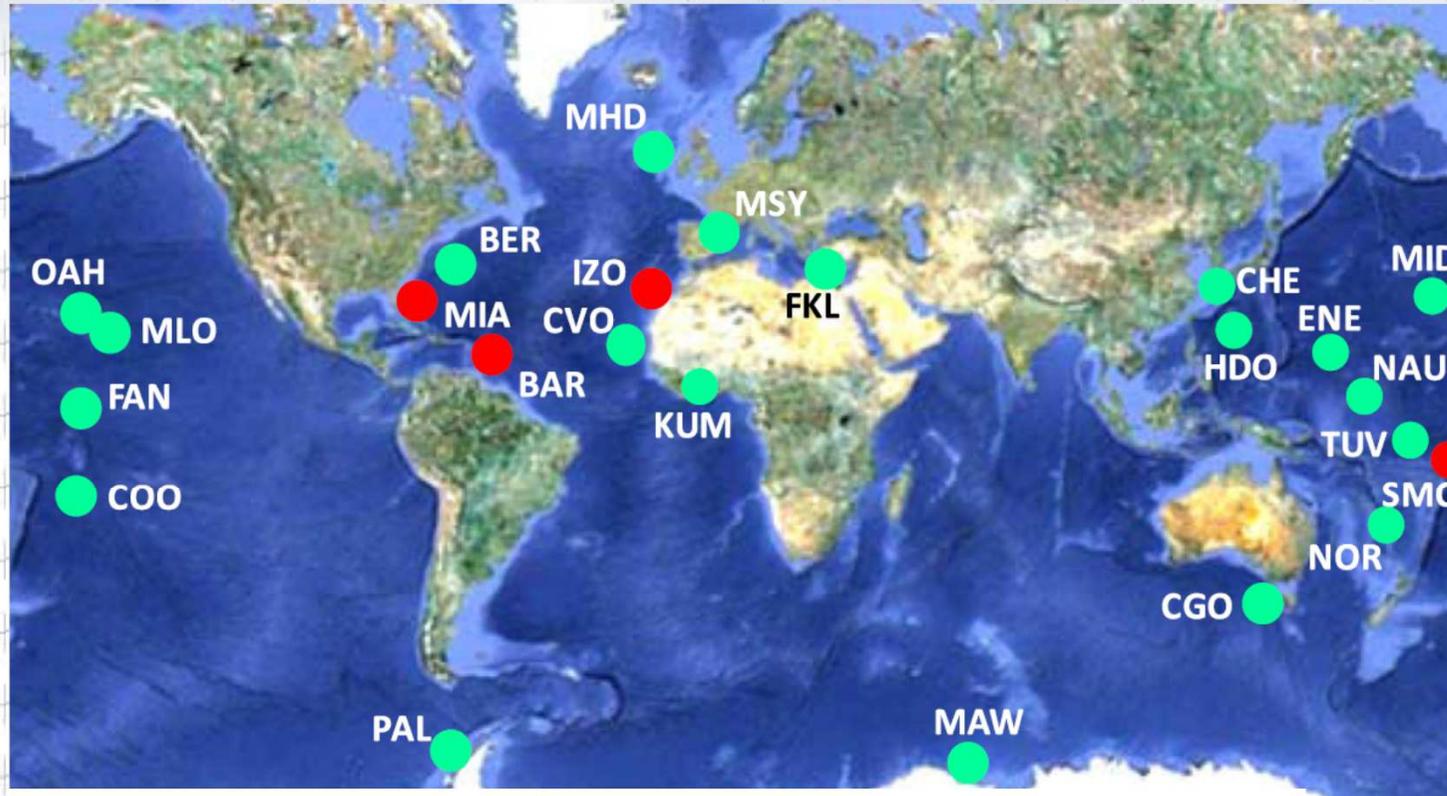
Extinction vertical distribution

Test new inexpensive developments

Color index from simple multi-wavelength radiometers

Total sky-cameras

Long term monitoring dust background-observatories:



- at least 4 years
- Active during the last 20 years

Review Article

Aeolian Research Aeolian Research 6 (2012) 55–74

A review of methods for long term in situ characterization of aerosol dust

Sergio Rodríguez^{a,*}, Andrés Alastuey^b, Xavier Querol^b

There is a big gap of ground-based observations in Middle East: in-situ PM, sunphotometers and lidars/ceilometers

Synergies between air quality and dust monitoring should be found. This is of critical importance for big cities in Northern Africa and Middle East

3rd Training Course on WMO SDS-WAS products (satellite and ground observation and modelling of atmospheric dust)
Muscat-Oman, December 8-12, 2013



Ground observations of mineral dust

Emilio Cuevas [ecuevasa@aemet.es] (remote sensing)

Sergio Rodríguez [srodriguezg@aemet.es] (in-situ)

Izaña Atmospheric Research Centre, AEMET

La Marina 20

Santa Cruz de Tenerife, 38071

Canary Islands, Spain

<http://izana.aemet.es>